



### Supplement of

### **Optical properties of atmospheric fine particles near Beijing during the HOPE-J**<sup>3</sup>**A campaign**

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### **1** S1 System performance

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Fig. S1 Example data of the transmitted intensity measured with the CCD spectrometer and the scattering signal measured with the PMT of the cavity-enhanced abldoemeter during the experimental period. The cavity was flushed with particle-free air every hour to acquire the  $I_0(\lambda)$ spectrum. No obvious drift in the LED light intensity was observed even after 6 hours of measurement, indicating the high stability of the instrument under these operating conditions.

#### 1 S2 Chemical apportionment of aerosol optical properties

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Chemical apportionment of light extinction of  $PM_{1,0}$  was determined with a revised IMPROVE 3 (Interagency Monitoring of Protected Visual Environments) algorithm (Pitchford et al., 2007). 4 Although the IMPROVE algorithm is a simplified predictor of extinction, it is nevertheless a useful 5 tool to estimate the contribution of different particle components to haze levels (Pitchford et al., 6 2007). Light extinction at  $\lambda = 550$  nm can be estimated by multiplying the mass concentrations by 7 component-specific mass extinction efficiencies (MEEs) of each of seven major components: sulfate 8 9 (assumed to be ammonium sulfate), nitrate (assumed to be ammonium nitrate), organic mass (OM, based on the measured mass concentration of organic carbon, OC), elemental carbon (EC), fine soil, 10 sea salt (chlorine, Cl), and coarse mass (the differences between PM<sub>10</sub> and PM<sub>2.5</sub> mass concentration). 11 It can be expressed by the following (Pitchford et al., 2007): 12

$$\begin{aligned} \alpha_{ext,550nm} &\approx 2.2 \times f_s(\text{RH}) \times [\text{Small Sulfate}] + 4.8 \times f_L(\text{RH}) \times [\text{Large Sulfate}] \\ &+ 2.4 \times f_s(\text{RH}) \times [\text{Small Nitrate}] + 5.1 \times f_L(\text{RH}) \times [\text{Large Nitrate}] \\ &+ 2.8 \times [\text{Small Organic Mass}] + 6.1 \times [\text{Large Organic Mass}] \\ &+ 10 \times [\text{Elemental Carbon}] + 1 \times [\text{Fine Soil}] \\ &+ 1.7 \times f_{ss}(\text{RH}) \times [\text{Sea Salt}] + 0.6 \times [\text{Coarse Mass}] \\ &+ \text{Rayleigh scattering (site specific}) + 0.33 \times [\text{NO}_2(\text{ppb})] \end{aligned}$$
(S1)

where f(RH) is the water growth factor of inorganic components,  $f_S(RH)$  and  $f_L(RH)$  are the water growth factors for the small and large particle size modes for sulfate and nitrate, respectively, and  $f_{SS}(RH)$  is the hygroscopic growth factor for sea salt.

The large and small parts are defined by the IMPROVE formula as (Pitchford et al., 2007; Cao
et al., 2012):

19  $[Large X] = [Total X]^2/20$ , for  $[Total X] < 20 \ \mu g \ m^{-3}$ 

20 [Large X] = [Total X], for [Total X] 
$$\ge 20 \ \mu g \ m^{-3}$$

21 [Small X] = [Total X] - [Large X]

where X = sulfate, nitrate or OM. The concentration of ammonium sulfate ([(NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub>]) was 1.375 times the sulfate concentration ([SO<sub>4</sub><sup>2-</sup>]), and the ammonium nitrate ([NH<sub>4</sub>NO<sub>3</sub>]) was 1.29 times the nitrate concentrations ([NO<sub>3</sub><sup>-</sup>]). The OM concentration was estimated by multiplying the reported OC concentration by a factor of 1.6 (Turpin and Lim, 2001). The sea salt mass concentration was estimated by multiplying the Cl<sup>-</sup> mass concentration by a factor of 1.8. The ammonium cation was not used directly in the IMPROVE algorithm. It was assumed to be fully neutralized by  $SO_4^{2-}$  and  $NO_3^{-}$  and treated as ammonium sulfate ((NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub>) and ammonium nitrate(NH<sub>4</sub>NO<sub>3</sub>), respectively.

