



# Ozone source apportionment during peak summer events over southwestern Europe

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**Abstract.** It is well established that in Europe, high O<sub>3</sub> concentrations are most pronounced in southern/Mediterranean countries due to the more favourable climatological conditions for its formation. However, the contribution of the different sources of precursors to O<sub>3</sub> formation within each country relative to the imported (regional and hemispheric) O<sub>3</sub> is poorly quantified. This lack of quantitative knowledge prevents local authorities from effectively designing plans that reduce the exceedances of the O<sub>3</sub> target value set by the European air quality directive. O<sub>3</sub> source attribution is a challenge because the concentration at each location and time results not only from local biogenic and anthropogenic precursors, but also from the transport of O<sub>3</sub> and precursors from neighbouring regions, O<sub>3</sub> regional and hemispheric transport and stratospheric O<sub>3</sub> injections. The main goal of this study is to provide a first quantitative estimation of the contribution of the main anthropogenic activity sectors to peak O<sub>3</sub> events in Spain relative to the contribution of imported (regional and hemispheric) O<sub>3</sub>. We also assess the potential of our source apportionment method to improve O<sub>3</sub> modelling. Our study applies and thoroughly evaluates a countrywide O<sub>3</sub> source apportionment method implemented in the CALIOPE air quality forecast system for Spain at high resolution (4 × 4 km<sup>2</sup>) over a 10-day period characterized by typical summer conditions in the Iberian Peninsula (IP). The method tags both O<sub>3</sub> and its gas precursor emissions from source sectors within one simulation, and each tagged species is subject to the typical physico-chemical pro-

cesses (advection, vertical mixing, deposition, emission and chemistry) as the actual conditions remain unperturbed. We quantify the individual contributions of the largest NO<sub>x</sub> local sources to high O<sub>3</sub> concentrations compared with the contribution of imported O<sub>3</sub>. We show, for the first time, that imported O<sub>3</sub> is the largest input to the ground-level O<sub>3</sub> concentration in the IP, accounting for 46 %–68 % of the daily mean O<sub>3</sub> concentration during exceedances of the European target value. The hourly imported O<sub>3</sub> increases during typical northwestern advections (70 %–90 %, 60–80 µg m<sup>-3</sup>), and decreases during typical stagnant conditions (30 %–40 %, 30–60 µg m<sup>-3</sup>) due to the local NO titration. During stagnant conditions, the local anthropogenic precursors control the O<sub>3</sub> peaks in areas downwind of the main urban and industrial regions (up to 40 % in hourly peaks). We also show that ground-level O<sub>3</sub> concentrations are strongly affected by vertical mixing of O<sub>3</sub>-rich layers present in the free troposphere, which result from local/regional layering and accumulation, and continental/hemispheric transport. Indeed, vertical mixing largely explains the presence of imported O<sub>3</sub> at ground level in the IP. Our results demonstrate the need for detailed quantification of the local and remote contributions to high O<sub>3</sub> concentrations for local O<sub>3</sub> management, and show O<sub>3</sub> source apportionment to be an essential analysis prior to the design of O<sub>3</sub> mitigation plans in any non-attainment area. Achieving the European O<sub>3</sub> objectives in southern Europe requires not only ad hoc local actions but also decided national and European-wide strategies.

## 1 Introduction

Tropospheric ozone ( $O_3$ ) is an air pollutant of major public concern as it harms human health (WHO, 2013) and sensitive vegetation (Booker et al., 2009), and contributes to climate change (Jacob and Winner, 2009).  $O_3$  is formed in the atmosphere through non-linear photochemical reactions among carbon monoxide (CO), peroxy radicals generated by the photochemical oxidation of volatile organic compounds (VOCs) and nitrogen oxides ( $NO_x$ ) (Crutzen, 1974). Therefore, meteorological stagnation, high temperatures, high solar radiation and low precipitation enhance tropospheric  $O_3$  formation (Demuzere et al., 2009; Otero et al., 2016). Atmospheric circulation also controls the short- and long-range transport of  $O_3$  affecting its lifetime in the atmosphere (Monks et al., 2015). For example, the transport of precursors emitted in urban and industrialized areas may cause  $O_3$  production downwind (Holloway et al., 2003).

According to the European Environmental Agency (EEA) around 95 %–98 % of the population in Europe were exposed to  $O_3$  concentrations that exceeded the guidelines of the World Health Organization (WHO) during 2013–2015 period (EEA, 2017). These guidelines establish a maximum daily 8 h averaged (MDA8)  $O_3$  concentration of  $100 \mu\text{g m}^{-3}$  never to be exceeded. The European air quality directive (2008/50/EC) is less restrictive as it sets an  $O_3$  target value of  $120 \mu\text{g m}^{-3}$  for the MDA8 concentration, which can be exceeded up to 25 days per calendar year averaged over 3 years.

Southern European countries around the Mediterranean Basin are particularly exposed to exceedances of the  $O_3$  target value in summer due to the influence of frequent anticyclonic and clear-sky conditions that favour photochemical  $O_3$  formation in the troposphere (EEA, 2017). In addition, its geographic location also makes the basin a receptor of the long-range transport of pollution from Europe, Asia and even North America (Lelieveld et al., 2002; Gerasopoulos, 2005). The importance of long-range transport on surface  $O_3$  has been studied in the Mediterranean Basin, indicating that the emission sources within the basin have a dominating influence on surface  $O_3$ , whereas remote sources are more important than local sources for  $O_3$  mixing ratios at higher altitudes (Richards et al., 2013; Safieddine et al., 2014). Recent studies have suggest that the upper  $O_3$ -rich air masses could increase the surface  $O_3$  concentration in the Mediterranean Basin (Kalabokas et al., 2017; Querol et al., 2018). Further detailed and quantitative studies on the mechanism linking the upper  $O_3$ -rich layer with increases in the ground-level  $O_3$  concentration in episodes require further clarification particularly regarding the contribution of  $O_3$  transported at regional and hemispheric scales.

Several studies in the Iberian Peninsula (IP) have addressed the causes of  $O_3$  episodes by looking at the circulation of air masses (Millán, 2014, and references therein). In the Atlantic region, the blocking anticyclones over western Europe favour the inter-regional transport of  $O_3$  in the

area and its accumulation for several days during the most severe episodes (Alonso et al., 2000; Gangoiti et al., 2002, 2006; Valdenebro et al., 2011; Saavedra et al., 2012; Monteiro et al., 2016). Conversely, on the Mediterranean coast, the typical summer synoptic meteorological conditions with a lack of strong synoptic advection, combined with the orographic characteristics and the sea and land breezes, favour episodes where high levels of  $O_3$  are accumulated by recirculation of air masses loaded with  $O_3$  precursors (Millán et al., 1997, 2000; Toll and Baldasano, 2000; Gangoiti et al., 2001; Pérez et al., 2004; Jiménez et al., 2006; Gonçalves et al., 2009; Millán, 2014; Querol et al., 2017, 2018). The coupling between synoptic and mesoscale processes governing the levels of  $O_3$  in the western Mediterranean Basin need further research in order to understand the  $O_3$  intercontinental contribution. Furthermore, based on our understanding there is a lack of research quantifying the contribution of the activity sources to the  $O_3$  local formation during peak events in this region.

$O_3$  analyses in the western Mediterranean Basin show that regional background  $O_3$  levels have remained high without significant changes (EEA, 2017; EMEP-CCC, 2016; Querol et al., 2016). However, they have increased at traffic and urban background sites (EEA, 2017; Querol et al., 2016; Sicard et al., 2016; Saiz-Lopez et al., 2017). The reasons behind the urban  $O_3$  upward trend are not clear yet due to the complex VOC– $NO_x$  regime; part of the  $O_3$  increase may have resulted from the reduction of NO emissions relative to  $NO_2$  and therefore to a lower NO titration effect in VOC-limited regimes. The most intense  $O_3$  events in the last decade, measured by the number of exceedances of the  $O_3$  target value are recorded over areas downwind of large urban and industrial hot spots (Monterio et al., 2012; Querol et al., 2016; EEA, 2016). Overall, a number of these types of  $O_3$  events occur in June–July and during summer heat waves (i.e. 2003 and 2015).

According to the European air quality directive, in zones exceeding the  $O_3$  target value, EU member states must develop plans to attain compliance by reducing the emission of  $O_3$  precursors. Abatement of the tropospheric  $O_3$  concentration in the western Mediterranean Basin has been insufficient thus far (Querol et al., 2018). Effective planning requires an accurate quantitative knowledge of the sources of these precursors and their respective contributions to the exceedances of the  $O_3$  target value (Querol et al., 2016; Borrego et al., 2016). However, source attribution of surface  $O_3$  concentrations remains a challenge, because the concentration at each location and time results not only from local biogenic and anthropogenic precursors, but also from the transport of  $O_3$  and its precursors from neighbouring regions,  $O_3$  hemispheric transport (UNECE, 2010) and stratospheric  $O_3$  injections (Monks et al., 2015).

At present, there are no methods based on observations that distinguish the origin of  $O_3$ . Despite their inherent uncertainties, chemical transport models (CTMs) allow for the

apportionment of the contribution of any source (by sector and/or region) to O<sub>3</sub> concentrations. The most widely used approach is the “brute force” method, which consists of running an ensemble of simulations zeroing out the sources one by one and then comparing them with a baseline simulation that accounts for all of the sources. Several O<sub>3</sub> source apportionment studies at a European scale have applied the brute force method to quantify the contribution of one or two emission sectors. For example, road transport emissions using the EMEP model (Reis et al., 2000), biogenic and anthropogenic emissions using the Polyphemus model (Sartelet et al., 2012), transport-related emissions including road transport, shipping and aviation using the WRF-CMAQ model (TRANSPHORM, 2014), and ship emissions with CAMx (Aksoyoglu et al., 2016). Brute force is simple to implement, as it does not require additional coding in the CTM. However, as it quantifies the contribution of each source based on its absence, it does not reproduce actual atmospheric conditions; therefore, it is susceptible to inaccuracies in the prediction of O<sub>3</sub> peaks under non-linear regimes (Cohan and Napelenok, 2011). Actually, brute force is not suitable for retrieving source contributions when the relationship between emissions and concentrations is non-linear, but it is useful for analysing the concentration responses to emission abatement scenarios (Clappier et al., 2017).

Recently, CTMs include algorithms that tag multiple pollutants by source (region and/or sector) all the way through the pollutant’s lifetime, from emission to deposition. This integrated source apportionment approach has several advantages. First, it allows for the identification of the main sources contributing to high O<sub>3</sub> levels under actual atmospheric conditions, which is a preliminary step towards designing refined and efficient emission abatement scenarios. Second, as we show below, it supports enhanced model evaluation and therefore potential model improvements by identifying problems in emission estimates (sectors or regions) or chemical boundary conditions. The Integrated Source Apportionment Method (ISAM) within the Community Multiscale Air Quality (CMAQ) model has shown promising results for O<sub>3</sub> tagging, exhibiting less noise in locations where brute force results are demonstrably inaccurate (Kwok et al., 2013, 2015). Recent ISAM experiments have quantified that the contribution of traffic in the cities of Madrid and Barcelona to the daily O<sub>3</sub> peaks downwind of the urban areas is particularly significant (up to 80–100 µg m<sup>-3</sup>) (Valverde et al., 2016a). O<sub>3</sub> tagging methods are also included in other regional and global models applied over Europe (Karamchandani et al., 2017; Butler et al., 2018).

The integrated source apportionment tools combined with high-resolution emission and meteorological models can help unravelling the sources responsible for peak summer events of O<sub>3</sub> in the western Mediterranean Basin. Quantifying the contribution of emission sources during acute O<sub>3</sub> episodes is a prerequisite for the design of future mitigation strategies in the region. In this framework, the main goal

of this study is to provide a first quantitative estimation of the contribution of the main anthropogenic activity sectors compared to the imported concentration (regional and hemispheric) to peak O<sub>3</sub> events in Spain. We also assess the potential of our source apportionment method to improve O<sub>3</sub> modelling. Our study applies, for the first time, a countrywide O<sub>3</sub> source apportionment at high resolution over the IP during the period between 21 and 31 July 2012, which is representative of the typical summer synoptic conditions in the region. We use the CMAQ-ISAM within the CALIOPE air quality forecast system for Spain (<http://www.bsc.es/caliope/es>, last access: 4 April 2019), which runs at a horizontal resolution of 4 × 4 km<sup>2</sup> over the IP. The system is fed by the HERMESv2.0 emission model, which provides disaggregated emissions based on local information and state-of-the-art bottom-up approaches for the most prevalent pollution sectors.

The paper is organized as follows. In Sect. 2 we introduce the CALIOPE system, the set-up of ISAM and the HERMESv2.0 emission model for O<sub>3</sub> source apportionment studies, and the methodology used to quantify the evaluation of the model. In Sect. 3 we demonstrate the representativeness of the selected episode, we evaluate the model and we provide an analysis of the source-sector contribution to Spanish O<sub>3</sub> under the different synoptic patterns occurring during the study period. In Sect. 4, we discuss our findings, the regulatory implications and future research.

## 2 Methodology

### 2.1 Air quality model

We used the CALIOPE air quality modelling system (<http://www.bsc.es/caliope>) to simulate the O<sub>3</sub> dynamics over the IP during the selected episode. CALIOPE is described elsewhere (Baldasano et al., 2008, 2011; Pay et al., 2010, 2012; and reference therein). The system consists of the HERMESv2.0 emission model (Guevara et al., 2013), the WRF-ARWv3.6 meteorological model (Skamarock and Klemp, 2008), the CMAQ v5.0.2 chemical transport model (Byun and Schere, 2006) and the BSC-DREAM8bv2 mineral dust model (Basart et al., 2012). CALIOPE first runs over Europe at a 12 × 12 km<sup>2</sup> horizontal resolution (EU12 domain) and then over the IP at a 4 × 4 km<sup>2</sup> resolution (IP4 domain) (Fig. S1 in the Supplement). In the present work, the system is configured with 38 sigma layers up to 50 hPa, both for WRF and CMAQ. The planetary boundary layer (PBL) is characterized by approximately 11 layers, and the bottom layer’s depth is ∼ 39 m. The EU12 domain uses meteorological initial and boundary conditions from the final analyses provided by the National Centers of Environmental Prediction (FNL/NCEP) at a 0.5° × 0.5° resolution. The first 12 h of each meteorological run are treated as cold start, and the next 23 h are provided to the chemical transport model.

Boundary conditions for reactive gases and aerosols come from the global MOZART-4/GEOS-5 model at  $1.9^\circ \times 2.5^\circ$  horizontal resolution (Emmons et al., 2010). CMAQ uses the CB05 gas-phase mechanism with active chlorine chemistry, an updated toluene mechanism (CB05TUCL; Whitten et al., 2010; Sarwar et al., 2012) and the sixth-generation CMAQ aerosol mechanism including sea salt, aqueous/cloud chemistry and the ISORROPIA II thermodynamic equilibrium module (AERO6; Reff et al., 2009; Appel et al., 2013). Table S1 in the Supplement depicts the remaining CALIOPE configuration options.

