Supplement of

# Experimental determination of Henry's law constants of difluoromethane (HFC-32) and the salting-out effects in aqueous salt solutions relevant to seawater 

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S1. Equilibrium time for the PRV-HS method experiments


Figure S1. Relative areas of GC-MS peaks for $\mathrm{CH}_{2} \mathrm{~F}_{2}$ versus headspace time duration for equilibration of $9.0 \mathrm{~cm}^{3}$ of aqueous $\mathrm{CH}_{2} \mathrm{~F}_{2}$ at 353 K .

## S2. An example of the IGS method experiments

Figure S 2 shows an example of time profile of $P_{t}$ and how to calculate the $k_{1}$ value for the IGS method experiments. The $k_{1}$ value at each time was calculated by fitting nearest three data of $P_{t}$ for each time. The average of the $k_{1}$ values is given as the $k_{1}$ value for the experimental run. Two standard deviation of the $k_{1}$ values gives errors of the $k_{1}$ value for the experimental run.


Figure S2. An IGS experimental result for $V=0.350 \mathrm{dm}^{3}$ and $F=3.32 \times 10^{-4} \mathrm{dm}^{\mathbf{3}} \mathrm{s}^{\mathbf{- 1}}$ at $\mathbf{2 5} 5^{\circ} \mathrm{C}$. (upper panel) time profile of $P_{t}$; (lower panel) values of $k_{1}$ calculated by fitting nearest three data of $\boldsymbol{P}_{t}$ for each time with respect to Eq. (1).

## S3. Results of the PRV-HS method experiments

Figure S 3 illustrates the results of a PRV-HS experiment at 313 K . In panel A, peak area $\left(S_{i j}\right)$ is plotted against the volume of the $\mathrm{CH}_{2} \mathrm{~F}_{2}$ gas mixture added $\left(v_{j}\right)$ for $V_{i}=9.0,7.5,6.0,4.5,3.0$, and $1.5 \mathrm{~cm}^{3}$. For each $V_{i}$, the data form a straight line intersecting the origin, indicating that $S_{i j}$ is proportional to $v_{j}$ for vials with the same value of $V_{i}$. The slope $\left(L_{i}\right)$ of each

5 line is obtained by linear regression with respect to Eq. (8), and the reciprocal of the slope $\left(L_{i}^{-1}\right)$ is plotted against the phase ratio $\left(V_{i} / V_{0}\right)$ in panel B of Fig. S3. Plots of $L_{i}^{-1}$ and $V_{i} / V_{0}$ obey Eq. (9). Table S1 lists the values of $L_{i}^{-1}$, the slopes and the intercepts for linear regression with respect to Eq. (9), and the $K_{\mathrm{H}}(T)$ values calculated from the slopes and the intercepts. Two measurements of $K_{\mathrm{H}}(T)$ were carried out at each temperature.

Furthermore, the $K_{\mathrm{H}}(T)$ values, along with errors of them at $95 \%$ confidence level, were also estimated by non-liner
10 fitting of the two datasets simultaneously at each temperature by use of Eq. (11) (Fig. S4). The $K_{\mathrm{H}}(T)$ values and their errors thus estimated are plotted in Fig. 2 and are listed in Table S1.

Table S1. $L_{i}$ values for various $V_{i} / V_{0}$ ratios at various temperatures, slopes and intercepts for linear regression with respect to Eq. $(10), K_{H}(T)$ values calculated from the slopes and intercepts, and $K_{H}(T)$ values and the errors at $95 \%$ confidence level estimated by non-linear fitting the two datasets simultaneously at each temperature (Fig. S4) with respect to Eq. (11).

