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# Does afforestation deteriorate haze pollution in Beijing–Tianjin–Hebei (BTH), China?

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Abstract. Although aggressive emission control strategies have been implemented recently in the Beijing-Tianjin-Hebei area (BTH), China, pervasive and persistent haze still frequently engulfs the region during wintertime. Afforestation in BTH, primarily concentrated in the Taihang and Yan Mountains, has constituted one of the controversial factors exacerbating the haze pollution due to its slowdown of the surface wind speed. We report here an increasing trend of forest cover in BTH during 2001-2013 based on long-term satellite measurements and the impact of the afforestation on the fine-particle (PM<sub>2.5</sub>) level. Simulations using the Weather Research and Forecast model with chemistry reveal that afforestation in BTH since 2001 has generally been deteriorating the haze pollution in BTH to some degree, enhancing PM<sub>2.5</sub> concentrations by up to 6 % on average. Complete afforestation or deforestation in the Taihang and Yan Mountains would increase or decrease the PM2.5 level within 15 % in BTH. Our model results also suggest that implementing a large ventilation corridor system would not be effective or beneficial to mitigate the haze pollution in Beijing.

# 1 Introduction

Heavy haze with extremely high levels of fine particles  $(PM_{2.5})$ , caused by rapid growth of industrialization, urbanization, and transportation, frequently covers northern China during wintertime, particularly in the Beijing–Tianjin–Hebei area (BTH). The haze pollution in BTH remarkably impairs visibility and potentially causes severe health defects (Lim et al., 2013; Wang and Hao, 2012). The Chinese State Council issued the "Air Pollution Prevention and Control Action Plan" (APPCAP) in September 2013 with the aim of improving China's air quality within 5 years and reducing  $PM_{2.5}$  by up to 25% by 2017 relative to 2012 levels. Although aggressive emission control strategies have been undertaken since the initiation and implementation of the Action Plan, widespread and persistent haze still often engulfs BTH.

Aside from emissions, meteorological conditions play a key role in the haze pollution, affecting the formation, transformation, diffusion, transport, and removal of PM<sub>2.5</sub> in the atmosphere (Bei et al., 2012, 2017). Multifarious measurements have provided cumulative evidence that the widespread slowdown of surface wind speeds has occurred globally and in China since the 1980s (Chen et al., 2013; McVicar et al., 2012), which facilitates the pollutant accumulation to deteriorate air quality (Zhao et al., 2013; Sun et al., 2016). An increase in the surface roughness induced by increased vegetative biomass has been proposed to be responsible for the surface-level stilling to some degree (Wu et al., 2016b; Vautard et al., 2010). Consequently, there has been ongoing debate in China on whether the afforestation program contributes enough to mitigate the haze formation in BTH (China Forestry Network, 2016a, 2017).

Deforestation and its potential to severe droughts and massive floods has raised serious concerns in China since the 1970s, fostering the largest afforestation project in the world (Liu et al., 2008). Six key afforestation programs have been implemented since 2001, and "the Green Great Wall" of China has been established in northern China (Duan et al., 2011). A remarkable forest growth has been reported in the northwest of BTH from 2000 to 2010 (Li et al., 2016), which has the potential to increase the surface roughness and decrease the surface wind speed (Wu et al., 2016a; Bichet et al., 2012), and could potentially aggravate the haze pollution. In addition, previous studies have shown that the afforestation is beneficial for the atmosphere to remove  $O_3$ ,  $NO_x$ ,  $SO_2$ , and  $PM_{2.5}$  through the dry deposition process (Zhang et al., 2015, 2017; Huang et al., 2016). Hence, a large artificial ventilation corridor system has been proposed, highly anticipated to ventilate Beijing (China Forestry Network, 2014, 2016b, c).

In the present study, we report an analysis of long-term satellite measurements of the land cover change in BTH and quantitatively evaluate the impacts of the afforestation on the haze pollution in BTH using the WRF-CHEM model. We have further evaluated the effect of the proposed large artificial ventilation corridor system on the haze mitigation in Beijing. The model configuration and methodology are provided in Sect. 2. Data analysis and model results are presented in Sect. 3, and conclusions are given in Sect. 4.