For the IP4 domain, HERMESv2.0 estimates emissions for Spain with a temporal and spatial resolution of 1 h and up to  $1 \text{ km} \times 1 \text{ km}$ , according to the Selected Nomenclature for Air Pollution (SNAP), which are then aggregated to a  $4 \times 4 \text{ km}^2$  resolution (Guevara et al., 2013). HERMESv2.0 is suitable for source apportionment studies thanks to its level of detail in the calculation of the emission fluxes by source (Guevara et al., 2014). The model is currently based on data from 2009, which was the closest year with updated information on local emission activities in HERMESv2.0 at the time that this work started. For neighbouring countries and international shipping activities, HERMESv2.0 uses the annual gridded national emission inventory provided by the European Monitoring and Evaluation Programme (EMEP) disaggregated to a  $4 \times 4 \text{ km}^2$  resolution using a SNAP-sector-dependent spatial, temporal and speciation treatment (Ferreira et al., 2013).

HERMESv2.0 integrates the Model of Emissions of Gas and Aerosols from Nature (MEGANv2.0.4; Guenther et al., 2006) to estimate VOCs and  $\text{NO}_x$  emissions from vegetation, which play a major role in  $\text{O}_3$  photochemistry, using temperature and solar radiation from the WRF model. Note that we configured MEGAN to compute VOC emissions from cultivated crops; the agriculture emission module in HERMESv2.0 estimates the VOCs from manure management and field burning of agricultural residues. In this study, we have updated MEGANv2.0.4 with emission factors from MEGANv2.1 (<http://lar.wsu.edu/megan/guides.html>, last access: 4 April 2019). In Sect. 2 of the Supplement, we provide a comparison with measurements from the DAURE campaign (Pandolfi et al., 2014) that shows the reasonably good behaviour of our modelled isoprene.

Urban VOC emissions could be a relevant source of  $\text{O}_3$ . Over Spanish urban areas, HERMESv2.0 estimates VOC emissions from road transport and the use of solvents (Fig. 1) following bottom-up approaches (Guevara et al., 2013). However, uncertainties in the estimation of urban VOC emission inventories, as stated recently by several works (Pan et al., 2015; Liu et al., 2017; McDonald et al., 2018; Lewis, 2018) makes the urban VOC contribution to tropospheric  $\text{O}_3$  uncertain. In order to overcome this problem, continuous monitoring of urban VOCs should be performed in Spanish cities, following the example of other regions in which  $\text{O}_3$  is also a major problem such as Mexico City (Jaimes-Palomera

et al., 2016). In addition, the use of formaldehyde satellite observations to constrain urban VOC emissions could also be pointed out as a future task to improve the representativeness of urban emission inventories (Zhu et al., 2014).

## 2.2 Ozone source apportionment method

We applied ISAM to quantify contributions from different SNAP categories to the surface  $\text{O}_3$  over the IP. The ISAM  $\text{O}_3$  tagging method is a mass-balance technique that tags both  $\text{O}_3$  and its gas precursor emissions ( $\text{NO}_x$  and VOC) from each source sector within one simulation (Kwok et al., 2013, 2015). Each tagged species undertakes typical physico-chemical processes (advection, vertical mixing, deposition, emission and chemistry) without perturbing the actual conditions. The  $\text{O}_3$  rate of change for each tag in any grid cell is calculated as follows (Eq. 1):

$$\frac{dC_{\text{tag}}}{dt} = P_{\text{tag}} - D \frac{C_{\text{tag}}}{\sum_{\text{tag}} C}, \quad (1)$$

where  $C_{\text{tag}}$  represents the  $\text{O}_3$  concentration related to a tagged source of interest,  $P_{\text{tag}}$  is the chemical production rate of  $\text{O}_3$  formed by the precursors emitted for each tag and  $D$  is the total chemical destruction rate of  $\text{O}_3$  in this grid cell. Different ratios of  $\text{NO}_x/\text{VOC}$  cause the formation of  $\text{O}_3$  in each grid cell, which is either controlled by  $\text{NO}_x$ - or VOC-limited conditions. ISAM uses the  $\text{H}_2\text{O}_2/\text{HNO}_3$  ratio to determine whether  $\text{O}_3$  is  $\text{NO}_x$ - or VOC-sensitive (above or below 0.35, respectively) (Zhang et al., 2009). The bulk  $\text{O}_3$  concentration in each model grid cell ( $P_{\text{bulk}}$ ) is equal to the sum of  $\text{O}_3$  tracers that were produced under either  $\text{NO}_x$ - or VOC-sensitive conditions (Eq. 2),

$$P_{\text{bulk}} = \sum_{\text{tag}} P_{\text{tag}} = \sum_{\text{tag}} P_{\text{tag}}^{\text{N}} + \sum_{\text{tag}} P_{\text{tag}}^{\text{V}}, \quad (2)$$

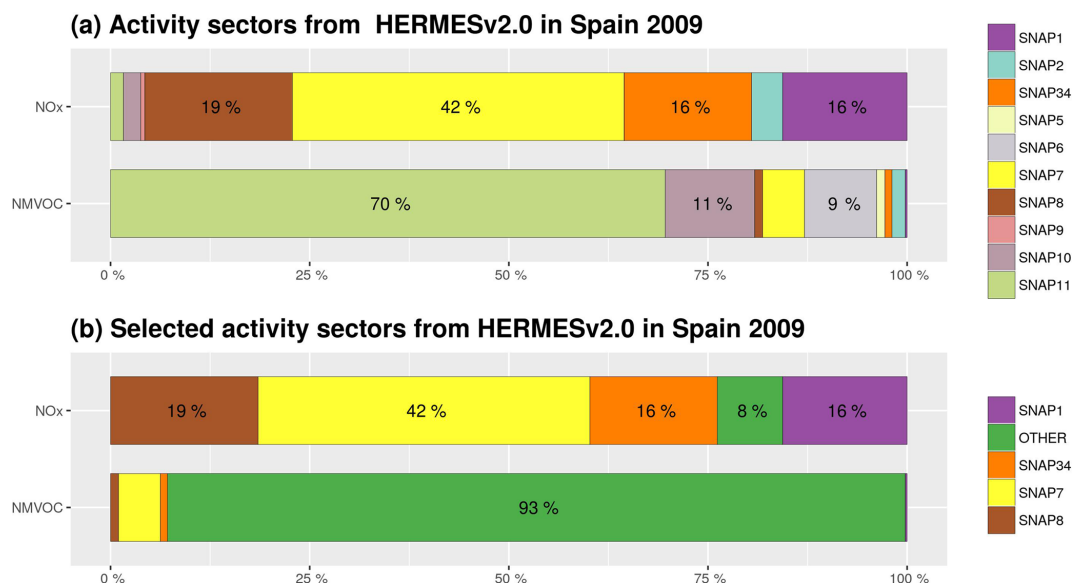
where  $P_{\text{tag}}^{\text{N}}$  and  $P_{\text{tag}}^{\text{V}}$  are the  $\text{O}_3$  produced under  $\text{NO}_x$ - and VOC-limited conditions, respectively, according to Eqs. (3) and (4):

$$P_{\text{tag}}^{\text{N,new}} = P_{\text{tag}}^{\text{N,old}} + P_{\text{bulk}}^{\text{new}} \frac{\sum_x \text{NO}_{x,\text{tag}}}{\sum_{\text{tag}} \sum_x \text{NO}_{x,\text{tag}}}, \quad (3)$$

$$P_{\text{tag}}^{\text{V,new}} = P_{\text{tag}}^{\text{V,old}} + P_{\text{bulk}}^{\text{new}} \frac{\sum_y \text{VOC}_{y,\text{tag}} \times \text{MIR}_y}{\sum_{\text{tag}} \sum_y \text{VOC}_{y,\text{tag}} \times \text{MIR}_y}. \quad (4)$$

In Eqs. (3) and (4),  $\text{NO}_{x,\text{tag}}$  and  $\text{VOC}_{j,\text{tag}}$  are the respective concentrations of the  $x$  nitrogen and  $y$  VOC species in CB05 that participate in the photochemical  $\text{O}_3$  formation for each source sector tag and grid cell, and  $\text{MIR}_y$  is the maximum incremental reactivity factor of each  $y$  species of VOC emitted by each source-sector tag, corresponding to the  $\text{O}_3$  generating potential of each single VOC species (Carter, 1994).





**Figure 1.** Percentage of the contribution of emissions to total annual emissions by SNAP sector calculated by HERMES for Spain 2009 (a) and for the selected SNAP sector accounting for more than 90 % of NO<sub>x</sub> total emission to be tracked with ISAM (b). “OTHER” compiles the SNAP categories 2 (residential combustion), 5 (fugitive emissions from fuels), 6 (solvent use), 9 (waste management), 10 (agriculture) and 11 (other sources). (NMVOC refers to non-methane volatile organic compounds.)

### 2.3 Ozone tagged species

Table 1 summarizes the O<sub>3</sub> tagged sources in the present study, and Fig. 1a depicts the HERMESv2.0 model estimates of the contribution from each SNAP category to the total emissions of O<sub>3</sub> precursors in Spain. The largest NO<sub>x</sub> sources are road transport (SNAP7, 42 %), non-road transport (SNAP8, 19 %), manufacturing industries (SNAP34, 16 %) and energy production (SNAP1, 16 %). VOCs are dominated by biogenic sources (SNAP11, 70 %) and to a lesser extent by the agricultural sector (SNAP10, 11 %), solvent and other product uses (SNAP6, 9 %), and road transport (SNAP7, < 7 %). The selected (tagged) SNAP categories in this study are the energy, industrial, road transport and non-road transport sectors (Fig. 1b), which account for 92 % of the total NO<sub>x</sub> emissions in Spain. An additional tracer (OTHER) gathers the remaining emission categories that were not explicitly tracked (i.e. SNAP2, 5, 6, 9, 10 and 11).

In addition to the selected sources, we tracked the contributions of the chemical boundary conditions (BCON) and the initial conditions (ICON). BCON represents both the O<sub>3</sub> directly transported through the IP4 domain boundaries and the formation of O<sub>3</sub> resulting from precursors that are also transported through the boundaries. BCON O<sub>3</sub> comes from the EU12 parent domain, which includes the O<sub>3</sub> produced in Europe and the O<sub>3</sub> transported at a global scale (both tropospheric and stratospheric O<sub>3</sub>) provided by the MOZART-4/GEOS model (Fig. S1 in the Supplement). In the following, we name BCON O<sub>3</sub> as the imported O<sub>3</sub> to the IP4 do-

main. Tagging the initial O<sub>3</sub> allows for the quantification of the number of spin-up days to minimize the impact of model initialization. For the present run, we required 6 days of spin-up to set the contribution of initial conditions to less than 1 % of the net hourly O<sub>3</sub> concentration over 95 % of the available O<sub>3</sub> stations.

### 2.4 Evaluation method

We evaluate the simulated concentrations against air quality measurements from the Spanish monitoring stations that are part of the European Environment Information and Observation Network (EIONET; <https://www.eionet.europa.eu/>, last access: 4 April 2019). The EIONET network provides a relatively dense geographical coverage of the Spanish territory. During the 21–31 July episode, we used the measurements from 347 stations for O<sub>3</sub> and 357 stations for NO<sub>2</sub> with a temporal coverage above 85 % on an hourly basis. Figure S2 in the Supplement shows the distribution of the stations for O<sub>3</sub> and NO<sub>2</sub>.

The evaluation based on discrete statistics includes the correlation coefficient ( $r$ ), the mean bias (MB), the normalized mean bias (NMB) and the root mean square error (RMSE) (Appendix B). We used the “openair” package (Carslaw and Ropkins, 2012) for R (v3.3.2; R Core Team, 2016) to compute the statistics. We calculate statistics on an hourly basis for O<sub>3</sub> and NO<sub>2</sub>, as well as for the regulatory MDA8 in the case of O<sub>3</sub>. The evaluation also takes the station type, following the categories established by the EEA into account (i.e.

**Table 1.** Description of the O<sub>3</sub> tagged sources in the present study.

ISAM tag*	Emission by SNAP category	Description
SNAP1	SNAP1	SNAP1: energy industry
SNAP34	SNAP34	SNAP34: manufacturing industries (combustion and processes)
SNAP7	SNAP7	SNAP7: road transport, exhaust and non-exhaust
SNAP8	SNAP8	SNAP8: non-road transport (international shipping, airport and agricultural machinery)
OTHER	SNAP2 + SNAP5 + SNAP6 + SNAP9 + SNAP10 + SNAP11	SNAP2: residential and commercial/institutional combustion; SNAP5: fugitive emissions from fuels; SNAP6: product use including solvents; SNAP9: waste management; SNAP10: agriculture; and SNAP11: other sinks
BCON	–	Chemical boundary conditions of the IP4 domain from the EU12 simulation which includes the contribution from Europe and the international contribution from MOZART-4. O <sub>3</sub> external contribution
ICON	–	Initial chemical condition of the IP4 domain

\* Each ISAM tag is applied to O<sub>3</sub> and its precursor species in the CB05 (NO<sub>x</sub> and VOCs). NO<sub>x</sub> species contributing to O<sub>3</sub> formation involve (9 species): NO, NO<sub>2</sub>, nitrogen trioxide (NO<sub>3</sub>), dinitrogen pentoxide (N<sub>2</sub>O<sub>5</sub>), nitrous acid (HONO), peroxyacyl nitrates (PAN), higher peroxyacyl nitrates (PANX), peroxyacetic acid (PAA) and organic nitrates (NTR). VOC species contributing to O<sub>3</sub> formation include (14 species): acetaldehyde (ALD2), higher aldehydes (ALDX), ethene (ETH), ethane (ETHA), ethanol (ETOH), formaldehyde (FORM), internal olefin (IOLE), isoprene (ISOP), methanol (MEOH), olefin (OLE), paraffin (PAR), monoterpene (TERP), toluene (TOL) and xylene (XYL).

rural background, suburban background, urban background, industrial and traffic).

There are no direct evaluation methods for apportioned pollutants. Instead, we designed a diagnostic plot for source apportionment analysis at each individual receptor, including a time series of measured and observed O<sub>3</sub> and NO<sub>2</sub> concentrations as well as the simulated tagged sources. In addition, this plot includes the simulated wind speed and direction. These plots are helpful as they compare the modelled O<sub>3</sub> and NO<sub>2</sub> with the observations, while highlighting the sources and circulation patterns at least partly responsible for the model behaviour. This work will only discuss the source apportionment plots at key O<sub>3</sub> receptor regions in detail, given the high number of stations (260) that simultaneously measure O<sub>3</sub> and NO<sub>2</sub>.

Evaluation results are discussed together with the source apportionment results. On the one hand, the interpretation of the source apportionment results benefits from model evaluation. On the other hand, the source apportionment results support enhanced model evaluation as it allows for the identification of potential errors in emission estimates for specific sectors and/or in the chemical boundary conditions.