| $\begin{gathered} T \\ (\mathrm{~K}) \end{gathered}$ | $L_{i}$ (a.u.) ${ }^{\text {a }}$ |  |  |  |  |  | Eq. (10) <br> Intercept | $\begin{aligned} & \text { Eq. (10) } \\ & \text { Slope } \end{aligned}$ | $K_{\mathrm{H}}\left(\mathrm{M} \mathrm{atm}^{-1}\right)$ |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | $V_{\mathrm{i}} / V=0.421$ | 0.351 | 0.280 | 0.210 | 0.140 | 0.070 |  |  | Eq. (10) | Eq. $(11)^{\text {b, c }}$ | Eq. (13) ${ }^{\text {b }}$ |
| 353 | $3.226 \pm 0.002$ | $3.270 \pm 0.026$ | $3.330 \pm 0.004$ | $3.391 \pm 0.008$ | $3.462 \pm 0.014$ | $3.526 \pm 0.009$ | 3.581 | -0.870 | 0.026 | $0.027$ | 0.028 |
|  | $2.044 \pm 0.006$ | $2.050 \pm 0.012$ | $2.112 \pm 0.010$ | $2.132 \pm 0.009$ | $2.186 \pm 0.021$ | $2.209 \pm 0.011$ | 2.248 | -0.513 | 0.027 | $\begin{gathered} \pm 0.002 \\ ( \pm 0.003) \\ \hline \end{gathered}$ | $\pm 0.003$ |
| 343 | $3.000 \pm 0.018$ | $3.025 \pm 0.009$ | $3.070 \pm 0.008$ | $3.089 \pm 0.015$ | $3.117 \pm 0.015$ | $3.148 \pm 0.018$ | 3.179 | -0.423 | 0.031 |  |  |
|  | $1.949 \pm 0.004$ | $1.955 \pm 0.005$ | $1.968 \pm 0.003$ | $1.998 \pm 0.004$ | $2.020 \pm 0.002$ | $2.030 \pm 0.009$ | 2.050 | -0.258 | 0.031 | $\begin{gathered} \pm 0.001 \\ ( \pm 0.002) \end{gathered}$ | $\pm 0.002$ |
| 333 | $3.247 \pm 0.018$ | $3.234 \pm 0.018$ | $3.243 \pm 0.015$ | $3.241 \pm 0.010$ | $3.247 \pm 0.009$ | $3.223 \pm 0.013$ | 3.231 | 0.034 | 0.037 | 0.036 |  |
|  | $3.080 \pm 0.009$ | $3.044 \pm 0.006$ | $3.082 \pm 0.005$ | $3.127 \pm 0.009$ | $3.113 \pm 0.008$ | $3.134 \pm 0.014$ | 3.149 | -0.213 | 0.034 | $\begin{gathered} \pm 0.003 \\ ( \pm 0.004) \end{gathered}$ | $\pm 0.002$ |
| 323 | $3.208 \pm 0.011$ | $3.190 \pm 0.008$ | $3.133 \pm 0.010$ | $3.134 \pm 0.011$ | $3.092 \pm 0.008$ | $3.093 \pm 0.006$ | 3.055 | 0.355 | 0.042 | 0.043 |  |
|  | $3.357 \pm 0.010$ | $3.289 \pm 0.014$ | $3.275 \pm 0.005$ | $3.233 \pm 0.004$ | $3.226 \pm 0.016$ | $3.160 \pm 0.001$ | 3.135 | 0.496 | 0.044 | $\begin{gathered} \pm 0.002 \\ ( \pm 0.004) \end{gathered}$ | $\pm 0.001$ |
| 313 | $3.245 \pm 0.018$ | $3.185 \pm 0.013$ | $3.100 \pm 0.015$ | $3.022 \pm 0.012$ | $2.995 \pm 0.012$ | $2.915 \pm 0.011$ | 2.848 | 0.935 | 0.052 | 0.052 |  |
|  | $2.162 \pm 0.031$ | $2.134 \pm 0.010$ | $2.060 \pm 0.014$ | $2.029 \pm 0.018$ | $1.992 \pm 0.010$ | $1.925 \pm 0.018$ | 1.896 | 0.612 | 0.052 | $\begin{gathered} \pm 0.003 \\ ( \pm 0.005) \end{gathered}$ | $\pm 0.001$ |

a. Errors are $2 \sigma$ for the regression only.; b. Errors are those at $95 \%$ confidence level for the regression only.; c. Number in parenthesis represents both errors at $95 \%$ confidence level for the regression and potential systematic bias ( $\pm 4 \%$ ).


