



Comparison of VLT/X-shooter OH and O₂ rotational temperatures with consideration of TIMED/SABER emission and temperature profiles

Stefan Noll¹, Wolfgang Kausch^{2,1}, Stefan Kimeswenger^{3,1}, Stefanie Unterguggenberger¹, and Amy M. Jones⁴

¹Institute for Astro- and Particle Physics, University of Innsbruck, Technikerstr. 25/8, 6020 Innsbruck, Austria

²Department of Astrophysics, University of Vienna, Türkenschanzstrasse 17, 1180 Vienna, Austria

³Instituto de Astronomía, Universidad Católica del Norte, Avenida Angamos 0610, Antofagasta, Chile

⁴Max Planck Institute for Astrophysics, Karl-Schwarzschild-Str. 1, 85748 Garching, Germany

Correspondence to: S. Noll (stefan.noll@uibk.ac.at)

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Abstract. Rotational temperatures T_{rot} derived from lines of the same OH band are an important method to study the dynamics and long-term trends in the mesopause region near 87 km. To measure realistic temperatures, the rotational level populations have to be in local thermodynamic equilibrium (LTE). However, this might not be fulfilled, especially at high emission altitudes. In order to quantify possible non-LTE contributions to the OH T_{rot} as a function of the upper vibrational level v' , we studied a sample of 343 echelle spectra taken with the X-shooter spectrograph at the Very Large Telescope at Cerro Paranal in Chile. These data allowed us to analyse 25 OH bands in each spectrum. Moreover, we could measure lines of O₂ $b(0-1)$, which peaks at about 94 to 95 km, and O₂ $a(0-0)$ with an emission peak at about 90 km. The latter altitude is reached in the second half of the night after a rise of several km because of the decay of a daytime population of excited O₂. Since the radiative lifetimes for the upper levels of the two O₂ bands are relatively long, the derived T_{rot} are not significantly affected by non-LTE contributions. These bands are well suited for a comparison with OH if the differences in the emission profiles are corrected. For different sample averages, we made these corrections by using OH emission, O₂ $a(0-0)$ emission, and CO₂-based temperature profile data from the multi-channel radiometer SABER on the TIMED satellite. The procedure relies on differences of profile-weighted SABER temperatures. For an O₂ $a(0-0)$ -based reference profile at 90 km, we found a good agreement of the O₂ with the SABER-related temperatures, whereas the

OH temperatures, especially for the high and even v' , showed significant excesses with a maximum of more than 10 K for $v' = 8$. The exact value depends on the selected lines and molecular parameters. We could also find a nocturnal trend towards higher non-LTE effects, particularly for high v' . The amplitude of these variations can be about 2 K or less, which tends to be significantly smaller than the total amount of the non-LTE contributions. The variations revealed may be important for dynamical studies based on T_{rot} derived from OH bands with high v' .

1 Introduction

Studies of long-term trends as well as short-term variations of the temperature in the mesopause region are often based on so-called rotational temperatures T_{rot} derived from emission lines of the hydroxyl (OH) airglow (e.g. Beig et al., 2003, 2008; Khomich et al., 2008; Reisin et al., 2014), which is mostly caused by the chemical reaction of hydrogen (H) and ozone (O₃) (Bates and Nicolet, 1950) at a typical height of 87 km (e.g. Baker and Stair, 1988). Ground-based instruments usually measure the intensities of a few lines with different rotational upper levels N' of a single band (e.g. Beig et al., 2003; Schmidt et al., 2013). Assuming a Boltzmann distribution of the populations at the different rotational levels, a temperature can be derived, which is reliable if the frequency of thermalising collisions is high compared to the

natural radiative lifetime of a state. The latter ranges from about 58 to 5 ms for the relevant upper vibrational levels $v' = 1$ to 9 of the OH electronic ground state (Xu et al., 2012). For the OH peak emission height, this appears to be sufficiently long to establish the assumed Boltzmann distribution (Khomich et al., 2008).

However, at higher altitudes with lower air density, collisions may be too infrequent, especially for the higher v' . Hence, the measured T_{rot} can deviate from the true temperatures due to the contribution of a population of OH molecules, where the condition of a local thermodynamic equilibrium (LTE) for the rotational levels is not fulfilled. Comparisons of T_{rot} from various bands differing in v' show clear signatures of such non-LTE effects (Cosby and Slanger, 2007; Noll et al., 2015). The measured T_{rot} tend to increase with v' and even v' appear to indicate higher T_{rot} than the adjacent odd v' . For this reason, the highest T_{rot} are found for OH bands with $v' = 8$. As discussed by Noll et al. (2015), this can only be explained by a significant impact of a nearly nascent vibrational and rotational level distribution related to the hydrogen–ozone reaction, which mostly populates $v' = 9$ and 8 (Charters et al., 1971; Ohoyama et al., 1985; Adler-Golden, 1997). The differences in the results for even and odd v' can then be understood by the higher initial T_{rot} for $v' = 8$ (Llewellyn and Long, 1978) and the high probability of radiative transitions with $\Delta v = 2$ (Mies, 1974; Rousselot et al., 2000; Khomich et al., 2008). The T_{rot} non-LTE contributions also appear to vary with time, as observed by Noll et al. (2015). At Cerro Paranal (24.6° S), they tend to enhance during the night when data of several years are averaged. This is probably caused by a rising OH emission layer, which can be inferred from a decrease of the measured OH intensity (Yee et al., 1997; Melo et al., 1999; Liu and Shepherd, 2006; von Savigny, 2015). Apart from v' , the measured T_{rot} are also strongly affected by the choice of lines. Higher N' tend to show stronger non-LTE contributions (Pendleton et al., 1989, 1993; Perminov and Semenov, 1992; Cosby and Slanger, 2007). For this reason, only lines of the first three or four rotational levels are usually used for T_{rot} derivations (e.g. Beig et al., 2003; Perminov et al., 2007; Cosby and Slanger, 2007; Schmidt et al., 2013). However, even this restriction is not sufficient to obtain line-independent T_{rot} . The OH electronic ground level has two substates: $X^2\Pi_{3/2}$ and $X^2\Pi_{1/2}$. Population differences of several per cent were found by Noll et al. (2015) for rotational levels of the same substate and for comparisons between the two substates.

The approach to estimate temperatures in the mesopause region by means of airglow lines with different rotational upper levels can also be applied to bands of molecular oxygen (O_2). However, long ground-based time series are only available for $\text{O}_2(b^1\Sigma_g^+ - X^3\Sigma_g^-)(0-1)$ at about 0.865 μm (e.g. Khomich et al., 2008; Reisin et al., 2014), which is sufficiently bright, hardly blended by OH lines, and not affected by self-absorption in the lower atmosphere like the stronger (0-0) band. Since $\text{O}_2b(0-1)$ remains unresolved with the stan-

dard instruments in use, a simulation of the temperature dependence of the band shape is required (e.g. Scheer and Reisin, 2001; Shiokawa et al., 2007; Khomich et al., 2008). In contrast to OH, the radiative lifetime τ of $\text{O}_2b(0-1)$ is relatively long. Using the Einstein- A coefficients from the HITRAN2012 molecular line database (Rothman et al., 2013), τ is about 210 s. This suggests that $\text{O}_2b(0-1)$ T_{rot} are not significantly affected by non-LTE contributions. The $b^1\Sigma_g^+$ state is essentially populated by atomic oxygen recombination and subsequent collisions (Greer et al., 1981), which results in an emission profile with a peak at about 94 to 95 km (derived from the (0-0) band; Watanabe et al., 1981; Greer et al., 1986; Yee et al., 1997); i.e. $\text{O}_2b(0-1)$ is sensitive to a different part of the mesopause region compared to OH. With full widths at half maximum (FWHM) of 8 to 9 km (Baker and Stair, 1988; Yee et al., 1997) for both kinds of emission, there is only a minor overlap of their profiles.

For a comparison with OH, transitions from the $a^1\Delta_g$ state of O_2 to the ground state $X^3\Sigma_g^-$ are more promising since the measured nighttime emission peak heights of the (0-0) band at 1.27 μm are close to those of OH (McDade et al., 1987; López-González et al., 1989; Lipatov, 2013). During most of the night, they are at about 90 km. At the beginning of the night, the effective emission heights can be several km lower. This behaviour is caused by the contribution of a daytime-related $a^1\Delta_g$ excitation by ozone photolysis to the nighttime-related atomic oxygen recombination (Vallance Jones and Gattinger, 1963; McDade et al., 1987; López-González et al., 1989, 1992). The former is relevant for the night since the spontaneous radiative lifetime of $a^1\Delta_g$ is about 75 min due to an Einstein- A coefficient of about $2.2 \times 10^{-4} \text{ s}^{-1}$ for the (0-0) band (Lafferty et al., 1998; Newman et al., 1999; Gamache and Goldman, 2001). The daytime emission peaks at about 50 km (Evans et al., 1968; Howell et al., 1990) and effectively rises in altitude after sunset up to the nighttime peak due to the faster decay of the $a^1\Delta_g$ population by collisional quenching at lower altitudes. Although non-LTE effects are negligible, only Mulligan and Galligan (1995) studied T_{rot} based on the bright $\text{O}_2a(0-0)$ band. Their sample consists of ground-based observations with a Fourier transform spectrometer at Maynooth, Ireland (53.2° N), and covers 91 twilight periods over 18 months. The T_{rot} measurements required simulations of the band for different ambient temperatures due to the strong line blending at their resolution. Moreover, $\text{O}_2a(0-0)$ is strongly affected by self-absorption, which required careful radiative transfer calculations. These efforts might explain the lack of $\text{O}_2a(0-0)$ T_{rot} in the literature. As an alternative, $\text{O}_2a(0-1)$ can be used. Although this band is fairly weak and partly blended with strong OH lines, it was studied by Perminov and Lipatov (2013) using spectroscopic data taken at Zvenigorod, Russia (55.7° N), from 2010 to 2011. The band emission had to be modelled to retrieve T_{rot} .

In this paper, we will compare T_{rot} derived from 25 OH bands, $\text{O}_2b(0-1)$, and $\text{O}_2a(0-0)$ to quantify the non-LTE effects related to the OH measurements. For this purpose, we will use 343 observations from the X-shooter echelle spectrograph (Vernet et al., 2011) at the Very Large Telescope (VLT) of the European Southern Observatory (ESO) at Cerro Paranal in Chile (2635 m, 24°38' S, 70°24' W). The OH data were already discussed by Noll et al. (2015). As X-shooter covers the wavelength range from 0.3 to 2.5 μm with a resolving power of at least $\lambda/\Delta\lambda \approx 3300$, the different OH and O_2 bands can be studied in parallel and without band simulations for the T_{rot} retrieval. The resolution is sufficient to measure individual lines even in $\text{O}_2a(0-0)$.

We will not directly compare the derived T_{rot} since the studied bands probe different parts of the mesopause temperature profile due to the differences in the altitude distributions of the emission. In particular, the significant changes in the emission profile of $\text{O}_2a(0-0)$ (López-González et al., 1992; Gao et al., 2011; Lipatov, 2013) have to be corrected. For this reason, we will use volume emission rate (VER) and temperature profiles from satellite observations to derive temperature corrections depending on the VER differences of the bands to be compared. For this purpose, data of the ten-channel broad-band radiometer SABER (Russell et al., 1999) on the TIMED satellite launched in 2001 are most suited. The limb sounder measures profiles of the crucial $\text{O}_2a(0-0)$ band and two OH-related channels at 1.64 and 2.06 μm , which probe different upper vibrational levels v' (4/5 and 8/9; see Baker et al., 2007). The latter is important since the OH emission peak heights depend on v' (Adler-Golden, 1997; Xu et al., 2012; von Savigny et al., 2012). The temperature profiles are derived from the SABER CO_2 channels under the consideration of non-LTE effects (Mertens et al., 2001, 2004; Remsberg et al., 2008; Rezac et al., 2015). TIMED has a slowly precessing polar orbit, which allows a 24 h coverage at low and mid-latitudes within about 60 days (Russell et al., 1999). Since this results in only two narrow local time intervals for a given site and day of year, our study will focus on average X-shooter and SABER data for certain periods.

The paper is structured as follows. We will start with the description of the X-shooter and SABER data used for this study (Sect. 2). Then, we will describe line measurements and T_{rot} retrieval from the X-shooter data, the derivation of SABER-based profiles suitable for the measured airglow bands, and the correction of the corresponding T_{rot} for the differences in the VER profiles (Sect. 3). In Sect. 4, we will discuss the resulting estimates of the non-LTE effects related to the OH T_{rot} and their change with nighttime and season. Finally, we will draw our conclusions (Sect. 5).

2 Data set

2.1 X-shooter

The X-shooter instrument (Vernet et al., 2011) of the VLT is an echelle slit spectrograph that covers the very wide wavelength range from 0.3 to 2.5 μm by three arms called UVB (0.30 to 0.59 μm), VIS (0.53 to 1.02 μm), and NIR (0.99 to 2.48 μm). This allows one to study various airglow bands simultaneously, as demonstrated by Noll et al. (2015), who investigated 25 OH($v' - v''$) bands between 0.58 and 2.24 μm with v' from 2 to 9. Since X-shooter is an instrument for observing astronomical targets, airglow studies have to rely on archival, unprocessed spectra, which differ in terms of line of sight, exposure time, and slit-dependent spectral resolution. For this study, we have used the same data set that was selected and processed by Noll et al. (2015). The extended analysis of these data, therefore, focuses on the $\text{O}_2b(0-1)$ and $\text{O}_2a(0-0)$ bands (see Sect. 3.1).

Noll et al. (2015) reduced archival data from October 2009 (start of the archive) to March 2013 using version V2.0.0 of the ESO public pipeline (Modigliani et al., 2010). The resulting wavelength-calibrated two-dimensional (2-D) spectra were then collapsed into a 1-D spectrum by applying a median along the spatial direction. For the flux calibration, response curves were created based on data of spectrophotometric standard stars (Moehler et al., 2014) corrected for molecular absorption (Smette et al., 2015; Kausch et al., 2015) and atmospheric extinction (Patat et al., 2011). Due to the instability of the flat-field calibration lamps for the correction of the pixel-to-pixel sensitivity variations in the 2-D echelle spectra, the fluxes of long wavelengths beyond 1.9 μm including the OH bands (8-6) and (9-7) are relatively uncertain. For 77 % of the data, this could be improved by means of spectra of standard stars with known spectral shapes, i.e. Rayleigh–Jeans slopes.

From the processed data, a sample of 343 VIS- and NIR-arm spectra was selected. The selection criteria were a minimum exposure time of 3 min, a maximum slit width of 1.5'' corresponding to $\lambda/\Delta\lambda \approx 5400$ for the VIS arm and 3300 for the NIR arm, and a difference in start, end, and length of the exposure between VIS and NIR arm of less than 5 min. The latter was necessary since the observations in the different X-shooter arms can be performed independently. Moreover, 21 % of the combined VIS- and NIR-arm spectra were rejected due to unreliable T_{rot} in at least one of the 25 investigated bands. This can be caused by low data quality, an erroneous data reduction, or critical contaminations by residuals of the observed astronomical object. 36 spectra of the final sample could only be used up to 2.1 μm , which excludes OH(9-7), because a so-called K -blocking filter was applied (Vernet et al., 2011). The 343 selected spectra show a good coverage of nighttimes and seasons and, therefore, appear to be representative of Cerro Paranal. For more details, see Noll et al. (2015).