The revised IMPROVE algorithm above was parameterized for atmospheric aerosol extinction at  $\lambda = 550$  nm. In this study, the optical properties of PM<sub>1.0</sub> were measured at  $\lambda = 470$  nm. The IMPROVE algorithm needs to be refined to better represent the chemical apportionment of light extinction for PM<sub>1.0</sub> particles. Since dry PM<sub>1.0</sub> aerosols were measured, the coarse mass and the hygroscopic increase of inorganic component were ignored for the inputs of the IMPROVE algorithm. The fine soil component was also left out in apportioning the extinction coefficient due to its small fraction of PM<sub>1.0</sub> particles during wintertime in Beijing (Sun et al., 2014).

10 Under the assumptions of the IMPROVE algorithm (in which particles are treated as separate 11 entities), the difference in the wavelength will only affect the dry mass extinction efficiency terms. 12 We modified the MEE terms in equation (S1) of each individual particle components with a scaling 13 factor so as to be directly comparable to measurements at  $\lambda = 470$  nm. The scaling factor was the 14 ratio of the MEEs of each species at  $\lambda = 470$  nm and  $\lambda = 550$  nm calculated with Mie theory from the 15 literature reported complex refractive index (Table S1) and the measured mean number size 16 distribution.

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Table S1 Densities and complex refractive indexes of different types of aerosols. (adapted from
Cheng et al., 2008)

Chemical Species	Density (g cm <sup>-3</sup> )	Complex Refractive Index			
$(NH_4)_2SO_4$	1.748	1.54 - 10 <sup>-7</sup> i			
NH <sub>4</sub> NO <sub>3</sub>	1.725	1.54 - 10 <sup>-7</sup> i			
OM	1.4	1.55 - 0.001 i			
Sea salt	2.0	1.54 - 10 <sup>-7</sup> i			
EC	1.5	1.80 - 0.54 i			

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Under the assumption that particles were completely separated (externally mixed) and can be adequately estimated as separate terms of each single chemical component, *j*, the extinction coefficient of component *j* can be calculated from (Bohren and Huffman, 1998; Petersson et al., 2004; Hand and Malm, 2007):

$$\alpha_{ext,j} = \int_{0}^{\infty} \frac{\pi}{4} D_p^2 Q_{ext}(m_j, D_p, \lambda) N(D_p) dD_p$$
(S2)

where  $N(D_p)$  is the number of particles with mean diameter  $D_p$  per unit volume in the size bin  $dD_p$ , 1 and m = n + ik is the complex refractive index.  $Q_{ext}$  is the extinction efficiency and can be calculated 2 3 with Mie theory for spherical particles.

The mass extinction efficiency(MEE) is a function of the Mie extinction efficiency, particle 4 diameter and species density ( $\rho_i$ ) (Hand and Malm, 2007): 5

6 
$$MEE_{j} = \frac{3}{2} \frac{Q_{ext}(m_{j}, D_{p}, \lambda)}{\rho_{j} D_{p}}$$
(S3)

7 If the mass size distribution of each species *j* does not vary with total mass concentration, the light extinction coefficient of species *j* can be calculated with : 8

9 
$$\alpha_{ext,i} = MEE_iM_i$$
 (S4)

10 where  $M_i$  is the mass concentration per unit volume.

If the mass size distribution of each species varies with the total mass concentration, the 11 following equation can be used for the MEE calculation (Cheng et al., 2008; Cheng et al., 2015). 12

13  
$$MEE_{j} = \frac{\sum_{bin=1}^{D_{bin}} MEE(j, D_{bin})M_{j,bin}}{\sum_{bin=1}^{D_{bin}} M_{j,bin}}$$
(S5)

The total extinction is a linear combination of all the species : 14

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 $\alpha_{ext} = \sum_{i} MEE_{i}M_{j}$ Since the size-segregated mass concentration of each chemical composition was not available in this work. Equation S5 was not used. We used equation (S3) (with literature reported complex refractive index, table S1, and the measured mean size distribution) for the calculation of the MEE

(S6)

values at  $\lambda = 470$  nm and  $\lambda = 550$  nm. 19

Under the assumptions of the IMPROVE algorithm (in which particles are treated as separate 20 entities), the difference in the wavelength will only affect the dry mass extinction efficiency (MEE) 21 terms. We modified the MEE terms (equation S1) of each individual particle components with a 22 scaling factor so as to be directly comparable to measurements at  $\lambda = 470$  nm. The scaling factor 23

