

Supplement of Atmos. Chem. Phys., 19, 6497–6507, 2019  
<https://doi.org/10.5194/acp-19-6497-2019-supplement>  
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*Supplement of*

## **Effect of sea salt aerosol on tropospheric bromine chemistry**

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**Text S1.** Updated parameterization of first-order loss of gas-phase HOBr in cloud

In the model, HOBr(g) is uptaken by cloud water (clouds droplets and ice crystal quasi-liquid surfaces), and then reacts with Br<sup>-</sup>/Cl<sup>-</sup>/HSO<sub>3</sub><sup>-</sup>/SO<sub>3</sub><sup>2-</sup> in the aqueous phase.



Here we number the four reactions in the same order as in the manuscript. Chen et al. (2017) computed the overall reactive uptake coefficient of HOBr in cloud ( $\gamma$ ) using following parameterization:

$$\gamma = \left( \frac{1}{\gamma_d} + \frac{1}{a_b} + \frac{1}{\Gamma} \right)^{-1} \quad (1)$$

$$\Gamma = \Gamma_1 + \Gamma_4 + \Gamma_{5a} + \Gamma_{5b} \quad (2)$$

where  $\gamma_d$  is gas-phase diffusion coefficient (unitless), calculated following Schmidt et al. (2016) and Sherwen et al. (2016);  $a_b$  (0.6, Ammann et al., 2013) is mass accommodation coefficient (unitless); and  $\Gamma$  is bulk reaction coefficient (unitless).  $\Gamma$  is the sum of reaction coefficients of the four reactions ( $\Gamma_1$ ,  $\Gamma_4$ ,  $\Gamma_{5a}$ , and  $\Gamma_{5b}$ ), which are calculated as:

$$\Gamma_1 = 4HRT \sqrt{D_l k_1 [\text{H}^+][\text{Br}^-]} \frac{f_r}{c} \quad (3)$$

$$\Gamma_4 = 4HRT \sqrt{D_l k_4 [\text{H}^+][\text{Cl}^-]} \frac{f_r}{c} \quad (4)$$

$$\Gamma_{5a} = 4HRT \sqrt{D_l k_{5a} [\text{HSO}_3^-]} \frac{f_r}{c} \quad (5)$$

$$\Gamma_{5b} = 4HRT \sqrt{D_l k_{5b} [\text{SO}_3^{2-}]} \frac{f_r}{c} \quad (6)$$

where  $H$  ( $1300 \text{ M atm}^{-1} \times \exp[6014 \text{ K}^{-1} \times (1/T - 1/298.15)]$ , Sander, 2015) is the Henry's law constant of HOBr;  $T$  is air temperature (K);  $R$  ( $8.31 \times 10^{-2} \text{ L bar mol}^{-1} \text{ K}^{-1}$ ) is the universal gas constant;  $D_l$  ( $1.4 \times 10^{-5} \text{ cm}^2 \text{ s}^{-1}$ , Ammann et al., 2013) is the liquid phase diffusion coefficient for HOBr;  $f_r$  and  $c$  is the reacto-diffusive correction term (Ammann et al., 2013) and the mean thermal velocity ( $\text{cm s}^{-1}$ ), respectively, calculated following Schmidt et al. (2016) and Sherwen et al. (2016);  $k_1$  ( $1.6 \times 10^{10} \text{ M}^{-2} \text{ s}^{-1}$ , Beckwith et al., 1996),  $k_4$  ( $2.3 \times 10^{10} \text{ M}^{-2} \text{ s}^{-1}$ , Liu and Margerum, 2001),  $k_{5a}$  ( $3.2 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$ , Liu, 2000), and  $k_{5b}$  ( $5.0 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$ , Troy and Margerum, 1991) is the reaction rate coefficient for (R1), (R4), (R5a), and (R5b), respectively.