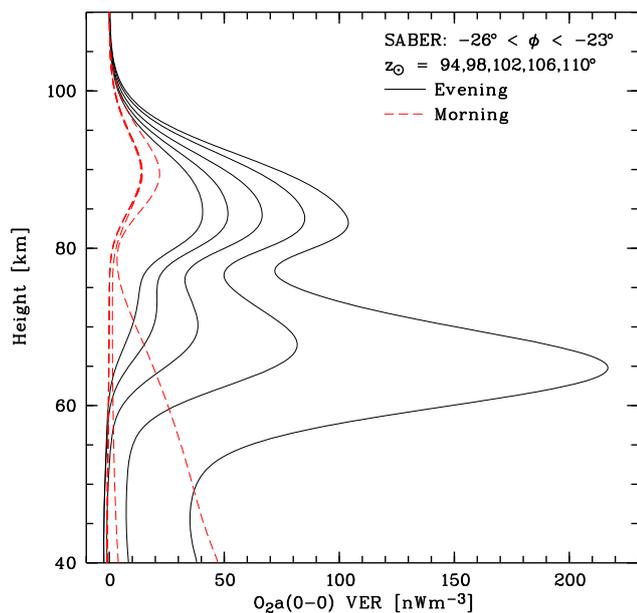


Figure 1. Average SABER volume emission rate profiles of $O_2a(0-0)$ during evening (solid lines) and morning twilight (dashed lines) for the latitude range from 23 to 26° S and the years from 2009 to 2013. The plot shows profiles for five different solar zenith distances z_{\odot} from 94 to 110°. The VERs decrease with increasing z_{\odot} .

2.2 SABER

Our analysis of VER and temperature profiles is based on version 2.0 of the limb sounding products of the SABER multi-channel photometer on TIMED (Russell et al., 1999). We used VER profiles from the 1.27 μm channel of $O_2a(0-0)$ and the two OH-related channels centred on 1.64 and 2.06 μm . We took the “unfiltered” data products, which are corrected VERs that also consider the emission of the targeted molecular band(s) outside the filter (Mlynczak et al., 2005). For $O_2a(0-0)$, the correction factor is 1.28 (Mlynczak et al., 2007). Finally, we used the SABER kinetic temperature T_{kin} profiles. In version 2.0, they are based on the limb radiance measurements in the 4.3 and 15 μm broad-band channels and a complex retrieval algorithm involving non-LTE radiative transfer calculations (Rezac et al., 2015). For our analysis, we could use the VER and T_{kin} profiles between 40 and 110 km because only for this altitude range all data are available. The lower and upper limits are related to the VERs and T_{kin} , respectively. We mapped all profiles (having a vertical sampling of about 0.4 km) to a regular grid with a step size of 0.2 km.

To cover a similar period as the X-shooter sample (see Sect. 2.1), we only considered SABER data of the years 2009 to 2013. The X-shooter data set is a nighttime sample with a minimum solar zenith distance z_{\odot} of 104°. For this reason, we only selected SABER data with a minimum z_{\odot} of 100°. With this limit, it is guaranteed that the $O_2a(0-0)$ main emis-

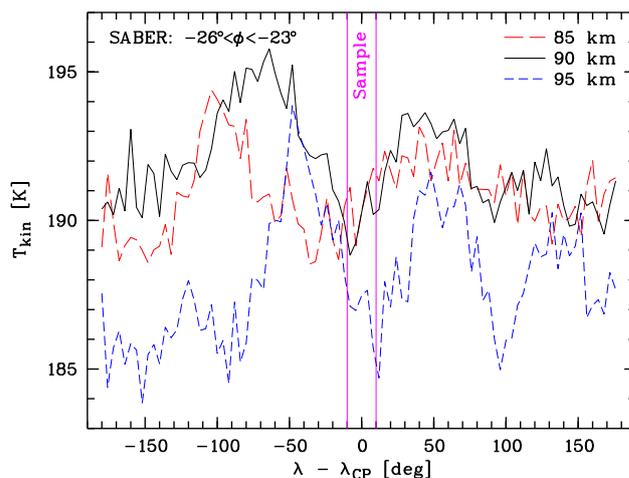


Figure 2. SABER kinetic temperature T_{kin} in K as a function of the longitude distance to Cerro Paranal $\lambda - \lambda_{\text{CP}}$ in degrees for the altitudes 85 (long dashed), 90 (solid), and 95 km (short dashed) and bin sizes of 4° and 1 km. The averages were derived from SABER data for the latitudes between 23 to 26° S, the years from 2009 to 2013, and zenith distances of at least 100°. The final SABER sample selection criterion $|\lambda - \lambda_{\text{CP}}| \leq 10^\circ$ is visualised by vertical bars.

sion peak is located in the mesopause region; i.e. it is typical of nighttime conditions. The fast change of the $O_2a(0-0)$ VER profile at twilight is illustrated in Fig. 1. As discussed in Sect. 1, the daytime population related to ozone photolysis with a maximum around the stratopause decays after sunset. The late nighttime profile remains relatively constant until the middle atmosphere is illuminated by the sun at dawn.

Since airglow emissions and temperatures significantly depend on geographic location (e.g. Yee et al., 1997; Melo et al., 1999; Hays et al., 2003; Marsh et al., 2006; Xu et al., 2010) due to differences in the solar irradiation and activity of tides, planetary and gravity waves, we selected only SABER data within the range from 23 to 26° S. Cerro Paranal is located at 24.6° S. With this criterion and the exclusion of 14 profiles which were suspicious in terms of the VER peak positions for the two OH channels, the resulting sample includes 27 796 measurements. We further restricted the data set by also considering longitudinal variations. Figure 2 shows the SABER T_{kin} for the altitude levels 85, 90, and 95 km as a function of longitudinal distance $\lambda - \lambda_{\text{CP}}$ to Cerro Paranal (70.4° W). On average, there are variations of several K and changes of the variability pattern depending on altitude. For this reason, we limited the sample by the criterion $|\lambda - \lambda_{\text{CP}}| \leq 10^\circ$, which resulted in 1685 profiles. A further reduction of the longitude range is not prudent because then there could be too few profiles for representative results for time and/or seasonal averages. We also checked the results presented in the subsequent sections with respect to the longitude limit (see Sect. 3.3). Indeed, the agreement of X-shooter and SABER variations does not signifi-

cantly improve with more restrictive limits. Also note that the SABER tangent point can easily move by more than 100 km during a profile scan. Moreover, the TIMED yaw cycle of 60 days (Russell et al., 1999) causes large gaps in the distribution of local times and days of year. For odd months (e.g. March), the period around midnight is well covered at Cerro Paranal. Even months show a better evening/morning coverage. Therefore, limitations in the correlation of X-shooter and SABER data are unavoidable.

3 Analysis

This study mainly focuses on the quantification of non-LTE contributions to OH T_{rot} . The following section describes the different steps of this analysis, which relies on X-shooter-related OH and O₂ T_{rot} measurements and the SABER-related correction of temperature differences by different band-specific emission profiles.

3.1 Line intensities and rotational temperatures

The wavelength range, spectral resolution, and quality of the X-shooter spectra (Sect. 2.1) allowed us to derive T_{rot} from line intensities of 25 OH bands, O₂b(0-1), and O₂a(0-0).

3.1.1 OH

The T_{rot} related to OH are based on the line selection and intensity measurements of Noll et al. (2015). The selection only includes lines originating from rotational levels $N' \leq 4$ of both electronic substates $X^2\Pi_{3/2}$ and $X^2\Pi_{1/2}$ to minimise the non-LTE effects that increase with N' (Cosby and Slinger, 2007; Khomich et al., 2008). Moreover, it focuses on P -branch lines (related to an increase of the total angular momentum by 1) since these lines are well separated in the X-shooter spectra. The OH Λ doublets (e.g. Rousselot et al., 2000) are not resolved. Due to reasons like blending by air-glow lines of other bands or strong absorption in the lower atmosphere (e.g. by water vapour), the number of suitable P -branch lines varied between 2 and 8 for the 25 considered OH bands. In the case of OH(6-4), an additional R -branch line was measured in order to have a minimum of three lines for quality checks. The measured intensities of all the lines were corrected for molecular absorption in the lower atmosphere by means of a transmission curve for Cerro Paranal computed with the radiative transfer code LBLRTM (Line-By-Line Radiative Transfer Model) (Clough et al., 2005) by Noll et al. (2012). The absorption by water vapour was adapted based on data for the precipitable water vapour (PWV) retrieved from standard star spectra (Kausch et al., 2015). For the OH line shape, Doppler-broadened Gaussian profiles related to a temperature of 200 K were assumed.

The measurement of T_{rot} is based on the assumption of a Boltzmann distribution of the N' -related populations, which are derived from $I/(g'A)$; i.e. the measured intensity

I divided by the N' -dependent statistical weight g' (number of sublevels) times the Einstein- A coefficient for the line. In this case, the natural logarithm of this population of a hyper-fine structure level of N' , which we call y , is a linear function of the level energy E' . The line slope is then proportional to the reciprocal of T_{rot} (Meinel, 1950; Mies, 1974). Noll et al. (2015) used the molecular parameters E' , A , and g' from the HITRAN2012 database (Rothman et al., 2013), which are based on Goldman et al. (1998) for the investigated lines (Rothman et al., 2009), and from van der Loo and Groenenboom (2007, 2008). We will mostly focus on the HITRAN-based parameters since they appear to provide more reliable T_{rot} .

Noll et al. (2015) corrected the resulting T_{rot} for the non-LTE effects caused by the different band-dependent line sets (see Sect. 1) by introducing a reference line set consisting of the three P_1 - ($X^2\Pi_{3/2}$) and P_2 -branch lines ($X^2\Pi_{1/2}$) with $N' \leq 3$. The temperature changes for the transition from the actual to the reference line set were calculated based on the sample mean T_{rot} of the 16 most reliable bands. The resulting set-specific ΔT_{rot} were then applied to the temperatures for the individual observations. For this study, we used a different reference line set only including the three P_1 -branch lines. The corresponding results are easier to interpret since the mixing of P_1 - and P_2 -branch lines causes a negative non-LTE effect; i.e. T_{rot} decreases by the lower relative population of the $X^2\Pi_{1/2}$ state, which has higher E' than $X^2\Pi_{3/2}$. We calculated the correction to the new reference set based on the four very reliable OH bands (4-1), (5-2), (6-2), and (7-4) (cf. Noll et al., 2015). This choice is not critical since the slight ν' dependence of the energy differences for the rotational levels has only a minor impact on the resulting ΔT_{rot} . The T_{rot} only based on the P_1 branch are 3.0 ± 0.3 K higher than those given by Noll et al. (2015) if the molecular parameters are taken from HITRAN. In the case of molecular data from van der Loo and Groenenboom (2008), ΔT_{rot} is 2.6 ± 0.4 K. If we focused on the fainter P_2 -branch lines, the corrections would be 4.3 ± 1.0 and 5.9 ± 2.5 K, respectively.

Our study is based on mean T_{rot} derived from bands with the same ν' , which show very similar intensity and T_{rot} variations. We took the ν' -specific T_{rot} from Noll et al. (2015) (corrected for the different reference line set) for each of the 343 sample spectra. Note that the $\nu' = 9$ mean values of 36 spectra with reduced wavelength range do not include OH(9-7) (see Sect. 2.1). The effect on the sample mean T_{rot} is negligible. The very faint OH(8-2) band is only considered by a fixed ΔT_{rot} for $\nu' = 8$ derived from the sample mean (see Noll et al., 2015). Excluding errors in the molecular parameters, the uncertainties of the T_{rot} (ν') based on the sample mean intensities are between 1.0 K for $\nu' = 8$ and 2.0 K for $\nu' = 2$. These values are mainly based on the T_{rot} variation for fixed ν' and are related to HITRAN data. For molecular data from van der Loo and Groenenboom (2008), the corresponding uncertainties range from 1.8 K for $\nu' = 9$ to 3.6 K for $\nu' = 2$. The T_{rot} uncertainties for individual observations

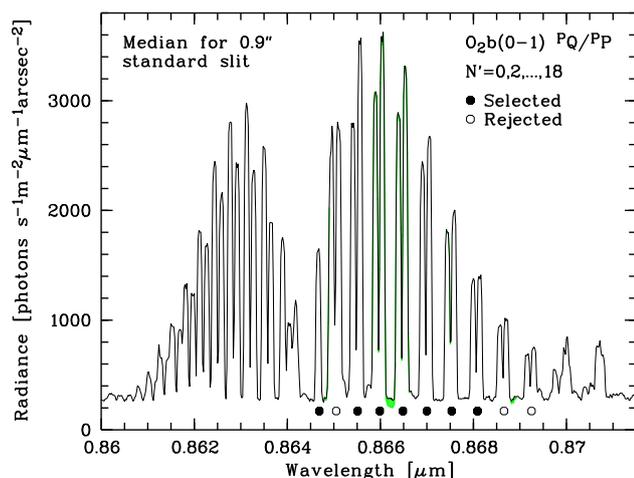


Figure 3. Median X-shooter spectrum of $O_2b(0-1)$ for the 189 VIS-arm sample spectra taken with the standard slit width of $0.9''$. The original spectrum was corrected for solar absorption lines in the continuum (green areas). Pairs of $^P Q$ and $^P P$ lines with even upper rotational levels N' from 0 to 18 are marked by circles. Note that $^P Q(0)$ does not exist. The filled symbols indicate lines that were used for the derivation of T_{rot} .

are higher than the stated values due to measurement errors depending on the signal-to-noise ratio.

3.1.2 $O_2b(0-1)$

As discussed in Sect. 1, $O_2b(0-1)$ at $0.865 \mu\text{m}$ is the most popular non-OH molecular band for T_{rot} retrievals in the mesopause region. It can also be measured in the X-shooter spectra. With a minimum resolving power of 5400 in the VIS arm, the individual lines of $O_2b(0-1)$ are not separated. However, in the P branch, blended lines originate from the same upper rotational level of $b^1\Sigma_g^+(v=0)$ (see Fig. 3). This allows T_{rot} to be directly measured without the requirement for T_{rot} -dependent synthetic spectra and data fitting, which is the standard approach (e.g. Scheer and Reisin, 2001; Shiokawa et al., 2007; Khomich et al., 2008). Consequently, the systematic uncertainties of the T_{rot} measurement are expected to be significantly lower. The two blended sub-branches are $^P Q$ and $^P P$ (e.g. Gamache et al., 1998), where the rotational level N increases by 1 and the total angular momentum J changes by 0 or +1 depending on the spin orientations in the lower state $X^3\Sigma_g^-(v=1)$. Due to the symmetry of the O_2 molecule, the upper state rotational levels N' have to be even. The selection rules also forbid $^P Q(0)$; i.e. there is only a single P -branch line for $N' = 0$.