Figure S3. Headspace GC-MS measurements for six series of test samples containing water ( $V_{i}$ in $\mathbf{c m}^{\mathbf{3}}$ ) to which a $\mathbf{C H}_{\mathbf{2}} \mathrm{F}_{\mathbf{2}}$-air mixture was added ( $v_{j}$ in $\mathrm{cm}^{3}$ ) at $\mathbf{3 1 3} \mathrm{K}$. (a) Plot of peak area $\left(S_{i j}\right)$ versus $v_{j}$ for test samples containing volume $V_{i}$ of water. Slope $\left(L_{i}\right)$ was obtained by linear fitting of the data to Eq. (8) for samples of the same $V_{i}$. (b) Plot of $L_{i}^{-1}$ versus $V_{i} / V_{0}$ fitted to Eq. (10).


Figure S4. Plot of $L_{i}$ versus $V_{i} / V_{0}$ for the PRV-HS measurements at each temperature. Bold curves represent the simultaneous fitting of the two datasets at each temperature by Eq. (11).

S4. Determination of salting-out effects in artificial seawater


Figure S5. Plots of values of $F /\left(k_{1} R T V\right)$ against $F$ at each temperature for $0.35 \mathbf{d m}^{3}$ of a-seawater at $\mathbf{4 . 4 5 2 \%}$. Error bars represent $2 \sigma$ due to errors of values of $k_{1}$ as described in Sect. S2. Grey symbols represent the data excluded for calculating the average.


Figure S6. Plots of values of $F /\left(k_{1} R T V\right)$ against $F$ at each temperature for $0.35 \mathrm{dm}^{\mathbf{3}}$ of a-seawater at $\mathbf{8 . 9 2 1 \%}$. Error bars represent $\mathbf{2 \sigma}$ due to errors of values of $k_{1}$ as described in Sect. S2. Grey symbols represent the data excluded for calculating the average.


Figure S7. Plots of values of $F /\left(k_{1} R T V\right)$ against $F$ at each temperature for $0.35 \mathrm{dm}^{3}$ of a-seawater at $\mathbf{2 1 . 5 2 0 \%}$. Error bars represent $2 \sigma$ due to errors of values of $k_{1}$ as described in Sect. S2. Grey symbols represent the data excluded for calculating the average.


Figure S8. Plots of values of $F /\left(k_{1} R T V\right)$ against $F$ at each temperature for $0.35 \mathrm{dm}^{\mathbf{3}}$ of a-seawater at $51.534 \%$. Error bars represent $2 \sigma$ due to errors of values of $k_{1}$ as described in Sect. S2. Grey symbols represent the data excluded for calculating the average.


Figure S9. $\log$-log plots for $\ln \left(K_{H}(T) / K_{\text {eq }}{ }^{s}(T)\right)$ vs. salinity in a-seawater at each temperature. Bold lines represent the fitting obtained by a liner regression. Errors are those at $\mathbf{9 5 \%}$ confidence level for the regression only.

5 Table S2. Values of $\boldsymbol{k}_{s}$ (Eq. (17)) and comparison of values of $K_{\text {eq }}{ }^{S}$ calculated at each temperature by Eq. (17) with those by Eq. (22).