24 (Scaling factor<sub>j</sub> = 
$$\frac{MEE_{j,\lambda=470nm}}{MEE_{j,\lambda=550nm}}$$
) was the ratio of the MEEs of each species at  $\lambda = 470$ 

nm and  $\lambda = 550$  nm, calculated with equation (S3) under different pollution days was shown in Fig. 25

S2. As an average result, the calculated dry mass extinction efficiencies of inorganic mass (including
 sulfate, nitrate and sea salt), organic mass and element carbon at λ = 470 nm were 1.31, 1.30 and
 1.08 times larger than that at λ = 550 nm.

- 6.0 6.0 Ammonium nitra Clear Slightly Po (a) (b) Mass extinction efficiency (m<sup>2</sup> g<sup>-1</sup>) Mass extinction efficiency (m<sup>2</sup> g<sup>-1</sup>) 5.5 5.5 5.0 5.0 4.5 4.5 4.0 4.0 3.5 3.5 450 460 470 480 490 500 510 520 530 540 550 450 460 470 480 490 500 510 520 530 540 550 Wavlength (nm) Wavlength (nm) 7.5 5.5 (c) Organic Mass (d) Sea Salt Mass extinction efficiency (m<sup>2</sup> g<sup>-1</sup>) Clear Clea Mass extinction efficiency  $(m^2 g^{-1})$ Slightly Polluted Polluted Slightly Polluted 7.0 5.0 6.5 45 6.0 4.0 5.5 3.5 5.0 4.5 450 460 470 480 490 500 510 3.0 530 540 520 550 450 460 470 480 490 500 510 . 520 530 540 550 Wavlength (nm) Wavlength (nm) 9.2 (e) nent Carbon Ck Mass extinction efficiency (m<sup>2</sup> g<sup>-1</sup>) iahtiv P 9.0 8.8 8.6 8.4 8.2 8.0 460 470 540 450 480 490 500 510 520 530 550 Wavlength (nm)
- 4



Fig. S2 Calculated wavelength dependent mass extinction efficiencies of inorganic mass (including
sulfate, nitrate and sea salt), organic mass and element carbon.

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10 The dry mass extinction efficiencies of inorganic mass (including sulfate, nitrate and sea salt), 11 organic mass, and elemental carbon for the input of the IMPROVE formula at  $\lambda = 470$  nm should 12 accordingly be scaled by factors of 1.31, 1.30 and 1.08, respectively. The modified IMPROVE 13 function for PM1.0 extinction at  $\lambda = 470$  nm can be rewritten as following:

$$\alpha_{ext,470nm,PM_{1.0}} \approx 2.88 \times [Small Sulfate] + 6.29 \times [Large Sulfate] +3.14 \times [Small Nitrate] + 6.68 \times [Large Nitrate] +3.64 \times [Small Organic Mass] + 7.93 \times [Large Organic Mass] +10.8 \times [Elemental Carbon] + 2.23 \times [Sea Salt]$$
(S7)

The  $PM_{1.0}$  mass concentration can be reconstructed as the sum of its major chemical components (Pitchford et al., 2007):

4 Reconstructed mass\_
$$PM_{1.0} = [(NH_4)_2SO_4] + [NH_4NO_3] + [SS] + [OM] + [EC]$$
  
=  $1.375 \times [SO_4^{2-}] + 1.29 \times [NO_3^{-}] + 1.8 \times [Cl^{-}] + 1.6 \times [OC] + [EC]$  (S8)