For the T_{rot} determination, the intensities of the P -branch line pairs with N' from 0 to 18 were measured (see Fig. 3). The lines with N' higher than 18 are already relatively faint and strongly blended with OH lines. Before the line intensities could be determined, the continuum had to be measured and subtracted. Since an important part of the continuum is

caused by zodiacal light (sunlight scattered by interplanetary dust particles) and/or scattered moonlight (e.g. Noll et al., 2012), the strong solar Ca^+ absorption line at 866.2 nm is critical for the continuum window between the $N' = 6$ and 8 lines (see Fig. 3). For this reason, the solar flux atlas of Wallace et al. (2011) was used to calculate a continuum correction, which required the adaption of the solar transmission to the corresponding resolution of the X-shooter spectra and the determination of the variable fractional contribution of a solar-type spectrum to the continuum at $O_2b(0-1)$. The latter was derived from the depth of the Ca^+ absorption in the X-shooter data compared to the solar spectrum and is important since there is also an airglow continuum related to reactions involving NO (Khomich et al., 2008; Noll et al., 2012; Semenov et al., 2014). After the removal of the solar lines, the median continuum level was measured in nine fixed windows, which were optimised for the setups with the lowest spectral resolution, i.e. a slit width of $1.5''$. The range between the $N' = 2$ and 6 lines could not be used for the continuum determination due to significant contributions by OH lines. The resulting continuum fluxes were then linearly interpolated and subtracted from the X-shooter spectra. Finally, the intensities of the line pairs were integrated within limits which were adapted to the spectral resolution. Since the P branch of $O_2b(0-1)$ is not affected by molecular absorption (cf. Sect. 3.1.1), the corresponding correction could be neglected. As the observations were carried out at different zenith distances, the intensities were corrected to be representative of the zenith by applying the equation of van Rhijn (1921) (see Noll et al., 2015).

The resulting sample mean populations of the rotational levels $N' = 0$ to 18 are shown in Fig. 4. Similar to OH described in Sects. 3.1.1 and Noll et al. (2015), they correspond to $I/(g'A)$, where I and A are sums for the line pairs. The molecular parameters were taken from HITRAN2012 (Rothman et al., 2013), which is essentially based on Gamache et al. (1998) for $O_2b(0-1)$. Most y values show a good correlation with E' . The outliers can be explained by contamination from OH lines. This result supports the assumption that non-LTE effects are not relevant for the rotational level population. The radiative lifetime of about 210 s derived from the HITRAN data appears to be sufficiently long.

The optimal line set for the calculation of T_{rot} turned out to consist of $^P P(0)$ and the P -branch line pairs with N' from 4 to 14. For the mean populations shown in Fig. 4, the corresponding systematic T_{rot} uncertainty is only 1.8 K. This line set could be applied to 326 spectra. Due to instrumental issues or contamination by astronomical objects, 15 fits were significantly better by using only six instead of seven lines/line pairs. In two cases, only five emission features were considered. The optimal line sets were determined by an automatic iterative procedure. In order to avoid systematic differences depending on the line set, the T_{rot} based on less than seven data points were corrected. For this, the reliable spectra were used to derive the differences between the T_{rot} for

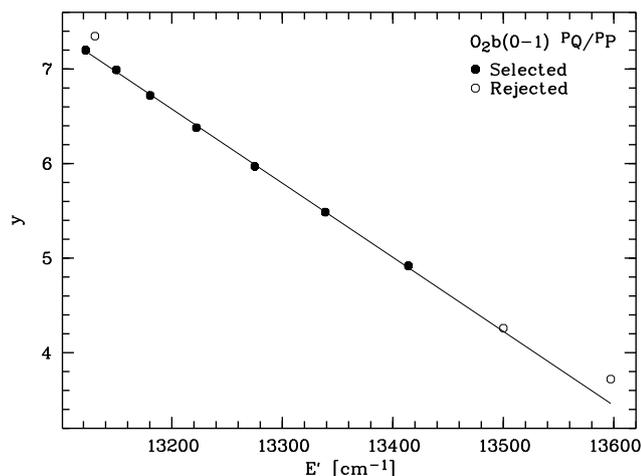


Figure 4. Derivation of $O_2b(0-1)$ T_{rot} under the assumption of a Boltzmann distribution for the different rotational levels. The abscissa shows the HITRAN-based energy of a given level in cm^{-1} . The ordinate displays the natural logarithm of the HITRAN-based population of a hyperfine-structure level relative to 1 Rayleigh second. Resulting from PQ and PP line intensities, the mean populations of the even rotational levels with N' from 0 to 18 for the full X-shooter sample are shown. Filled symbols mark reliable data points that were used for the derivation of T_{rot} (see Fig. 3), which is illustrated by the solid regression line.

the full and each reduced line set. The mean amount of the correction was 0.5 K.

We applied further corrections to the measured T_{rot} . First, we investigated possible systematic errors related to the widths of our continuum windows, which also had to be suitable for the low-resolution data. For 23 reliable high-resolution spectra taken with the $0.7''$ slit, we could significantly increase the number of continuum pixels. The resulting temperatures were about 0.4 K lower than those obtained based on the standard windows. As a consequence, we decreased the T_{rot} of the whole sample by the same amount. Second, we corrected the influence of the slit width on T_{rot} by adapting the resolution of the 23 $0.7''$ spectra to those taken with a $0.9''$, $1.2''$, or $1.5''$ slit and deriving the T_{rot} differences. This resulted in temperature increases of about 0.2, 0.4, and 1.0 K, respectively. All these corrections are relatively small, which makes us confident that our T_{rot} for $O_2b(0-1)$ are robust.

3.1.3 $O_2a(0-0)$

The X-shooter NIR arm covers the strong $O_2a(0-0)$ band at $1.27\ \mu\text{m}$ (see Fig. 5), which contains nine branches related to the dominating magnetic dipole transitions (e.g. Laferty et al., 1998; Newman et al., 1999; Rousselot et al., 2000; Leshchishina et al., 2010). With the minimum resolving power of 3300 in the X-shooter NIR arm, the lines of two branches are sufficiently separated to allow individual inten-

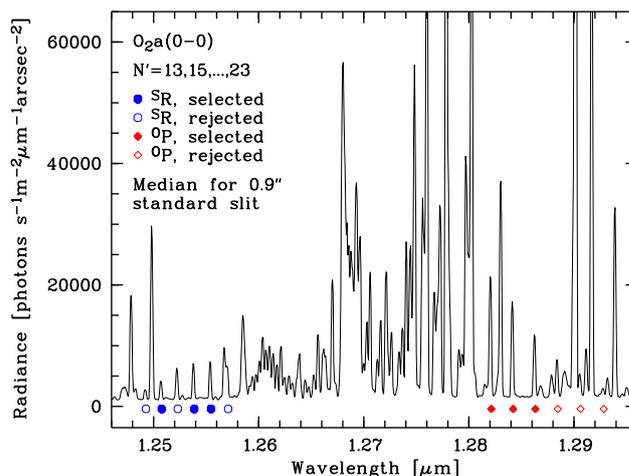


Figure 5. Median X-shooter spectrum of $O_2a(0-0)$ for the 177 NIR-arm sample spectra taken with the standard slit width of $0.9''$. SR and OP lines with odd upper rotational levels N' from 13 to 23 are marked by circles and diamonds, respectively. The filled symbols indicate lines that were used for the derivation of T_{rot} .

sity measurements. This could not be performed by Mulligan and Galligan (1995), who have done the only known $O_2a(0-0)$ T_{rot} measurements (see Sect. 1). The suitable branches are SR and OP (see Fig. 5), which have the same upper levels but differ by a change of N by -2 and $+2$, respectively. To fulfil the selection rules for ΔJ , these ΔN have to be accompanied by an opposite change of the spin ΔS by 1. Only transitions with odd N' originate from $a^1\Delta_g$. For the line measurements, we focused on intermediate N' between 13 and 23, which are sufficiently strong and not too close to the highly blended branches in the band centre. The line intensities were measured in a similar way as described in Sect. 3.1.2 for $O_2b(0-1)$. The effect of solar lines on the continuum could be neglected. However, since the lower state $X^3\Sigma_g^-(v=0)$ of $O_2a(0-0)$ is the ground state, the resulting line intensities had to be corrected for the strong self-absorption at low altitudes. The procedure was similar to the approach for OH described in Sects. 3.1.1 and Noll et al. (2015). Since the transmission curves for Cerro Paranal with a resolving power of about 1×10^6 calculated by Noll et al. (2012) were not accurate enough for $O_2a(0-0)$, we recalculated the O_2 absorption with the maximum resolving power of 4×10^6 inherent to the radiative transfer code LBLRTM (Clough et al., 2005). For a nominal temperature of 190 K for the $O_2a(0-0)$ emission layer, we obtained zenith transmissions between 0.14 and 0.80 for the measured SR lines and between 0.57 and 0.96 for the corresponding OP lines. The use of a fixed temperature for the Doppler broadening calculation has only a minor impact on the resulting T_{rot} . We performed a second iteration of the transmission correction using the measured T_{rot} . The mean T_{rot} change compared to the first iteration was only about 0.1 K.

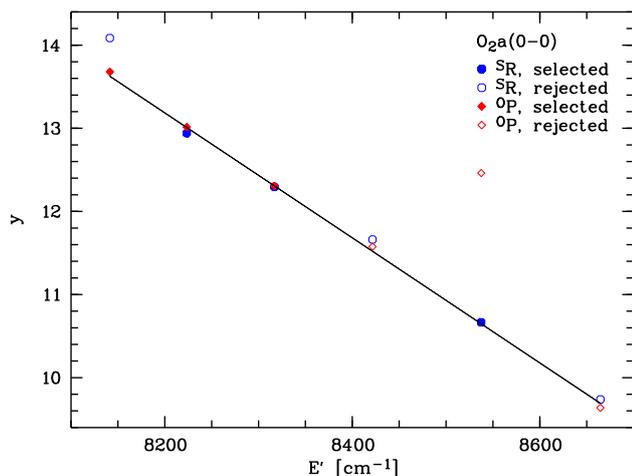


Figure 6. Derivation of $O_2a(0-0) T_{\text{rot}}$. The plot is similar to Fig. 4. Resulting from $^S R$ (circles) and $^O P$ (diamonds) line intensities, the mean populations of the odd rotational levels with N' from 13 to 23 for the full X-shooter sample are shown. Filled symbols mark reliable data points that were used for the derivation of T_{rot} (see Fig. 5), which is illustrated by the solid regression line.

Figure 6 shows the sample mean populations for the upper states of the measured $O_2a(0-0)$ lines. The populations were calculated by means of the molecular parameters of the HITRAN2012 database (Rothman et al., 2013), which are based on several studies (Gamache et al., 1998; Newman et al., 1999; Washenfelder et al., 2006; Leshchishina et al., 2010). Most lines indicate a tight linear relation of y and E' . The good agreement of the populations derived from $^S R$ and $^O P$ lines with the same N' is also convincing. The few outliers can be explained by blends with O_2 lines from other branches and OH lines. This result supports the expected good thermalisation of the rotational level populations of $a^1\Delta_g(v=0)$ due to the long radiative lifetime. The HITRAN data indicate an Einstein-A coefficient of $2.28 \times 10^{-4} \text{ s}^{-1}$ (cf. Sect. 1), which corresponds to $\tau \approx 73$ min.

We can directly constrain τ from our data. Since the daytime production of $a^1\Delta_g$ by ozone photolysis stops at sunset, the subsequent decrease of the $O_2a(0-0)$ intensity is an indicator of τ (e.g. Mulligan and Galligan, 1995; Pendleton et al., 1996; Nair and Yee, 2009). Since there are also $a^1\Delta_g$ losses by collisions, the derived lifetime is only a lower limit. Figure 7 shows the $O_2a(0-0)$ intensities normalised to the mean after midnight as a function of the time after sunset for our X-shooter data set. The steep intensity decrease at the beginning of the night is evident. We fit this decrease by an exponential function plus 1. The latter reflects the population related to the nighttime $a^1\Delta_g$ excitation, which is considered to be constant. This assumption is reasonable since the average nighttime variations are expected to be much weaker than those of the daytime-related population. The resulting lower lifetime limit is 46.1 ± 1.6 min, which is within the range of

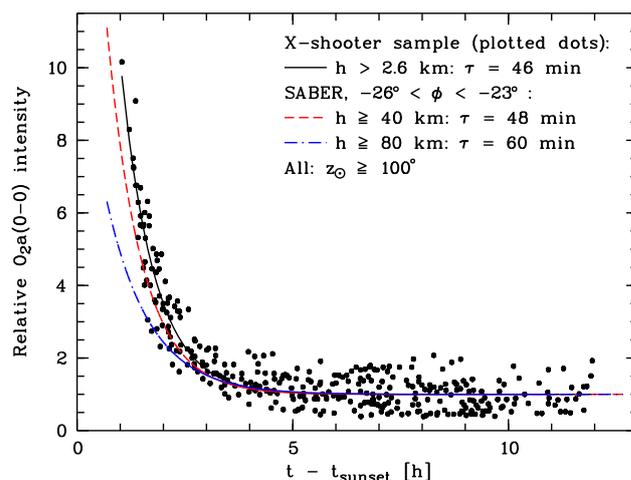


Figure 7. $O_2a(0-0)$ intensity relative to mean after midnight as a function of the time since sunset for the X-shooter sample (dots). The solid line shows the best fit of a constant and an exponential function to the data. The plot also displays the corresponding fit results for the SABER sample without longitude restriction and lower intensity integration limits of 40 (dashed) and 80 km (dash dotted). For the three fits, the resulting radiative lifetimes τ of the daytime population are shown in the legend.

other ground-based measurements (Mulligan and Galligan, 1995; Pendleton et al., 1996). We can also derive τ from our SABER data set (Sect. 2.2). For the latitude range from 23 to 26° S, z_0 greater than 100° , and a minimum height of 40 km, we obtained about 48 min, which is close to the X-shooter value. The SABER best-fit curve in Fig. 7 is shifted to earlier times due to the sparse twilight sampling of the X-shooter data with a minimum z_0 of 105° . This could also contribute to the small difference of the τ values. As illustrated by Fig. 1, the decay of the daytime population is faster at lower altitudes, where collisions are more frequent. Therefore, the resulting τ depends on the considered altitude range. For a lower limit of 80 km, we obtained about 60 min (see Fig. 7). Also based on SABER data, Nair and Yee (2009) derived an Einstein-A coefficient consistent with the theoretical value for a height of 85 km.

As shown by Figs. 5 and 6, the optimal line set for the T_{rot} retrieval consists of $^O P$ (13), $^S R$ (15), $^O P$ (15), $^S R$ (17), $^O P$ (17), and $^S R$ (21). In principle, the two lines with $N' = 23$ could also have been included since the T_{rot} derived from the mean populations in Fig. 6 only differs by 0.1 K from the result for our standard line set. However, they are relatively faint, which causes significant uncertainties if the T_{rot} are obtained from individual spectra. Using the mean populations and the extended line set, the systematic T_{rot} uncertainty from the regression analysis is about 2.1 K.