| Temperature $\left({ }^{\circ} \mathrm{C}\right)$ | $\begin{gathered} k_{\mathrm{s}} \\ \left(\%{ }^{-1}\right) \end{gathered}$ | [ $K_{\text {eq }}{ }^{\text {S }}$ from Eq. (17) $] /\left[K_{\text {eq }}{ }^{\text {S }}\right.$ from Eq. (22) $]$ |  |  | $\left[K_{\text {eq }}{ }^{\text {S }}\right.$ at $\left.30 \% 0\right] /\left[K_{\text {eq }}{ }^{\text {S }}\right.$ at 40\%o $]$ |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | at $30 \%$ | at 35\% | at 40\% | Eq. (17) | Eq. (22) |
| 3.0 | 0.00811 | 1.027 | 1.008 | 0.988 | 1.084 | 1.043 |
| 5.8 | 0.00785 | 1.033 | 1.014 | 0.995 | 1.082 | 1.042 |
| 10.5 | 0.00768 | 1.033 | 1.016 | 0.997 | 1.080 | 1.042 |
| 15.5 | 0.00718 | 1.044 | 1.028 | 1.012 | 1.074 | 1.041 |
| 20.3 | 0.00728 | 1.037 | 1.020 | 1.003 | 1.076 | 1.040 |
| 25.0 | 0.00704 | 1.040 | 1.024 | 1.008 | 1.073 | 1.039 |
| 29.9 | 0.00731 | 1.027 | 1.010 | 0.992 | 1.076 | 1.039 |
| 34.8 | 0.00713 | 1.029 | 1.012 | 0.995 | 1.074 | 1.038 |
| 39.5 | 0.00709 | 1.026 | 1.010 | 0.992 | 1.073 | 1.038 |

## S5. Discussion of potential reason for this salting-out effect of $\mathbf{C H}_{2} \mathbf{F}_{\mathbf{2}}$ solubility in a-seawater (deviation from Sechenov relationship)

The reason that the salting-out effect of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ solubility in a-seawater depends on $S^{0.5}$ is not clear. Specific properties of $\mathrm{CH}_{2} \mathrm{~F}_{2}$-small molecular volume, which results in small work of cavity creation (Graziano, 2004; 2008), and large solute- solvent attractive potential energy in water and a-seawater- may cause deviation from Sechenov relationship. This possibility may be discussed here.

I calculate Ben-Naim standard Gibbs energy $\Delta G^{*}$, enthalpy $\Delta H^{*}$, and entropy $\Delta S^{*}$ changes for dissolution of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ in water because these values correspond to the values for the transfer from a fixed position in the gas phase to a fixed position in water. Values of $\Delta G^{*}, \Delta H^{*}$, and $\Delta S^{*}$ are calculated on the basis of the Ostwald solubility coefficient, $L(T)$, as follows.

$$
\begin{align*}
& \ln (L(T))=\ln \left(R T K_{\mathrm{eq}}{ }^{S}(T)\right)  \tag{B1}\\
& \Delta G^{\cdot}=R^{\prime} T \ln (L(T))  \tag{B2}\\
& \Delta H^{\cdot}=-\frac{\partial}{\partial(1 / T)}\left(\frac{\Delta G^{\cdot}}{T}\right)  \tag{B3}\\
& \Delta S^{*}=\frac{\Delta H^{\prime}-\Delta G^{\cdot}}{T} \tag{B4}
\end{align*}
$$

where both $R$ and $R^{\prime}$ represent gas constant but their units are different: $R=0.0821 \mathrm{in} \mathrm{atm} \mathrm{dm}^{3} \mathrm{~K}^{-1} \mathrm{~mol}^{-1} ; R^{\prime}=8.314 \mathrm{in} \mathrm{J} \mathrm{K}$ $\mathrm{mol}^{-1}$.

Combining Eqs. (B1), (B2), (B3), and (B4) with Eqs. (14) and (15), $\Delta G^{*}\left(\mathrm{~kJ} \mathrm{~mol}^{-1}\right), \Delta H^{*}\left(\mathrm{~kJ} \mathrm{~mol}{ }^{-1}\right)$, and $\Delta S^{*}\left(\mathrm{~J} \mathrm{~mol}^{-1}\right.$ $\mathrm{K}^{-1}$ ) are represented by $\Delta G_{\text {sol }}$ and $\Delta H_{\text {sol }}$ as follows:

$$
\begin{align*}
& \Delta G^{\prime}=\Delta G_{\text {sol }}+R^{\prime} T \ln (R T)  \tag{B5}\\
& \Delta H^{\prime}=\Delta H_{\text {sol }}+R^{\prime} T  \tag{B6}\\
& \Delta S^{\cdot}=\frac{\Delta H_{\text {sol }}-\Delta G_{\text {sol }}}{T}+R^{\prime}-R^{\prime} \ln (R T) \tag{B7}
\end{align*}
$$