However, such parameterization subjects to the assumption that physical uptake of HOBr, including diffusion in the gas phase, accommodation on the liquid surface, and diffusion in the aqueous phase, occurs independently and

separately for each pathway. We argue that such assumption does not hold. Instead, aqueous  $\text{HOBr}(\text{aq}) + \text{Br}^-/\text{Cl}^-/\text{HSO}_3^-/\text{SO}_3^{2-}$  reactions should all follow the same physical uptake of HOBr. We express the first-order aqueous-phase loss of  $\text{HOBr}(\text{aq})$  as a sum of the four pathways. We update the parameterization of  $\text{HOBr}(\text{aq}) + \text{Br}^-/\text{Cl}^-/\text{HSO}_3^-/\text{SO}_3^{2-}$  reaction system as described in Chen et al. (2017) by computing bulk reaction coefficient  $\Gamma$  using the sum ( $R^I$ ) of the first order loss rate coefficient of  $\text{HOBr}(\text{aq})$  via (R1), (R4), (R5a), and (R5b).

$$\Gamma = 4HRT\sqrt{D_l R^I} \frac{f_r}{c} \quad (7)$$

$$R^I = k_1[\text{H}^+][\text{Br}^-] + k_4[\text{H}^+][\text{Cl}^-] + k_{5a}[\text{HSO}_3^-] + k_{5b}[\text{SO}_3^{2-}] \quad (8)$$

Then the overall reactive uptake coefficient of HOBr in cloud ( $\gamma$ ) is computed using equation (1). The updated parameterization tends to have greater impact on the bulk reaction coefficient ( $\Gamma$ ) and thus on  $\gamma$  at a lower cloud pH. Following Chen et al. (2017), we consider the condition with a temperature of 283 K, cloud liquid water content of  $0.2 \text{ g m}^{-3}$ , cloud droplet radius of  $10 \mu\text{m}$ , gas-phase  $\text{SO}_2$ , HCl and HBr concentrations to be 100 ppt, 1 ppt and 1 ppt, respectively. For cloud pH of 3, we estimate that  $\Gamma$  is 0.11 following Chen et al. (2017), but 0.06 (decreased by 41%) with the updated parameterization. Accordingly,  $\gamma$  decreases from 0.018 to 0.016 (by 10%). For cloud pH of 6,  $\Gamma$  decreases from 6.4 to 5.1 (by 20%) with the updated parameterization, but  $\gamma$  remains almost the same (0.021).

The overall first-order loss rate of  $\text{HOBr}(\text{g})$   $R_{\text{HOBr}}$  ( $\text{molecules cm}^{-3} \text{ s}^{-1}$ ) is computed as:

$$R_{\text{HOBr}} = \frac{1}{4} c \gamma S [\text{HOBr}] \quad (9)$$

Where  $S$  is the total surface area concentration of cloud droplets and ice crystals ( $\text{cm}^2 \text{ cm}^{-3}$ ). We distribute the loss by the above four pathways on the basis of their relative aqueous rates.

For (R1) and (R4), we also consider the dependence of  $\text{Br}_2$  production yield ( $Y$ ) on  $[\text{Br}^-]$  and  $[\text{Cl}^-]$  following Fickert et al. (1999), Chen et al. (2017), and Wang et al. (2019).

$$Y = 0.41 \log_{10} \left( \frac{[\text{Br}^-]}{[\text{Cl}^-]} \right) + 2.25 \quad \text{for } [\text{Br}^-] / [\text{Cl}^-] < 5 \times 10^{-4} \quad (10)$$

$$Y = 0.90 \quad \text{for } [\text{Br}^-] / [\text{Cl}^-] \geq 5 \times 10^{-4} \quad (11)$$

Finally, we compute the loss rate of  $\text{HOBr}(\text{g})$  via (R1), or the production rate of  $\text{Br}_2(\text{g})$ ,  $R_{\text{Br}_2}$  as:

$$R_{\text{Br}_2} = \frac{k_1[\text{H}^+][\text{Br}^-] + k_4[\text{H}^+][\text{Cl}^-]}{R^I} Y R_{\text{HOBr}} \quad (12)$$

The loss rate of  $\text{HOBr}(\text{g})$  via (R4), or the production rate of  $\text{BrCl}(\text{g})$ ,  $R_{\text{BrCl}}$  is:

$$R_{\text{BrCl}} = \frac{k_1[\text{H}^+][\text{Br}^-] + k_4[\text{H}^+][\text{Cl}^-]}{R^I} (1 - Y) R_{\text{HOBr}} \quad (13)$$