## 2 Data, model, and methodology

## 2.1 MODIS data

The data utilized in the study are the annual land cover product, MCD12Q1, derived from the Terra and Aqua Moderate Resolution Imaging Spectroradiometer (MODIS) observations since 2001 (Friedl et al., 2002). The product has been widely used in studies of atmospheric science, hydrology, ecology, and land change science (Gerten et al., 2004; Guenther et al., 2006; Reichstein et al., 2007; Turner et al., 2007). Wu et al. (2008) have compared four global land cover datasets across China, concluding that the MODIS land cover product is the most representative over China with the minimal bias from the China's National Land Cover Dataset. The MCD12Q1 (Version 5.1) IGBP (International Geosphere-Biosphere Programme) scheme with a spatial resolution of 500 m is utilized to explore the variability of the land cover fraction (LCF) from 2001 to 2013 in BTH and assimilated into the WRF-CHEM model. The high-resolution land cover product is generated using a supervised classification algorithm in conjunction with a revised database of high-quality land cover training sites (Friedl et al., 2002). The accuracy of the IGBP layer of MCD12Q1 is estimated to be 72.3-77.4 % globally, with a 95 % confidence interval (Friedl et al., 2002, 2010). Great efforts have been made to evaluate the accuracies of the global land cover datasets over China. The overall accuracy of MCD12Q1 in China is estimated to be 55.9-68.9% (Bai et al., 2015; Yang et al., 2017), which could be



Figure 1. (a) The model domain, region of interest (ROI), and monitoring sites. (b) The topography and monitoring sites in January 2014. The circles represent the centers of cities with ambient monitoring sites, and the size of circles denotes the number of monitoring sites in the cities. The boundary of BTH region is highlighted with bright lines. The Yan (Yanshan Mt.) and Taihang Mountains are also displayed.

increased to about 70 % when ignoring the differences of five forests.

## 2.2 WRF-CHEM model and configurations

We use a specific version of the WRF-CHEM model (Grell et al., 2005) to investigate the impacts of the afforestation on the haze pollution in BTH. The model includes a new flexible gas phase chemical module and the CMAQ/Models3 aerosol module developed by US EPA (Binkowski and Roselle, 2003). The wet deposition of chemical species follows the CMAQ method. The dry deposition parameterization follows Wesely (1989), and the dry deposition velocity of aerosols and trace gases is calculated as a function of the local meteorology and land use. The photolysis rates are calculated by an FTUV (fast radiation transfer model) (Li et al., 2005). The inorganic aerosols are predicted using ISORROPIA Version 1.7 (http://nenes.eas.gatech.edu/ISORROPIA/, last access: 30 June 2018) (Nenes et al., 1998). The secondary organic aerosol (SOA) is predicted using a non-traditional SOA module, including the VBS (volatility basis set) modeling approach and SOA contributions from glyoxal and methylglyoxal. Detailed information about the WRF-CHEM model can be found in previous studies (Li et al., 2010, 2011a, b, 2012).

Model spin-up time

Simulation regions	Beijing–Tianjin–Hebei–Shandong		
Simulation period	1 December 2013–31 January 2014		
Domain size	$200 \times 200$		
Domain center	38° N, 116° E		
Horizontal resolution	$6 \text{ km} \times 6 \text{ km}$		
Vertical resolution	35 vertical levels with a stretched vertical grid with spacing ranging		
	from 30 m near the surface, to 500 m at 2.5 km and 1 km above 14 km		
Microphysics scheme	WSM 6-class graupel scheme (Hong and Lim, 2006)		
Boundary layer scheme	MYJ TKE scheme (Janjić, 2002)		
Surface layer scheme	MYJ surface scheme (Janjić, 2002)		
Land-surface scheme	Unified Noah land-surface model (Chen and Dudhia, 2001)		
Longwave radiation scheme	Goddard longwave scheme (Chou and Suarez, 2001)		
Shortwave radiation scheme	Goddard shortwave scheme (Chou and Suarez, 1999)		
Meteorological boundary and initial conditions	NCEP $1^{\circ} \times 1^{\circ}$ reanalysis data		
Chemical initial and boundary conditions	MOZART 6 h output (Horowitz et al., 2003)		
Anthropogenic emission inventory	SAPRC-99 chemical mechanism emissions (Zhang et al., 2009)		
Biogenic emission inventory	MEGAN model developed by Guenther et al. (2006)		