Values of $\Delta G^{*}, \Delta H^{*}$, and $\Delta S^{*}$ calculated at 298 K are listed in Table S3. Table S3 also lists values of $\Delta G^{*}, \Delta H^{*}$, and $\Delta S^{\circ}$ reported for $\mathrm{CH}_{3} \mathrm{~F}$ and $\mathrm{C}_{2} \mathrm{H}_{6}$ (Graziano, 2004) and $\mathrm{CH}_{4}$ (Graziano, 2008) at 298 K . The chemicals, which having a methyl group, in Table S 3 are classified into two groups $\left(\mathrm{CH}_{2} \mathrm{~F}_{2}\right.$ and $\mathrm{CH}_{3} \mathrm{~F} ; \mathrm{CH}_{4}$ and $\left.\mathrm{C}_{2} \mathrm{H}_{6}\right)$ according to $\Delta G^{*}$.

Table S3 lists values of $\Delta G_{\mathrm{c}}, E_{a}$ and $\Delta H^{h}$ deduced using a scaled particle theory (Granziano, 2004; 2008). $\Delta G_{c}$ is the work of cavity creation to insert a solute in a solvent. $E_{a}$ is a solute-solvent attractive potential energy and accounts for the solute-solvent interactions consisting of dispersion, dipole-induced dipole, and dipole-dipole contributions. $\Delta H^{h}$ is enthalpy of solvent molecules reorganization caused by solute insertion. The solvent reorganization mainly involves a rearrangement of H-bonds.
$\Delta G_{\mathrm{c}}$ is entropic in nature in all liquids, being a measure of the excluded volume effect due to a reduction in the spatial configurations accessible to liquid molecules upon cavity creation. Hence, $\mathrm{C}_{2} \mathrm{H}_{6}$ has larger value of $\Delta G_{\mathrm{c}}$ than $\mathrm{CH}_{3} \mathrm{~F}$ and $\mathrm{CH}_{4}$. $\Delta G_{\mathrm{c}}, E_{a}$, and $\Delta H^{h}$ are related to $\Delta G^{*}$ and $\Delta H^{*}$ as follows (Graziano, 2008):

$$
\begin{align*}
& \Delta G^{\cdot}=\Delta G_{\mathrm{c}}+E_{a}  \tag{B8}\\
& \Delta H^{\cdot}=E_{a}+\Delta H^{h} \tag{B9}
\end{align*}
$$

Table S3 thus suggests that smaller value of $\Delta G^{*}$ of $\mathrm{CH}_{3} \mathrm{~F}$ than $\mathrm{CH}_{4}$ is due to large solute-solvent attractive potential energy $\left(-E_{a}\right)$ of $\mathrm{CH}_{3} \mathrm{~F}$.

Table S3. Ben-Naim standard hydration Gibbs energy $\Delta G^{*}$, enthalpy $\Delta H^{*}$, and entropy $\Delta S^{*}$ changes for dissolution of $\mathbf{C H}_{2} \mathrm{~F}_{2}$ at 298 $K$ determined here and the corresponding values and values of $\Delta G_{c}, E_{a}$ and $\Delta H^{h}$ reported for $\mathrm{CH}_{3} F$ and $C_{2} H_{6}$ (Granziano, 2004) and $\mathrm{CH}_{4}$ (Graziano, 2008).

|  | $\Delta G^{*}$ <br> $\left(\mathrm{~kJ} \mathrm{~mol}^{-1}\right)$ | $\Delta H^{\cdot}$ <br> $\left(\mathrm{kJ} \mathrm{mol}^{-1}\right)$ | $\Delta S^{*}$ <br> $\left(\mathrm{~J} \mathrm{~K}^{-1} \mathrm{~mol}^{-1}\right)$ | $\Delta G_{\mathrm{c}}$ <br> $\left(\mathrm{kJ} \mathrm{mol}^{-1}\right)$ | $E_{a}$ <br> $\left(\mathrm{~kJ} \mathrm{~mol}^{-1}\right)$ | $\Delta H^{h}$ <br> $\left(\mathrm{~kJ} \mathrm{~mol}^{-1}\right)$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\mathrm{CH}_{2} \mathrm{~F}_{2}$ | -1.1 | -14.7 | -45.4 |  |  |  |
| $\mathrm{CH}_{3} \mathrm{~F}$ | -0.9 | -15.8 | -50.0 | 23.3 | -24.3 | 8.5 |
| $\mathrm{CH}_{4}$ | 8.4 | -10.9 | -64.7 | 22.9 | -14.5 | 3.7 |
| $\mathrm{C}_{2} \mathrm{H}_{6}$ | 7.7 | -17.5 | -84.5 | 28.4 | -20.7 | 3.2 |