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Studies in other Chinese megacities, such as Shanghai (Cheng et al., 2015) and Guangzhou (Tao 6 7 et al., 2014), and data from US monitoring sites show that the revised IMPROVE algorithm 8 underestimates the PM2.5 extinction under high aerosol loading but overestimates the values under low aerosol loading. The underestimation and overestimation ratios in different studies ranged from 9 -11 to -26 % and +25 to +54 %, respectively (Cheng et al., 2015). In this work,  $\alpha_{ep,470}$  of PM<sub>1.0</sub> 10 particles was reconstructed using the modified IMPROVE algorithm based on the measured 11 concentrations of each composition (Fig. S3a), which correlated well with the measured  $\alpha_{ep,470}$  ( $R^2 =$ 12 0.96) during this campaign (Fig. S3c). With modified IMPROVE function (Eq. S7), the agreement 13 between the measured and calculated  $PM_{1,0}$  extinctions is good (with a slope of  $1.04 \pm 0.04$ ) when 14 the measured extinction coefficient is lower than 300 Mm<sup>-1</sup> (as shown in the insert of Fig. S3c). 15 When the observed extinction coefficients are larger than 300 Mm<sup>-1</sup>, the reconstructed values of the 16 modified IMPROVE algorithm were 16% lower than observed values (calculated from the average 17 of the ratios of the measured extinction to the reconstructed extinction for all points  $> 300 \text{ Mm}^{-1}$ ). 18 The modified IMPROVE algorithm for PM<sub>1.0</sub> at  $\lambda = 470$  nm represents the chemical apportionment 19 of light extinction quite well. The reconstructed  $PM_{1,0}$  mass concentration (Fig. S3b) using the 20 modified IMPROVE algorithm was well correlated with the measured PM<sub>1.0</sub> mass concentration (the 21 summation of the concentrations of eight water-soluble ion compositions and carbon concentration 22 (including elemental carbon, [EC], and organic mass,  $1.6 \times [OC]$ )) (R<sup>2</sup> = 0.99, slope = 1.00, intercept 23 = 0.28) (Fig. S3d), indicating that the modified IMPROVE algorithm can be used to estimate the 24 chemical apportionment for extinction in this campaign. 25



Fig. S3 (a) and (b) : The plot of measured and reconstruction value of the extinction and mass concentration, respectively; (c) and (d) : Scatter plot of the measured extinction coefficient at  $\lambda = 470$ nm and PM<sub>1.0</sub> mass concentration against the reconstructed values with the modified IMPROVE formula. Insert of (c) shows the linear regression between measured and reconstructed extinction coefficient under lower aerosol load condition (with extinction smaller than 300 Mm<sup>-1</sup>).

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#### **S3** Relationships between aerosol optical properties and wind direction

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The local wind direction and wind speed were measured on the roof of building with a Gill 3 MetPak-II weather station. Fig. S4 presents the relationship of  $\alpha_{sp,470}$ ,  $\alpha_{ap,470}$  and  $\omega_{470}$  to local wind 4 directions and wind speed during the campaign. On these graphs, Beijing centre would be at ~206°. 5 Fig. S4a and S4b display the wind direction versus wind speed, with  $\alpha_{sp,470}$  and  $\alpha_{ap,470}$  as the color 6 scale. From November 2014 to January 2015, the winds were mostly from the southeast and 7 southwest. There was no obvious correlation between extensive optical properties and wind direction. 8 The average wind speed was 1.2 m s<sup>-1</sup> during the field campaign. When the instantaneous wind 9 speeds were higher than the average wind speed, the values of  $\alpha_{sp,470}$  and  $\alpha_{ap,470}$  decreased. The 10 average values of  $\alpha_{sp,470}$  for v < 1.2 m s<sup>-1</sup> and v > 1.2 m s<sup>-1</sup> were 198 Mm<sup>-1</sup> and 54 Mm<sup>-1</sup>, respectively. 11 Similarly, the average values of  $\alpha_{ap,470}$  for v < 1.2 m s<sup>-1</sup> and v > 1.2 m s<sup>-1</sup> were 41 Mm<sup>-1</sup> and 11 Mm<sup>-1</sup>, 12 respectively. When v > 1.2 m s<sup>-1</sup>, lower values of  $\alpha_{sp,470}$  (< 100Mm<sup>-1</sup>) and  $\alpha_{ap,470}$  (< 50Mm<sup>-1</sup>) 13 occurred more frequently when the local wind came from 225 - 270°, which indicated that the air 14 parcel was relatively clean in the wind direction. When v < 1.2 m s<sup>-1</sup>, values of  $\alpha_{sp,470}$  and  $\alpha_{ap,470}$ 15 occurred similar frequently and ranged widely when the local wind same from the south. Fig. S4c 16 shows the relationship between the wind direction and  $\omega_{470}$ , with the color scale as wind speed. The 17 average value of  $\omega_{470}$  ranged from 0.7 to 0.9 and was not strongly correlated with the wind direction. 18 However, higher values  $\omega_{470}$  (0.9 - 1.0) occurred more frequently when the local wind came from 19 270 - 160°, which indicated that the air parcel in this wind direction contained less light absorbing 20 particulate matter. 21