### 3 Results

#### 3.1 Description of the ozone episode

Our first estimation of the origin of peak O<sub>3</sub> events in Spain focuses on the episode from 21 to 31 July 2012. Figure 2a illustrates the relevance of the episode showing the observed MDA8 O<sub>3</sub> concentrations trends at the Spanish EIONET stations during the (extended) summers (i.e. from April to September) from 2000 to 2012, in addition to the concen-

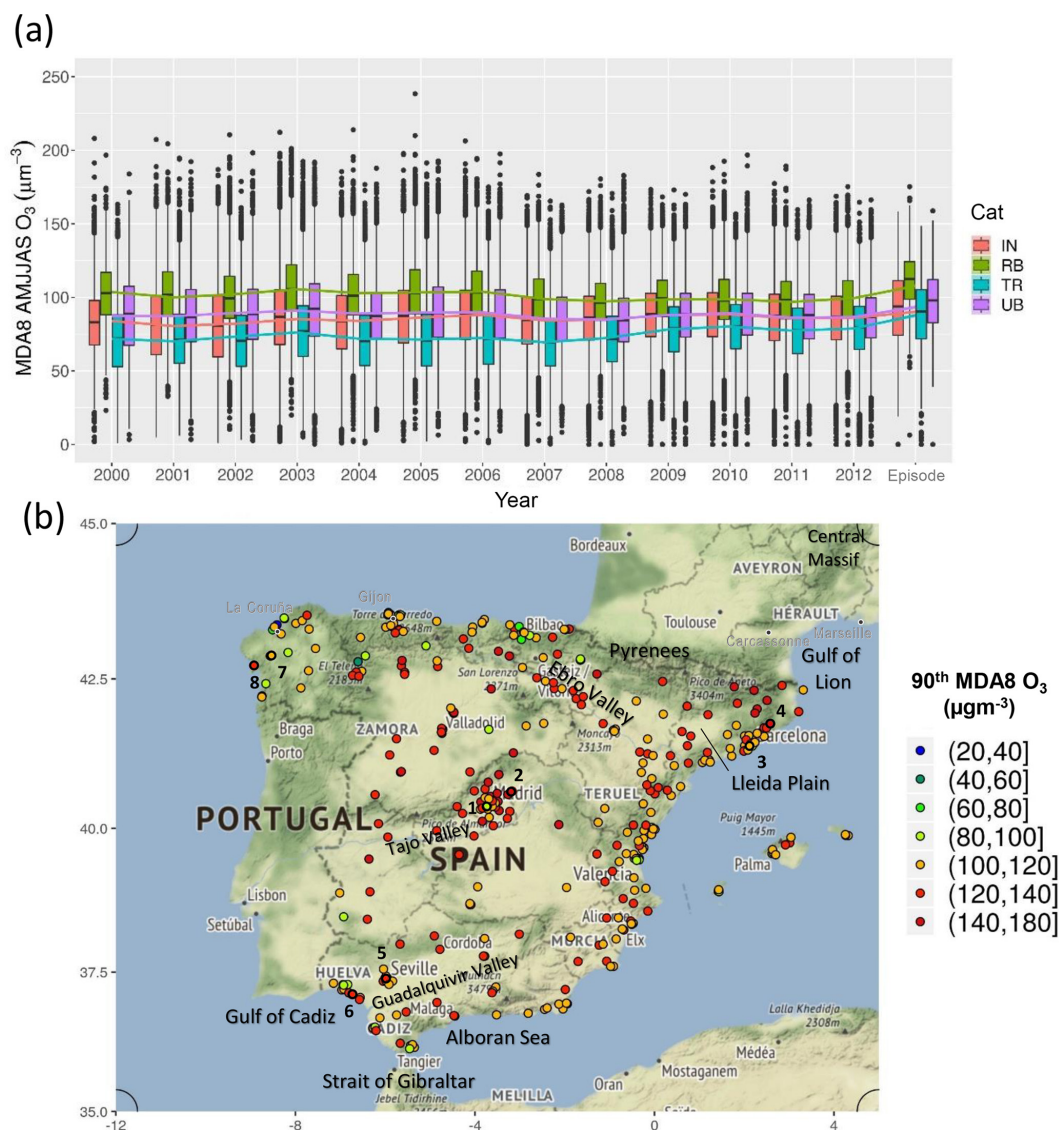
trations recorded during the episode. Although the selected episode is not the most severe between 2000 and 2012 at a national scale, it comprises a period with high MDA8 O<sub>3</sub> concentrations measured at rural background stations, (the 75th percentile of those values was above the target value) similar to the particularly severe summer of 2003 (Solberg et al., 2008).

This episode is also interesting because it was widespread and affected large areas of Europe (EEA, 2013). During this period alone, 33 % and 12 % of the total number of exceedances for the information threshold and the target value in 2012, respectively, were measured. The O<sub>3</sub> regional context of the episode allows us to study the influence of the imported O<sub>3</sub> to Spain.

The maps of the 90th percentile of the measured MDA8 O<sub>3</sub> concentrations over Spain (Fig. 2b) show high concentration spots throughout the domain. The exceedances of the target value were found in the regions surrounding large urban areas (Madrid, Barcelona, Valencia and Seville) and along Spanish valleys (i.e. Ebro Valley and Guadalquivir Valley).

There were more than 100 exceedances of the O<sub>3</sub> target value on most of the days during the episode, with relative maxima on 25, 28 and 31 July attributed to the change in the synoptic conditions (Fig. S3 in the Supplement). Figure 3 shows the meteorological patterns (2 m temperature, 10 m wind, precipitation, mean sea level pressure and geopotential height at 500 hPa) modelled by WRF-ARW during the 3 distinctive days over the outer EU12 domain.

Our characterization of the study period is based on the circulation type classification proposed in Valverde et al. (2014), who applied an objective synoptic classification method over the period from 1983 to 2012, specifically designed to study air quality dynamics over the IP. Stagnant



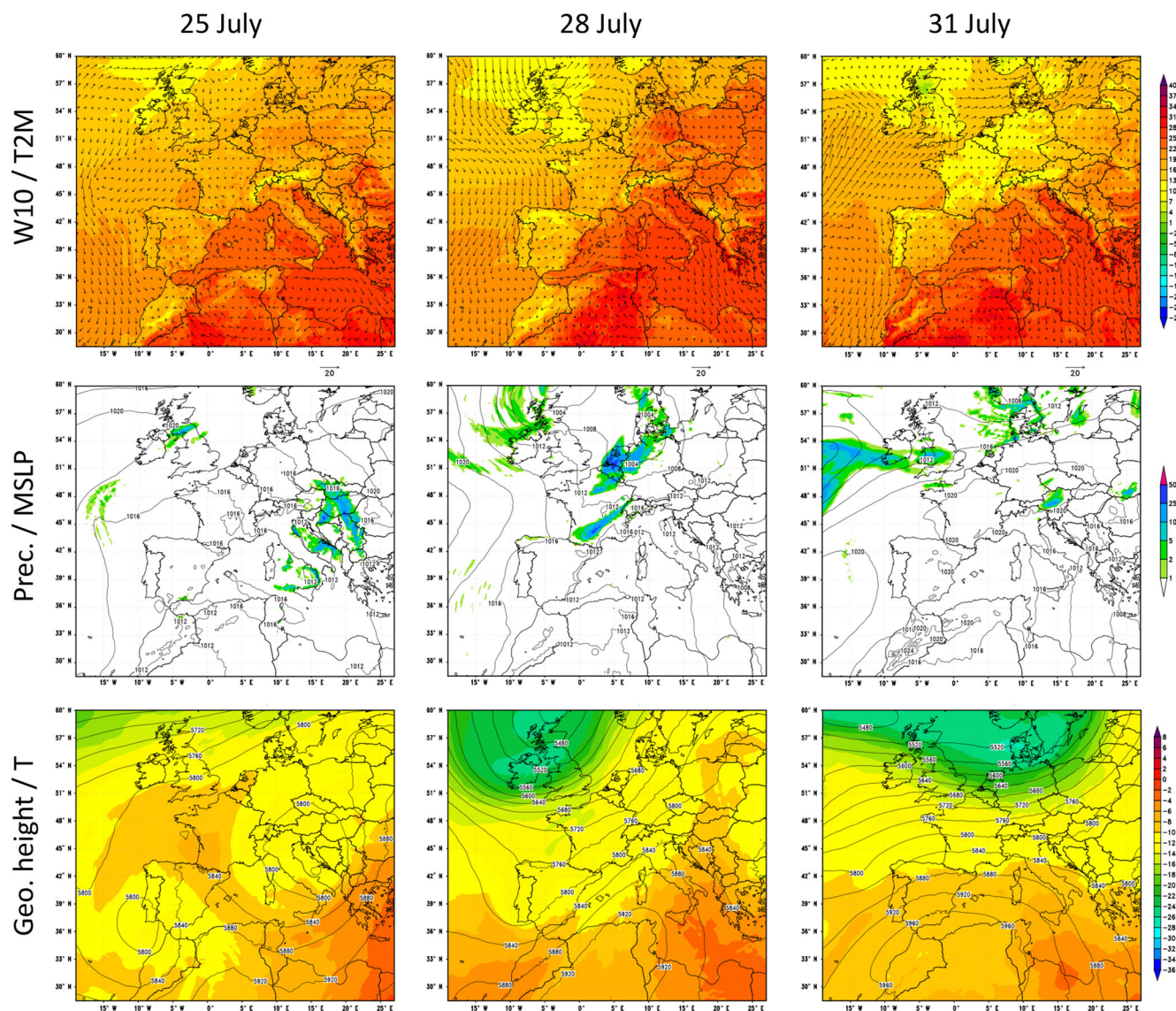
**Figure 2.** (a) Temporal distribution of the MDA8 O<sub>3</sub> concentration during the extended summer (from April to September, AMJJAS) at the Spanish EIONET stations for the 2000–2012 period and the episode (from 21 to 31 July 2012) by station type: IN (industrial), RB (rural background), TR (traffic) and UB (urban background). (b) The 90th percentile of the MDA8 O<sub>3</sub> concentration at the Spanish EIONET stations during the episode. Numbers indicate the stations cited in Sect. 3.4.

conditions and northwestern advections are the most frequent summer synoptic circulation patterns over the IP, occurring on ~ 44 % of the days in a year (Jorba et al., 2004; Valverde et al., 2014). Stagnant conditions are characterized by reduced surface pressure gradients and weak synoptic winds, intense solar radiation, and the development of the Iberian thermal low (ITL). The ITL forces the convergence of surface winds from the coastal areas towards the central plateau enhancing sea breezes and mountain–valley winds and subsidence over the western Mediterranean Basin, as described by Millán et al. (1997, 2000) and Millán (2014). In contrast, northwestern advections (NWad) transport air masses from the Atlantic towards the north and west of the IP and they

are characterized by atmospheric instability and intense ventilation. Periods of accumulation and venting of pollutants follow the same sequence of pressure ridging and troughing respectively, of the lower and middle troposphere of the IP during the warm season (Querol et al., 2017, 2018). According to the circulation type classification in Valverde et al. (2014), the selected episode started with the development of the ITL (21–25 July), followed by a NWad-venting period (26–29 July) and ended with the development of another ITL (30–31 July).

Figure 4 shows the 90th percentile (p90) of the simulated hourly O<sub>3</sub> and NO<sub>2</sub> concentrations corresponding to the 3 distinctive days with the relative maxima of exceedances.



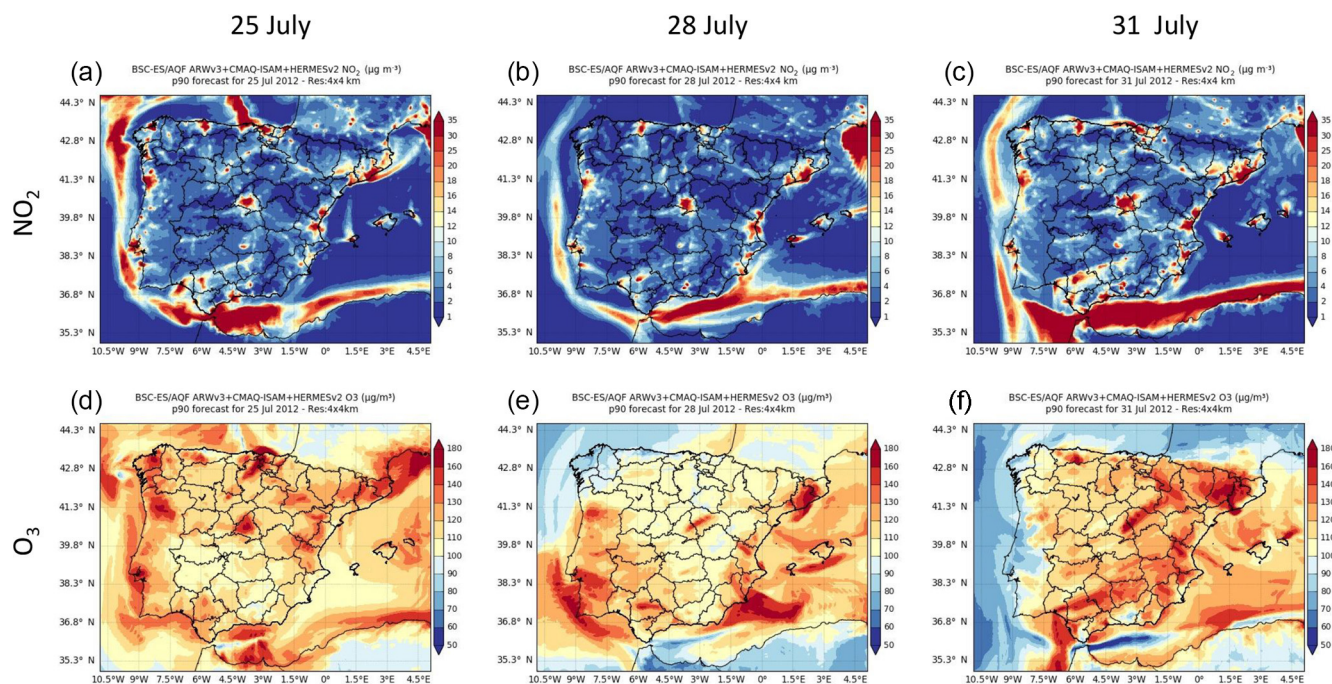


**Figure 3.** WRF-ARW meteorological fields at 06:00 UTC for 25, 28 and 31 July in the EU12 domain: 10 m wind speed (W10,  $\text{ms}^{-1}$ ), 2 m temperature (T2M,  $^{\circ}\text{C}$ ), 6 h accumulated precipitation (Prec., mm), mean sea level pressure (MSLP, hPa), 500 hPa geopotential height in contours (Geo. height, m), and 500 hPa temperature in shaded colours ( $T$ ,  $^{\circ}\text{C}$ ).

In the northern Spanish Mediterranean areas, intense  $\text{O}_3$  episodes often affect the plains and valleys located 60 km north of the Barcelona metropolitan area (BMA) in summer (Toll and Baldasano, 2000; Gonçalves et al., 2009; Valverde et al., 2016a; Querol et al., 2017). High  $\text{NO}_x$  concentrations from the BMA combined with high biogenic VOC levels are driven inland by mesoscale processes (sea breezes and mountain–valley winds). This happened on 31 July when the highest p90 of the hourly  $\text{O}_3$  concentrations ( $160\text{--}180\text{ }\mu\text{g m}^{-3}$ ) in Spain occurred over the north and northwest regions of the BMA. Occasionally, as occurred on 25 July, anticyclonic winds over the western Mediterranean Sea deflect the sea-breeze flow enriched with precursors from the

BMA towards the Gulf of Lion where it reaches the highest p90 of the hourly  $\text{O}_3$  concentrations ( $160\text{--}180\text{ }\mu\text{g m}^{-3}$ ) in the IP Mediterranean region. Eastern Spanish Mediterranean areas show similar  $\text{O}_3$  dynamics, with inland regions depicting the highest  $\text{O}_3$  peaks ( $140\text{--}160\text{ }\mu\text{g m}^{-3}$ ) when stagnant conditions cover the central and eastern IP.