Graziano (2008) definitively explained the salting-out of $\mathrm{CH}_{4}$ by sodium chloride at molecular level on the basis of a scaled particle theory. He explained that $\Delta G_{\mathrm{c}}$ increase was linearly related to the increase in the volume packing density of the solutions $\left(\xi_{3}\right)$ with adding NaCl . Such an increase of $\Delta G_{\mathrm{c}}$ is probably the case for salting-out of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ by a-seawater observed in this study. He also explained that $E_{a}$ was linearly related to the increase in $\xi_{3}$ assuming that a fraction of the dipole-induced dipole attractions could be taken into account by the parameterization of the dispersion contribution.

I think the possibility that $E_{\mathrm{a}}$ may be nonlinearly related to the increase in $\xi_{3}$ because of dipole-dipole interaction between $\mathrm{CH}_{2} \mathrm{~F}_{2}$ and solvents. Temperature dependence in Eq. (22) suggests that salting-out effect of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ by a-seawater is enthalpic. Eqs. (22) and (B9) thus suggests that the salting-out of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ is mostly related to change in $E_{a} . \mathrm{CH}_{2} \mathrm{~F}_{2}$ has relatively small value of $\Delta G_{c}$ because of its small molecular volume compared to other chemicals such as $\mathrm{C}_{2} \mathrm{H}_{6}$. Accordingly, $\Delta G^{\cdot}$, that is, solubility of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ would depend on $E_{a}$ rather than $\Delta G_{\mathrm{c}}$. Therefore, I think that specific properties of $\mathrm{CH}_{2} \mathrm{~F}_{2}-$ small molecular volume, which results in small work of cavity creation (Graziano, 2004; 2008), and large solute-solvent attractive potential energy in water and a-seawater- may cause deviation from Sechenov relationship.

## References

Graziano, G.: Case study of enthalpy-entropy noncompensation. Journal of Chemical Physics, 120, 4467-4471, doi: 10.1063/1.1644094, 2004.

Graziano, G.: Salting out of methane by sodium chloride: A scaled particle theory study. Journal of Chemical Physics, 129, 084506, doi: 10.1063/1.2972979, 2008.

S6. Estimated results (Sect. 3.3) for monthly amount of $\mathbf{C H}_{2} \mathbf{F}_{2}$ dissolved in the ocean mixed layer at solubility equilibrium with the atmospheric $\mathrm{CH}_{2} \mathrm{~F}_{2}(1 \mathrm{patm})$ and the depth distribution of the $\mathrm{CH}_{2} \mathrm{~F}_{2}$ dissolved in each semihemisphere

Table S4. Monthly amount of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ dissolved in the ocean mixed layer at solubility equilibrium with the atmospheric $\mathbf{C H}_{2} \mathrm{~F}_{2}$ (partial pressure, 1 patm) and the depth distribution of the $\mathbf{C H}_{2} \mathbf{F}_{2}$ dissolved in the southern semi-hemisphere ( $90^{\circ} \mathbf{S - 3 0} \mathbf{~ S}$ ).

