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Relative humidity and O_3 concentration as two prerequisites for sulfate formation

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Supplementary information

S1. Methodology for estimation of the mass concentrations of $PM_{2.5}$ components

S1.1 Organic matter

The mass concentration of organic matter (OM) was calculated from organic carbon (OC) measurements by multiplying OC by a factor that represents the mass contributions of other elements, such as oxygen, hydrogen, and nitrogen. The OM/OC ratio varies from 1.4 to 2.2 and is expected to increase as aerosols age (El-Zanan et al., 2005). We chose a factor of 1.6 to calculate OM in Beijing following advice in the literature (Xing et al., 2013).

S1.2 Minerals

The total mass concentration of minerals, referred to as "minerals", can be estimated by the following equation (Chan et al., 1997):

$$[\text{minerals}] = 2.2[\text{Al}] + 2.49[\text{Si}] + 1.63[\text{Ca}] + 2.42[\text{Fe}] + 1.94[\text{Ti}] \ , \quad \text{(Eq. 1)}$$

where [*x*] represents the mass concentration of species *x*. According to Zhang et al. (2003), on average Al accounted 7 % of total mineral dust mass concentrations in North, Northwest, and West China. Mineral concentrations can thus also be estimated by Eq. 2:

$$[\text{minerals}] = [\text{Al}]/0.07 \ , \quad \text{(Eq. 2)}$$

We calculated [minerals] with the two methods above and found no significant differences (Fig. S1). Equation 2 was therefore employed to calculate [minerals] in this study.

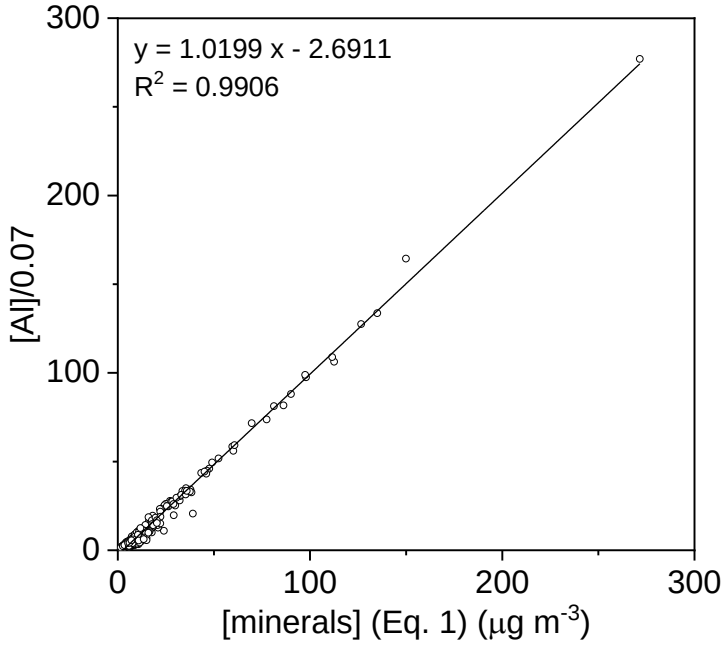

Figure S1. Comparison of the two methods for the calculation of [minerals].

S1.3 Trace element oxides

The enrichment factors (EFs) of trace element oxides (TEOs) can be used to determine whether natural or anthropogenic sources dominated our observations. The EF value of element *i* was defined as follows:

$$EF_i = \frac{[X_i/X_{ref}]_{sample}}{[X_i/X_{ref}]_{crust}} , \quad \text{(Eq. 3)}$$

where $[X_i/X_{ref}]_{sample}$ is the mass concentration ratio of element *i* to the reference element in our samples and $[X_i/X_{ref}]_{crust}$ is the mass concentration ratio of element *i* to the reference element in average crust (Hans Wedepohl, 1995). Al was used as the reference element in this study. The EFs of each element are depicted in Fig. S2.