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Fig. S4 Local wind direction and wind speed plots for the campaign: (a) wind direction versus wind speed (m s<sup>-1</sup>), with  $\alpha_{sp,470}$  (Mm<sup>-1</sup>) as the color scale, (b) wind direction versus wind speed (m s<sup>-1</sup>), with  $\alpha_{ap,470}$ , (Mm<sup>-1</sup>) as the color scale, and (c) wind direction versus  $\omega_{470}$ , with the color scale as wind speed (m s<sup>-1</sup>).

### 1 S4 Frequency distribution and diurnal variations of dry PM1.0 optical properties

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With increasing pollutant level, the extensive optical properties ( $\alpha_{ep,470}$ ,  $\alpha_{sp,470}$  and  $\alpha_{ap,470}$ ) 3 increased strongly, in accordance with the expected, strong dependence of particle size on light 4 scattering. In contrast, changes in the intensive optical properties ( $\omega_{470}$ ) were more modest. The 5 6 optical measurement data are presented as histograms of the relative frequency of occurrence of  $\alpha_{ep.470}$ ,  $\alpha_{sp.470}$ ,  $\alpha_{ap.470}$  and  $\omega_{470}$  (Fig. S5). Approximately 71% of extinction and scattering coefficients 7 values were lower than 200 Mm<sup>-1</sup> and nearly 80% of absorption coefficient values were located in 8 the range of 1.5–50 Mm<sup>-1</sup>. Approximately 90% of the  $\omega_{470}$  values fell into the range of 0.70–0.97. 9 Compared with polluted days, the frequency distribution of the SSA showed similar patterns, and the 10 11 average values were similar on clear and slightly polluted days.

The diurnal variations of hourly averaged extinction, scattering, absorption coefficient and SSA 12 on clear, slightly polluted and polluted days are presented in Fig. S6. Broadly similar patterns were 13 observed for the extensive optical properties for different pollutant levels, although diurnal 14 15 variability was weak under polluted conditions owing to the high number of pre-existing particles and to the weakened diurnal changes in the boundary layer height that typically prevails under these 16 conditions (Yang et al., 2015; Gao et al., 2016). Extinction ( $\alpha_{ep,470}$ ) and scattering ( $\alpha_{sp,470}$ ) tended to 17 be lower during daytime and higher at night, in accordance with a lower nocturnal boundary layer 18 height (Ma et al., 2011). Emissions associated with morning rush hour are apparent in the aerosol 19 optical properties, especially in clean and slightly polluted conditions.  $\alpha_{ep,470}$  and  $\alpha_{sp,470}$  increased 20 slowly in the morning (07:00–09:00 LT) to peak values at 09:00 LT, indicating significant emission 21 and formation of particles during this period; these properties then decreased slowly until about 22 14:00 LT. The maximum values of  $\alpha_{ap,470}$  and lowest SSA values occurred during the traffic rush 23 hour and are therefore attributed to direct emissions of light absorbing species from vehicles. We 24 note that the increase in  $\alpha_{ap,470}$  from 06:00 to the maximum at 08:00 to 09:00 LT varied from 10 to 25 20 Mm<sup>-1</sup> and was quite consistent across different pollutant days. This observation suggests that the 26 number and type of particles emitted during this time period is not strongly influenced by 27 pre-existing pollutant levels, but mainly determined by relatively constant daily emissions from 28 traffic. 29



Fig S5 The frequency distributions of dry  $PM_{1.0}$  optical properties at  $\lambda = 470$  nm observed during the campaign. (a) extinction coefficient, (b) scattering coefficient, (c) absorption coefficient, (d) single scattering albedo on clear, slightly polluted and polluted days.



Fig. S6 Diurnal variations of hourly averaged (a) extinction coefficient, (b) scattering coefficient, (c) absorption coefficient and (d) SSA at  $\lambda =$ 

4 470 nm on clear, slightly polluted and polluted days. The error bars are 5th and 95th percentiles and the limits of the boxes represent 25th and

<sup>5 75</sup>th percentiles.