In the centre of the IP, intense  $\text{O}_3$  episodes occurred during the development of the ITL, where the affected area depends on the synoptic conditions (Querol et al., 2018). Under the absence of synoptic forcing (e.g. 25 July), the MMA had the highest p90 of the hourly  $\text{O}_3$  concentrations ( $\sim 140\text{--}160\text{ }\mu\text{g m}^{-3}$ ). In contrast, when mountain–valley winds are reinforced with synoptic westerlies (e.g. 31 July; Fig. 3) the



**Figure 4.** Ground-based concentration maps (in  $\mu\text{g m}^{-3}$ ) for NO<sub>2</sub> (a, b, c) and O<sub>3</sub> (d, e, f) corresponding to the 90th percentile of the average hourly concentrations on 25 (a, d), 28 (b, e) and 31 (c, f) July 2012.

urban NO<sub>x</sub> plume is channelled along the mountain ranges in Madrid towards the northeast and the highest p90 values of the hourly O<sub>3</sub> concentrations are found along the valley ( $\sim 140\text{--}160 \mu\text{g m}^{-3}$ ).

In the north and northeast of the IP, the p90 values of the hourly O<sub>3</sub> concentrations show a significant increase when the blocking anticyclone over western Europe is combined with the development of the ITL (e.g. 25 July). The stagnant conditions favour the accumulation of O<sub>3</sub> precursors around main cities and industrial areas and enhance the local O<sub>3</sub> formation.

The NWad pattern (e.g. 28 July) significantly decreases the p90 of the hourly O<sub>3</sub> concentrations in the centre and north of the IP. The northwesterly winds decrease the temperature and therefore the O<sub>3</sub> formation. As a consequence, O<sub>3</sub> levels are reduced in the plumes from the BMA and the MMA, although they are still significant in the latter. Overall, the p90 of the hourly O<sub>3</sub> concentrations during the NWadv pattern was  $\sim 100 \mu\text{g m}^{-3}$  in most background areas. In contrast, during the ITL it was above  $120 \mu\text{g m}^{-3}$ .

### 3.2 Statistical evaluation

CALIOPE has been evaluated in detail elsewhere (Pay et al., 2014, and references therein). Furthermore, the system has been evaluated using the DELTA Tool (Thunis and Cuvelier, 2014) developed by the Forum for Air Quality Modelling in Europe to support and harmonize the model evaluation in the frame of the air quality directive. Valverde et al. (2016a, b)

used DELTA Tool v4.0 and showed that CALIOPE accomplishes the quality objectives as defined in the air quality directive for 78 % of the NO<sub>2</sub> and 91 % of the O<sub>3</sub> monitoring stations during the summer 2012. Here, we evaluate the updated version of CALIOPE using ISAM to quantify the system's ability to reproduce O<sub>3</sub> and NO<sub>2</sub> concentrations during the selected episode. Table 2 compiles the quartiles of the statistics calculated by station type.

The model slightly overestimates the average hourly and MDA8 O<sub>3</sub> concentrations with MB values of +12 and +6  $\mu\text{g m}^{-3}$ , respectively. The  $r$  is above 0.6 at more than 50 % of the stations and above 0.7 at 25 % of them. The MB values for the average hourly and MDA8 O<sub>3</sub> concentrations are lower at RB stations ( $\pm 4 \mu\text{g m}^{-3}$ ) than at IN, TR and UB stations (between +6 and +16  $\mu\text{g m}^{-3}$ ) at 50 % of the stations. As expected, the highest number of exceedances of the O<sub>3</sub> target value was recorded at RB stations (260 exceedances) followed by IN stations (204 exceedances).

At RB stations, average hourly O<sub>3</sub> is overestimated (+4  $\mu\text{g m}^{-3}$ ) and MDA8 O<sub>3</sub> is underestimated (−4  $\mu\text{g m}^{-3}$ ), which indicates that nighttime O<sub>3</sub> is overestimated. The nighttime overestimation is a common feature of CTMs and it is typically attributed to the underestimation of the O<sub>3</sub> titration by NO (Bessagnet et al., 2016; Sharma et al., 2017).

CALIOPE underestimates the average hourly NO<sub>2</sub> concentrations with −7  $\mu\text{g m}^{-3}$  at TR stations and −2  $\mu\text{g m}^{-3}$  at RB stations. This partly explains the high overestimation of the average hourly and MDA8 O<sub>3</sub> concentration at TR and UB stations, as well as the systematic overestimation



**Table 2.** Statistics for average hourly O<sub>3</sub>, MDA8 O<sub>3</sub> and average hourly NO<sub>2</sub> concentrations in the episode as a function of the station type. Exceedances indicate the number of exceedances of the European air quality directive standards for hourly O<sub>3</sub> (180 µg m<sup>-3</sup>), MDA8 O<sub>3</sub> (120 µg m<sup>-3</sup>) and average hourly NO<sub>2</sub> (200 µg m<sup>-3</sup>). *N* indicates the number of monitoring stations used in the statistical calculation. MO and MM depict the measured and modelled mean concentrations, respectively. Statistics are calculated by considering more than 75 % of the hours in a day, as established by Directive 2008/50/EC. The statistics correspond to the 50th (25th, 75th) quantiles by station. Type indicates the station categories in the calculation of statistics: all of the stations (ALL), industrial (IN), traffic (TR), urban background (UB), suburban background (SB) and rural background (RB) stations.

Pollutant	Type	Exceedances: observed/modelled	<i>N</i>	MO (µg m <sup>-3</sup> )	MM (µg m <sup>-3</sup> )	MB (µg m <sup>-3</sup> )	NMB (%)	RMSE (µg m <sup>-3</sup> )	<i>r</i>
Hourly O <sub>3</sub>	ALL	26/216	348	77.3 (66.7, 86.5)	89.9 (83.5, 95.6)	12.6 (4.4, 20.2)	16.8 (5.1, 29.2)	26.7 (21.5, 32.1)	0.65 (0.57, 0.72)
	IN	5/23	106	74.1 (62.2, 83.2)	82.2 (83.4, 94.5)	14.1 (4.5, 21.5)	19.8 (5.5, 34.7)	26.8 (21.1, 32.8)	0.66 (0.56, 0.73)
	TR	0/58	70	68.4 (57.1, 76.7)	83.8 (74.4, 89.2)	15.9 (8, 21.8)	23.4 (10.1, 38.5)	28.8 (24.5, 33.9)	0.63 (0.54, 0.70)
	UB	0/78	56	74.8 (64.1, 80.0)	89.4 (81.4, 94.5)	15.6 (7.8, 21.3)	22.1 (10.2, 31.1)	28.2 (24.5, 33.5)	0.65 (0.60, 0.70)
	SB	4/48	44	83.2 (78.1, 87.2)	93.5 (88.2, 97.3)	10 (4.6, 14.6)	12.9 (5.4, 17.5)	25.1 (21.4, 29.3)	0.66 (0.60, 0.72)
	RB	17/9	66	91.2 (80.9, 97.1)	96.2 (92.6, 99.6)	4.5 (-3.3, 14.7)	4.8 (-3.2, 17.2)	21.2 (18.3, 28.0)	0.67 (0.57, 0.74)
MDA8 O <sub>3</sub>	ALL	751/822	348	101 (91.0, 113.4)	106.2 (101.8, 113.2)	5.7 (-3.7, 16.9)	5.5 (-3.3, 17.4)	17.9 (13.5, 25.4)	0.64 (0.39, 0.78)
	IN	204/187	106	96.7 (88.9, 109.5)	104.8 (99.4, 109.2)	5.7 (-2.9, 16.5)	5.7 (-2.9, 17.8)	17.0 (13.0, 24.5)	0.68 (0.48, 0.82)
	TR	62/145	70	92.7 (83.6, 99.9)	104.1 (98.3, 110.2)	13.3 (4.6, 23.8)	15.1 (4.6, 27.2)	21.2 (15.2, 31.2)	0.54 (0.27, 0.76)
	UB	86/174	56	100.0 (89.3, 105.5)	107.6 (102.2, 122.8)	14.1 (2.1, 21.7)	14.8 (2.1, 22.5)	22.6 (15.6, 28)	0.59 (0.34, 0.77)
	SB	139/131	44	110.7 (99.8, 118.5)	110.0 (105.5, 119.2)	3.1 (-6.2, 10.5)	2.9 (-5.4, 10.1)	16.9 (13.6, 22)	0.5 (0.31, 0.67)
	RB	260/185	66	113.1 (104.9, 119.9)	108.6 (104.5, 113.2)	-3.7 (-11.1, 5.5)	-3.3 (-9.1, 5.1)	15 (11.8, 20.2)	0.72 (0.60, 0.84)
Hourly NO <sub>2</sub>	ALL	3/0	358	14.0 (8.4, 21.2)	9.9 (4.4, 16.2)	-4.1 (-8.4, -0.5)	-31.9 (-54.2, -4.9)	12.8 (8.2, 17.6)	0.43 (0.29, 0.55)
	IN	0/0	120	11.5 (7.9, 16.3)	8.8 (3.8, 14.7)	-3.0 (-6.6, 0.2)	-28.6 (-55.0, 1.1)	11.0 (7.9, 15.6)	0.40 (0.26, 0.50)
	TR	3/0	95	22.9 (17.7, 28.7)	14.1 (8.7, 21.6)	-7.4 (-13.9, -2.6)	-35.2 (-58.1, -14.3)	18.4 (14.1, 23.1)	0.42 (0.28, 0.57)
	UB	0/0	63	17.0 (13, 21.2)	13.2 (7.8, 18.4)	-4.0 (-8, -0.3)	-26.9 (-47.4, -2)	14.1 (10.5, 17.3)	0.49 (0.36, 0.63)
	SB	0/0	33	12 (8.2, 14.6)	7.6 (4.8, 11.1)	-2.8 (-6.5, 0.8)	-34.4 (-48.6, 7.5)	6.4 (7.9, 13.6)	0.5 (0.41, 0.65)
	RB	0/0	41	4.3 (3.5, 9.3)	2.8 (1.6, 4.0)	-2.0 (-5.3, -0.3)	-49.5 (-72.3, -9.7)	4.4 (2.7, 7.1)	0.34 (0.24, 0.43)

of average hourly  $\text{O}_3$  concentration at nighttime (due to a lack of  $\text{O}_3$  titration by  $\text{NO}$ ). The average hourly  $\text{NO}_2$  concentration at TR stations features the highest  $r$  (with 25 % of stations above 0.6), which proves the reasonably accurate representation of temporal emission in urban areas by the HERMESv2.0 model (Guevara et al., 2014; Baldasano et al., 2011). In contrast, the RMSE is highest at TR stations, which results from the underestimation of  $\text{NO}_2$  peaks during traffic rush hours. Underestimation of  $\text{NO}_2$  traffic peaks is a common problem in Eulerian mesoscale models (Pay et al., 2014), as emission heterogeneity is lost in the grid cell-averaging process, which is especially critical in urban areas. Next generation microscale models will potentially solve this problem (Lateb et al., 2016). Besides the dilution of the emission in the grid, meteorology may also play an important role in the low performance of  $\text{NO}_2$  and  $\text{O}_3$  in hotspot areas. Several inter-comparison studies (e.g. EURODELTA and AQMEII) agree on the limitations of models to simulate meteorological variables that affect the average hourly  $\text{NO}_2$  temporal variability, which controls model performance for  $\text{O}_3$  in high- $\text{NO}_x$  environments and their downwind areas (Bessagnet et al., 2016).

Figure 5 classifies the average hourly and MDA8  $\text{O}_3$  concentrations at the air quality stations into four MB categories that account for 93 % of the stations. The best performance for  $\text{O}_3$  (type B) is found at the 28 % of stations located in the areas surrounding the MMA, the BMA and most of the northern Mediterranean stations, which is consistent with the highest  $r$  ( $0.6 < r < 0.9$ ) found in the centre and north of the IP (Fig. S4 in the Supplement). The highest  $\text{O}_3$  overestimations (type D) are present at the 36 % of stations that are mainly located in highly industrialized areas in Spain (Guadalquivir Valley, Strait of Gibraltar, Valencia) and inside the MMA. The next sections analyse the origin of these  $\text{O}_3$  biases using the source apportionment time series.

The comparison with previous CALIOPE studies (Baldasano et al., 2011; Pay et al., 2014) indicates that  $r$  is in the same range for  $\text{O}_3$  (0.6–0.7) and  $\text{NO}_2$  (0.4–0.5) at individual stations; the same applies for the RMSE (15–29 and 10–20  $\mu\text{g O}_3 \text{ m}^{-3}$ ). Modelled  $\text{O}_3$  shows better performance at traffic stations in large cities, as stations influenced by road transport emissions (i.e. high- $\text{NO}_x$  environments) are better characterized with a more pronounced daily variability (Baldasano et al., 2011). At the European scale, several model inter-comparisons (Giornado et al., 2015; Bessagnet et al., 2016) have indicated that  $\text{O}_3$  concentrations in summer agree with the surface observations with  $r$  values between 0.5 and 0.6.  $\text{NO}_2$  hourly variability is underestimated overall due to uncertainties in the emission estimates, meteorological inputs and model resolution. These studies highlight the limitations of models with respect to simulating meteorological variables that affect the  $\text{NO}_2$  hourly variability, and therefore the model performance for  $\text{O}_3$  in high- $\text{NO}_x$  environments and their downwind areas.

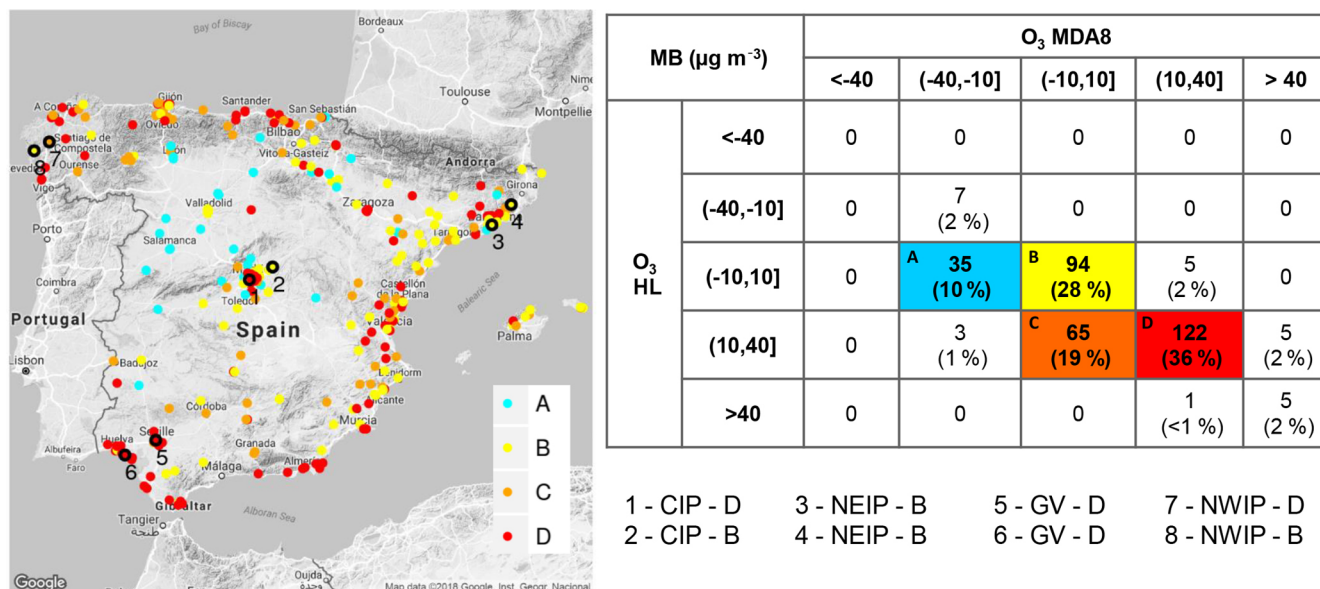
Section S4 in the Supplement discusses the meteorological evaluation results and their impact on pollutant concentrations. Not surprisingly, temperature shows the best behaviour when compared with observations (Table S2 in the Supplement). The modelled wind speed is overestimated, particularly during nighttime (Fig. S5 in the Supplement), coincident with low-level wind speed. The nighttime overestimation of wind is a source of error in modelled  $\text{NO}_2$  and  $\text{O}_3$  nighttime concentrations (Vautard et al., 2012; Bessagnet et al., 2016).