28 h

High PM<sub>2.5</sub> pollution episodes from 1 December 2013 to 31 January 2014 in the North China Plain (NCP) have been simulated using the WRF-CHEM model. The model simulation domain is shown in Fig. 1, and detailed model configurations can be found in Table 1. The chemical initial and boundary conditions are interpolated from the 6h output of MOZART-4 (Emmons et al., 2010; Horowitz et al., 2003). MOZART-4 is driven by meteorology fields from the NASA GMAO GEOS-5 model, using anthropogenic emissions based on the Streets et al. (2006) inventory and fire emissions from FINN-v1 (Wiedinmyer et al., 2011). The model has been evaluated comprehensively with several sets of observations, reproducing the tropospheric chemical composition well (Emmons et al., 2010). The model results have been successfully and widely used as the initial and lateral boundary conditions for chemical transport models. The anthropogenic emission inventory used in the present study is developed by Zhang et al. (2009), with the base year of 2013, including contributions from agriculture, industry, power, residential, and transportation sources. Figure S1 shows the emission distribution of OC (organic carbon), VOCs (volatile organic compounds),  $NO_x$ , and  $SO_2$  in the simulation domain. The high emissions of OC, VOCs,  $NO_x$ , and  $SO_2$  are generally concentrated in the plain region of BTH and Shandong province, the downwind area of afforestation.

The hourly near-surface CO, SO<sub>2</sub>, NO<sub>2</sub>, O<sub>3</sub>, and PM<sub>2.5</sub> mass concentrations released by China's Ministry of Environmental Protection are used to validate the model simulations and are accessible from the following website: http: //www.aqistudy.cn/ (last access: 30 June 2018).

We use the normalized mean bias (NMB), the index of agreement (IOA), and the correlation coefficient (R) to assess the WRF-CHEM model performance in simulating air pollutants against measurements.

3.7

$$NMB = \frac{\sum_{i=1}^{N} (P_i - O_i)}{\sum_{i=1}^{N} O_i},$$
(1)

$$IOA = 1 - \frac{\sum_{i=1}^{N} (P_i - O_i)^2}{\sum_{i=1}^{N} (|P_i - \overline{O}| + |O_i - \overline{O}|)^2},$$
(2)

$$R = \frac{\sum_{i=1}^{N} (P_i - \overline{P}) (O_i - \overline{O})}{\left[\sum_{i=1}^{N} (P_i - \overline{P})^2 \sum_{i=1}^{N} (O_i - \overline{O})^2\right]^{\frac{1}{2}}},$$
(3)

where  $P_i$  and  $O_i$  are the calculated and observed air pollutant concentrations, respectively. N is the total number of the predictions used for comparisons, and  $\overline{P}$  and  $\overline{O}$  represent the average of predictions and observations, respectively. The IOA ranges from 0 to 1, with 1 showing perfect agreement of the prediction with the observation. The *R* ranges from -1 to 1, with 1 implicating perfect spatial consistency of observations and predictions.

## 2.3 MCD12Q1 data assimilation to the WRF-CHEM model

The IGBP layer in MCD12Q1 is suitable for the WRF-CHEM IGBP land cover scheme, which consists of 11 natural vegetation classes, 3 developed and mosaicked land classes, and 3 non-vegetated land classes. Table S1 displays the comparison of land cover classification between the WRF-CHEM model and MCD12Q1. We use the gridded LCF of each category to assimilate the MCD12Q1 satellite data to the WRF-CHEM model.

$$LCF_{i,j,k} = \frac{Area_{i,j,k}}{Area_{i,j}},$$
(4)

where *i* and *j* are grid cell indices of the WRF-CHEM model domain, Area<sub>*i*,*j*,*k*</sub> stands for the total area of each land cover category *k* within grid cell (i, j), and Area<sub>*i*,*j*</sub> is the area of grid cell (i, j). The LCF<sub>*i*,*j*,*k*</sub> ranges from 0 to 1.