The resulting T_{rot} were corrected for different systematic effects. For this purpose, the 47 spectra taken with the narrow $0.4''$ and $0.6''$ slits were used. First, we checked the

quality of the crucial O_2 self-absorption correction. This was achieved by multiplying the optical depth of all lines by a line-independent but variable factor and recalculating T_{rot} to find the results with the lowest regression uncertainties. For the 47 test spectra, this procedure resulted in an optimal factor of 1.03 ± 0.03 and a T_{rot} decrease by 1.1 ± 0.8 K. This small correction confirms the reliability of our radiative transfer calculations. Then, similar corrections as for $O_2b(0-1)$ (Sect. 3.1.2) were performed. With continuum windows optimised for the narrow slits instead of the wide $1.5''$ slit, we obtained a ΔT_{rot} of -0.3 K. The effect of the spectral resolution on the line measurements resulted in T_{rot} corrections of $+1.2$, $+0.4$, and $+1.4$ K for the $0.9''$, $1.2''$, and $1.5''$ slits. Taking the $0.6''$ slit as a reference, the T_{rot} of the three spectra obtained with a $0.4''$ slit were also corrected and increased by 3.5 K. This relatively large correction is probably related to the starting resolution of underlying faint lines. In general, the applied corrections are small and partly cancel each other out. The systematic T_{rot} uncertainty slightly increases from 2.1 to 2.2 K when the uncertainties of the T_{rot} corrections are also considered.

3.2 Emission profiles

In order to compare the T_{rot} derived from the discussed bands (Sect. 3.1), the differing height distributions of the emission have to be considered since the measured bands probe different parts of the mesopause temperature profile. Since the ground-based X-shooter data do not provide this information, we obtained VER and temperature profiles representative of Cerro Paranal from the satellite-based SABER data (Russell et al., 1999).

3.2.1 OH

As described in Sect. 2.2, SABER has two OH-related photometric channels centred on 1.64 and $2.06 \mu\text{m}$ (Russell et al., 1999), which essentially cover the (4-2), (5-3) and (8-6), (9-7) bands, respectively (see Baker et al., 2007). The SABER OH profiles cannot be used directly since the OH emission height depends on v' (Adler-Golden, 1997; Xu et al., 2012; von Savigny et al., 2012), which ranges from 2 to 9 for our X-shooter data (Sect. 3.1.1). For this reason, we had to derive v' -related profiles from the SABER data.

It is reasonable to assume that the v' -dependent emission peak altitudes are equidistant (cf. von Savigny et al., 2012) (see also Sect. 3.4). Even if there are significant deviations from this assumption, the maximum errors should only be of the order of a few 100 m, which is still relatively small compared to the typical FWHM of 8 to 9 km of the OH VER profiles (e.g. Baker and Stair, 1988). Then, the v' -dependent emission peak altitudes can be derived by means of effective v' for the two SABER channels. We achieved this by calculating OH line spectra from the vibrational level populations and T_{rot} given by Noll et al. (2015) for HITRAN data

and convolving them with the SABER filter curves (Baker et al., 2007). In this way, we obtained fractional contributions of the different OH bands and, finally, the effective v' , which are 4.57 and 8.29 for the 1.64 and $2.06 \mu\text{m}$ channels, respectively. Although we neglected the higher relative populations at high N' compared to those related to T_{rot} (e.g. Cosby and Slanger, 2007) and did not consider the variability of the relative populations (Noll et al., 2015), the resulting values are sufficiently robust for our purpose. The corresponding altitude uncertainty should be well below 100 m. From $\Delta v' = 3.72$ and the emission peak altitudes h_{peak} of both OH channels, we derived a mean altitude difference of 0.37 km for $\Delta v' = 1$ and our Cerro Paranal SABER sample consisting of 1685 measurements (Sect. 2.2). This is in good agreement with data from other studies (von Savigny et al., 2012). Since h_{peak} is not very accurate due to a step size of 0.2 km (Sect. 2.2), we used the centre of the range limited by the interpolated half maximum positions h_{cen} as a more robust measure to calculate the v' -dependent profile altitudes. In this case, the mean $\Delta v' = 1$ altitude difference is 0.39 km with a standard deviation σ of 0.15 km.

The VER profiles of the two OH channels (see Fig. 8) show slightly different shapes. For the Cerro Paranal SABER sample, the mean FWHM are 8.1 and 9.1 km for the 1.64 and $2.06 \mu\text{m}$ channels, respectively. The corresponding σ are 1.9 and 2.1 km. For a better comparability of the v' -dependent profiles and the related temperatures (see Sect. 3.3), we used merged profiles. For all SABER measurements, we shifted both profiles normalised to the same integrated intensity to the calculated h_{cen} for each v' and averaged them. The resulting mean profiles for the Cerro Paranal SABER sample are shown in Fig. 8. They have a FWHM of 8.6 km with a σ of 1.9 km. The average h_{cen} range from 86.2 km ($v' = 2$) to 89.0 km ($v' = 9$) with a scatter between 1.3 and 1.4 km. The profiles are asymmetric. The mean h_{cen} are about 0.6 km ($\sigma = 1.1$ km) higher than the corresponding h_{peak} . For the effective height h_{eff} , i.e. the VER-weighted altitude, the average difference is 1.2 km. The calculation of h_{eff} excluded altitudes with negative VERs and $|h - h_{\text{cen}}| > 15$ km (see also Sect. 3.3).

3.2.2 $O_2b(0-1)$

SABER does not provide VER profiles for $O_2b(0-1)$. For this reason, we assumed a simple Gaussian profile, which appears to be justified by the relatively small measured asymmetries (see Khomich et al., 2008). Based on values from the literature (see Sect. 1), where the measurements are usually based on $O_2b(0-0)$ (e.g. Yee et al., 1997), peak height and FWHM were set to 94.5 and 9 km (see Fig. 8). Since our $O_2b(0-1)$ profile is fixed, we checked the influence of variations of h_{peak} and FWHM on the derived temperatures (see Sect. 3.3). For this analysis, we assumed the same variations of these parameters as described in Sect. 3.2.1 for OH. As a result, we found for individual profiles an average σ of 2.0 K. How-

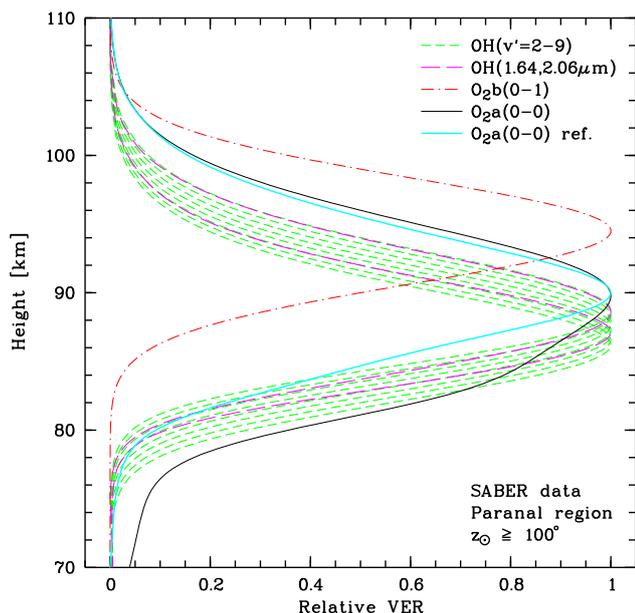


Figure 8. Mean emission profiles of different molecular transitions (see legend) for the Cerro Paranal SABER sample. The volume emission rates scaled to the emission peak are given as a function of the height in km. The emission peak height of the OH profiles increases with v' and the mean profile of the 2.06 μm SABER channel peaks higher than the 1.64 μm one.

ever, for the sample mean (where transient waves vanish), σ is only 0.5 K. Since we are mainly interested in average properties, the exact choice of h_{peak} and FWHM for $\text{O}_2b(0-1)$ is not critical for our study.

3.2.3 $\text{O}_2a(0-0)$

SABER has a dedicated channel for $\text{O}_2a(0-0)$ (see Sect. 2.2). This was one of the reasons for selecting SABER for this study since this band shows strong variations of the emission profile. In Fig. 1, this was illustrated for twilight conditions. As our X-shooter data set is a nighttime sample, we show the average emission profiles for the Cerro Paranal SABER sample at five different time intervals in Fig. 9. For all periods, most of the emission originates close to the OH emission layer (see also Fig. 8), which is crucial for a reliable T_{rot} comparison (see Sect. 3.3). However, the emission peak indicates a clear transition from about 84 to 90 km in the first half of the night. The early-night peak is obviously caused by a daytime ozone (and its photolysis) maximum in the upper mesosphere, whereas the late-night peak appears to be related to atomic oxygen recombination and subsequent collisions (e.g. López-González et al., 1989; Mulligan and Galligan, 1995; Kaufmann et al., 2003) (see also Sect. 1).

The nocturnal variations of $\text{O}_2a(0-0)$ are illustrated in Fig. 10. Here, the effective heights h_{eff} are given compared with those of $\text{OH}(v' = 8)$. The calculation of the $\text{O}_2a(0-0)$

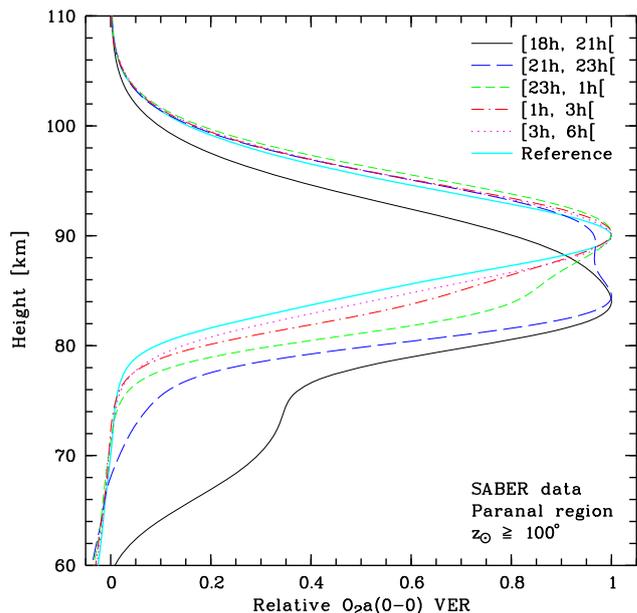


Figure 9. $\text{O}_2a(0-0)$ VER profiles scaled to the emission peak for the Cerro Paranal SABER sample. Mean profiles for five different nocturnal local time (LT) periods are shown (see legend). For comparison, the $\text{O}_2a(0-0)$ -based reference profile is also displayed (light solid line).

h_{eff} excluded altitudes with negative VERs, which essentially appear in the lower mesosphere (see Fig. 9). Additional height limits as for OH (see Sect. 3.2.1) were not applied because of the strong profile variations. The general profile limits of 40 and 110 km (Sect. 2.2) are not an issue for the nighttime. Figure 10 shows a relatively large scatter of 2.4 km in the $\text{O}_2a(0-0)$ h_{eff} at the beginning of the night, which is not seen for $\text{OH}(v' = 8)$ ($\sigma = 1.1$ km). In the course of the night, the deviation (4.4 km) and scatter become smaller until a convincing linear relation with $\text{OH}(v' = 8)$ is established. The correlation coefficient for the last period is $r = 0.62$, which is significantly higher than $r = 0.35$ for the first period. The increase of h_{eff} for $\text{O}_2a(0-0)$ is partly compensated by a simultaneous rise of the OH layer of 1.1 km between the first and last period. The smallest discrepancy and scatter of 0.2 and 1.1 km for $\text{O}_2a(0-0)$ are found for the fourth nighttime interval. The result for the discrepancy will change with a different $\text{OH}(v')$. The analysis of the h_{eff} data confirms our expectation that an OH and $\text{O}_2a(0-0)$ T_{rot} comparison is possible with only minor corrections for a wide range of nighttimes (see Sect. 3.3). A correction is not negligible even for equal h_{eff} since the profile shapes and FWHM can be very different (see Fig. 8). The average emission profile FWHM for $\text{O}_2a(0-0)$ and OH are 12.9 and 8.7 km, respectively. There is still a difference of 3.6 km (corresponding to 11.7 and 8.0 km) for the fourth nighttime period, which shows the best agreement in terms of h_{eff} . The emission profile widths for $\text{O}_2a(0-0)$ and OH decrease during the night. The difference is 4.6 and

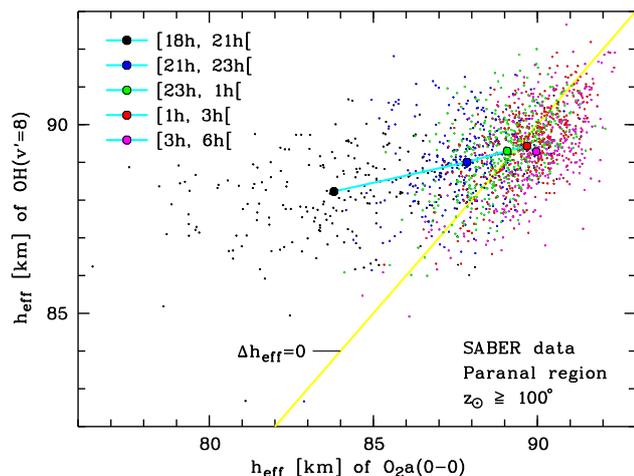


Figure 10. Comparison of effective emission heights h_{eff} in km for OH($v' = 8$) and O₂a(0-0) for the Cerro Paranal SABER sample. The mean values for five different nighttime LT periods are shown by big circles of different colours connected by lines (see legend). The colours of the individual data points correspond to those of the related mean values.

2.3 km, respectively, if the FWHM of the first and last period are compared.

The emission profiles of O₂a(0-0) and the negligible non-LTE contributions to the corresponding T_{rot} make this band ideal for an analysis of the OH-related T_{rot} . However, the strong profile variations do not allow one to assign a reference altitude for the O₂a(0-0)-related temperatures. Therefore, we also introduce a reference profile, which is based on O₂a(0-0) profiles observed in the second half of the night. Since the corresponding median h_{peak} and FWHM are 89.8 and 10.9 km, respectively, we selected 90.0 and 11.0 km as reference values. Observed SABER VER profiles with similar properties were then averaged. As the Cerro Paranal SABER sample is too small for this, we used the larger sample without longitude restriction (see Sect. 2.2). We considered all 297 profiles with h_{peak} between 88 and 90 km and FWHM between 10.8 and 11.2 km. To avoid any broadening of the averaged profile, the limits for the FWHM were more restrictive. In the case of h_{peak} , this effect could be prevented by shifting the profiles to the reference height before the averaging. Figures 8 and 9 show the resulting reference profile, which is not significantly affected by the decaying daytime $a^1\Delta_g$ population. The profile is slightly asymmetric with $h_{\text{cen}} = 90.1$ km and $h_{\text{eff}} = 90.4$ km.