|  | Amount $\left(\mathrm{Gg} \mathrm{patm}^{-1}\right)$ | Distribution of the amount of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ dissolved in the ocean mixed layer with respect to the ocean mixed layer depth (\%) |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | 10-100 m | 100-200 m | 200-300 m | 300-400 m | 400-500 m | 500-600m |
| January | 0.0169 | 94.9 | 2.9 | 1.0 | 0.5 | 0.3 | 0.3 |
| February | 0.0201 | 92.1 | 3.6 | 2.9 | 1.0 | 0.3 | 0.0 |
| March | 0.0255 | 87.8 | 9.2 | 1.7 | 0.7 | 0.2 | 0.4 |
| April | 0.0338 | 66.5 | 31.8 | 1.1 | 0.2 | 0.1 | 0.2 |
| May | 0.0409 | 48.5 | 48.1 | 2.2 | 0.8 | 0.3 | 0.0 |
| June | 0.0510 | 26.8 | 62.7 | 8.0 | 1.7 | 0.8 | 0.1 |
| July | 0.0571 | 14.1 | 69.3 | 12.2 | 3.3 | 0.9 | 0.1 |
| August | 0.0640 | 8.5 | 65.8 | 17.0 | 6.2 | 2.3 | 0.2 |
| September | 0.0609 | 13.5 | 61.0 | 14.6 | 8.2 | 2.7 | 0.0 |
| October | 0.0504 | 24.7 | 58.6 | 12.1 | 2.9 | 1.4 | 0.3 |
| November | 0.0335 | 60.4 | 30.5 | 4.6 | 2.2 | 2.3 | 0.1 |
| December | 0.0196 | 95.1 | 4.3 | 0.4 | 0.2 | 0.0 | 0.0 |

Table S5. Monthly amount of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ dissolved in the ocean mixed layer at solubility equilibrium with the atmospheric $\mathbf{C H}_{2} \mathrm{~F}_{2}$ (partial pressure, $1 \mathbf{p a t m}$ ) and the depth distribution of the $\mathrm{CH}_{2} \mathrm{~F}_{2}$ dissolved in the southern semi-hemisphere ( $\mathbf{3 0}{ }^{\circ} \mathrm{S}-\mathbf{0}^{\circ} \mathrm{S}$ ).

|  | Amount <br> $\left(\mathrm{Gg} \mathrm{patm}^{-1}\right)$ | Distribution of the amount of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ dissolved in the ocean mixed layer <br> with respect to the ocean mixed layer depth $(\%)$ |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | $10-100 \mathrm{~m}$ | $100-200 \mathrm{~m}$ | $200-300 \mathrm{~m}$ | $300-400 \mathrm{~m}$ | $400-500 \mathrm{~m}$ | $500-600 \mathrm{~m}$ |
| January | 0.0084 | 99.6 | 0.4 | 0 | 0 | 0 | 0 |
| February | 0.0084 | 99.7 | 0.3 | 0 | 0 | 0 | 0 |
| March | 0.0089 | 100.0 | 0 | 0 | 0 | 0 | 0 |
| April | 0.0106 | 100.0 | 0 | 0 | 0 | 0 | 0 |
| May | 0.0131 | 100.0 | 0 | 0 | 0 | 0 | 0 |
| June | 0.0163 | 97.1 | 2.9 | 0 | 0 | 0 | 0 |
| July | 0.0189 | 80.1 | 19.9 | 0 | 0 | 0 | 0 |
| August | 0.0193 | 73.1 | 26.9 | 0 | 0 | 0 | 0 |
| September | 0.0165 | 82.2 | 17.8 | 0 | 0 | 0 | 0 |
| October | 0.0124 | 94.6 | 5.4 | 0.1 | 0 | 0 | 0 |
| November | 0.0097 | 99.9 | 100.0 | 0 | 0 | 0 | 0 |
| December | 0.0087 |  |  | 0 | 0 | 0 |  |

Table S6. Monthly amount of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ dissolved in the ocean mixed layer at solubility equilibrium with the atmospheric $\mathbf{C H}_{2} \mathrm{~F}_{\mathbf{2}}$ (partial pressure, $1 \mathbf{p a t m}$ ) and the depth distribution of the $\mathbf{C H}_{2} \mathbf{F}_{2}$ dissolved in the northern semi-hemisphere ( $0^{\circ} \mathbf{N - 3 0}{ }^{\circ} \mathbf{N}$ ).