To evaluate the afforestation impacts on the haze pollution in BTH, we have used and modified the coupled unified Noah land-surface model (LSM), which was developed based on Oregon State University LSM (Chen and Dudhia, 2001). Noah is able to reasonably reproduce the observed diurnal variation of sensible heat fluxes and surface skin temperature. Also, it is capable of capturing the diurnal and seasonal evolution in evaporation and soil moisture (Chen et al., 1996). Despite some remaining issues, Noah has been chosen for further refinement and implementation in NCEP regional and global coupled weather and climate models because of its relative simplicity and adequate performance (Mitchell, 2005, update 2017). The surface roughness length (SFz0) in Noah is calculated based on the dominant land cover category (https://ral.ucar.edu/solutions/products/ unified-noah-lsm/, last access: 30 June 2018).

$$SFz0 = (5)$$

$$\int SFz0_{\min} \qquad G_{T} \le G_{\min}$$

$$\begin{cases} \text{OI} \ \text{ZOmin} & \text{OI} \ \text{ZOmin} \\ (1 - G_{\text{f}}) \times \text{SFz0}_{\text{min}}, & G_{\text{min}} \leq G_{\text{T}} \leq G_{\text{max}} \\ + G_{\text{f}} \times \text{SFz0}_{\text{max}} & G_{\text{T}} \gg G_{\text{max}}, \end{cases}$$

where SFz0<sub>min</sub> and SFz0<sub>max</sub> are the minimum and maximum SFz0 for each category.  $G_{\rm f}$  is the area fractional coverage of green vegetation, and  $G_{\rm T}$ ,  $G_{\rm min}$ , and  $G_{\rm max}$  are the threshold, minimal, and maximal value of  $G_{\rm f}$ , respectively.  $G_{\rm T}$ , SFz0<sub>min</sub>, and SFz0<sub>max</sub> are listed in Table S2.

In order to more precisely simulate surface stress within the sub-grid scale in heterogeneous terrain, the effective roughness length has been extensively studied, especially in the 1990s. Claussen (1990) has defined the effective roughness length as a value of the area average of the roughness length in heterogeneous terrain. The effective roughness length relies upon the blending height (Wieringa, 1986; Mason, 1988; Wood and Mason, 1991; Philip, 1996; Mahrt, 1996), at which the flow is approximately in equilibrium with underlying surface conditions and independent of horizontal position (Ma and Daggupaty, 1998). We have modified the Noah SFz0 calculation using the spatial average of the vegetation roughness length.

$$SFz0 = \sum_{k} LCF_k \times SFz0_k \tag{6}$$

 $SFzO_k$  denotes the gridded area fraction of land cover category k and is alculated by Eq. (5).



Figure 2. Land cover change from 2001 to 2013. Spatial distributions of (a) forests, (b) shrublands, (c) croplands, and (d) grasslands.

**Table 2.** Land cover change over Beijing and BTH from 2001 to2013.

Land cover categories	Land cover description	Beijing	BTH
1–5	Forests	$14.9\% \\ -12.6\% \\ -0.1\% \\ -2.0\% \\ -0.2\%$	7.2 %
6–7	Shrublands		-3.9 %
12/14	Croplands		1.9 %
8–10	Grasslands		-5.1 %
Others	–		-0.1 %

#### 3 Results and discussions

#### 3.1 Land cover change in BTH

The land cover in BTH and Beijing exhibits appreciable variation from 2001 to 2013 (Fig. 2 and Table 2). In BTH, forests and croplands have increased by 7.2% and 1.9%, while shrublands and grasslands/savannas have decreased by 3.9 % and 5.1%, respectively. In Beijing, forests have increased by 14.9%, while shrublands have decreased by 12.6%. Apparently, the forest LCF has increased substantially in western and northern BTH, concentrated in the Taihang and Yan Mountains, with an increase up to 50 %. This result is consistent with the previous study of Li et al. (2016), which has reported a remarkable forest growth in the northwest of NCP from 2000 to 2010. As such, a "Green Great Wall" has been established (Fig. 2a), which has reportedly protected the southeastern BTH from the dust pollution (Liu et al., 2008; Duan et al., 2011; Parungo et al., 2013). The land cover change, particularly the evident forest growth, is primarily attributed to China's national afforestation programs aiming to increase the forest coverage and to conserve soil and water, including the Grain for Green Project, the Three-North Shelter Forest System Project (Phase IV), and the Natural Forest Conservation Program (Yin et al., 2010; Cao et al., 2011).