## 1 S5 Comparison of the measured aerosol optical properties to those at other 2 locations

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The Huairou site is a new suburban site at which aerosol optical properties have not previously 4 been reported. To put our observations at Huairou in context, the scattering and absorption 5 6 coefficients and SSA observed in this campaign are compared in Table S2 to those at other locations 7 (urban, suburban, and rural sites). As would be expected given the high concentrations of particulate matter in much of China, the mean  $\alpha_{sp,470}$  value at Huairou was considerably higher than values 8 observed in America and Europe, including the Los Angeles basin measurements in Pasadena 9 10 (Thompson et al., 2012) and the urban site of Granada (Titos et al., 2012). Within China, the Huairou values for scattering and absorption were higher than in Shanghai (Li et al., 2013) and similar to the 11 urban site of Guangzhou in China (Garland et al., 2008). Compared with other non-urban polluted 12 13 sites in China, both  $\alpha_{sp,470}$  and  $\alpha_{ap,470}$  at Huairou were lower than Xinken (Cheng et al., 2008), Yufa (Garland et al., 2009) and much lower than Xianghe (Li et al., 2007). Moreover, the  $\alpha_{sp,470}$  values 14 were comparable to those observed at Shangdianzi, an atmospheric background site located  $\sim 150$  km 15 northeast of the urban center of Beijing (Yan et al., 2008). The average value of  $\alpha_{ap,470}$  at Huairou 16 was lower than those seen at other urban and suburban locations in China, with the exceptions of 17 Guangzhou (reflecting the lower SSA values observed in Huairou) and the rural site of Shangdianzi. 18

The regional differences in SSA can be considered in terms of the different sources of particles, 19 including local primary emissions, transport emissions and secondary aerosol formations. The 20 average value of  $\omega_{470}$  at Huairou was  $0.80 \pm 0.08$ , which was lower than suburban sites in Xinken 21 22  $(0.83 \pm 0.05)$ , Xianghe (0.81-0.85) and Yufa  $(0.86 \pm 0.07)$ , and rural sites in Shangdianzi  $(0.88 \pm 0.07)$ 0.05), and Pasadena (0.92  $\pm$  0.08). Compared to urban sites, the Huairou SSA was similar to 23 observations in Beijing  $(0.80 \pm 0.09)$  (He et al., 2009), but considerably higher than in Shanghai 24  $(0.70 \pm 0.07)$  (Li et al., 2013) and Granada  $(0.71 \pm 0.07)$  (Titos et al., 2012). The lower SSA values 25 probably arose from the higher contribution of vehicular emissions in Shanghai (Zhou et al., 2009) 26 and both traffic emissions and a higher mass fraction of light absorbing particles caused by fuel-oil 27 combustion in Granada (Titos et al., 2012). 28

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Table S2. List of the mean values of aerosol scattering, absorption coefficients, and single scattering albedo in this campaign and recently reported values from references.