### 3.3 Source-sector ozone contributions during peak episodes

Figure 6 shows the p90 of the average hourly  $\text{O}_3$  concentration over the IP tagged by source type (Table 1) for different days (25, 28 and 31 July). (Fig. S6 in the Supplement shows similar plots for  $\text{NO}_2$ .) The imported  $\text{O}_3$  is by far the largest contributor showing a p90 ranging from 70 to 120  $\mu\text{g m}^{-3}$  in the east/north/centre of the IP on 25/28/31 July, respectively. The imported  $\text{O}_3$  enters the study domain through the IP4 domain boundaries and it can only be transported, scavenged, deposited or depleted by  $\text{O}_3$  precursors. Therefore, areas with low imported  $\text{O}_3$  concentrations ( $< 50 \mu\text{g m}^{-3}$ ) are good indicators of (1) the accumulation of specific  $\text{O}_3$  precursors that deplete imported  $\text{O}_3$ , and (2) the subsequent  $\text{O}_3$  photochemical production that occur mostly under stagnant conditions and around the largest industrial/urban areas. The p90 of the hourly imported  $\text{O}_3$  concentration shows the lowest values under two different conditions and in regions. On 25 July in the northwestern IP and Portugal, stagnant conditions allow the accumulation of pollutants that titrate the imported  $\text{O}_3$  concentrations down to 30–70  $\mu\text{g m}^{-3}$ . At the same time  $\text{O}_3$  is locally produced downwind of major northern cities due to traffic emissions (La Coruña, Gijón, Bilbao) (60–120  $\mu\text{g m}^{-3}$ ), shipping activities (up to 40  $\mu\text{g m}^{-3}$ ) and the generation of energy and industrial processes (10–20  $\mu\text{g m}^{-3}$ ). On 31 July in the northeast of the IP, the pollutants transported from the Gulf of Lion and Catalonia towards the Mediterranean act as a sink of imported  $\text{O}_3$  reducing its concentration down to 60  $\mu\text{g m}^{-3}$ . As a result, there is local  $\text{O}_3$  formation up to 120–160  $\mu\text{g m}^{-3}$  along the Ebro Valley and the Lleida Plain. In a source attribution study over northern Portugal, Borrego et al. (2016) also found a reduction of imported  $\text{O}_3$  and subsequent  $\text{O}_3$  formation by local sources under similar meteorological conditions.

Following imported  $\text{O}_3$ , the largest contributor to  $\text{O}_3$  is the road transport sector. Downwind of major urban areas in Spain (i.e. Madrid, Barcelona, Bilbao, Seville and Valencia), on-road traffic contributed as much as 60–120  $\mu\text{g m}^{-3}$  to the p90 of the hourly  $\text{O}_3$  concentrations, and affected different areas depending on the synoptic/mesoscale regimes (Fig. 6). In the north of the BMA, the p90 of the hourly  $\text{O}_3$  concentration from the road transport sector reaches its maximum when stagnant conditions affect the centre and eastern IP





**Figure 5.** Air quality stations classified by both mean bias (MB, in  $\mu\text{g m}^{-3}$ ) for average hourly and MDA8 O<sub>3</sub> at the Spanish EIONET stations and lumped by category (A, B, C and D). Numbered black circles indicate the stations under study in central IP (CIP; stations 1 and 2), northeastern IP (NEIP; station 3 and 4), Guadalquivir Valley (GV; stations 5 and 6) and northwestern IP (NWIP; stations 7 and 8).

(e.g. 31 July). As noted above, mesoscale winds carry traffic O<sub>3</sub> precursors from the BMA inland, channelled by north–south valleys towards the intra-mountain plain in the north. Over the MMA, the p90 of the hourly O<sub>3</sub> concentration from the road transport sector showed a maximum when the ITL was combined with the synoptic westerlies (e.g. 31 July), carrying high O<sub>3</sub> as far as the Ebro Valley, as shown in Fig. 4.

Regarding the contribution from the non-road transport sector, the Atlantic regions of the IP show the highest p90 of the average hourly O<sub>3</sub> concentration ( $25\text{--}40\mu\text{g m}^{-3}$ ) on 25 July. The stagnant conditions favoured the accumulation of precursors from the Atlantic shipping route and the formation of O<sub>3</sub> within the region. The Spanish Mediterranean region shows the highest p90 of the average hourly O<sub>3</sub> concentrations from the non-road transport sector close to the south-eastern coasts of the IP ( $\sim 180\mu\text{g m}^{-3}$ ) when the westerlies in the Strait of Gibraltar inject precursors from international shipping into the Mediterranean Basin (e.g. 28 and 31 July). Note that during days with high p90 of the average hourly O<sub>3</sub> concentration from non-road transport, the imported O<sub>3</sub> concentration shows the lowest p90 due to the NO titration effect over emission areas.

The elevated point source emission sectors (i.e. energy and industry) contributed less to O<sub>3</sub> than the traffic sector, but their contributions were significant reaching  $15\text{--}25\mu\text{g m}^{-3}$  of the p90 of the average hourly O<sub>3</sub> concentrations (Fig. 6). The north and northeast of the IP, the Mediterranean coast and the Guadalquivir Valley are the most affected regions under stagnant conditions.

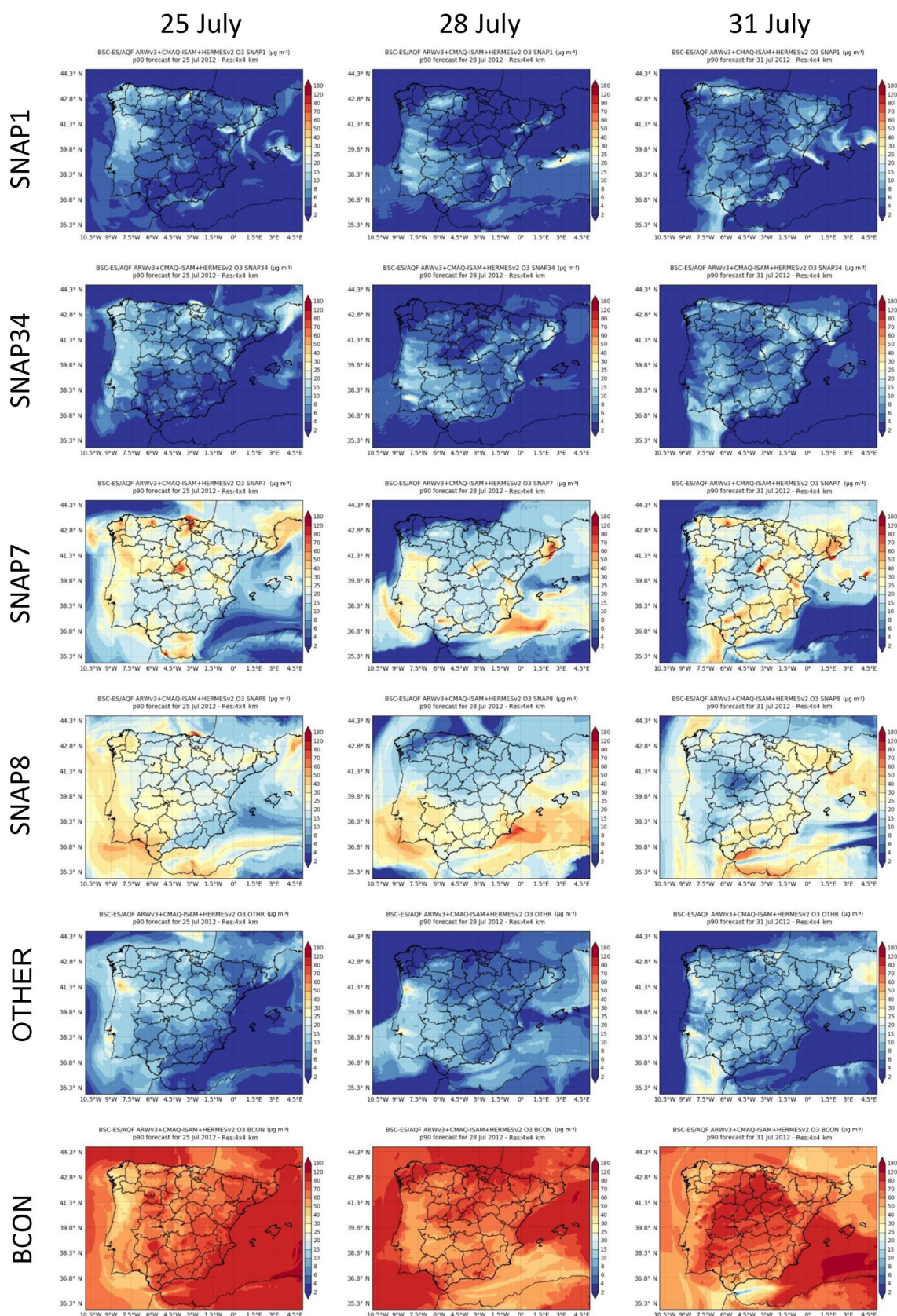
The contribution of the remaining sectors (OTHER) to the p90 of the average hourly O<sub>3</sub> concentrations was simi-

lar to that of the elevated point sources ( $15\text{--}25\mu\text{g m}^{-3}$ ), but it reached up to  $30\mu\text{g m}^{-3}$  in areas downwind of Oporto and Lisbon (Fig. 6). OTHER includes the formation of O<sub>3</sub> from the remaining anthropogenic and biogenic sources (accounting for less than 8 % of total NO<sub>x</sub> emissions, but 93 % of total VOC). The high OTHER concentration around the biggest cities in Portugal may be related to precursors emitted by the residential sector (SNAP2 and SNAP9) and biogenic emissions, as found in other source apportionment studies over Portugal (Borrego et al., 2016; Karamchandani et al., 2017).

### 3.4 Regionalization of source-sector contributions

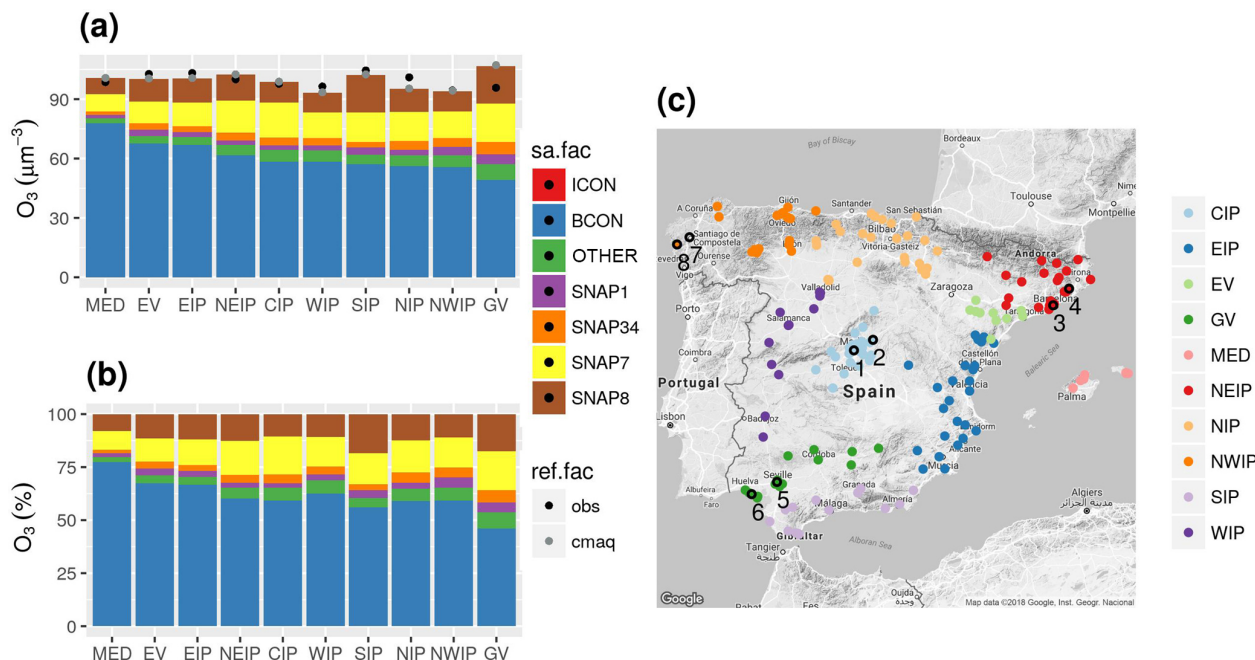
We have identified 10 O<sub>3</sub> receptor regions with similar characteristics in terms of meteorological and geographical patterns, O<sub>3</sub> dynamics and main source contributors (Figs. 4 and 6). The receptor regions defined in our work are consistent with Diéguez et al. (2014) and Querol et al. (2016), who proposed a similar regionalization based on observations from air quality stations. Figure 7c shows the location of the air quality stations belonging to each receptor region corresponding to the centre of the IP (CIP), the east of the IP (EIP), the Ebro Valley (EV), the Guadalquivir Valley (GV), the Mediterranean Sea (MED), the northeast of the IP (NEIP), the north of the IP (NIP), the northwest of the IP (NWIP), the south of the IP (SIP) and the west of the IP (WIP).

Figure 7a shows the absolute O<sub>3</sub> contribution of each tagged source at air quality stations by region along with the modelled and observed daily mean concentration during exceedances of  $120\mu\text{g m}^{-3}$  of the observed MDA8 ozone. Note



**Figure 6.** Tagged O<sub>3</sub> concentrations (in  $\mu\text{g m}^{-3}$ ) corresponding to the 90th percentile (p90) of the average hourly concentrations: SNAP1, SNAP34, SNAP7, SNAP8, OTHER and BCON for 25 July (first column), 28 July (second column) and 31 July (third column) in 2012.





**Figure 7.** Daily mean contribution in  $\mu\text{g m}^{-3}$  (a) and in percentage (b) of tagged sources of O<sub>3</sub> during exceedances of the observed  $120 \mu\text{g m}^{-3}$  for MDA8 O<sub>3</sub> averaged by the identified receptor regions (c). Black and grey dots represent observed and modelled daily mean concentrations during exceedances of  $120 \mu\text{g m}^{-3}$  of the observed MDA8 O<sub>3</sub>. Regions correspond to the centre of the IP (CIP), the east of the IP (EIP), the Ebro Valley (EV), the Guadalquivir Valley (GV), the Mediterranean Sea (MED), the northeast of the IP (NEIP), the north of the IP (NIP), the northwest of the IP (NWIP), the south of the IP (SIP) and the west of the IP (WIP). Numbered black circles indicate the stations under study CI (1–2), NEIP (3–4), GV (5–6) and NWIP (7–8).

that differences between sectors are more evident when normalizing (Fig. 7b). (Table S3 in the Supplement compiles the numerical values of Fig. 7.) Figure 7 indicates that during exceedances of the MDA8 target value there is a good agreement ( $r = 0.79$ ) between the sum of apportioned O<sub>3</sub> and the observed concentrations over the receptor regions.