3.3 Temperature corrections

The final step of the preparation of T_{rot} data for a comparison is the correction of differences in the emission profiles. This requires some knowledge of the true temperature profile in the mesopause region. SABER provides this kind of data (T_{kin}), which are based on CO₂-related measurements

and non-LTE radiative transfer calculations (see Sect. 2.2). For the derivation of the sample-averaged temperature differences for slightly deviating emission profiles, the T_{kin} uncertainties are probably small. However, we also want to directly compare X-shooter T_{rot} and SABER T_{kin} . For this reason, the temperature uncertainties have to be known in detail. According to Rezac et al. (2015) and J. M. Russell III (personal communication, 2015), the total uncertainties for single profiles are about 5 K at 90 km and tend to increase with altitude. Since we are interested in average properties, we can neglect the statistical errors, which leaves systematic uncertainties of 3 to 4 K in the relevant altitude range. This may further decrease if it is considered that systematic errors can also partly cancel out (Rezac et al., 2015). For this reason, we assume a relevant uncertainty of 2.5 K as input for Sect. 4. As an additional check, we compared the preferred version 2.0 data to those of version 1.07, which is possible for SABER data taken until the end of 2012. The old T_{kin} show deviations from the new ones, which are between +2.0 K for OH($v' = 2$) and −0.2 K for O₂b(0-1). Assuming that the new T_{kin} are more correct, the changes are consistent with the assumed errors.

To obtain effective temperatures T_{eff} from the SABER T_{kin} for the different bands, we used a similar procedure as described for h_{eff} (Sects. 3.2.1 and 3.2.3); i.e. the T_{kin} of the different altitudes were weighted depending on the related VERs for each emission layer. Altitudes with negative VERs were rejected. Moreover, weights were set to zero for altitude distances of more than 15 km from the OH h_{cen} .

The difference between the resulting T_{eff} of a given band and a desired reference band ΔT_{eff} indicates the change of the effective true temperature by deviating altitudes, widths, and shapes of the corresponding emissions profiles. For the derivation of possible non-LTE effects, these ΔT_{eff} have to be subtracted from the measured T_{rot} . The corrections cannot directly be applied since both quantities originate from different data sets, i.e. the Cerro Paranal SABER sample with 1685 profiles (Sect. 2.2) and the X-shooter sample with 343 spectra (Sect. 2.1). For this reason, we needed to obtain SABER-specific parameters such as T_{eff} and h_{eff} for each X-shooter spectrum. We achieved this by smoothing the SABER data in the two dimensions: day of year (DOY) and time. As a smoothing filter, we used a 2-D Gaussian with σ of half a month (15.2 days) and half an hour. The Gaussians were calculated across the New Year boundary if required. For each X-shooter data point, we could then compute weights for the different SABER data points and, finally, weighted averages for the SABER-related parameters.

A suitable smoothing procedure is crucial since both data sets have large gaps. In particular, the SABER data for our selected area around Cerro Paranal only cover a very narrow range of times for a certain DOY (see Sect. 2.2). It is desirable to average as many SABER data as possible to obtain reliable mean values. However, if this requires adding data taken at very different conditions, this might cancel out

variability patterns that are required for realistic corrections. The SABER data set is already a compromise since it includes data that were taken several 100 km away from Cerro Paranal. Hence, we checked the influence of the smoothing procedure on the resulting weighted parameters. Focusing on the X-shooter sample mean T_{eff} for the eight OH, two O₂, and one reference emission layers (see Sect. 3.2), we found a mean change of about 0.1 K for a bisection of the Gaussian σ . A doubling of the limiting periods resulted in a corresponding mean shift of about 0.2 K. Moreover, a different smoothing function (e.g. an exponential function with a flat top for distances smaller than σ) had little influence on the results. Finally, the introduction of the solar activity (measured by the solar radio flux) as a third dimension did not significantly change the resulting T_{eff} either. Therefore, we conclude that the details of the smoothing procedure for the SABER data are not critical for our analysis.

The main contributions to the weighted parameter averages mainly originate from a relatively small fraction of the SABER data. Within the 1σ perimeter of the 2-D Gaussian around the Cerro Paranal data points, the number of SABER profiles ranges from 0 to 41 for our applied procedure. The mean is 12. If a weight sum is used (defining weight = 1 in the centre of the Gaussian), values between 4 and 49 with a mean of 23 are found. These numbers suggest that profile-to-profile variations limit the accuracy of the resulting averages. This is confirmed by typical T_{eff} mean errors between 1.3 K (O₂a(0-0)) and 1.9 K (O₂b(0-1)) for individual X-shooter observations. These uncertainties are comparable to the systematic T_{rot} measurement errors (see Sect. 3.1). However, for sample averages, the statistical T_{eff} -related errors strongly decrease. The full sample mean errors are only about 0.1 K. The uncertainties are further reduced if the ΔT_{eff} for the T_{rot} correction are taken into account. In this case, the individual mean errors range from 0.6 K (O₂a(0-0)) to 1.3 K (O₂b(0-1)) for the reference profile at 90 km described in Sect. 3.2.3. The uncertainties for the sample mean are negligible.

The most critical source of error is probably a significant deviation of the T_{eff} climatologies from those related to the Cerro Paranal X-shooter data. This is relevant since the set of SABER profiles used covers a relatively wide area comprising 3° in latitude and 20° in longitude (see Sect. 2.2). Moreover, there are large gaps in the local time coverage for a given month. In order to investigate this, we used the five nighttime periods first introduced in Fig. 9. For the OH(v') and two O₂ bands, we then measured the sample-averaged difference of the X-shooter-related T_{rot} and the SABER-related T_{eff} for each time interval. Finally, mean values and standard deviations were calculated for the five data points. This provides us with a rough estimate of the agreement of the X-shooter and SABER nocturnal variability patterns. The resulting standard deviations range from 0.7 K for OH($v' = 9$) to 2.1 K for OH($v' = 2$). The average for all emission layers is 1.5 K. We can repeat this calculation

by taking the temperature differences for two bands, which is needed for the emission profile correction. For the emission of O₂a(0-0) as reference profile, the standard deviations now range from 0.2 K for OH($v' = 3$) to 3.2 K for O₂b(0-1), which reflects the relatively low effective height of the reference layer. Our preferred reference profile at 90 km cannot be used in this context since T_{rot} measurements do not exist. For this reason, we calculated the standard deviations for all possible band combinations and performed a regression analysis depending on the related height differences Δh_{eff} . Then, the climatology-related uncertainties can be estimated only based on differences of emission altitudes. This approach assumes that h_{eff} is the major driver for the changes in the variability pattern, which is likely since the widths of the emission layers do not differ a lot. Moreover, a linear change of the variability with Δh_{eff} is required, which is more difficult to fulfil if the height differences are relatively large. Significant variations of the OH non-LTE effects could influence the fit (see Sect. 4). For our X-shooter sample, we obtain a slope of 0.37 K km⁻¹ and an offset of 0.47 K. The latter includes influences like FWHM differences. The fit is robust since the uncertainties of both values are lower by more than an order of magnitude. For a change from a band-specific emission profile to our reference profile, the linear fit results in systematic errors in the temperature correction by the different climatologies for the SABER and X-shooter data between 0.7 K for OH($v' = 9$) and 2.0 K for O₂b(0-1). For the discussion of seasonal variations in Sect. 4.3, we also performed the whole analysis by using double month periods, which have a similar length as the TIMED yaw cycle of about 60 days (Russell et al., 1999). Independent of the month grouping starting with December/January or January/February, we obtained slightly higher errors between 1.1 K for OH($v' = 9$) and 2.3 K for O₂b(0-1). Our estimates might be lower than the real systematic uncertainties due to the low bin number, which is limited by the small data sets and gaps in the time coverage (cf. Sect. 4.3).

The relatively large distances between Cerro Paranal and the tangent points where the SABER profiles were taken are an important reason for the deviating variability. With the approach described above, we can check the robustness of our area selection limits. Focusing on the $|\lambda - \lambda_{\text{CP}}| \leq 10^\circ$ criterion (Sect. 2.2), we also studied 5 and 20° limits. For the correction of the differences between the emission profiles of the measured bands and the reference profile, we find average temperature uncertainties of 1.2, 1.3, and 1.4 K for the 5, 10, and 20° limits, respectively. The averages were derived from the results for the eight OH and two O₂ emission layers. The errors show the expected increase with growing longitude interval. However, the changes are small enough that the selection of the limit does not need to be very accurate. A tight longitude range is not recommended since this would significantly shrink the SABER sample and, hence, increase the statistical errors.

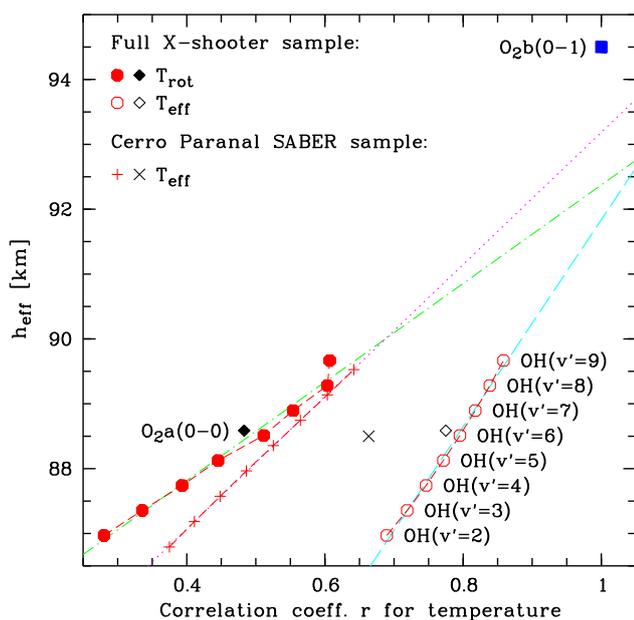


Figure 11. Effective emission height h_{eff} in km from SABER data vs. the correlation coefficient r for the correlation of the T_{rot} (filled symbols), T_{eff} for the full X-shooter sample (open symbols), and T_{eff} for the Cerro Paranal SABER sample (crosses) of OH(v') and O₂a(0-0) with those of O₂b(0-1). The dash-dotted (T_{rot} , X-shooter sample), long-dashed (T_{eff} , X-shooter sample), and dotted line (T_{eff} , SABER sample) show fits of the OH data, which are connected by short dashed lines.

3.4 Emission heights and temperature correlations

In the same way as described in Sect. 3.3 for T_{eff} , the band-specific h_{eff} (Sect. 3.2.1) were derived for each X-shooter observation. It is possible to check the reliability of these data by comparing average h_{eff} with correlation coefficients r for the correlation of the corresponding T_{rot} with the T_{rot} for a reference band derived from the same data set (Sect. 3.1). Assuming a gradual change of the temperature variability pattern with height, the T_{rot} -based r should be a function of the altitude difference of the emission layers used for the correlation analysis. The required altitude-dependent temperature variations can be caused by waves like tides (e.g. Marsh et al., 2006). A complication could be v' -dependent variations of the non-LTE contributions to the measured OH T_{rot} (see Sect. 4). As r is sensitive to measurement uncertainties, we only selected the most reliable OH band for each v' , i.e. we used (2-0), (3-1), (4-2), (5-3), (6-2), (7-4), (8-5), and (9-7). The results for the latter band are only based on 240 instead of 343 spectra due to a lack of wavelength coverage and issues with the flux calibration (see Sect. 2.1).

Figure 11 shows the mean h_{eff} vs. r for the X-shooter sample with O₂b(0-1) as a reference band ($r = 1$). There is a striking linear relation for the OH data, which supports our assumption of equidistant emission layers (Sect. 3.2.1). The

slope of a linear fit is 7.6 ± 0.4 km. The deviation of OH(9-7) may be explained by the different sample and the measurement uncertainties. O₂a(0-0) is very close to the fit and best agrees with OH(6-2). This is another argument for using this band for temperature comparisons with OH data. O₂b(0-1) deviates by about 2 K from the extrapolated fit, which is most probably caused by a non-linear relation of h_{eff} and r for the large emission height differences. The mesopause temperature minimum, which is close to the O₂b(0-1) emission layer (see Sect. 4), might have an impact. Our interpretations are supported by the r that we derived from the SABER-related T_{eff} for the X-shooter sample (Fig. 11). There is an almost perfect linear relation of h_{eff} and r for these data, which is in good agreement with the found equidistance of the OH layers and the position of O₂a(0-0) in relation to the OH bands. The position of O₂b(0-1) deviates in a similar way from a linear fit of the OH data with a slope of 16.0 ± 0.4 km. The correlations for T_{eff} are significantly stronger than for T_{rot} . The relatively low signal-to-noise ratio of the O₂b(0-1) data compared to those of the other bands can have a small impact. However, a direct correlation for e.g. $v' = 3$ and 8 revealed $r = 0.86$ for T_{rot} and 0.97 for T_{eff} . Possible variations of the OH T_{rot} non-LTE contributions do not appear to be sufficiently strong since the r difference for O₂a(0-0) is in good agreement with those of OH (cf. Sect. 4.2). Therefore, the best explanation for the r discrepancies is the weighted averaging of the SABER measurements to derive representative values for the X-shooter data set (Sect. 3.3). This effect can be estimated by calculating r for the T_{eff} of the original Cerro Paranal SABER sample with 1685 profiles (Sect. 2.2). As shown in Fig. 11, the OH r values for these data with a slope of 10.2 ± 0.1 km are very close to those for T_{rot} ; i.e. the r differences are indeed related to the averaging procedure. The position of O₂a(0-0) deviates from the trend for OH, which is probably caused by differences in the X-shooter and SABER sample properties. The latter also contributes to the remaining differences between the OH data points. Moreover, the limited vertical resolution of the SABER profiles can cause higher r for T_{eff} than for T_{rot} . The sampling of about 0.4 km (Sect. 2.2) is similar to the height difference of adjacent OH(v') emission layers (Sect. 3.2.1). Therefore, small-scale variations cannot be probed by SABER.

4 Results and discussion

After the derivation of T_{rot} corrected for the emission layer differences (Sect. 3), we can finally compare these temperatures and quantify the OH-related non-LTE effects depending on v' and time.