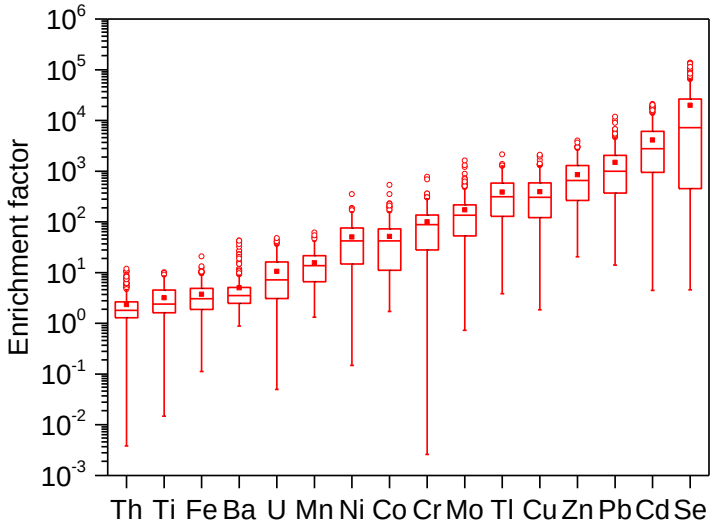

Figure S2. Elemental enrichment factors (EFs) of our samples. The boxes represent, from top to bottom, the 75th, 50th, and 25th percentiles for each element. The whiskers, solid red squares, and open red circles represent 1.5 times the interquartile range (IQR), mean values, and outlier data points, respectively.

If the EF was < 5, the element was considered to originate mainly from natural sources; if 5 < EF< 20, the element originated from both natural and anthropogenic sources; if EF > 20, the element originated mainly from anthropogenic sources. According to Zhang et al. (2013), the mass concentrations of TEOs can be estimated by multiplied a correction factor to represent the contribution of oxygen. For elements originating from anthropogenic sources only, a factor of 1 was applied, whereas for elements of both natural and anthropogenic origin, a factor of 0.5 was applied to represent the anthropogenic part. As multiple forms of metal oxides were identified, which were hard to quantify, a multiplicative factor of 1.3 was used when considering the metal abundance. The mass concentration of TEOs was calculated as described in Zhang et al. (2013):

$$[\mathrm{TEOs}] = 1.3 \times [0.5 \times (\mathrm{Ba} + \mathrm{Mn} + \mathrm{U}) + (\mathrm{Ni} + \mathrm{Co} + \mathrm{Cr} + \mathrm{Mo} + \mathrm{Tl} + \mathrm{Cu} + \mathrm{Zn} + \mathrm{Pb} + \mathrm{Cd} + \mathrm{Se})] \ , \quad \text{(Eq. 4)}$$

S1.4 Aerosol water content

Aerosol water content (AWC) was calculated using the ISORROPIA-II thermodynamic model (http://isorropia.eas.gatech.edu).

The Na^+–K^+–Ca^{2+}–Mg^{2+}–NH_4^+–SO_4^{2-}–NO_3^-–Cl^-–H_2O aerosol system was applied in reverse mode (Fountoukis and Nenes, 2007; Nenes et al., 1998).

S2 Results and discussion

S2.1 Sulfate formation mechanism

Sulfate can be formed through the oxidation of SO_2 by OH radicals in the gas phase (Stockwell and Calvert, 1983), through the oxidation of dissolved SO_2 by various oxidants (e.g., O_3, H_2O_2, NO_2, and O_2) in the aqueous phase (Seinfeld and Pandis, 2006), which may be transition metal ions (TMIs)-catalysed, or through heterogeneous reaction on the surface of sea-salt or dust aerosols (Gurciullo et al., 1999; Usher, 2002).

The rate of the SO_2 + OH reaction can be expressed as:

$$R_{SO_2+OH} = k_0[SO_2(g)][OH(g)] \ , \quad \text{(Eq. 5)}$$

where k_0 is the rate constant and [x] represents the concentration of species x. The production rate of sulfate through OH radical oxidation can be expressed as:

$$P_{OH} = \frac{3600 \times 96 \times p \times R_{SO_2+OH}}{RT} \ , \quad \text{(Eq. 6)}$$

where 3600 is a time conversion factor (s h^{-1}), 96 is the molar mass of SO_4^{2-} (g mol^{-1}), p is atmospheric pressure (kPa), R is the gas constant (8.31 Pa m^3 mol^{-1} K^{-1}), and T is the temperature (K).