The MED region shows the highest imported O<sub>3</sub> contribution (76 %) because it is relatively far from important anthropogenic NO<sub>x</sub> + VOC sources in the IP. Under the ITL influence (25 and 31 July), MED received air masses enriched with on-road traffic precursors from southern France and the NEIP, which enhanced O<sub>3</sub> production up to 7 %. Shipping emissions in the MED region contributed up to 8 % of the total O<sub>3</sub>.

After MED, there is a cluster of regions along the Spanish Mediterranean coast (i.e. NEIP, EV and EIP) that show imported O<sub>3</sub> contributions between 60 % and 68 % of the daily mean O<sub>3</sub> under exceedances. This is explained by their proximity to the eastern boundary and the frequent mesoscale phenomena enhancing the recirculation and accumulation of imported O<sub>3</sub> along the Spanish Mediterranean coast. The contribution of road and non-road transport is similar ( $\sim 11$  %–16 %) because these regions have both important roads and maritime trade routes. Note that the SIP region, which is also located on the Spanish Mediterranean coast, shows a daily mean imported O<sub>3</sub> concentration lower

than other regions along the Spanish Mediterranean coast ( $\sim 57$  %) and the highest non-road transport contribution in Spain (19 %). The main sink of imported O<sub>3</sub> are precursors resulting from dense shipping traffic through the Strait of Gibraltar, which have a substantial impact on the O<sub>3</sub> production downwind (either in the Alboran Sea or the Gulf of Cadiz).

The regions including the largest metropolitan areas in Spain are the CIP (Madrid) and the NEIP (Barcelona). Both regions show an imported O<sub>3</sub> contribution of  $\sim 60$  % and a similar contribution from the road transport sector (18 % and 16 %, respectively). However, the NEIP shows a slightly higher contribution from non-road transport (13 vs. 10 %) due to the influence of international shipping near coastal areas.

The northern and northwestern regions of the IP (NIP and NWIP) had relatively lower imported O<sub>3</sub> contributions (56 %–59 %). The contribution of non-road transport was  $\sim 10$  %–12 %, slightly lower than in the Mediterranean coast, and that of road transport was also significant ( $\sim 14$  %–15 %). The contribution from the industrial sector was one of the highest in the country ( $\sim 5$  %) which was probably related to the influence of the large industrial facilities located in several areas of the north of Spain. The contribution from the energy sector in the NWIP region was the highest in

Spain ( $\sim 5\%$ ) due to emissions from large coal-fired power plants located in the area.

The WIP had the lowest daily mean  $\text{O}_3$  concentrations during days exceeding the  $\text{O}_3$  target value ( $93.5\text{ }\mu\text{g m}^{-3}$ ) and a high imported  $\text{O}_3$  contribution ( $\sim 63\%$ ).  $\text{NO}_x$  emissions in the WIP region are moderate (Fig. 4), which could explain the low daily mean  $\text{O}_3$  concentration. There is a significant contribution from traffic (14 % for road transport and 11 % for non-road transport) and industrial and energetic sectors (7 %) to the daily mean  $\text{O}_3$  concentrations. These anthropogenic contributions suggest that  $\text{O}_3$  in the WIP may be produced by precursors transported from the surrounding cities (Porto, Lisbon and Madrid) and the highly industrialized areas in the NWIP and the NIP (Fig. 6).

The Guadalquivir Valley had the lowest imported  $\text{O}_3$  contribution in the IP ( $\sim 46\%$ ) and the highest daily  $\text{O}_3$  concentration during days of exceedance. The on-road traffic was the highest anthropogenic contributor to  $\text{O}_3$  ( $\sim 18\%$ ) due to the emissions from three major cities (Seville, Huelva and Cordoba). Although  $\text{O}_3$  in Huelva may be overestimated (as discussed later), shipping is the second most important contributor to  $\text{O}_3$  in the Guadalquivir Valley ( $\sim 17\%$ ) which is probably linked to the important fluvial transport along the river. (The Guadalquivir River is one of the most important routes for merchandise transport in Europe.) In fact, the non-road transport sector is the highest contributor ( $\sim 17\%$ – $19\%$ ) in southern Spain, both in the Guadalquivir Valley and the SIP, which is also due to the dense maritime routes across the Strait of Gibraltar.

The following sections analyse the source apportionment results at regions with a high on-road traffic contribution (i.e. CIP and NEIP) and a high contribution from industry and energy production (i.e. NWIP and Guadalquivir Valley).

### 3.4.1 The centre of the Iberian Peninsula

Figure 8 shows the source apportionment time series of the average hourly  $\text{O}_3$  and  $\text{NO}_2$  concentrations at two stations, an urban station in Madrid (station 1 in Fig. 8a), and another station located in Guadalajara (station 2 in Fig. 8b), which is a medium size city affected by Madrid's urban plume. At the urban station, the model reproduces the  $\text{O}_3$  traffic cycle ( $r = 0.66$  and  $\text{MB} = 22.5\text{ }\mu\text{g m}^{-3}$ ) featuring the typical low  $\text{O}_3$  concentrations ( $< 40\text{ }\mu\text{g m}^{-3}$ ) in the early morning and in the afternoon due to  $\text{O}_3$  titration (Fig. 8a). However,  $\text{O}_3$  was overestimated (MB type D) during daytime peaks due to the overestimation of the  $\text{NO}_2$  morning peaks during stagnant conditions, coincident with the highest road transport contribution for both pollutants. The results point towards a poor representation of the meteorological condition in the city during stagnant conditions as shown in the meteorological evaluation (Sect. 4 in the Supplement).

At the urban station downwind (Fig. 8b), modelled  $\text{O}_3$  is positively biased (MB type B) due to the underestimation of  $\text{NO}_2$ . Note that the uncertainty in  $\text{NO}_2$  traffic emissions in

medium-sized cities is larger than in the largest urban areas (i.e. Madrid and Barcelona) because data are generally unavailable and emissions are estimated based on population density (Baldasano et al., 2011).

The imported  $\text{O}_3$  is the main contributor at both stations, but  $\text{O}_3$  formation due to traffic increases significantly during peaks at both stations. The highest  $\text{O}_3$  concentrations ( $\sim 160\text{ }\mu\text{g m}^{-3}$ ) are modelled when westerly winds channelled along the Tajo Valley carry the polluted air masses in a northeasterly direction. This results in an  $\text{O}_3$  contribution of  $\sim 70\text{ }\mu\text{g m}^{-3}$  from the road transport sector in downwind areas (see wind vectors in Fig. 8b on 28 and 31 July). The  $\text{O}_3$  contribution from the industrial sector (whose precursors could come from facilities in the south of the MMA, Fig. 4) reinforces the  $\text{O}_3$  peaks up to  $\sim 10\text{ }\mu\text{g m}^{-3}$ , whereas the contribution from non-road transport systematically increases the background  $\text{O}_3$  concentration by  $\sim 15\text{ }\mu\text{g m}^{-3}$ . The  $\text{O}_3$  contribution from non-road transport in this region may arise mainly from Madrid's airports and the agricultural machinery operating in the surrounding rural areas.

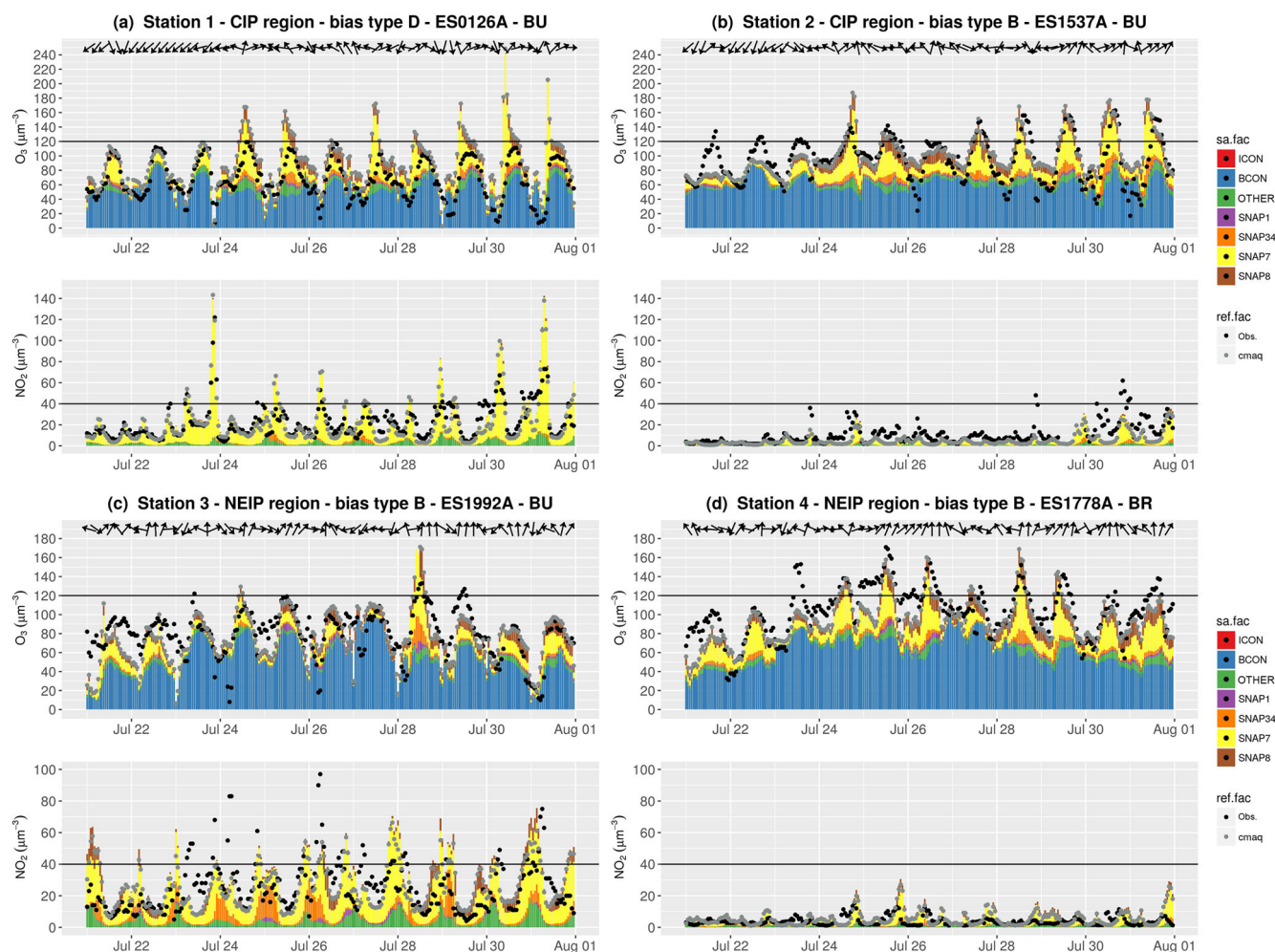
### 3.4.2 The northeast of the Iberian Peninsula

Figure 8c and d show the source apportionment time series at two stations in the NEIP, an urban station in Barcelona (station 3) and a remote rural downwind area (station 4). Not surprisingly, at the urban station,  $\text{NO}_2$  levels of up to  $100\text{ }\mu\text{g m}^{-3}$  affect  $\text{O}_3$  concentrations by titration during traffic peaks. In contrast, the rural station downwind depicts a higher  $\text{O}_3$  and lower  $\text{NO}_2$  concentration than the urban station. Absolute  $\text{O}_3$  biases at both stations are  $\sim 10\text{ }\mu\text{g m}^{-3}$  (MB type B).

$\text{O}_3$  mostly results from import and from the NO titration effect due to local road transport and industrial sources (Fig. 8c). However, the  $\text{O}_3$  diurnal cycle in the urban areas of the NEIP is less marked than in the CIP due to the persistently high  $\text{O}_3$  concentration at night ( $\sim 60\text{ }\mu\text{g m}^{-3}$ ). The breezes and mountain–valley winds contribute to the accumulation and recirculation of pollutants in this region.

At the rural station, modelled  $\text{O}_3$  peaks ( $> 120\text{ }\mu\text{g m}^{-3}$ ) are in a good agreement with observations (Fig. 8d), which suggests that the model generally reproduces the main transport paths, photochemical processes and relative contributions from different sources. Imported  $\text{O}_3$  is one of the main contributors to ground-level  $\text{O}_3$  (from 40 to  $100\text{ }\mu\text{g m}^{-3}$ ), but during peaks the on-road traffic contribution sharply increases up to  $80\text{ }\mu\text{g m}^{-3}$ .

The  $\text{O}_3$  concentration from the road transport sector arriving at rural areas in the NEIP, mainly originates from Barcelona and its surroundings as a result of the afternoon sea breezes (see wind vectors in Fig. 8c and d). However, under specific meteorological patterns these winds also carry precursors from other cities located in the northwestern Mediterranean Basin. Other authors (Gangoiti et al., 2001) have hypothesized that the high  $\text{O}_3$  concentration in the



**Figure 8.** Source apportionment time series for  $\text{O}_3$  and  $\text{NO}_2$  concentrations (in  $\mu\text{g m}^{-3}$ ) in the episode at the selected stations in the centre of the IP (CIP) region (a, b), and in the northeast of the IP (NEIP) (c, d). Colour bars (sa.fac) indicate the  $\text{O}_3$  tags, and black and grey dots (ref.fac) indicate the respective observed and modelled concentrations. Black horizontal lines represent the  $\text{O}_3$  target value ( $120 \mu\text{g m}^{-3}$ ) and the  $\text{NO}_2$  limit value ( $40 \mu\text{g m}^{-3}$ ) as a reference. The locations of the stations are shown in Fig. 7 using the corresponding numbers.

western Mediterranean Basin is influenced by transport from France via the Carcassonne gap. The present experiment cannot quantify the contribution of French cities to the  $\text{O}_3$  concentration over the NEIP, but future studies could explicitly tag the emission from the French regions.