#### **3.2 Model performance**

We have first assimilated into the WRF-CHEM model the MCD12Q1 product of 2013 and performed the numerical simulation of haze pollution episodes from 1 December 2013 to 31 January 2014. To enable discussion, we have defined the simulation with the 2013 land cover as the reference case (hereafter referred to as REF case), and results from the reference simulation are compared to observations in BTH.

Considering the key role of meteorological fields in determining the formation, transformation, diffusion, transport, and removal of the air pollutants (Bei et al., 2017), Fig. S2 presents the comparison of the simulated wind speed and direction, and planetary boundary layer height with the reanalysis data from ECMWF (European Centre for Medium-range Weather Forecasts) at monitoring sites. The predicted temporal variations of the three meteorological parameters are generally in agreement with the reanalysis data, with the IOAs exceeding 0.80 and the absolute NMB less than 25 %.

Figure 3 presents the calculated and observed temporal profiles of near-surface air pollutant concentrations averaged at monitoring sites in BTH during the simulation period, including PM<sub>2.5</sub>, O<sub>3</sub>, NO<sub>2</sub>, SO<sub>2</sub>, and CO. The WRF-CHEM model generally reproduces the haze pollution episodes well. For example, all the haze events during the period are captured successfully (Fig. 3a), with an IOA of 0.90 and a NMB of 2.1 % for PM2.5 mass concentrations. The model reasonably yields O3 variations compared to observations, with an IOA of 0.80, but underestimates O<sub>3</sub> concentrations, with a NMB of -15.9% (Fig. 3b). In winter, the insolation is weak in the north of China, unfavorable for the O<sub>3</sub> photochemical production, so the O<sub>3</sub> level is substantially influenced by the boundary conditions (Li et al., 2017; Wu et al., 2017). Hence, one of possible reasons for the  $O_3$  underestimation might be from the uncertainty in the O<sub>3</sub> boundary conditions. The simulated temporal variations of NO2 mass concentrations are well consistent with the observation, and the IOA and NMB are 0.91 and 0.6%, respectively. The SO<sub>2</sub> and CO temporal variations are also reasonably replicated against observations, with IOAs of 0.82 and 0.84, respectively.

Figure 4 shows the spatial comparison of calculated and observed PM<sub>2.5</sub> concentrations. Generally, the average predicted PM<sub>2.5</sub> spatial patterns agree well with the observations at the monitoring sites in BTH during the whole period (Fig. 4b) and each month (Fig. 4c and d), with *R* values exceeding 0.85, indicating good agreement of simulations with observations. The observed PM<sub>2.5</sub> concentrations frequently exceed 150  $\mu$ g m<sup>-3</sup> in BTH, showing the frequent occurrence of heavy haze pollution events. The model gen-



Figure 3. Comparisons of observed (black dots) and simulated (solid red lines) diurnal profiles of near-surface hourly mass concentrations of (a)  $PM_{2.5}$ , (b)  $O_3$ , (c)  $NO_2$ , (d)  $SO_2$ , and (d) CO averaged at monitoring sites in BTH from 1 December 2013 to 31 January 2014.

erally yields the observed high  $PM_{2.5}$  concentrations in BTH and their surrounding areas, although the model underestimation or overestimation still exists. Additionally, compared to observations, the model also performs well in simulating the spatial pattern of haze episodes with various timescales ranging from 8 to 16 days (Fig. S3).

The good agreements of the simulated mass concentrations of air pollutants with observations at monitoring sites in BTH show that the emission inventory used in the present study and simulated wind fields are generally reasonable, providing a reliable base for the further assessment. It is worth noting that, although the predicted meteorological parameters are generally consistent with the reanalysis data from ECMWF at monitoring sites, other factors still affect the meteorological field simulations and cause biases to compensate some of the deficiencies of the WRF-CHEM model, such as overestimation of surface wind speeds.