|  | Amount <br> $\left(\mathrm{Gg} \mathrm{patm}^{-1}\right)$ | Distribution of the amount of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ dissolved in the ocean mixed layer <br> with respect to the ocean mixed layer depth (\%) |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | $10-100 \mathrm{~m}$ | $100-200 \mathrm{~m}$ | $200-300 \mathrm{~m}$ | $300-400 \mathrm{~m}$ | $400-500 \mathrm{~m}$ | $500-600 \mathrm{~m}$ |
| January | 0.0132 | 96.4 | 3.6 | 0 | 0 | 0 | 0 |
| February | 0.0126 | 95.9 | 4.1 | 0 | 0 | 0 | 0 |
| March | 0.0107 | 98.7 | 1.3 | 0 | 0 | 0 | 0 |
| April | 0.0087 | 99.8 | 0.2 | 0 | 0 | 0 | 0 |
| May | 0.0079 | 100.0 | 0 | 0 | 0 | 0 | 0 |
| June | 0.0080 | 100.0 | 0 | 0 | 0 | 0 | 0 |
| July | 0.0084 | 100.0 | 0 | 0 | 0 | 0 | 0 |
| August | 0.0082 | 100.0 | 0 | 0 | 0 | 0 | 0 |
| September | 0.0080 | 100.0 | 0 | 0 | 0 | 0 | 0 |
| October | 0.0086 | 100.0 | 0 | 0 | 0 | 0 | 0 |
| November | 0.0100 | 100.0 | 0 | 0 | 0 | 0 | 0 |
| December | 0.0118 | 100.0 |  | 0 | 0 | 0 |  |

5 Table S7. Monthly amount of $\mathbf{C H}_{2} \mathbf{F}_{2}$ dissolved in the ocean mixed layer at solubility equilibrium with the atmospheric $\mathbf{C H}_{2} \mathbf{F}_{2}$ (partial pressure, $1 \mathbf{p a t m}$ ) and the depth distribution of the $\mathrm{CH}_{2} \mathrm{~F}_{2}$ dissolved in the northern semi-hemisphere ( $\mathbf{3 0 ^ { \circ }} \mathbf{N - 9 0}{ }^{\circ} \mathbf{N}$ ).

|  | Amount <br> $\left(\right.$ Gg patm $\left.^{-1}\right)$ | Distribution of the amount of $\mathrm{CH}_{2} \mathrm{~F}_{2}$ dissolved in the ocean mixed layer <br> with respect to the ocean mixed layer depth $(\%)$ |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  | $10-100 \mathrm{~m}$ | $100-200 \mathrm{~m}$ | $200-300 \mathrm{~m}$ | $300-400 \mathrm{~m}$ | $400-500 \mathrm{~m}$ | $500-600 \mathrm{~m}$ |
| January | 0.0205 | 41.3 | 50.1 | 7.0 | 1.4 | 0.2 | 0.0 |
| February | 0.0225 | 34.5 | 55.3 | 7.1 | 2.3 | 0.6 | 0.2 |
| March | 0.0208 | 49.7 | 42.3 | 4.9 | 1.7 | 0.7 | 0.6 |
| April | 0.0147 | 79.7 | 17.6 | 1.7 | 0.4 | 0.0 | 0.6 |
| May | 0.0081 | 90.1 | 9.9 | 0 | 0 | 0 | 0 |
| June | 0.0055 | 97.7 | 2.3 | 0 | 0 | 0 | 0 |
| July | 0.0045 | 96.6 | 3.4 | 0 | 0 | 0 | 0 |
| August | 0.0048 | 94.4 | 5.6 | 0 | 0 | 0 | 0 |
| September | 0.0059 | 97.7 | 2.3 | 0 | 0 | 0 | 0 |
| October | 0.0084 | 99.6 | 0.4 | 0 | 0 | 0 | 0 |
| November | 0.0121 | 89.6 | 10.4 | 0.1 | 0 | 0 | 0 |
| December | 0.0163 | 71.0 | 26.1 | 2.9 | 0 | 0 |  |