The loss rate of  $\text{HOBr}(\text{g})$  via (R5a)  $R_{\text{HSO}_3}$  is:

$$R_{\text{HSO}_3} = \frac{k_{5a}[\text{HSO}_3^-]}{R^I} R_{\text{HOBr}} \quad (14)$$

And the loss rate of HOBr(g) via (R5b)  $R_{\text{SO}_3}$  is:

$$R_{\text{SO}_3} = \frac{k_{5b}[\text{SO}_3^{2-}]}{R^I} R_{\text{HOBr}} \quad (15)$$

**Text S2.** Correction a registration error for SSA alkalinity in the standard GEOS-Chem code

In the standard GEOS-Chem code, SSA alkalinity was incorrectly registered as cloud pH in Headers/state\_chem\_mod.F. By doing so, the code generally assigned higher values (than the debromination threshold) for SSA alkalinity, thus largely prevented SSA debromination from happening. We corrected this error in our simulations.

Around line 1049 in state\_chem\_mod.F, we changed the following line

```
CALL Register_ChmField( am_I_Root, chmID, State_Chm%pHCloud, State_Chm, RC )
```

to

```
! Fix the typo
```

```
CALL Register_ChmField( am_I_Root, chmID, State_Chm%SSAlk, State_Chm, RC )
```

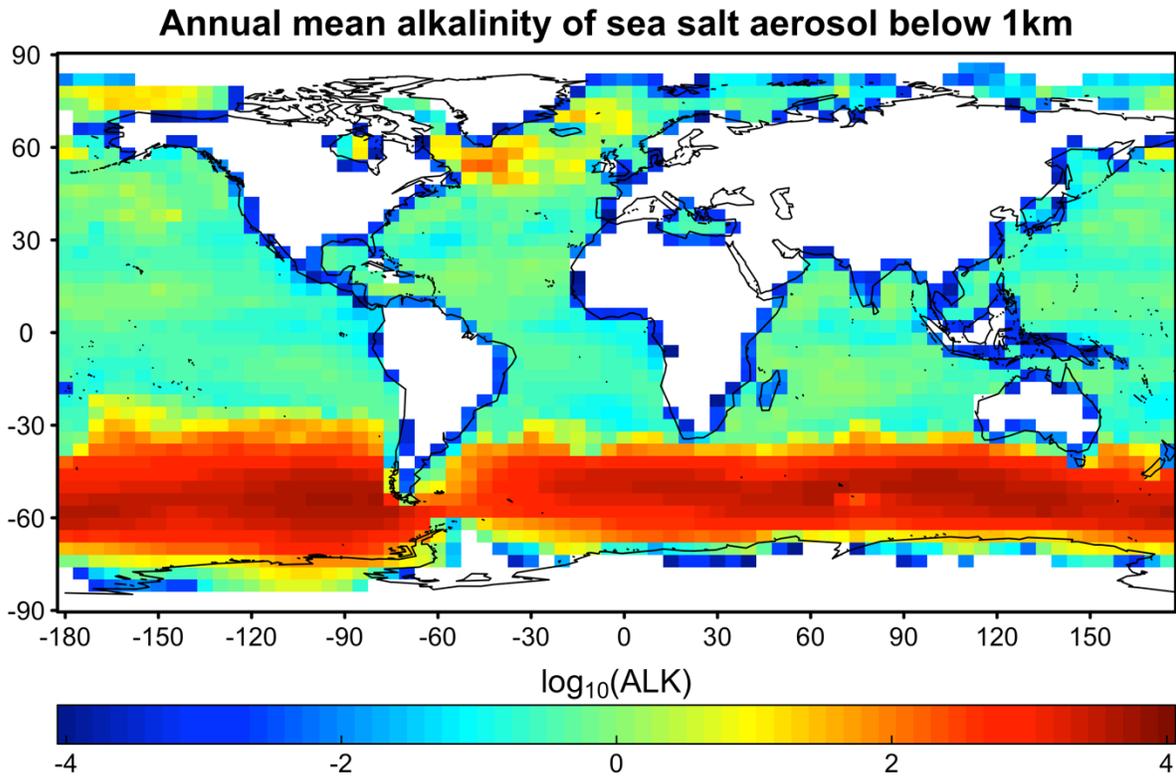


Figure S1. Annual mean alkalinity of sea salt aerosol (SSA) below 1km simulated by GEOS-Chem. Alkalinity (ALK) is in the unit of ppt. Results are shown as  $\log_{10}(\text{ALK})$ . White color means ALK is zero.