4.1 Temperature comparison for reference profile

Figure 12 gives an overview of the mean T_{rot} and their uncertainties for the eight OH(v'), O₂b(0-1), and O₂a(0-0) de-

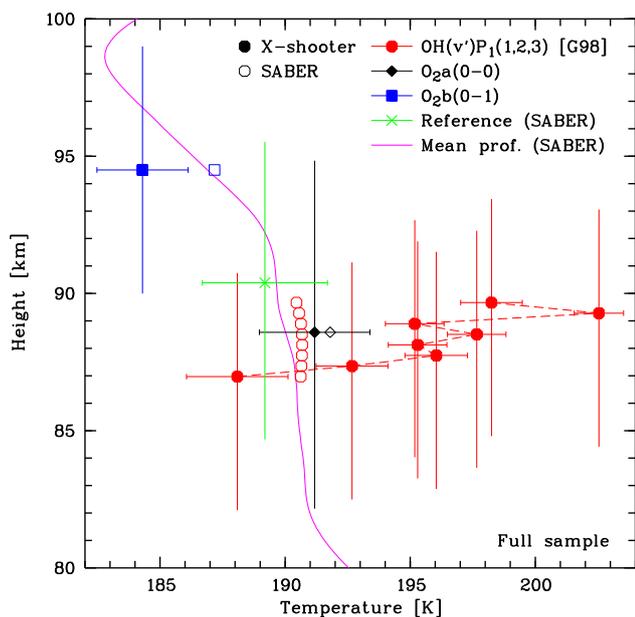


Figure 12. Comparison of temperatures in K and the corresponding heights in km for the full X-shooter sample. Solid symbols show X-shooter T_{rot} measurements and their uncertainties (excluding errors related to the molecular parameters where HITRAN was used) for OH from $v' = 2$ to 9 (cf. Fig. 11) using the first three P_1 lines (circles), $O_2a(0-0)$ (diamond), and $O_2b(0-1)$ (square). The effective emission heights were derived by means of SABER data. The bars in height direction mark the profile width at half maximum. The cross is related to the SABER T_{kin} for the $O_2a(0-0)$ -based reference profile. Open symbols indicate the T_{kin} results for the emission profiles of the bands measured by X-shooter. For these data, we assume temperature uncertainties of 2.5 K. The solid curve shows the mean T_{kin} profile for the X-shooter sample.

rived from the full X-shooter sample of 343 spectra (see Sect. 3.1). The T_{rot} are plotted in comparison to the corresponding SABER-based h_{eff} (Sect. 3.2) projected onto the X-shooter sample (Sect. 3.3). The FWHM of the emission profiles are indicated by vertical bars. The average $O_2a(0-0)$ profile is the widest with about 12.7 km. The figure shows a large overlap of the OH and $O_2a(0-0)$ emission profiles, as already demonstrated by Fig. 8. However, there are large differences in the measured OH T_{rot} , which range from 188.1 K for $v' = 2$ to 202.5 K for $v' = 8$. As discussed in Noll et al. (2015) and Sect. 1, the maximum at $v' = 8$ is typical of the $T_{\text{rot}}(v')$ pattern and can be explained by the contribution of a nearly nascent population, which shows a particularly hot rotational level population distribution for this v' . The mean T_{rot} of $O_2a(0-0)$ is at the lower end of the OH T_{rot} distribution with 191.2 ± 2.2 K. The $O_2b(0-1)$ T_{rot} of 184.3 ± 1.8 K is even lower. These temperatures can be compared to the SABER-based T_{eff} for the X-shooter sample with an estimated uncertainty of about 2.5 K (Sect. 3.3). While the T_{rot} and T_{eff} for each O_2 band agree within the errors, this is not the case for all the OH bands. The v' dependence of the tem-

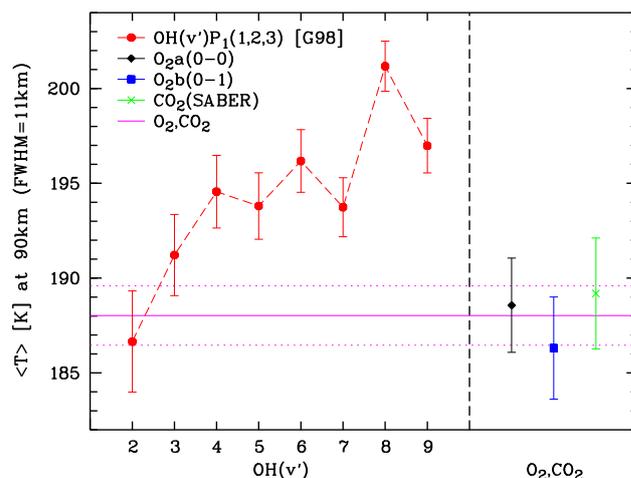


Figure 13. Comparison of mean temperatures in K for the $O_2a(0-0)$ -based reference profile and the full X-shooter sample. The v' -dependent OH T_{rot} (circles) are based on measurements of the first three P_1 lines and HITRAN-based molecular parameters (Goldman et al., 1998). The systematic errors related to the latter are not included in the plotted error bars. HITRAN-based molecular parameters were also used for the derivation of the T_{rot} from $O_2a(0-0)$ (diamond) and $O_2b(0-1)$ (square). The cross shows the CO_2 -based SABER T_{kin} for the reference profile and the full X-shooter sample. The mean and the uncertainties derived from the three non-OH temperatures are marked by horizontal lines.

perature is also very different. The OH T_{eff} only range from 190.5 to 190.7 K; i.e. the temperature gradient is almost zero. In Fig. 12, we also display the mean T_{kin} profile calculated for the X-shooter sample. Indeed, the mean temperature profile is relatively flat in the altitude range of the OH emission. Together with the large profile overlap, this explains the almost absent temperature gradient. This implies that the contribution of real temperature variations to the characteristic $T_{\text{rot}}(v')$ pattern is small. Non-LTE effects have to dominate. Only the relatively low T_{rot} of $O_2b(0-1)$ can be explained by the proximity of the mesopause temperature minimum.

Figure 12 also shows the h_{eff} , FWHM, and T_{eff} of the $O_2a(0-0)$ -based reference profile introduced in Sect. 3.2.3. T_{eff} is 189.2 ± 2.5 K. The different T_{rot} can be corrected to be representative of this profile that peaks at 90 km and has a FWHM of 11 km. This can be done by subtracting the difference between the T_{eff} related to a given band and the reference profile from the T_{rot} for the selected band (see Sect. 3.3). The resulting temperatures are shown in Fig. 13. For $O_2a(0-0)$ and $O_2b(0-1)$, we obtained 188.6 ± 2.5 K and 186.3 ± 2.7 K, respectively. The uncertainties are a combination of the measurement uncertainties already shown in Fig. 12 and the temperature correction errors discussed in Sect. 3.3. The two O_2 -related temperatures and the CO_2 -based SABER T_{eff} agree very well within their errors. Since these three temperatures are not affected by non-LTE effects, we can combine them, which results in an average tem-

perature of 188.0 ± 1.6 K. Clemesha et al. (2011) published a mean temperature profile based on sodium lidar measurements carried out at São José dos Campos in Brazil (23° S, 46° W) from 2007 to 2010. For our reference emission and their temperature profile, this implies a T_{eff} of 188 K, which is in remarkable agreement with our measurement despite their assumed absolute errors of about 5 K, differences in the time coverage (no January data), and the small but probably non-negligible differences in latitude and longitude. For the latter effect, see Fig. 2. The upper part of the lidar-based temperature profile is also in good agreement with the SABER one in Fig. 12. The lower part shows some discrepancies due to a different height and strength of the secondary minimum, which is at about 88 km. In the SABER case, there is only a nearly constant temperature down to about 81 km, which causes a lower onset of the steep temperature increase that is already visible in the Clemesha et al. (2011) profile above 80 km. The comparison with satellite and lidar data shows that the T_{rot} from $\text{O}_2a(0-0)$ and $\text{O}_2b(0-1)$ line measurements in X-shooter spectra provide reliable temperatures in the mesopause region.

The OH T_{rot} indicate significant non-LTE contributions in comparison to our LTE reference temperature of about 188 K. The temperatures range from 186.7 ± 2.7 K for $v' = 2$ to 201.2 ± 1.3 K for $v' = 8$. In the latter case, the difference is about 13 K. With the estimated errors, the $\Delta T_{\text{non-LTE}}$ of $v' = 8, 9$, and 6 are highly significant ($> 3\sigma$). The situation for $v' = 2$ and 3 is unclear. The OH bands (2-0) and (3-0), which by far show the lowest T_{rot} (Noll et al., 2015), are also affected by relatively large systematic uncertainties. OH(2-0) is partly absorbed by water vapour (more than any other band considered) and OH(3-0) is partly affected by line blending and close to the upper wavelength limit of the X-shooter VIS arm. Therefore, it cannot be excluded that the true temperatures of $v' = 2$ and 3 are similar to those of higher v' . The plotted temperatures are only valid if the T_{rot} are derived from the first three P_1 -branch lines. For example, the resulting $\Delta T_{\text{non-LTE}}$ and the significance of the non-LTE effects would increase if $P_1(4)$ was also used. In this case, the T_{rot} would be higher by about 1 K for all v' . The change to the first four P_2 -branch lines would even increase the OH T_{rot} by about 11 K. Moreover, the discussed OH T_{rot} are based on Einstein coefficients and level energies from the HITRAN database (Rothman et al., 2013), which are related to Goldman et al. (1998). The results change if we use the molecular parameters of van der Loo and Groenenboom (2007, 2008) (Fig. 14). The T_{rot} are more uncertain and tend to be lower than those based on HITRAN data. If we neglect the very uncertain data point for $v' = 2$ (see Noll et al., 2015), the mean shift is 4.2 K. Hence, only the $\Delta T_{\text{non-LTE}}$ of 10 K for $v' = 8$ remains significant with more than 3σ . Although the T_{rot} based on van der Loo and Groenenboom (2008) are less reliable than those based on Goldman et al. (1998), these results demonstrate that the uncertainties of the molecular parameters limit the accuracy of $\Delta T_{\text{non-LTE}}$ that can currently

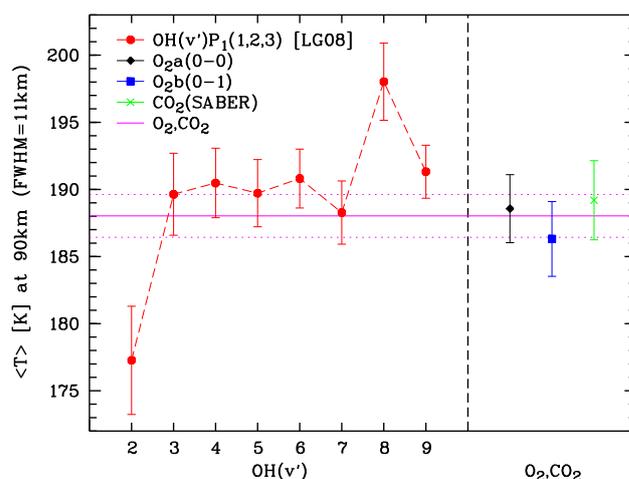


Figure 14. Same plot as Fig. 13 except for OH T_{rot} based on the molecular parameters of van der Loo and Groenenboom (2007, 2008).

be achieved. Nevertheless, $\Delta T_{\text{non-LTE}}$ of more than 10 K for $v' = 8$ are plausible if the first three P_1 -branch lines are used. Finally, the results show that the consideration of emission profile differences does not significantly change the characteristic $T_{\text{rot}}(v')$ pattern found by Cosby and Slanger (2007) and Noll et al. (2015). It can only be explained by non-LTE effects.

4.2 Nocturnal variations of temperature differences

Noll et al. (2015) studied the nocturnal variations of the OH T_{rot} by means of five nighttime periods, which we also used in Figs. 9 and 10. The analysis resulted in HITRAN-based T_{rot} differences (i.e. maximum minus minimum) between 4.8 K for $v' = 5$ and 7.0 K for $v' = 6$. In order to separate non-LTE from temperature profile variations, we used the observed $\text{O}_2a(0-0)$ profiles as reference for an emission profile correction. We did not take the reference profile at 90 km, which was the basis for the discussion in Sect. 4.1, since this reduces the uncertainties related to the temperature correction from about 1.2 to 0.9 K on average. At the beginning/end of the night, the uncertainties are higher for high/low v' . The resulting $\Delta T_{\text{non-LTE}}$ are displayed in Fig. 15. Compared to the uncorrected T_{rot} , the time-related variation for fixed v' significantly decreased. The $\Delta T_{\text{non-LTE}}$ differences range from 0.6 K for $v' = 3$ to 3.6 K for $v' = 9$. Hence, the average nocturnal variations of the non-LTE contributions to the OH T_{rot} are significantly lower than their actual values. Moreover, at least the low v' indicate that the T_{rot} variation is obviously dominated by the variability of the real temperature profile. However, it seems that only one-third of the $T_{\text{rot}}(v' = 9)$ variability can be explained in this way. Figure 15 indicates a nocturnal trend of increasing $\Delta T_{\text{non-LTE}}$. The trend appears to be stronger for higher v' . The significance is about 3σ for the differences between the first and last interval and is lim-

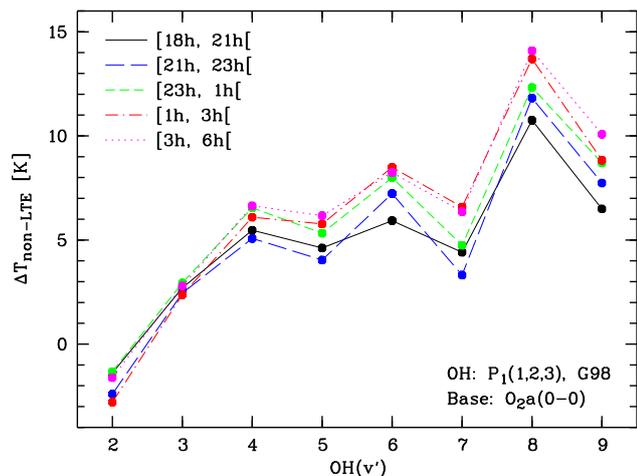


Figure 15. v' -dependent non-LTE contributions to OH T_{rot} based on the first three P_1 lines and HITRAN data (Goldman et al., 1998) for five nighttime LT periods derived from the full X-shooter sample. The LTE base temperature for each v' was derived from the corresponding $O_2a(0-0)$ T_{rot} corrected for the difference in the VER profiles. The uncertainties of the absolute $\Delta T_{\text{non-LTE}}$ are similar to those given in Fig. 13 (excluding the unknown uncertainties of the molecular parameters). Comparisons of $\Delta T_{\text{non-LTE}}$ for the different night-time periods and/or v' are more reliable. The errors should be of the order of 1 K, i.e. very likely smaller than 2 K.

ited by the statistical mean errors with an average of about 1.1 K and at least part of the temperature correction errors due to the changing $O_2a(0-0)$ emission profile.

Noll et al. (2015) discovered a significant increase of the T_{rot} differences of adjacent intermediate v' for the second compared to the first nighttime period. In Fig. 15, this causes a relatively strong $\Delta T_{\text{non-LTE}}$ increase for $v' = 6$ but a decrease for $v' = 5$ and 7. In order to better evaluate the reliability of these trends, we also calculated $\Delta T_{\text{non-LTE}}$ based on $OH(v' = 2)$ instead of $O_2a(0-0)$ emission profiles. Although the absolute $\Delta T_{\text{non-LTE}}$ are less trustworthy, and possible weak non-LTE T_{rot} variations for $v' = 2$ are neglected, this approach is justified for the comparisons of nighttime intervals and v' since it can further reduce the uncertainties related to the temperature correction. The emission profile differences between $OH(v' = 2)$ and the other OH bands are much smaller and less variable than those related to $O_2a(0-0)$. The resulting Fig. 16 does not show a $\Delta T_{\text{non-LTE}}$ decrease between the first and second nighttime period. However, significant increases appear to be limited to $v' = 6, 8,$ and 9. For $v' = 5$ and 7, the major change happens between the third and fourth period, i.e. after midnight. The last period indicates a decrease of $\Delta T_{\text{non-LTE}}$ for most v' . For $v' \geq 6$, the $\Delta T_{\text{non-LTE}}$ increases between the first and fourth period vary between 3.6 K ($v' = 7$) and 4.3 K ($v' = 8$). The significance is higher than for the results based on $O_2a(0-0)$ as reference band.