## 3.3 Effect of afforestation on haze pollution in BTH

Change in the land cover alters the surface roughness height (SFz0) that plays an important role in determining the surface level wind speed and energy exchange between the atmo-



**Figure 4.** Pattern comparisons of calculated and observed nearsurface  $PM_{2.5}$  mass concentrations. (a) Spatial correlation between calculated and observed  $PM_{2.5}$  concentrations during each month and the whole simulation period. Horizontal distribution of calculated (color contour) and observed (colored circles) average  $PM_{2.5}$ concentrations during (b) the whole simulation period, (c) December 2013, and (d) January 2014, along with the simulated wind fields (black arrows).

sphere and the land surface. Numerous studies have demonstrated that increasing SFz0 tends to decelerate the surface wind (Wu et al., 2016a, b), obstructing the dispersion of air pollutants (Sun et al., 2016; Zhao et al., 2013; Tie et al., 2015). In order to evaluate the impact of the afforestationinduced SFz0 change and resultant dynamical change (e.g., wind field) on the haze formation, a sensitivity experiment is designed, in which the MCD12Q1 product of 2001 is assimilated into the WRF-CHEM model to represent the land cover situations before the afforestation (hereafter referred to as SEN-AFF case).

Figure 5a and b display the SFz0 change and its correlation with forest LCF change from 2001 to 2013, respectively. The land cover change considerably alters the SFz0, particularly in the afforestation area, with a SFz0 increase ranging from 0.1 to 0.3 m. Apparently, the SFz0 exhibits a distinct increasing trend in western and northern BTH, concentrated in the Taihang and Yan Mountains, which is well consistent with the increase of the forest LCF. The SFz0 change is highly correlated with the forest LCF change, with a correlation coefficient of 0.91. Generally, the SFz0 is mainly dependent upon the LCF (Eq. 6) and sensitive to the forest change (Table S2). Therefore, afforestation constitutes the most important factor for the increase in the SFz0 in BTH.

It is worth noting that Jiménez and Dudhia (2012) have point out that there still exist large uncertainties in parame-



**Figure 5.** (a) SFz0 change from 2001 to 2013, and (b) its correlation with the forest LCF change; horizontal distribution of (c) absolute and (d) relative near-surface  $PM_{2.5}$  mass concentration changes caused by the afforestation. The wind field changes are shown in black arrows in (c) and (d).

terizing the air land interaction over complex terrain. Besides the vegetation effect on the roughness length, drag of subgrid features of topography needs to be considered. The parameterization of orographic flow over complex terrain is a challenging problem at the mesoscale numerical simulation. In early versions of the WRF model, a large bias in predicting surface winds over complex terrain has occurred due to the drag exerted by unresolved topography (Cheng and Steenburgh, 2005). Great efforts have been made to improve the simulation of orographic flow over complex terrain. The new parameterization scheme introduced in the WRF model since version 3.4.1 has corrected this high wind speed bias over plains and valleys (Mughal et al., 2017), and it has also corrected the low wind speed bias found over the mountains and hills (Jiménez and Dudhia, 2012).

Figure 5c and d illustrate the influence of the afforestation on the surface  $PM_{2.5}$  and wind field averaged during the simulation period (defined as (REF – SEN-AFF)). The prevailing westerly or northerly wind is decelerated in the western and northwestern BTH due to the increased SFz0 caused by the afforestation, with the wind speed decrease ranging from 0.3 to  $1.5 \text{ m s}^{-1}$ . The afforestation tends to deteriorate the haze pollution in BTH, particularly in the downwind area of the afforestation, with the period average  $PM_{2.5}$  enhancement reaching about 6–15 µg m<sup>-3</sup>, or 3–6%. The  $PM_{2.5}$  enhancement in Beijing is the most evident, corresponding to



**Figure 6.** Horizontal distribution of (a) the average near-surface  $PM_{2.5}$  mass concentration and (b) its change due to the afforestation during an intensified northerly/northwesterly event from 00:00 to 10:00 Beijing time on 18 January 2014. The wind field and its change are shown in black arrows.

the rapid growth of forests in the west and in the north of Beijing. Furthermore, during each episode, the afforestation generally tends to deteriorate the haze pollution in BTH, enhancing the  $PM_{2.5}$  concentration by about up to 3–6%, particularly in the downwind area of the afforestation (Fig. S4). On average, the difference of the simulated air pollutants and meteorological parameters between the REF and SEN-AFF case is not significant (Fig. 5c and d).

The occurrence of heavy haze pollution in BTH is generally associated with the weakening of northerly or northwesterly winds, which facilitates the accumulation of air pollutants in BTH. The afforestation in the western and northwestern BTH increases SFz0, further decelerating northerly or northwesterly winds and deteriorating the haze pollution. However, the afforestation only plays a marginal role in worsening the haze pollution and does not constitute the main cause of the heavy haze formation.