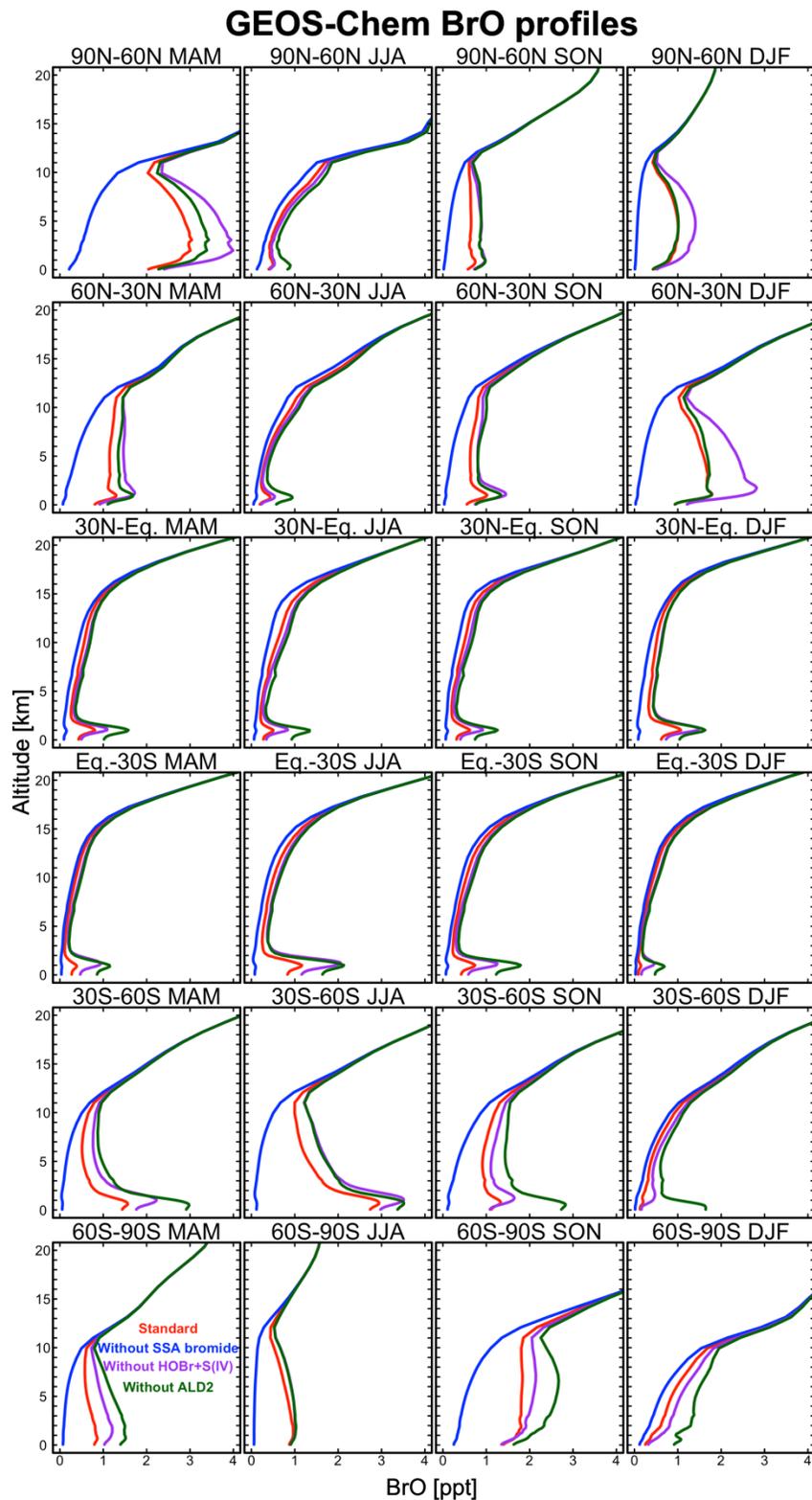


Figure S2. GEOS-Chem BrO vertical profiles averaged over different seasons and latitudinal bands. Model results are shown from our standard simulation including sea salt aerosol debromination (red lines), and sensitivity simulations not including SSA debromination (blue lines), HOBr + S(IV) reactions (purple lines), and oceanic acetaldehyde emissions (green lines).