SO_2 reacts with H_2O_2, O_3, NO_2, and O_2 (TMIs-catalysed) in the aqueous phase. The rates of the four main aqueous reactions are expressed as (He et al., 2018; Seinfeld and Pandis, 2006):

$$R_{SO_2+O_3} = (k_1[SO_2 \cdot H_2O] + k_2[HSO_3^-] + k_3[SO_3^{2-}])[O_3(aq)] \ , \quad \text{(Eq. 7)}$$

$$R_{SO_2+H_2O_2} = \frac{k_4[H^+][HSO_3^-][H_2O_2(aq)]}{1 + K[H^+]} \ , \quad \text{(Eq. 8)}$$

$$R_{SO_2+NO_2} = k_5[S(IV)][NO_2(aq)] \ , \quad \text{(Eq. 9)}$$

$$R_{SO_2+O_2} = k_6[H^+]^{-0.74}\,[S(IV)][Mn(II)][Fe(III)] \qquad (pH < 4.2) \ , \quad \text{(Eq. 10)}$$

$$R_{SO_2+O_2} = k_7\left[H^+\right]^{0.67}\,[S(IV)][Mn(II)][Fe(III)] \qquad (pH > 4.2)\ , \tag{Eq. 11}$$

The production rate of sulfate through aqueous oxidation routes can be expressed as:

$$P_{aqu(ox_i)} = 3600 \times 96 \times R_{SO_2+ox_i} \times \frac{LWC}{\rho_{H_2O}}\ , \tag{Eq. 12}$$

where k_n (n = 1–7) is the rate constant of each oxidation route, K = 13 M^{-1} at 298 K, LWC is the liquid water content (mg m^{-3}), ρ_{H2O} is the density of water (1 kg L^{-1}), and ox_i (i = O_3, H_2O_2, NO_2, and O_2) represents different oxidants.

The heterogeneous reaction rate $R_{het(ox_i)}$ can be expressed as (Jacob, 2000; Wang et al., 2012; Zheng et al., 2015):

$$R_{het(ox_i)} = k_{ox_i}[SO_2(g)]\ , \tag{Eq. 13}$$

where

$$k_{ox_i} = \left(\frac{d_p}{2D_i} + \frac{4}{v_i\,\gamma_i}\right)^{-1} S_p\ , \tag{Eq. 14}$$

d_p is the effective diameter of the particles (m), D_i is the gas phase molecular diffusion coefficient (m^2 s^{-1}), v_i is the mean molecular speed in the gas phase (m s^{-1}), and S_p is the aerosol surface area (m^2 m^{-3}). The uptake coefficient γ_i depends on RH:

$$\gamma_i = \begin{cases} \gamma_{low} & 0 < RH \leqslant 50\ \% \\ \gamma_{low} + \dfrac{(\gamma_{high} - \gamma_{low})(RH - 0.5)}{RH_{max} - 0.5} & 50\ \% < RH \leqslant RH_{max} \\ \gamma_{high} & RH_{max} < RH \leqslant 100\ \% \end{cases} \tag{Eq. 15}$$

where γ_{low} and γ_{high} can be obtained from Wang et al. (2012) and RH_{max} is the RH at which γ reaches γ_{high}. The rate of sulfate production via heterogeneous reactions $P_{het(ox_i)}$ can be expressed as:

$$P_{het(ox_i)} = \frac{3600 \times 96 \times p \times R_{het(ox_i)}}{RT}\ , \tag{Eq. 16}$$