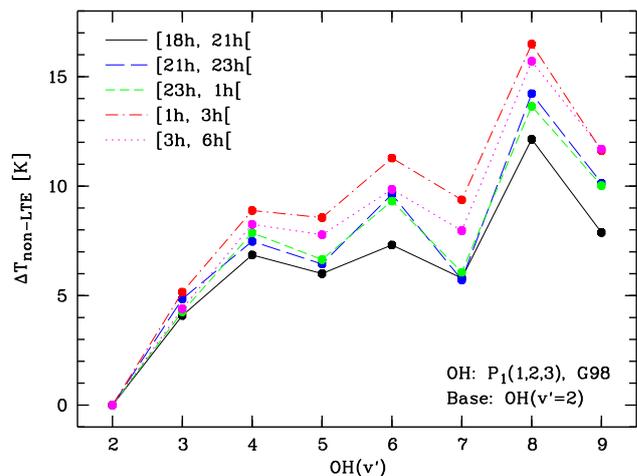


Figure 16. Same plot as Fig. 15 except for $OH(v' = 2)$ as base for the calculation of $\Delta T_{\text{non-LTE}}$. The errors for comparisons of different nighttime periods and/or v' are probably smaller than those of Fig. 15 since temperature correction uncertainties are less critical. The absolute $\Delta T_{\text{non-LTE}}$ are less reliable.

The found $\Delta T_{\text{non-LTE}}$ variations appear to be in agreement with those of the OH emission layer height (see Sect. 1). Up to the fourth nighttime period, the sample-averaged h_{eff} increases by 0.9 km ($v' = 9$) to 1.7 km ($v' = 2$). Afterwards the emission layers sink again except for $v' = 9$. See Fig. 10 for the h_{eff} trend related to $v' = 8$. As already discussed by Noll et al. (2015) based on T_{rot} differences and vibrational level populations, a higher emission layer is in a lower density environment, which decreases the frequency of thermalising collisions by O_2 (e.g. Adler-Golden, 1997; von Savigny et al., 2012; Xu et al., 2012) and, consequently, can increase the T_{rot} non-LTE effects. The SABER data show that the h_{eff} increase is more related to a VER decrease at low altitudes than an increase in the upper part of the emission layer. This trend mainly depends on changes in the distribution of atomic oxygen (which is required for the O_3 production) by chemical reactions and atmospheric dynamics (Lowe et al., 1996; Marsh et al., 2006; Nikoukar et al., 2007). In order to better estimate the effect of the VER profile change on the T_{rot} non-LTE effects, we calculated the fraction of OH emission above a certain altitude as a function of the nighttime period. For this, we used the original VER profiles of the two SABER channels at 1.64 and 2.06 μm corresponding to $v' = 4.6$ and 8.3, respectively (Sect. 3.2.1). For a minimum altitude of 95 km, we found an increase from 4.7 % (first period) to 7.3 % (fourth period) for $v' = 4.6$ and from 9.4 to 12.0 % for $v' = 8.3$. The differences are similar for both SABER channels but the absolute values are higher for $v' = 8.3$. The latter effect and the relative change of the emission fractions with time enhance with increasing minimum altitude. These results show the large impact of a relatively small rise of the emission layer on the highest parts of the

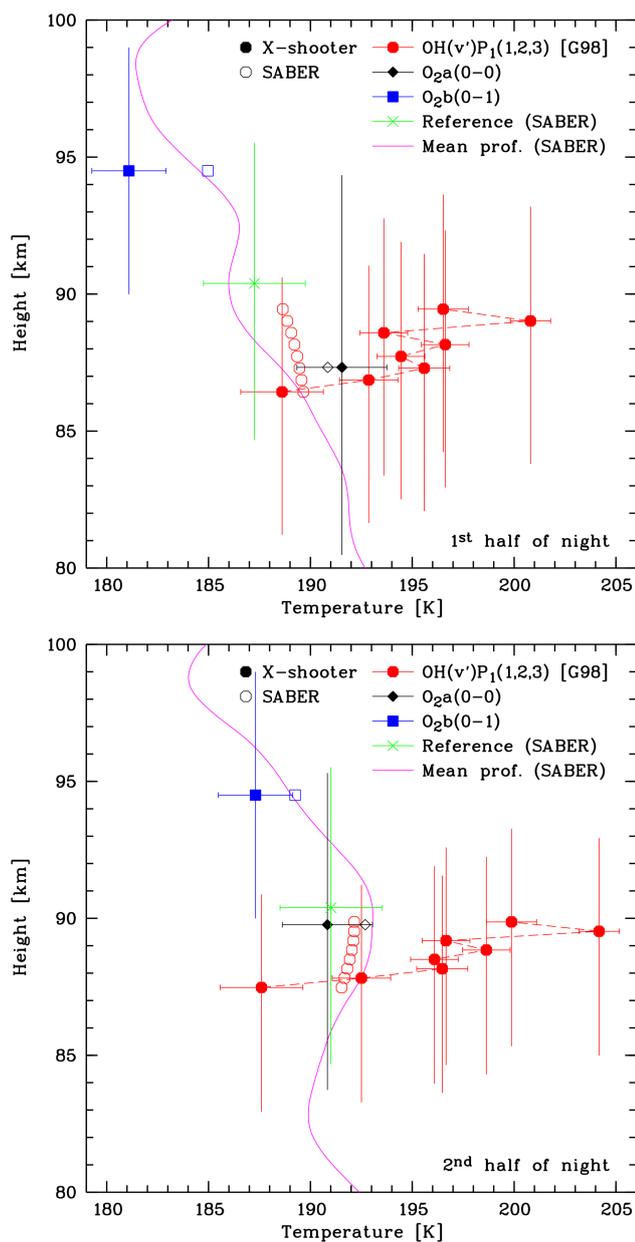


Figure 17. Same plot as Fig. 12 except for showing results for the first (upper panel) and the second half of the night (lower panel) separately.

VER profile. To fully understand Fig. 16, it will be necessary to model OH T_{rot} depending on v' and altitude, which has not been done so far. While there is some knowledge on the nascent population distribution over the rotational levels of the different v (Llewellyn and Long, 1978), it is not clear how this distribution is modified by collisional quenching.

Figure 17 illustrates the effect of the changing emission layers and temperature profile on the measured T_{rot} by comparing T_{rot} , T_{eff} , and T_{kin} for the first and second half of the night. The plots are similar to Fig. 12 discussed in Sect. 4.1.

The OH T_{rot} show a wider spread in the second half of the night. This can partly be explained by a significant change of the temperature profile. The SABER data reveal an average nocturnal trend from a single deep mesopause temperature minimum at the beginning to a double minimum, where the secondary trough is almost as deep as the primary one. The average altitudes of both minima are 98 to 99 km and between 80 and 85 km. Observations of a secondary minimum are not unusual (von Zahn et al., 1996; Shepherd et al., 2004; Friedman and Chu, 2007; Clemesha et al., 2011) and can be explained by tidal perturbations and tide-gravity wave interactions (Meriwether and Gardner, 2000; Meriwether and Gerrard, 2004). The appearance of a secondary mesopause minimum causes a temperature inversion at the main OH emission altitudes. The comparison of the second with the first half of the night, therefore, shows a T_{kin} increase above and decrease below 85 km. The OH T_{eff} are higher in the second period and this increase is stronger for higher v' due to the different temperature gradients. The amounts are 1.9 K for $v' = 2$ and 3.5 K for $v' = 9$. Nevertheless, a T_{rot} increase with v' remains, which is disturbed by the odd/even v' dichotomy, as indicated by Figs. 15 and 16. The O_2 -related T_{rot} follow the changes of the temperature profile convincingly. As expected, the emission height, profile width, and temperature for $\text{O}_2a(0-0)$ are much more similar to those for the reference profile (Sect. 3.2.3) in the second half of the night.

4.3 Seasonal variations of temperature differences

Noll et al. (2015) also studied the $T_{\text{rot}}(v')$ of the four meteorological seasons. The X-shooter data indicated higher temperatures for the equinoxes than for the solstices irrespective of v' . The maximum differences were between 7.4 K for $v' = 2$ and 3.3 K for $v' = 8$. After applying a temperature correction based on $\text{O}_2a(0-0)$, these differences decreased to 4.7 and 1.2 K, as shown in Fig. 18. These values imply stronger non-LTE effects near the equinoxes than near the solstices. Moreover, the effects seem to be stronger for lower v' . This is contradictory to the results and explanations discussed in Sect. 4.2. The apparent trend is most probably caused by an inefficient temperature correction. Although the reduction of the seasonal variations by the conversion from T_{rot} to $\Delta T_{\text{non-LTE}}$ was even the largest for low v' (2.7 vs. 2.1 K, see above), it was obviously insufficient to obtain reliable seasonal variations of the OH T_{rot} non-LTE contributions. This is probably related to the TIMED yaw cycle of 60 days, which causes large gaps in the seasonal coverage for fixed local times (Sects. 1 and 2.2). Therefore, the true seasonal variations could be stronger than those used for the temperature correction. At Cerro Paranal and similar latitudes, studying seasonal variations is more challenging than an analysis of the nocturnal variations since the combination of semi-annual and the weaker annual oscillations for OH T_{rot} (Takahashi et al., 1995; Gelinas et al., 2008) is more complex than

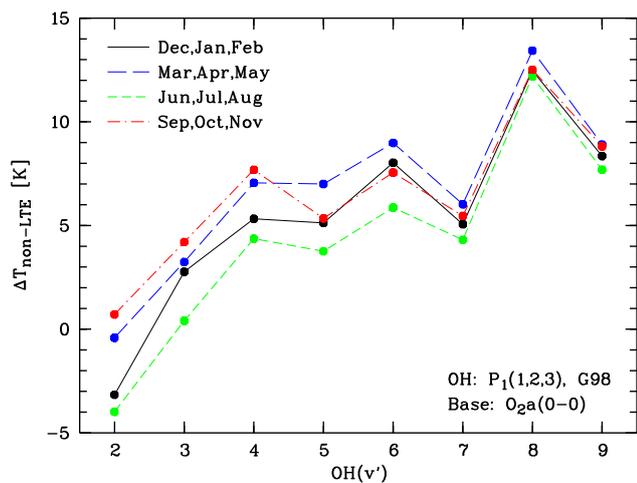


Figure 18. Same plot as Fig. 15 except for curves related to the four meteorological seasons instead of nighttime periods. The uncertainties for comparisons of $\Delta T_{\text{non-LTE}}$ for the different seasons and/or v' are somewhat higher. 2 K and more cannot be excluded.

the night trends (Takahashi et al., 1998; Gelinas et al., 2008; Noll et al., 2015), which are dominated by the solar migrating diurnal tide (Marsh et al., 2006; Smith, 2012). Hence, the limitations of our SABER and X-shooter data sets do not allow us to conclude on seasonal variations of the OH T_{rot} non-LTE effects.

5 Conclusions

Based on the VLT/X-shooter data set of Noll et al. (2015) with 343 spectra, we could compare their measurements of rotational temperatures T_{rot} from 25 OH bands covering $v' = 2$ to 9 with new results for the bands $\text{O}_2b(0-1)$ and $\text{O}_2a(0-0)$. For the latter band, it was just the second time that T_{rot} could be derived and the first time that it could be done with spectroscopically resolved lines. Since $\text{O}_2a(0-0)$ emits at similar altitudes as OH and the rotational level populations appear to be consistent with a Boltzmann distribution, the accurate T_{rot} for $\text{O}_2a(0-0)$ could be used to quantify the non-LTE contributions to the OH T_{rot} . To correct for the different emission profiles and the corresponding VER-weighted T_{eff} , we used OH, $\text{O}_2a(0-0)$, and CO_2 -based T_{kin} profile data from TIMED/SABER suitable for Cerro Paranal. The correction is particularly important for the beginning of the night when the $\text{O}_2a(0-0)$ emission peaks several km below the nighttime average due to a decaying daytime population from ozone photolysis. From the X-shooter data set, we obtained an effective population lifetime of about 46 min, which is in good agreement with SABER and other data. The required emission profiles for the different OH v' were derived from the two OH SABER channels under the assumption that emission peaks are equidistant for adjacent v' . This approach is supported by a comparison of the effective emission heights h_{eff}

with temperature correlation coefficients for the X-shooter and SABER samples, which showed nearly linear relations.

For an $\text{O}_2a(0-0)$ -based reference profile at 90 km, we compared the corrected temperature measurements and found that the results for $\text{O}_2b(0-1)$ (where a fixed emission profile at 94.5 km was assumed), $\text{O}_2a(0-0)$, and the SABER T_{kin} agree very well within the errors. The sample mean temperature at the reference altitude was 188.0 ± 1.6 K. Except for the uncertain $v' = 2$ and 3, the resulting OH-related temperatures were significantly higher. For the first three P_1 -branch lines and HITRAN molecular data, a maximum difference of about 13 K was achieved for $v' = 8$. The characteristic T_{rot} (v') pattern found by Cosby and Slanger (2007) and Noll et al. (2015) showing an odd/even v' dichotomy could also be seen in the corrected temperature data. The amount of non-LTE contributions to the OH T_{rot} $\Delta T_{\text{non-LTE}}$ strongly depends on the selected lines. A change of the line set affects all OH bands in a similar way. Apart from measurement and correction errors, the still uncertain molecular parameters limit the accuracy of the derived $\Delta T_{\text{non-LTE}}$ to a few K.

An analysis of the nocturnal variations of $\Delta T_{\text{non-LTE}}$ with five nighttime periods revealed that a significant fraction of the OH T_{rot} variability is caused by changes in the emission and mesopause kinetic temperature profiles. However, fractions above 50 % could only be found for low v' . At Cerro Paranal, the residual non-LTE effects tend to increase during the night and to be more variable at higher v' . The largest $\Delta T_{\text{non-LTE}}$ differences were found for the first (before 21:00 LT) and fourth period (between 01:00 and 03:00 LT), and $v' \geq 6$. They amount to about 4 K with an uncertainty of about 2 K. The amplitude of the variations appears to be only a minor fraction of the total amount of the non-LTE effects. The observed trends could be explained by the nocturnal rise of the OH emission layer, which we see in the SABER data up to the fourth nighttime period, and a corresponding reduction of the thermalising collisions due to the lower effective O_2 density. Higher v' could be more affected since the related emissions peak at higher altitudes. We also studied possible seasonal variations of $\Delta T_{\text{non-LTE}}$. However, the limitations of the SABER and X-shooter data sets did not allow us to find convincing trends.

In agreement with Noll et al. (2015), non-LTE effects are critical for absolute mesopause temperature estimates from OH T_{rot} and comparisons of bands with different v' . Dynamical studies based on T_{rot} derived from OH bands with high v' could also be significantly affected. In this respect, only low v' appear to be sufficiently safe.