### 3.4.3 Guadalquivir Valley

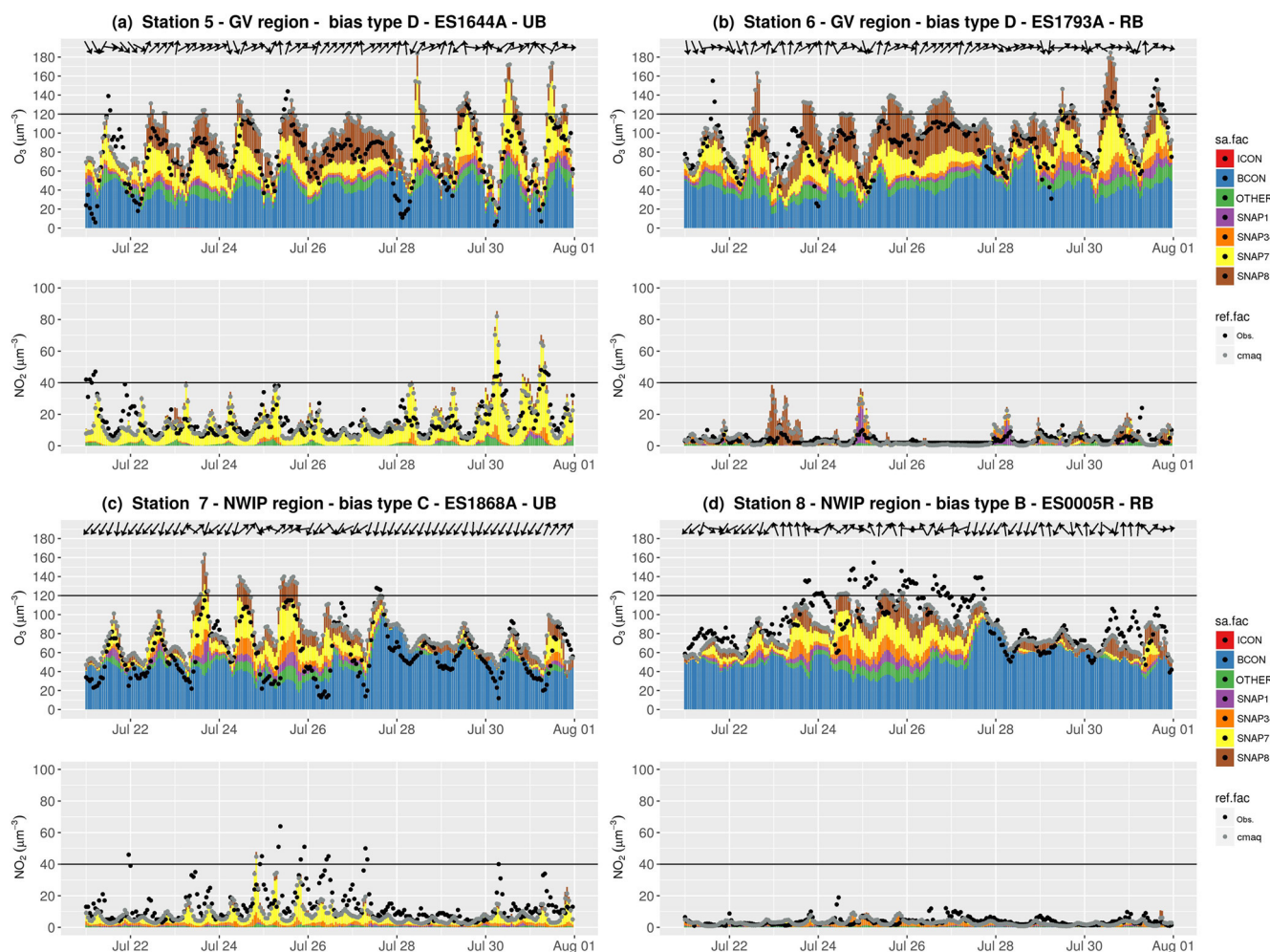
We have selected two stations along the Guadalquivir Valley, one in the urban area of Seville (station 5 in Fig. 9a), and one in a rural coastal area (station 6 in Fig. 9b). The contribution of non-road transport is due the influence of one of the largest Spanish harbours. The contribution of the energy sector to the  $\text{O}_3$  concentration is also noticed (e.g. 2 July). As expected, the urban station (Fig. 9a) shows a high  $\text{NO}_2$  concentration dominated by on-road traffic.  $\text{NO}_x$  from traffic is the main sink of  $\text{O}_3$  in the city and the model reproduces

the titration effect in agreement with observations (e.g. 28–31 July).

High  $\text{O}_3$  overestimations ( $10\text{--}30 \mu\text{g m}^{-3}$ ) at both stations (Fig. 9a and b) were detected during the 25–28 July period which correspond to intense and persistent southwesterly winds transporting air masses from the Atlantic Sea along the Guadalquivir Valley, as shown by the wind vectors in Fig. 9a and b. Although the model overestimates  $\text{O}_3$  concentrations, it reproduces the temporal variability.

Our results suggest that the non-road transport sector is a significant contributor along the Guadalquivir Valley during these days. The impact of shipping emission on  $\text{O}_3$  in the Guadalquivir Valley region is mainly evidenced by the relative high  $\text{NO}_x$  from ship exhaust (Fig. S3 in the Supplement). The  $\text{NO}_2$  time series at the coastal station (Fig. 9b) indicates that the model overestimates  $\text{NO}_2$  concentrations during hours when the  $\text{NO}_2$  contribution from non-road trans-





**Figure 9.** Source apportionment time series for  $\text{O}_3$  and  $\text{NO}_2$  concentrations (in  $\mu\text{g m}^{-3}$ ) in the episode at the selected stations in the Guadalquivir Valley (GV) (a, b), and in the northwest of the IP (NWIP) (c, d). Colour bars (sa.fac) indicate the  $\text{O}_3$  tags, and black and grey dots (ref.fac) indicate the respective observed and modelled concentrations. Black horizontal lines represent the  $\text{O}_3$  target value ( $120 \mu\text{g m}^{-3}$ ) and the  $\text{NO}_2$  limit value ( $40 \mu\text{g m}^{-3}$ ) as a reference. The locations of the stations are shown in Fig. 7 using the corresponding numbers.

port is highest. The unrealistic  $\text{NO}_2$  peaks from non-road transport suggest that shipping emissions are overestimated in the HERMESv2.0 model, which uses the EMEP gridded emission inventory at  $50 \text{ km} \times 50 \text{ km}$  horizontal resolution to estimate shipping emissions and spatially distributes them to the  $4 \times 4 \text{ km}^2$  IP4 domain using the marine routes reported by Wang et al. (2008).

A recent review on the state-of-the-art of marine traffic emissions (Russo et al., 2018) indicates that STEAM appears as the most reliable and detailed emissions inventory, as it is based on Automatic Identification System data and specific vessel information, with a resolution of  $2.5 \times 2.5 \text{ km}^2$  (Jalkanen et al., 2016). A comparative analysis indicates that EMEP gridded inventories are overestimated, in particular over hotspots on the Mediterranean shipping routes, and underestimated on secondary routes. This factor, added to the 15 % decrease of  $\text{NO}_x$  shipping emissions observed in

Europe between 2009 (HERMESv2.0 base year) and 2012 (EMEP CEIP, 2019) can at least partly explain the discrepancies observed.

#### 3.4.4 The northwest of the Iberian Peninsula

Figure 9c and d show the source apportionment time series at one urban (station 7) and one rural (station 8) background station in the NWIP. The urban station (Fig. 9c) located in Santiago de Compostela, a medium size city with  $\sim 100\,000$  inhabitants, shows a high  $\text{NO}_2$  concentration with a dominant contribution from the road transport sector. Traffic  $\text{NO}_2$  is the main sink of urban  $\text{O}_3$  via titration. Because  $\text{NO}_2$  is underestimated, especially during stagnant conditions (24–27 July),  $\text{O}_3$  concentrations are overestimated (MB type C).

Despite the  $\text{O}_3$  biases during stagnant conditions, the modelled  $\text{O}_3$  concentration is in general agreement with observa-

tions at the rural background station (Fig. 9d).  $\text{NO}_2$  is likely to be underestimated due to missing traffic emissions. As previously noted, traffic emissions are poorly constrained in small and medium-sized cities, due to a lack of detailed information. There is also additional uncertainty in the precursors emitted from the large coal power plants and industries in the region (Valverde et al., 2016b). Our study uses emissions for 2009, and it has been estimated that between 2009 and 2012 energy production in coal-fired power plants increased from 13.1 % to 19.4 % (IEA, 2019). This implies an increase in  $\text{NO}_x$  emissions from the power industry sector of around 19.5 % (EMEP CEIP, 2019).

The time series show that the model reproduces the observed  $\text{O}_3$  variability reasonably well under different synoptic conditions.  $\text{O}_3$  reaches its highest concentrations ( $\sim 100/150 \mu\text{g m}^{-3}$  in urban/rural areas) under stagnant conditions (24–27 July) when the contribution of anthropogenic sources from all activity sectors is highest (60 %–70 %).  $\text{O}_3$  concentrations decrease down to  $\sim 70 \mu\text{g m}^{-3}$  under northwest advective conditions (e.g. 28–30 July) when the imported  $\text{O}_3$  shows the highest contribution (80 %–90 %). Saavedra et al. (2012) found that stationary anticyclones over the NWIP play an important role in the occurrence of high  $\text{O}_3$  concentrations. Our results show that under these stagnant conditions  $\text{O}_3$  concentrations are largely due to in situ production (photochemistry) from on-road traffic, shipping, power plants and industry in almost the same proportion.

### 3.5 Imported ozone

Our results indicate that imported  $\text{O}_3$  represents the highest contribution to the ground-level  $\text{O}_3$  concentration in southwestern Europe. Imported  $\text{O}_3$  enters the IP4 domain through the boundaries; it includes the contribution of  $\text{O}_3$  from the EU12 domain, which in turn includes the contribution of hemispheric  $\text{O}_3$  from the MOZART-4 global model. The imported  $\text{O}_3$  contribution is as large as the background  $\text{O}_3$  regionally produced within the IP. Note that the small biases at rural background stations obtained in the evaluation section indicate an overall high performance of the modelled background  $\text{O}_3$  in the IP4 domain (Fig. S1 in the Supplement). Given the important implications and robustness of these results, we further analyse this contribution below. In particular, we aim to understand its high contribution within the IP, even far from the model domain boundaries.

Figure 10 shows the vertical cross-sections at 06:00, 12:00 and 18:00 UTC for  $\text{O}_3$  and  $\text{NO}_2$  at a constant latitude ( $40.38^\circ \text{N}$ ) on 25, 28 and 30 July. It helps to understand the vertical variability of both pollutants according to the PBL as schematized by Millán et al. (1996). The model predicts a pronounced  $\text{O}_3$  vertical gradient above a height of 4 km above sea level (a.s.l.), showing that  $\text{O}_3$  in the free troposphere is, to a large extent, imported to the IP4 domain (Safieddine et al., 2014). As expected,  $\text{NO}_2$  mixing ratios show a negative gradient with altitude as it is mainly emit-

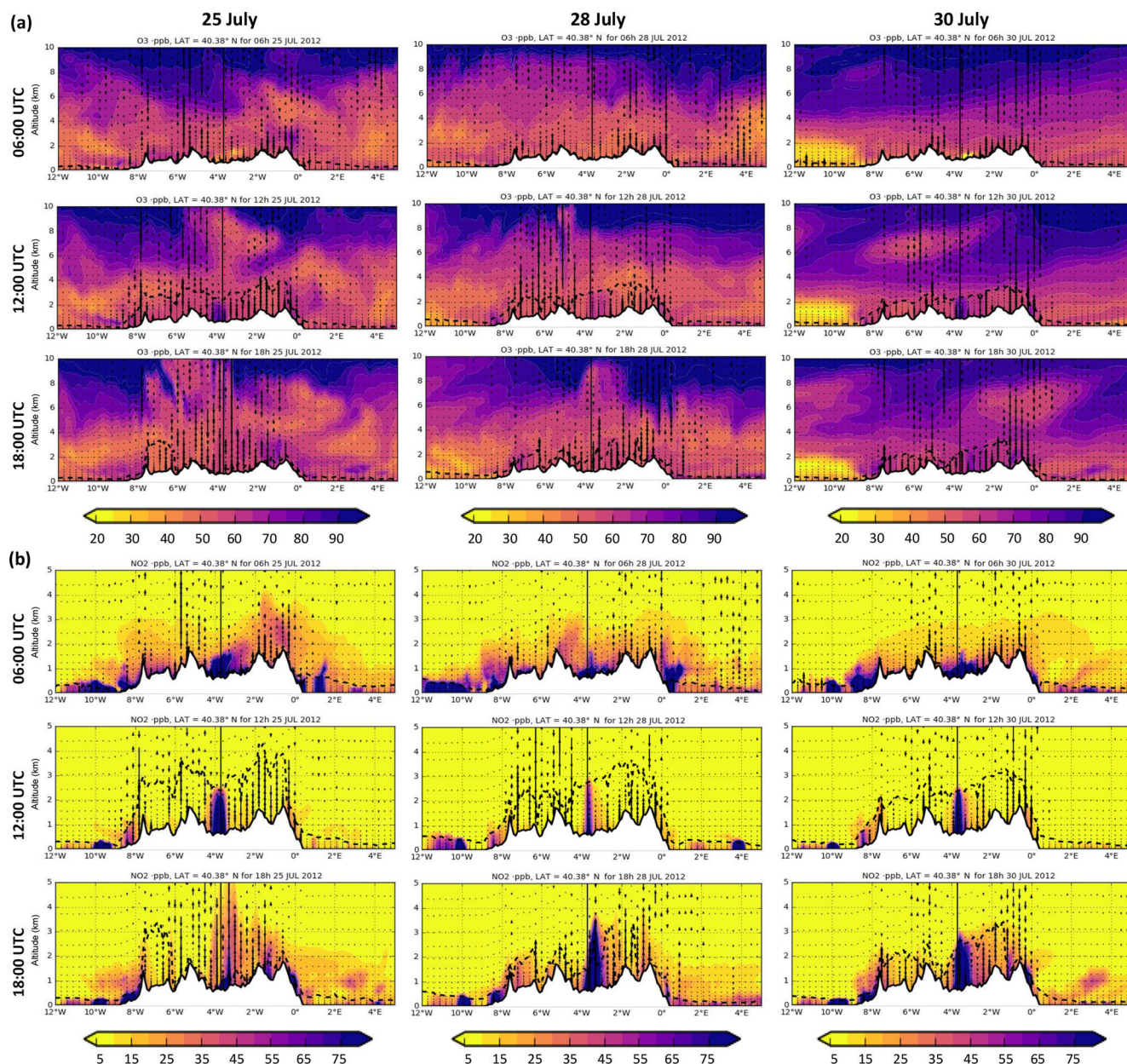
ted at the surface. In the morning, the sun starts to heat the ground up, producing convective thermals and forcing the growth of the mixing layer. At noon, the mixing height reaches its maximum: it is highest in the CIP (2–4 km) and decreases towards the coast ( $< 1 \text{ km}$ ) (Fig. 10). At the top of the mixing layer the  $\text{O}_3$ -enriched air aloft is entrained into the mixing layer, mixing with  $\text{O}_3$  and other pollutants produced locally within the mixing layer. When the mixing height decreases,  $\text{O}_3$  is left in the free troposphere forming high  $\text{O}_3$  residual layers (Gangoiti et al., 2001) that contribute to the regional transport. Over the following days, these residual layers (composed of imported  $\text{O}_3$  and local  $\text{O}_3$  produced within the domain over previous days) can be entrained by fumigation into the mixing layer to reach the surface. This fumigation effect, previously described in the eastern USA (Zhang and Rao, 1999; Langford et al., 2015) and in the western Mediterranean Basin (Kalabokas et al., 2017; Querol et al., 2018), leads to a rapid increase in  $\text{O}_3$  concentrations at ground level. The accumulation and recirculation of air masses is intensified along the eastern Mediterranean coast (Millán et al., 1996, 2000; Gangoiti et al., 2001; Querol et al., 2017) by the action of the breezes and mountain–valley winds. Furthermore, the small deposition velocity of  $\text{O}_3$  over the sea and its high atmospheric lifetime in the free troposphere contributes to enrich the  $\text{O}_3$  background concentration (ca. several weeks, Monks et al., 2015; Seinfeld and Pandis, 2016).