S2.2 Influencing parameters

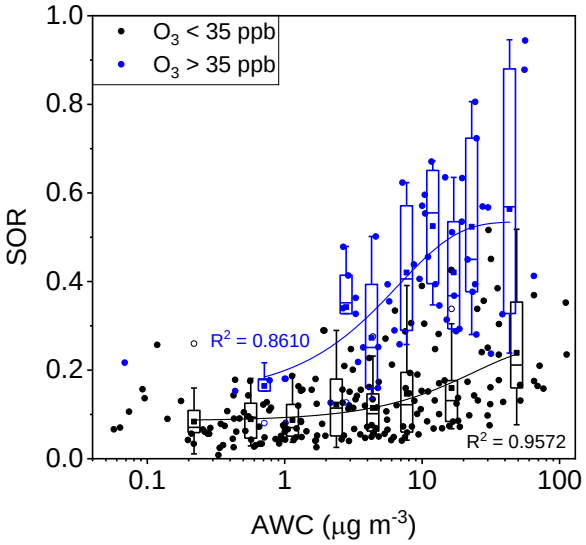

Figure S3. Plot of the SOR against aerosol water content (AWC) (note log scale), grouped by O_3 concentration. The solid blue circles represent O_3 > 35 ppb and the solid black circles represent O_3 < 35 ppb. The boxes represent, from top to bottom, the 75th, 50th, and 25th percentiles in each bin, which were also separated according to the 35 ppb O_3 concentration threshold; the bin widths were set such that there were an approximately equal number of data points in each bin. The whiskers, solid squares, and open circles represent 1.5 times the IQR, mean values, and outlier data points, respectively. The lines are best fits to the mean values based on a sigmoid function. Data for days with rain or snow were excluded from this plot.

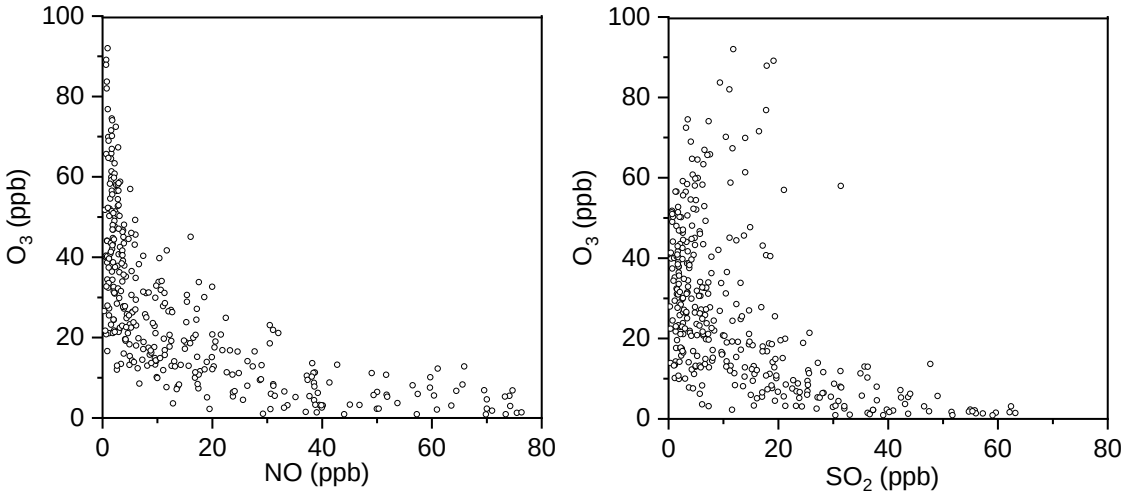

Figure S4. Plots of O_3 against the primary emission tracers NO and SO_2.