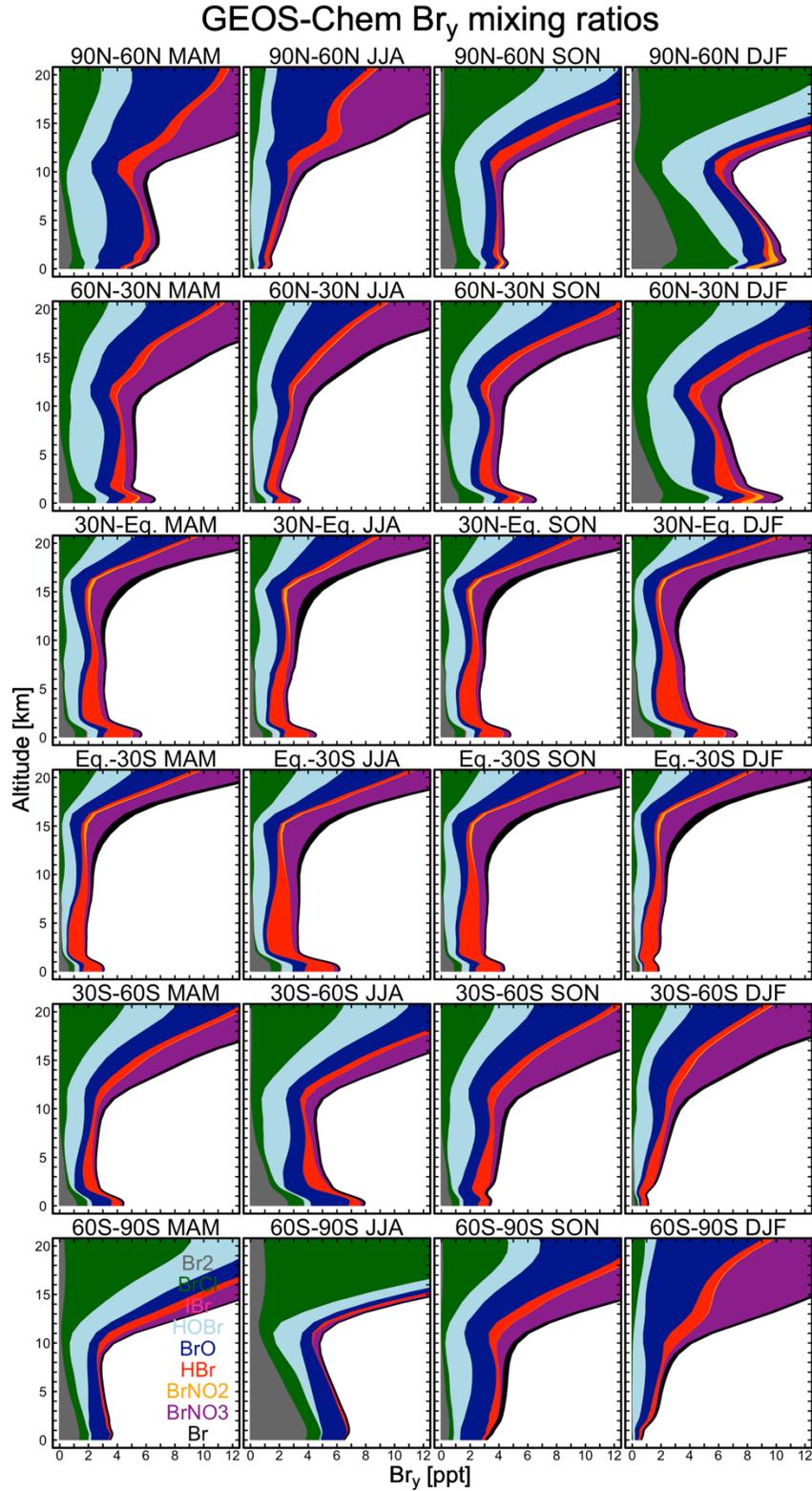


Figure S3. GEOS-Chem Br<sub>y</sub> vertical profiles over different seasons and latitudinal bands. Model results are shown from our standard simulation including sea salt aerosol debromination.