Apparently, during the haze development, when the northerly or northwesterly wind is weak or becomes calm, the SFz0 increase due to the afforestation contributes negligibly to the haze deterioration in BTH. However, once the northerly or northwesterly wind commences to strengthen but is not strong enough to evacuate the air pollutants in BTH, the SFz0 increase would play an appreciable role in sustaining high PM2.5 levels in the downwind area of the afforestation. Figure 6 presents the PM<sub>2.5</sub> contribution of the afforestation during the occurrence of a northerly gust on 18 January 2014. The intensified northerly wind cleanses the northern BTH, but the haze pollution is still very severe in the southern BTH. The afforestation considerably elevates the PM<sub>2.5</sub> concentration in southeastern BTH, particularly in Beijing and Tianjin, with the PM<sub>2.5</sub> contribution exceeding up to 15 % (Fig. 6b).

It is worth noting that the aerosol species (organic aerosol, sulfate, nitrate, ammonium, and elemental carbon) exhibit the same variation trend as the  $PM_{2.5}$  due to the afforestation



**Figure 7.** Impacts of complete deforestation/afforestation over Taihang and Yanshan Mountains on (a)–(c) SFz0 and (b)–(d) average near-surface PM<sub>2.5</sub> mass concentrations from 1 December 2013 to

31 January 2014, along with the wind field change (black arrows).

(Fig. S5). Apparently, the organic aerosol is the major contributor to the PM<sub>2.5</sub> variation due to the afforestation, followed by the sulfate and ammonium aerosol. The afforestation also increases emissions of the biogenic SOA (BSOA) precursors, such as isoprene and monoterpenes. However, due to the very low emissions of BSOA precursors during wintertime (Guenther et al., 2006, 2012), the BSOA contribution to PM<sub>2.5</sub> concentrations is insignificant, less than  $3 \mu g m^{-3}$  on average during the whole episodes (Fig. S6a). The average BSOA enhancement due to the afforestation is less than 0.5 % (Fig. S6b). Furthermore, in general, the afforestation has little effect on the boundary layer height, upward sensible heat flux (associated with turbulent mixing), and moisture (related to clouds) in BTH (Fig. S7).

To assess the upper limit of impacts of the afforestation on the PM<sub>2.5</sub> level in BTH, two additional experiments are conducted and compared to the REF case. The two experiments are one with complete deforestation and the other with complete afforestation in the Taihang and Yan Mountains (Fig. 7a and c). In the complete deforestation sensitivity case, the barren surface with SFz0 of 0.01 m is used to replace other land cover categories. In the complete afforestation case, the deciduous broadleaf forest category with SFz0 of 0.5 m is used to replace other land cover categories. As shown in Fig. 7, complete deforestation considerably decreases the PM<sub>2.5</sub> level in BTH, with the period average PM<sub>2.5</sub> reduction ranging from 5 to 18  $\mu$ g m<sup>-3</sup> generally, and in particular the PM<sub>2.5</sub> concentration in Beijing is reduced by more than 10  $\mu$ g m<sup>-3</sup>, due to the intensified northerly or northwesterly wind caused by the decrease of SFz0 in the Taihang and Yan Mountains. Complete afforestation deteriorates the haze pollution in BTH, and the haze pollution remains in the Taihang and Yan Mountains due to the weakened northerly or northwesterly wind. Additionally, the enhancement of  $PM_{2.5}$  concentrations in the foothills of the Taihang and Yan Mountains is obvious, varying from 10 to 25 µg m<sup>-3</sup> (Fig. 7d).

Interestingly, the afforestation deteriorates most of the haze pollution in Beijing (see Fig. 5). So it is anticipated that the proposed large ventilation corridor system could alleviate the haze pollution in Beijing (China Forestry Network, 2014, 2016b, c). Originally, the ventilation corridor system was devised to relieve the urban heat island effect and improve the thermal environmental conditions in urbanized regions. With the frequent occurrence of heavy haze in Beijing, the debatable system is expected to blow away the haze and bring blue skies to Beijing. In order to examine the effects of the wind corridor system, a sensitivity experiment is conducted based on the base case, in which three artificial ventilation corridors are designed in the northwest, north, and northeast of Beijing, with a width of 6 km (Fig. 8a). For all the grid cells within the corridors, the barren surface with SFz0 of 0.01 m is used to replace other land cover categories. In contrast to the expectation, our sensitivity results show that the  $PM_{2.5}$ reduction due to the designed ventilation corridor system is less than 1 % in Beijing (Fig. 8b). Note that the width of the ventilation corridor in the sensitivity study is 12 times the proposed one. Hence, the proposed large ventilation corridor system is not effective or beneficial to mitigate the haze pollution in Beijing.