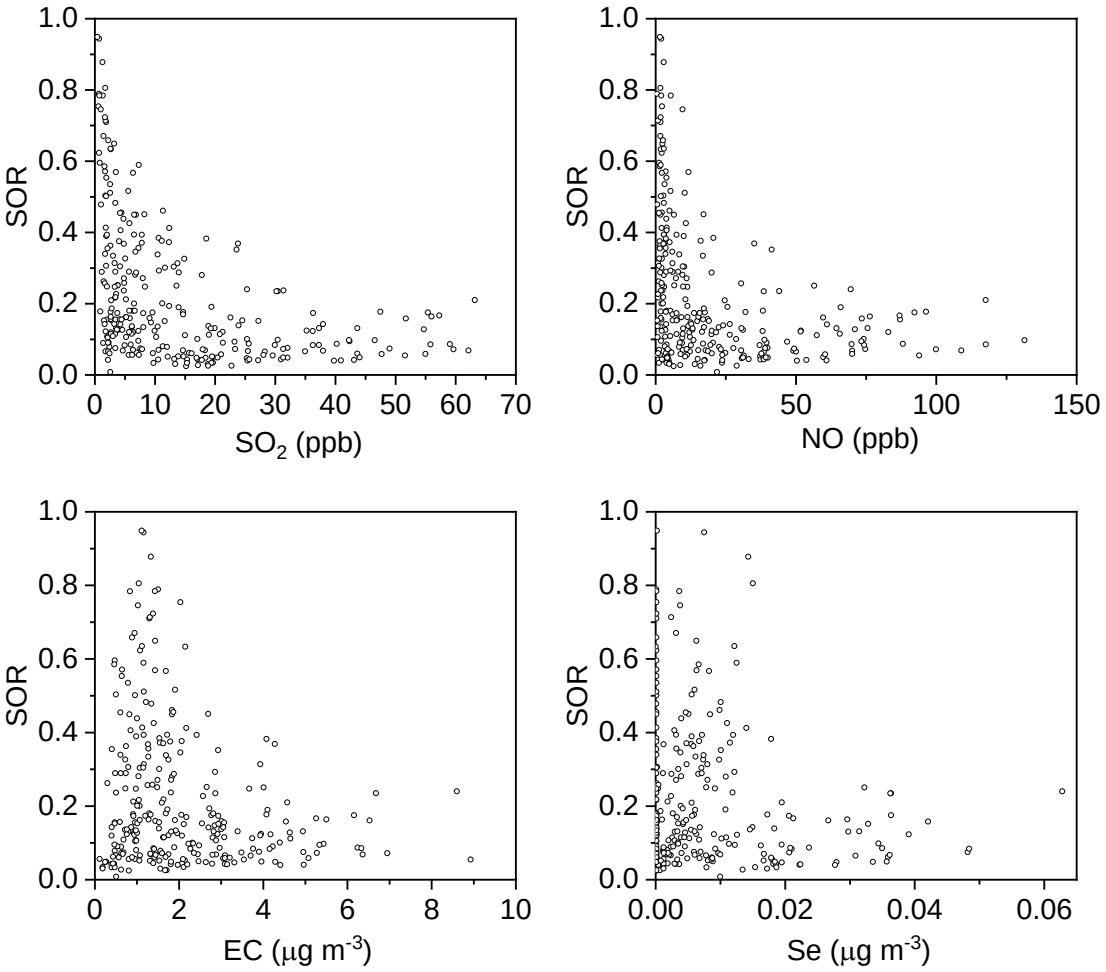

Figure S5. Plots of sulfur oxidation ratios (SORs) against the primary emission tracers SO_2, NO, EC, and Se.

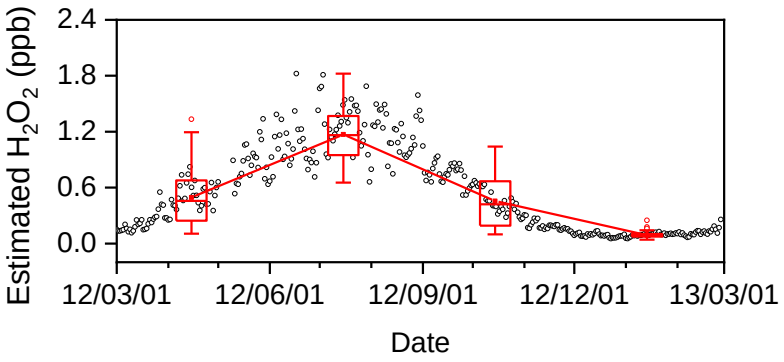

Figure S6. Time series of estimated H_2O_2 from March 1 2012 to February 28 2013. H_2O_2 was estimated from temperature (T) based on the fitting function $H_2O_2 = 0.1155e^{0.0846T}$ according to Fu (2014). The boxes represent, from top to bottom, the 75th, 50th, and 25th percentiles for each season. The whiskers, solid red squares, and open red circles represent 1.5 times the interquartile range (IQR), seasonal mean values, and outlier data points, respectively.