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Location	Date (MM/yy)	$\lambda$ (nm)	$\alpha_{sp} (\mathrm{Mm}^{-1})$	$\alpha_{ap} (\mathrm{Mm}^{-1})$	ω <sub>0</sub>	RH	Inlet	References
Granada, Spain (Urban)	03/2006 - 02/2007	$\alpha_{sp}:550$	61±25	24±9	0.71±0.07	<50%	PM <sub>10</sub>	Titos et al.(2012)
		$\alpha_{ap}$ : 550						
Guangzhou (Urban)	07/2006	$\alpha_{sp}$ : 550	151±103	34±27	$0.82 \pm 0.07$	<40%	$PM_{10}$	Garland et al.(2008)
		$\alpha_{ap}$ : 532						
Beijing (Urban)	01/2005 - 12/2006	$\alpha_{sp}$ : 525	255±243	45±39	$0.80 \pm 0.09$	<60%	TSP	He et al.(2009)
		$\alpha_{ap}$ : 532						
Shanghai (Urban)	04 - 05/2010	$\alpha_{sp}$ : 532#	102±74	44±35	$0.70 \pm 0.07$	41.2%	TSP	Li et al.(2013)
		$\alpha_{ap}:532$						
Xinken, PRD (Suburban)	10/2004 - 11/2004	$\alpha_{sp}$ : 550	333±138	70±42	$0.83 \pm 0.05$	<20%	$PM_{10}$	Cheng et al.(2008)
		$\alpha_{ap}: 630$						
Xianghe, Beijing (Suburban)	03/2005	$\alpha_{sp}$ : 550	468±472	65±75	0.81-0.85	<	TSP	Li et al.(2007)
		$\alpha_{ap}:550$				42.5%		
Yufa, Beijing (Suburban)	08/2006 - 09/2006	$\alpha_{sp}$ : 550	361±295	52±37	$0.86 \pm 0.07$	<32%	$PM_{10}$	Garland et al.(2009)
		$\alpha_{ap}:532$						
Huairou, Beijing (Suburban)	11/2014 - 01/2015	$\alpha_{sp}:470$	164±202	37±43	$0.80 \pm 0.08$	<15%	PM <sub>1.0</sub>	This work
		$\alpha_{ap}:470$						
Shangdianzi,	09/2003 - 01/2005	$\alpha_{sp}$ : 525	175±189	18±13	$0.88 {\pm} 0.05$	<60%	TSP	Yan et al.(2008)
Beijing (Rural)		$\alpha_{ap}:525$						
Pasadena, US (Rural)	05/2010 - 06/2010	$\alpha_{sp}$ : 532#	58±43	4±4	$0.92 \pm 0.08$	<50%	$PM_{1.0}$	Thompson et al.(2012)
		$\alpha_{ap}:532$						

# S6 Supporting information of the meteorological parameters and pollutant gases during the selected air pollution episode



Fig. S7 Temporal wind direction and speed, temperature, relative humidity and pressure of atmosphere air, PM<sub>2.5</sub>, CO, NO<sub>2</sub>, SO<sub>2</sub> and O<sub>3</sub> concentrations, and the measured aerosol extinction, scattering, and absorption coefficients during the air pollution episode.

### 2 S7 EC-tracer method for the estimation of secondary organic carbon

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EC is a good tracer of primary generated organic carbon (POC). Ambient OC/EC ratios larger than the OC/EC ratio of the primary ((OC/EC)<sub>pri</sub>) indicate the formation of secondary organic carbon (SOC). The concentrations of POC and SOC can be calculated with (Turpin and Huntzicker, 1995; Lim and Turpin, 2002; Lin et al., 2009):

$$OC_{pri} = EC \times (OC/EC)_{pri} + N$$
  

$$OC_{sec} = OC_{tot} - OC_{pri}$$
(S9)

9 where  $OC_{sec}$  is the SOC and  $OC_{tot}$  is the measured ambient OC concentration. The 10 concentration of POC ( $OC_{pri}$ ) could be calculated by the product of measured EC 11 concentration and the estimated (OC/EC)<sub>pri</sub>. N is the contribution of POC from 12 noncombustion sources or sample artifacts.

The estimation of (OC/EC)<sub>pri</sub> is based on the method used by Lim and Turpin (2002).
The scattering plot of OC and EC concentrations for the full campaign is shown in Fig. S8.
The lowest 20% percentile of ambient OC/EC ratios (shown in green dot points in Fig. S8)
were used for the determination of (OC/EC)<sub>pri</sub>. Time series of the estimated of POC and SOC mass concentrations are shown in Fig. S9.

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20 Fig. S8 Scatter plot of OC and EC concentrations. The green dot points were used to calculate

the primary OC/EC.





Fig. S9 Time series of (a) the measured OC mass concentrations, and the estimated primary
OC (POC) and secondary OC (SOC) mass concentrations determined with EC-tracer method.
(b) the percentage of POC and SOC in OC.

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### 7 S8 Modification of the IMPROVE with enhanced absorption of EC

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10 Fig. S10 Scatter plot of the measured extinction coefficient at  $\lambda = 470$  nm against the 11 reconstructed values with the new modified IMPROVE formula with MEE term of 23 m<sup>2</sup> g<sup>-1</sup>

12 at  $\lambda = 470$  nm.



4 Fig. S11 Average fractional contribution of each chemical composition to dry  $PM_{1.0}$ 5 extinction coefficient with respect to different pollution level with the new modified 6 IMPROVE formula.

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- 34