#### 4 Summary and conclusions

The annual land cover product, MCD12Q1, derived from the MODIS observations since 2001 has been used to analyze the land cover change in BTH. A considerable increasing trend of forests in the western and northwestern BTH has been identified, which is caused by China's national afforestation programs. Forests in BTH and Beijing have increased by 7.2 % and 14.9 %, respectively, from 2001 to 2013. The fast forest expansion has increased the surface roughness height, particularly in Beijing and its surrounding areas.

The MCD12Q1 product of 2013 has been assimilated into the WRF-CHEM model to represent the current land cover condition. Persistent haze pollution episodes in BTH from 1 December 2013 to January 2014 are simulated using the WRF-CHEM model. Generally, the WRF-CHEM model reproduces the temporal variations and spatial distributions of air pollutants reasonably well compared to observations at monitoring sites in BTH.

Sensitivity studies have demonstrated that the increase of the surface roughness height decreases the northwesterly or northerly wind speed in the western and northwestern BTH by about  $0.3-1.5 \text{ m s}^{-1}$ . The haze pollution is deteriorated



**Figure 8.** Impacts of an artificial large ventilation corridor system on (a) SFz0 and (b) average near-surface  $PM_{2.5}$  mass concentrations from 1 December 2013 to 31 January 2014, along with the wind field (black arrows).

in BTH to some degree, and  $PM_{2.5}$  concentrations are generally enhanced by less than 6% due to the afforestation. The heavy haze formation in BTH is generally associated with meteorological conditions when the northerly or northwesterly wind is weak. Once the northerly or northwesterly wind is strengthened during the haze development in BTH, afforestation plays a considerable role in maintaining high  $PM_{2.5}$  concentrations in the downwind of the afforestation area. Complete afforestation or deforestation in the Taihang and Yan Mountains would increase or decrease the  $PM_{2.5}$  level within 15% in BTH.

Additionally, our model results do not support that the proposed large ventilation corridor system is beneficial to alleviate the haze pollution in Beijing. Under the unfavorable synoptic situations, emissions mitigation is the sole optimum approach to mitigate the haze pollution in BTH.

It is worth noting that, in the present study, contributions of the surface roughness change induced by afforestation to the haze pollution are primarily evaluated using the WRF-CHEM model, but many other factors which directly or indirectly influence air quality are also modified by the land cover change, including surface moisture, terrestrial erosion, pollutants' dry deposition, planetary boundary layer thermal stability, etc. For example, changes in surface moisture and surface erosion impact the emissions of natural particles; changes in dry deposition directly influence the air quality in situ and indirectly the air quality downwind with occurrence of recirculation. Therefore, when changes in all those factors caused by land cover change are accounted for, the role of afforestation in air quality in situ might be uncertain. In the online WRF-CHEM model, besides the surface roughness, the impacts of afforestation on the heat flux, surface moisture, surface erosion, and dry deposition of air pollutants have also been considered. Considering that afforestation in BTH is mainly distributed in the mountain region, the surface roughness increases induced by afforestation obviously decrease surface wind speeds, facilitating accumulation of air

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pollutants in the downwind region and further deteriorating the haze pollution.

*Data availability.* The real-time O<sub>3</sub> and PM<sub>2.5</sub> observations are accessible for the public on the website http://106.37.208.233: 20035/ (China MEP, 2013a). One can also access the historic profile of observed ambient pollutants through visiting http://www. aqistudy.cn/ (China MEP, 2013b). The MODIS Land Cover products (MOD12Q1) are accessible for the public on the website https: //lpdaac.usgs.gov/dataset\_discovery/modis/modis\_products\_table.

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*Author contributions.* GL designed the study. XL and GL wrote the paper. XL, NB, JW, and XL performed model simulations. TF, SZ, JC, and ZA analyzed the data. All authors reviewed and commented on the paper.

*Competing interests.* The authors declare that they have no conflict of interest.

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