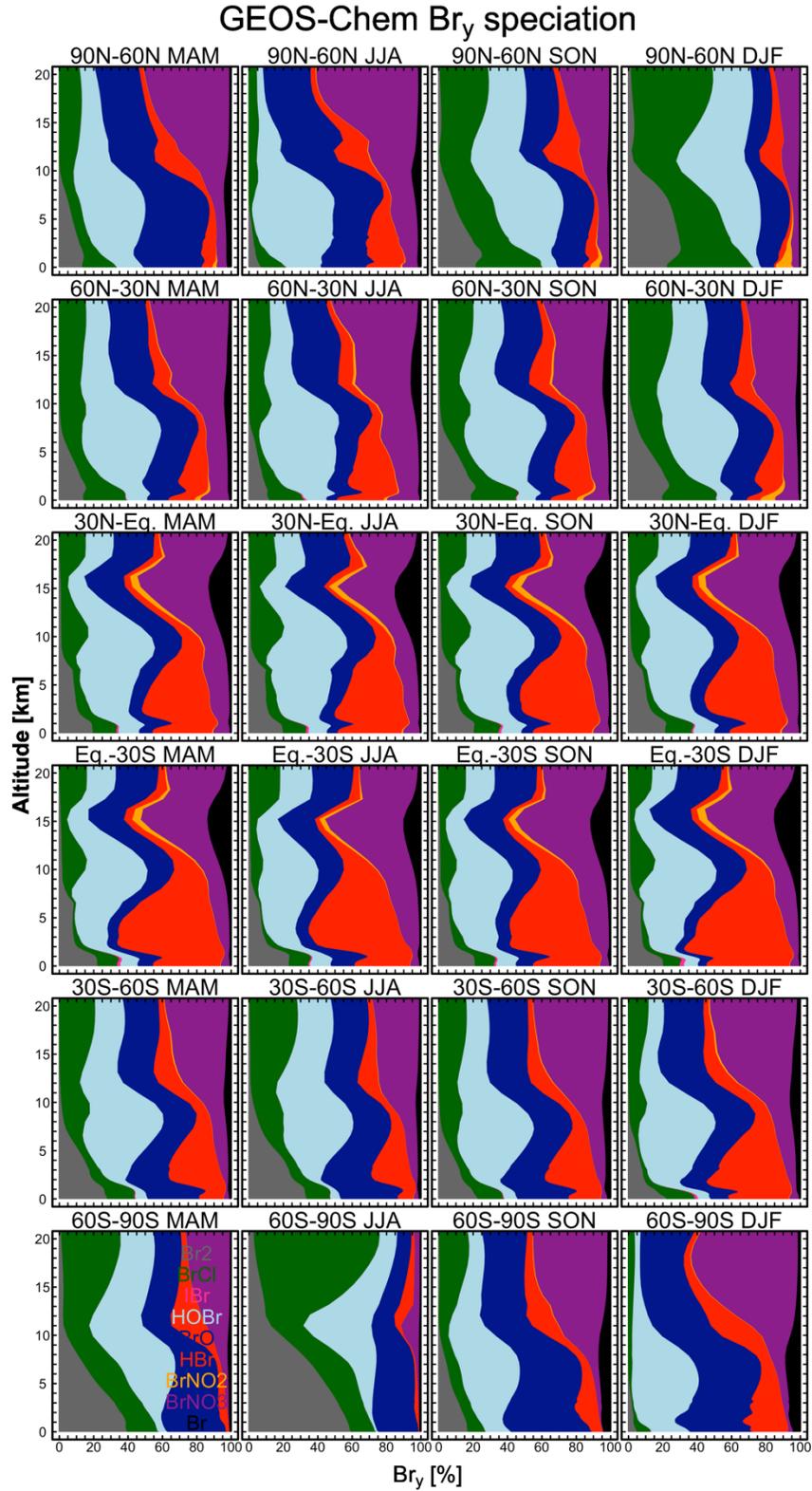


Figure S4. GEOS-Chem Br<sub>y</sub> speciation over different seasons and latitudinal bands. Model results are shown from our standard simulation including sea salt aerosol debromination.

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