The  $\text{O}_3$  fumigation effect was particularly intense from 30 to 31 July, when high  $\text{O}_3$  levels are found in the free troposphere compared with previous days (Fig. 10). The analysis of the  $\text{O}_3$  concentration map with the imported  $\text{O}_3$  contributions (Fig. 6) indicates that ground-based  $\text{O}_3$  is neither advected nor titrated; therefore, it can only result from vertical mixing. The high  $\text{O}_3$  mixing ratio in the free troposphere was mainly due to  $\text{O}_3$  advection entering the Atlantic boundary driven by westerlies (Fig. 3). We hypothesize that two events may have contributed to the increase of the  $\text{O}_3$  concentration by long-range transport: first, a low-pressure system in the British islands on 28 July could have transported significant amounts of  $\text{O}_3$  from the stratosphere to the free troposphere (Fig. S7 in the Supplement); second,  $\text{O}_3$  episodes generated in mid-July over the eastern USA predicted by the MOZART-4 model could have contributed to an increase in the  $\text{O}_3$  transported from North America to Europe by the action of the prevailing westerlies (Fig. S7 in the Supplement) associated with cyclonic systems along the “warm conveyor belt” (Pausata et al., 2012; Derwent et al., 2015).

## 4 Discussion and conclusions

Our study has provided a first estimation of the main sources responsible for high  $\text{O}_3$  concentrations in the western Mediterranean Basin during the 21–31 July 2012 period. We used the Integrated Source Apportionment Method (ISAM)





**Figure 10.** Cross section of modelled mixing ratios (in ppb) for O<sub>3</sub> (a) and NO<sub>2</sub> (b) at a constant latitude (latitude = 40.38°, Madrid city) of the daily Iberian thermal low circulation, equivalent to the conceptual scheme of Millán et al. (1996) for 25 July (first column), 28 July (second column) and 30 July (third column) at 06:00, 12:00 and 18:00 UTC. Dotted lines indicate the PBL height. Vertical arrows indicate the vertical wind. Up arrows depict positive winds. Note the different scales of the y axes between O<sub>3</sub> and NO<sub>2</sub>.

within the CALIOPE system to estimate the contribution of the main anthropogenic activity sectors to peak O<sub>3</sub> events in Spain compared with the imported O<sub>3</sub>. In addition, the use of ISAM has allowed an in-depth evaluation of the model.

The results demonstrate that the O<sub>3</sub> problem over the western Mediterranean Basin is local, regional and hemispheric. Long-range transport of O<sub>3</sub> from beyond the IP domain is the main contributor to the ground-level daily mean O<sub>3</sub> concentration (~ 45 %) during peak episodes. The imported O<sub>3</sub>

contribution ranges from 40 % during O<sub>3</sub> peaks to 80 % at night or during well-ventilated conditions. The absolute imported O<sub>3</sub> is higher in the northeast of the IP than in the central IP due to the recirculation and accumulation of pollutants along the Mediterranean. The high imported O<sub>3</sub> at the surface far away from the model boundaries is consistent with the high levels of O<sub>3</sub> in the free troposphere (resulting from local/regional layering and accumulation, and continen-

tal/hemispheric transport) along with intense vertical mixing during the day.

Our results support the European Commission (EC, 2004) in pointing out that the effectiveness of abatement strategies for achieving compliance with the European air quality standards in southern Europe might be compromised by the long-range transport of  $O_3$ . This is especially true in Mediterranean regions (i.e. NEIP, EV and EIP) where the contribution of imported  $O_3$  is particularly dominant (60 %–68 % of the daily mean  $O_3$  concentration) as a result of the accumulation and recirculation of pollutants over the Mediterranean Basin. In these areas, if the long-range transport of  $O_3$  is not reduced, the mean background level will not decrease, making it more vulnerable to exceedances of the  $O_3$  target values by enhanced local production under stagnant conditions.

During high  $O_3$  events, the imported  $O_3$  is added to the formation from local and regional anthropogenic sectors. Road transport is an important contributor to the  $O_3$  concentration in rural areas downwind of large cities in Spain; it contributed up to 16 %–18 % of the daily mean  $O_3$  concentration under exceedances of the target value for human health protection, and up to  $70 \mu\text{g m}^{-3}$  on an hourly basis downwind of Barcelona and Madrid.

The non-road transport sector (including international shipping, airports and agricultural machinery) is as significant as road transport inland (10 %–19 % of the daily mean  $O_3$  concentration during the peaks). There is a high influence of international shipping (13 %), affecting the coastal areas in the Mediterranean and the south of the IP (along the Strait of Gibraltar) with contributions of up to  $\sim 20$  and  $\sim 30 \mu\text{g m}^{-3}$ , respectively. Although the non-road transport contribution was found to be overestimated in coastal areas in the south of the IP in the present experiment, it cannot be neglected; furthermore, actions controlling international shipping should be considered as important as those related to road transport, especially in regions with big harbours (e.g. Huelva and Barcelona). Dalsøren et al. (2010) indicated that the annual  $O_3$  concentration is increasing annually between 1 and  $5 \mu\text{g m}^{-3}$  in areas impacted by shipping activities. Recent studies indicate that shipping emissions are projected to increase significantly due to increases in transportation demand and traffic. As the Strait of Gibraltar is the only shipping route connecting the Atlantic Ocean with the Mediterranean Basin, the regulation of these emissions is key in order to control  $O_3$  exceedances in Spain and the Mediterranean Basin. Shipping emissions can be regulated by each country within 400 km of coastlines, but policy-induced controls for offshore emissions are very dependent on the success of adopted and proposed regulations within the International Maritime Organization.

The energy and industrial sectors contribute  $\sim 6$  %–11 % of the daily mean  $O_3$  concentration during the peaks and over all the receptor regions. As they are usually injected at high altitudes, their contribution extends way beyond their surroundings. The energy combustion sector (3 %) and in-

dustrial and non-industrial combustion sectors (3 %) have a mean contribution of  $2$ – $4 \mu\text{g m}^{-3}$ , reaching to  $4$ – $6 \mu\text{g m}^{-3}$  in  $\text{NO}_x$ -limited areas (i.e. the western IP). In highly industrialized regions (i.e. Guadalquivir Basin and northwestern IP), abatement strategies affecting all sectors at a regional scale could contribute to decrease the local formation of  $O_3$  as the regional/local anthropogenic contribution can be greater than 50 % over several days.

In the Barcelona metropolitan area the contribution from energy and industrial sectors to the  $\text{NO}_2$  concentrations can be in the same range as the contribution from road transport ( $\sim 40$  %–60 %, Fig. 8c). In contrast, in areas downwind of Barcelona the contribution from energy and industrial sectors to  $O_3$  concentrations is relatively low compared with the contribution from road transport (Fig. 8d). The different contributions to the  $O_3$  concentration might be related to the different reactivity of VOCs for  $O_3$  formation. Each VOC emission source emits a different mix of VOCs, which contributes differently to photochemical ozone formation. For example, in the UK, Derwent et al. (2007) showed a higher photochemical  $O_3$  creation potential for road transport emissions than for production processes and combustion. Future national policy actions to control the emissions of VOCs should tackle the sources that contribute more to photochemical  $O_3$  formation.

The remaining sectors (i.e. SNAP 2, 5, 6, 9 and 11; see Table 1), are the fourth main contributors to the daily mean  $O_3$  concentration during the days exceeding the target value ( $\sim 2$  %–8 %). Future work should tag biogenic sources as an individual sector as they are the main contributor to VOC emissions in Spain (i.e.  $\sim 70$  % in 2009 according to the HERMESv2.0 model).

The air quality and meteorological evaluations indicate that uncertainties in our model are in the same range as the most recent inter-comparison studies using state-of-the-art air quality models. In addition, our model evaluation and the source apportionment results have allowed for a better understanding of the origin of model errors related to emission estimates. Our methodological choice was to use a detailed bottom-up emission inventory instead of a typical top-down regional emission inventory. Bottom-up emissions, estimated using source-specific emission factors and activity statistics, accurately characterize pollutant sources and allow for more realistic results to be obtained than those reported by top-down or regional emission inventories. To understand the impact of the use of 2009 data to study the year 2012, we revised the EMEP Centre on Emission Inventories and Projections (EMEP-CEIP), which collects and reviews the national emission inventories from parties to the Convention on Long-range Transboundary Air Pollution. Between 2009 and 2012, total  $\text{NO}_x$  and non-methane volatile organic compound (NMVOC) emissions in Spain decreased by  $-10.6$  % and  $-10.7$  %, respectively (EMEP CEIP, 2019). For  $\text{NO}_x$ , around 80 % of this reduction is linked to a reduction of road transport emissions, whereas in the case of NMVOCs  $\sim 50$  % of the reduction is due to a decrease in industrial emissions.

For our modelling study, we consider these differences as small and acceptable (and as not creating any major inconsistency). The difference of 10 %–15 % in emissions for certain precursors between 2009 and 2012 is within the typically larger ranges of uncertainty in emission inventories.

Another relevant and uncertain source of  $O_3$  is the VOC emitted in urban areas. Future research works should be devoted to the continuous monitoring of urban VOC and take advantage of satellite observations to improve speciation and spatial variability of urban VOC emissions.

We have identified two sources of uncertainty in the estimation of the imported  $O_3$ . First, it depends on both the performance of the CALIOPE system over EU12 and on MOZART-4 at a global scale. The small biases at rural background stations support an overall high performance baseline background  $O_3$  in the IP4 domain (Fig. S1 in the Supplement). Second, our set-up involves some uncertainty in the estimation of the imported contribution of  $O_3$ . In reality, there is a fraction of the imported  $O_3$  that may have been generated within the IP4 domain before the period of simulation (including the spin-up). We have assumed that this fraction is negligible and future works should check the extent to which this assumption is correct.

For regulatory applications, further source apportionment studies should target not only emissions from activity sectors, but also the source regions where the emission abatement strategies should be applied. In addition, future studies should preferentially cover multiple summer periods in order to improve representativeness. We note that our results cannot predict whether emission abatement will have either a positive or a negative effect in  $O_3$  changes due to the non-linearity of the  $O_3$  generation process. Subsequent source sensitivity analyses tailoring the identified main contribution sources could predict how  $O_3$  will respond to reductions in precursor emissions, which are essential to define the most efficient  $O_3$  abatement strategies in the western Mediterranean Basin.

Overall, we find that the imported  $O_3$  is the largest input to the ground-level  $O_3$  concentration in the IP during the episode studied. However, during stagnant conditions, the emission from local anthropogenic activities in the IP control the  $O_3$  peaks in areas downwind of the main urban and industrial regions. Furthermore, ground-level  $O_3$  concentrations are strongly affected by vertical downward mixing of  $O_3$ -rich layers in the free troposphere, which result from local/regional layering and accumulation, and continental/hemispheric transport. The importance of both imported and local contributions to the  $O_3$  peaks in the IP demonstrates the need for detailed quantification of both contributions to high  $O_3$  concentrations for local  $O_3$  management. Furthermore, the influence of local sources and topographical and meteorological conditions in the high  $O_3$  concentration indicate the importance of designing  $O_3$  abatement policies at the local scale.

This work has quantified the local and imported contributions to  $O_3$  during an episode in a particular area in southwestern Europe. In addition, we have provided a perspective regarding the potential use of source apportionment methods for regulatory studies in non-attainment regions. Further  $O_3$  source apportionment studies targeting other non-attainment regions in Europe are necessary prior to designing local mitigation measures that complement national and European-wide abatement efforts.

**Data availability.** Air quality measurements are available at the EIONET database (<https://www.eea.europa.eu/data-and-maps/data/aqereporting-8/T1/textbackslash#tab-figures-produced>, European Environmental Agency, Air Quality e-Reporting, last access 4 April 2019). The CMAQ code is available at <https://www.cmascenter.org/> (Community Modeling and Analysis System, last access 4 April, 2019). The WRF-ARW code is available at [http://www2.mmm.ucar.edu/wrf/users/download/get\\_source.html](http://www2.mmm.ucar.edu/wrf/users/download/get_source.html), (National Center for Atmospheric Research, last access: 4 April 2019). Download Mozart-4 outputs are available at <https://www2.acom.ucar.edu/gcm/mozart>, (National Center for Atmospheric Research, last access: 4 April 2019). The program that generates CMAQ boundary conditions from MOZART-4 output is available at <http://www.camx.com/getmedia/a0c2d710-6187-47bb-af3f-bd9b4a4d19bb/mozart2camx-12jul17>, (Ramboll Environ, last access: 4 April 2019). The HERMESv2 outputs are available upon request. The modelled data in this study are available upon request from the corresponding author.

## Appendix A

**Table A1.** Table 1a shows the acronyms used in the text.

BCON	Chemical boundary conditions to IP4, also referred to as the imported O <sub>3</sub> contribution
BMA	Barcelona metropolitan area
CIP	Central Iberian Peninsula
CMAQ	Community Multiscale Air Quality model
EIP	Eastern Iberian Peninsula
EU12	European domain at $12 \times 12$ km <sup>2</sup> horizontal resolution.
EV	Ebro Valley
GV	Guadalquivir Valley
IP	Iberian Peninsula
IP4	Iberian Peninsula domain at $4 \times 4$ km <sup>2</sup> horizontal resolution.
ISAM	Integrated Source Apportionment Method
MDA8	Maximum daily 8 h average concentration
MMA	Madrid metropolitan area
MED	Mediterranean Sea
NEIP	Northeast of the Iberian Peninsula
NIP	North of the Iberian Peninsula
NWIP	Northwest of the Iberian Peninsula
SIP	South of the Iberian Peninsula
SNAP1	Emission sector on combustion in energy
SNAP34	Emission sector on combustion and processes in industry
SNAP7	Emission sector on road transport, exhaust and non-exhaust
SNAP8	Emission sector on non-road transport (international shipping, airport and agricultural machinery)
WIP	West of the Iberian Peninsula
WRF	Weather Research and Forecasting model

## Appendix B

Definition of the discrete statistics used in the evaluation: correlation coefficient ( $r$ , Eq. B1), mean bias (MB, Eq. B2), normalized mean bias (NMB, Eq. B3) and root mean squared error (RMSE, Eq. B4). Where  $C_m(x, t)$  and  $C_o(x, t)$  are the respective modelled and observed concentrations at a location ( $x$ ) and time ( $t$ );  $N$  is the number of pairs of data.  $\overline{C_m}$  and  $\overline{C_o}$  are the modelled and observed mean concentrations over the whole period, respectively.

$$r = \frac{\sum_{i=1}^N (C_m(x, t) - \overline{C_m}) (C_o(x, t) - \overline{C_o})}{\sqrt{\sum_{i=1}^N (C_m(x, t) - \overline{C_m})^2} \sqrt{\sum_{i=1}^N (C_o(x, t) - \overline{C_o})^2}}, \quad (\text{B1})$$

$$\text{MB} = \frac{1}{N} \sum_{i=1}^N (C_m(x, t) - C_o(x, t)), \quad (\text{B2})$$

$$\begin{aligned} \text{NMB} &= \frac{\sum_{i=1}^N (C_m(x, t) - C_o(x, t))}{\sum_{i=1}^N (C_o(x, t))} \times 100 \\ &= \left( \frac{\overline{C_m}}{\overline{C_o}} - 1 \right) \times 100, \end{aligned} \quad (\text{B3})$$

$$\text{RMSE} = \frac{1}{N} \sqrt{\sum_{i=1}^N (C_m(x, t) - C_o(x, t))^2}. \quad (\text{B4})$$

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*Author contributions.* MTP designed the research. SN provided support regarding the ISAM set-up for this experiment. MTP performed the model simulations. All the authors analysed the data and discussed the results. MTP and CPGP wrote the paper.

*Competing interests.* The authors declare that they have no conflict of interest.

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