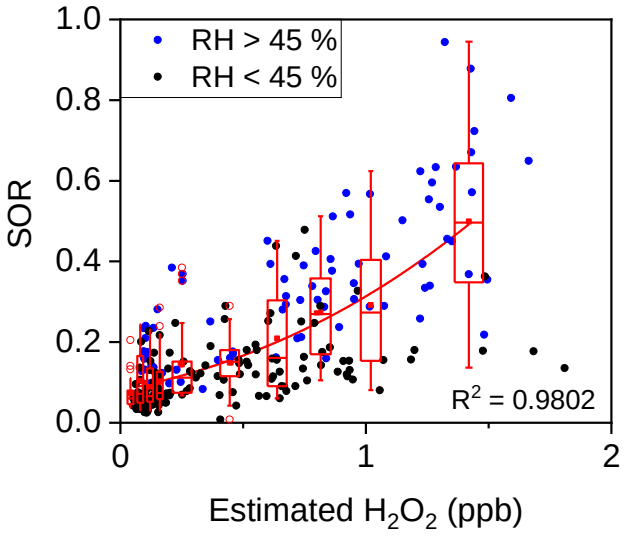

Figure S7. Plot of the SOR against estimated H_2O_2 grouped by RH. The solid blue circles represent RH > 45 % and the solid black circles represent RH < 45 %. The boxes represent, from top to bottom, the 75th, 50th, and 25th percentiles in each bin. The bin widths were set such that there were an approximately equal number of data points in each bin. The whiskers, solid squares, and open circles represent 1.5 times

the IQR, mean values, and outlier data points, respectively. The line are best fits to the mean values based on an exponential function. Data for days with rain were excluded from this plot.

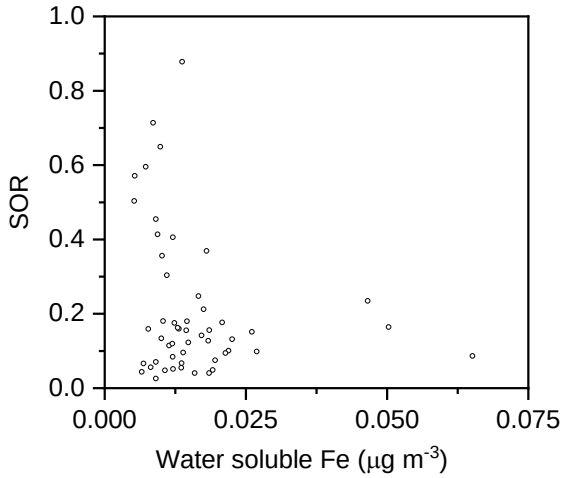

Figure S8. Plot of the SOR against water soluble Fe (54 samples selected every 6 days throughout the sampling period).

S2.3 Seasonal variations

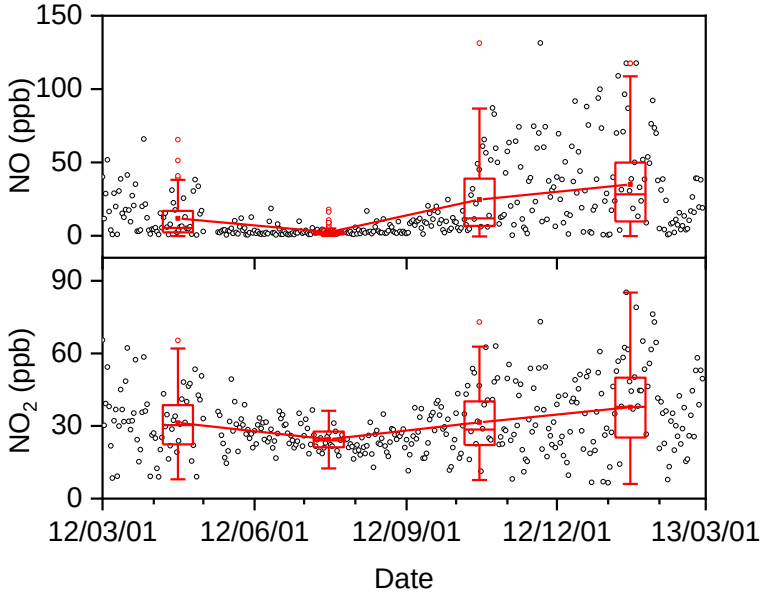

Figure S9. Time series of NO and NO_2 from March 1 2012 to February 28 2013 (open black circles). The boxes represent, from top to bottom, the 75th, 50th, and 25th percentiles for each season. The whiskers, solid red squares, and open red circles represent 1.5 times the IQR, seasonal mean values, and outlier data points, respectively.

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