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References

- Adler-Golden, S.: Kinetic parameters for OH nightglow modeling consistent with recent laboratory measurements, *J. Geophys. Res.*, 102, 19969–19976, doi:10.1029/97JA01622, 1997.
- Baker, D. J. and Stair Jr., A. T.: Rocket measurements of the altitude distributions of the hydroxyl airglow, *Phys. Scripta*, 37, 611–622, doi:10.1088/0031-8949/37/4/021, 1988.
- Baker, D. J., Thurgood, B. K., Harrison, W. K., Mlynczak, M. G., and Russell, J. M.: Equatorial enhancement of the nighttime OH mesospheric infrared airglow, *Phys. Scripta*, 75, 615–619, doi:10.1088/0031-8949/75/5/004, 2007.
- Bates, D. R. and Nicolet, M.: The photochemistry of atmospheric water vapor, *J. Geophys. Res.*, 55, 301–327, doi:10.1029/JZ055i003p00301, 1950.
- Beig, G., Keckhut, P., Lowe, R. P., Roble, R. G., Mlynczak, M. G., Scheer, J., Fomichev, V. I., Offermann, D., French, W. J. R., Shepherd, M. G., Semenov, A. I., Remsberg, E. E., She, C. Y., Lübken, F. J., Bremer, J., Clemesha, B. R., Stegman, J., Sigernes, F., and Fadnavis, S.: Review of mesospheric temperature trends, *Rev. Geophys.*, 41, RG1015, doi:10.1029/2002RG000121, 2003.
- Beig, G., Scheer, J., Mlynczak, M. G., and Keckhut, P.: Overview of the temperature response in the mesosphere and lower thermosphere to solar activity, *Rev. Geophys.*, 46, RG3002, doi:10.1029/2007RG000236, 2008.
- Charters, P. E., MacDonald, R. G., and Polanyi, J. C.: Formation of vibrationally excited OH by the reaction $H + O_3$, *Appl. Optics*, 10, 1747–1754, doi:10.1364/AO.10.001747, 1971.
- Clemesha, B., Simonich, D., and Batista, P.: Sodium lidar measurements of mesopause region temperatures at 23° S, *Adv. Space Res.*, 47, 1165–1171, doi:10.1016/j.asr.2010.11.030, 2011.
- Clough, S. A., Shephard, M. W., Mlawer, E. J., Delamere, J. S., Iacono, M. J., Cady-Pereira, K., Boukabara, S., and Brown, P. D.: Atmospheric radiative transfer modeling: a summary of the AER codes, *J. Quant. Spectrosc. Ra.*, 91, 233–244, doi:10.1016/j.jqsrt.2004.05.058, 2005.
- Cosby, P. C. and Slanger, T. G.: OH spectroscopy and chemistry investigated with astronomical sky spectra, *Can. J. Phys.*, 85, 77–99, 2007.
- Evans, W. F. J., Hunten, D. M., Llewellyn, E. J., and Vallance Jones, A.: Altitude profile of the infrared atmospheric system of oxygen in the dayglow, *J. Geophys. Res.*, 73, 2885–2896, doi:10.1029/JA073i009p02885, 1968.
- Friedman, J. S. and Chu, X.: Nocturnal temperature structure in the mesopause region over the Arecibo Observatory (18.35° N, 66.75° W): seasonal variations, *J. Geophys. Res.*, 112, D14107, doi:10.1029/2006JD008220, 2007.
- Gamache, R. R. and Goldman, A.: Einstein A coefficient, integrated band intensity, and population factors: application to the $a^1\Delta_g - X^3\Sigma_g^-(0-0)$ O₂ band, *J. Quant. Spectrosc. Ra.*, 69, 389–401, doi:10.1016/S0022-4073(00)00072-8, 2001.
- Gamache, R. R., Goldman, A., and Rothman, L. S.: Improved spectral parameters for the three most abundant isotopomers of the oxygen molecule, *J. Quant. Spectrosc. Ra.*, 59, 495–509, doi:10.1016/S0022-4073(97)00124-6, 1998.
- Gao, H., Xu, J. Y., Chen, G. M., Yuan, W., and Beletsky, A. B.: Global distributions of OH and O₂ (1.27 μm) nightglow emissions observed by TIMED satellite, *Sci. China Ser. E*, 54, 447–456, doi:10.1007/s11431-010-4236-5, 2011.
- Gelinas, L. J., Hecht, J. H., Walterscheid, R. L., Roble, R. G., and Woithe, J. M.: A seasonal study of mesospheric temperatures and emission intensities at Adelaide and Alice Springs, *J. Geophys. Res.*, 113, A01304, doi:10.1029/2007JA012587, 2008.
- Goldman, A., Schoenfeld, W. G., Goorvitch, D., Chackerian Jr., C., Dothe, H., Mélen, F., Abrams, M. C., and Selby, J. E. A.: Updated line parameters for OH $X^2\Pi - X^2\Pi$ (v' , v'') transitions, *J. Quant. Spectrosc. Ra.*, 59, 453–469, doi:10.1016/S0022-4073(97)00112-X, 1998.
- Greer, R. G. H., Llewellyn, E. J., Solheim, B. H., and Witt, G.: The excitation of O₂($b^1\Sigma_g^+$) in the nightglow, *Planet. Space Sci.*, 29, 383–389, doi:10.1016/0032-0633(81)90081-7, 1981.
- Greer, R. G. H., Murtagh, D. P., McDade, I. C., Dickinson, P. H. G., and Thomas, L.: ETON 1 – a data base pertinent to the study of energy transfer in the oxygen nightglow, *Planet. Space Sci.*, 34, 771–788, doi:10.1016/0032-0633(86)90074-7, 1986.
- Hays, P. B., Kafkalidis, J. F., Skinner, W. R., and Roble, R. G.: A global view of the molecular oxygen night airglow, *J. Geophys. Res.*, 108, 4646, doi:10.1029/2003JD003400, 2003.
- Howell, C. D., Michelangeli, D. V., Allen, M., Yung, Y. L., and Thomas, R. J.: SME observations of O₂($a^1\Delta_g$) nightglow – an assessment of the chemical production mechanisms, *Planet. Space Sci.*, 38, 529–537, doi:10.1016/0032-0633(90)90145-G, 1990.
- Kaufmann, M., Gusev, O. A., Grossmann, K. U., Martín-Torres, F. J., Marsh, D. R., and Kutepov, A. A.: Satellite observations of daytime and nighttime ozone in the mesosphere and lower thermosphere, *J. Geophys. Res.*, 108, 4272, doi:10.1029/2002JD002800, 2003.
- Kausch, W., Noll, S., Smette, A., Kimeswenger, S., Barden, M., Szyszka, C., Jones, A. M., Sana, H., Horst, H., and Kerber, F.: Molecfit: a general tool for telluric absorption correction, II. Quantitative evaluation on ESO-VLT/X-shooter spectra, *Astron. Astrophys.*, 576, A78, doi:10.1051/0004-6361/201423909, 2015.
- Khomich, V. Y., Semenov, A. I., and Shefov, N. N.: Airglow as an Indicator of Upper Atmospheric Structure and Dynamics, Springer, Berlin, Germany, 2008.
- Lafferty, W. J., Solodov, A. M., Lugez, C. L., and Fraser, G. T.: Rotational line strengths and self-pressure-broadening coefficients for the 1.27-μm, $a^1\Delta_g - X^3\Sigma_g^-, v = (0-0)$ band of O₂, *Appl. Optics*, 37, 2264–2270, doi:10.1364/AO.37.002264, 1998.

- Leshchishina, O., Kassi, S., Gordon, I. E., Rothman, L. S., Wang, L., and Campargue, A.: High sensitivity CRDS of the $a^1\Delta_g-X^3\Sigma_g^-$ band of oxygen near 1.27 μm : extended observations, quadrupole transitions, hot bands and minor isotopologues, *J. Quant. Spectrosc. Ra.*, 111, 2236–2245, doi:10.1016/j.jqsrt.2010.05.014, 2010.
- Lipatov, K. V.: Empirical model of variations in the IR atmospheric system of molecular oxygen: 2. emitting layer height, *Geomagn. Aeron.*, 53, 104–112, doi:10.1134/S001679321301012X, 2013.
- Liu, G. and Shepherd, G. G.: An empirical model for the altitude of the OH nightglow emission, *Geophys. Res. Lett.*, 33, L09805, doi:10.1029/2005GL025297, 2006.
- Llewellyn, E. J. and Long, B. H.: The OH Meinel bands in the airglow – the radiative lifetime, *Can. J. Phys.*, 56, 581–586, doi:10.1139/p78-076, 1978.
- López-González, M. J., López-Moreno, J. J., López-Valverde, M. A., and Rodrigo, R.: Behaviour of the O₂ infrared atmospheric (0-0) band in the middle atmosphere during evening twilight and at night, *Planet. Space Sci.*, 37, 61–72, doi:10.1016/0032-0633(89)90069-X, 1989.
- López-González, M. J., López-Moreno, J. J., and Rodrigo, R.: The altitude profile of the Infrared Atmospheric System of O₂ in twilight and early night – Derivation of ozone abundances, *Planet. Space Sci.*, 40, 1391–1397, doi:10.1016/0032-0633(92)90094-5, 1992.
- Lowe, R. P., Leblanc, L. M., and Gilbert, K. L.: WINDII/UARS observation of twilight behaviour of the hydroxyl airglow, at mid-latitude equinox, *J. Atmos. Terr. Phys.*, 58, 1863–1869, doi:10.1016/0021-9169(95)00178-6, 1996.
- Marsh, D. R., Smith, A. K., Mlynczak, M. G., and Russell, J. M.: SABER observations of the OH Meinel airglow variability near the mesopause, *J. Geophys. Res.*, 111, A10S05, doi:10.1029/2005JA011451, 2006.
- McDade, I. C., Llewellyn, E. J., Greer, R. G. H., and Murtagh, D. P.: Eton 6 – a rocket measurement of the O₂ infrared atmospheric (0-0) band in the nightglow, *Planet. Space Sci.*, 35, 1541–1552, doi:10.1016/0032-0633(87)90079-1, 1987.
- Meinel, A. B.: OH emission bands in the spectrum of the night sky. II, *Astrophys. J.*, 112, 120–130, doi:10.1086/145321, 1950.
- Melo, S. M. L., Lowe, R. P., and Takahashi, H.: The nocturnal behavior of the hydroxyl airglow at the equatorial and low latitudes as observed by WINDII: comparison with ground-based measurements, *J. Geophys. Res.*, 104, 24657–24666, doi:10.1029/1999JA900291, 1999.
- Meriwether, J. W. and Gardner, C. S.: A review of the mesosphere inversion layer phenomenon, *J. Geophys. Res.*, 105, 12405–12416, doi:10.1029/2000JD900163, 2000.
- Meriwether, J. W. and Gerrard, A. J.: Mesosphere inversion layers and stratosphere temperature enhancements, *Rev. Geophys.*, 42, RG3003, doi:10.1029/2003RG000133, 2004.
- Mertens, C. J., Mlynczak, M. G., López-Puertas, M., Wintersteiner, P. P., Picard, R. H., Winick, J. R., Gordley, L. L., and Russell III, J. M.: Retrieval of mesospheric and lower thermospheric kinetic temperature from measurements of CO₂ 15 μm Earth limb emission under non-LTE conditions, *Geophys. Res. Lett.*, 28, 1391–1394, doi:10.1029/2000GL012189, 2001.
- Mertens, C. J., Schmidlin, F. J., Goldberg, R. A., Remsberg, E. E., Pesnell, W. D., Russell III, J. M., Mlynczak, M. G., López-Puertas, M., Wintersteiner, P. P., Picard, R. H., Winick, J. R., and Gordley, L. L.: SABER observations of mesospheric temperatures and comparisons with falling sphere measurements taken during the 2002 summer MaCWAVE campaign, *Geophys. Res. Lett.*, 31, L03105, doi:10.1029/2003GL018605, 2004.
- Mies, F. H.: Calculated vibrational transition probabilities of OH($X^2\Pi$), *J. Mol. Spectrosc.*, 53, 150–188, doi:10.1016/0022-2852(74)90125-8, 1974.
- Mlynczak, M. G., Martin-Torres, F. J., Crowley, G., Kratz, D. P., Funke, B., Lu, G., Lopez-Puertas, M., Russell III, J. M., Kozyra, J., Mertens, C., Sharma, R., Gordley, L., Picard, R., Winick, J., and Paxton, L.: Energy transport in the thermosphere during the solar storms of April 2002, *J. Geophys. Res.*, 110, A12S25, doi:10.1029/2005JA011141, 2005.
- Mlynczak, M. G., Marshall, B. T., Martin-Torres, F. J., Russell III, J. M., Thompson, R. E., Remsberg, E. E., and Gordley, L. L.: Sounding of the atmosphere using broadband emission radiometry observations of daytime mesospheric O₂($^1\Delta$) 1.27 μm emission and derivation of ozone, atomic oxygen, and solar and chemical energy deposition rates, *J. Geophys. Res.*, 112, D15306, doi:10.1029/2006JD008355, 2007.
- Modigliani, A., Goldoni, P., Royer, F., Haigron, R., Guglielmi, L., François, P., Horrobin, M., Bristow, P., Vernet, J., Moehler, S., Kerber, F., Ballester, P., Mason, E., and Christensen, L.: The X-shooter pipeline, *SPIE Proc. Ser.*, 7737, 773728, doi:10.1117/12.857211, 2010.
- Moehler, S., Modigliani, A., Freudling, W., Giammichele, N., Gianninas, A., Gonneau, A., Kausch, W., Lançon, A., Noll, S., Rauch, T., and Vinther, J.: Flux calibration of medium-resolution spectra from 300 nm to 2500 nm: model reference spectra and telluric correction, *Astron. Astrophys.*, 568, doi:10.1051/0004-6361/201423790, 2014.
- Mulligan, F. J. and Galligan, J. M.: Mesopause temperatures calculated from the O₂($a^1\Delta_g$) twilight airglow emission recorded at Maynooth (53.2° N, 6.4° W), *Ann. Geophys.*, 13, 558–566, doi:10.1007/s00585-995-0558-1, 1995.
- Nair, H. and Yee, J.-H.: O₂($a^1\Delta_g$, $v=0$) chemical loss coefficients determined from SABER sunset measurements, *Geophys. Res. Lett.*, 36, L15829, doi:10.1029/2009GL039335, 2009.
- Newman, S. M., Lane, I. C., Orr-Ewing, A. J., Newnham, D. A., and Ballard, J.: Integrated absorption intensity and Einstein coefficients for the O₂ $a^1\Delta_g-X^3\Sigma_g^-(0,0)$ transition: a comparison of cavity ringdown and high resolution Fourier transform spectroscopy with a long-path absorption cell, *J. Chem. Phys.*, 110, 10749–10757, doi:10.1063/1.479018, 1999.
- Nikoukar, R., Swenson, G. R., Liu, A. Z., and Kamalabadi, F.: On the variability of mesospheric OH emission profiles, *J. Geophys. Res.*, 112, D19109, doi:10.1029/2007JD008601, 2007.
- Noll, S., Kausch, W., Barden, M., Jones, A. M., Szyszka, C., Kimeswenger, S., and Vinther, J.: An atmospheric radiation model for Cerro Paranal. I. The optical spectral range, *Astron. Astrophys.*, 543, doi:10.1051/0004-6361/201219040, 2012.
- Noll, S., Kausch, W., Kimeswenger, S., Unterguggenberger, S., and Jones, A. M.: OH populations and temperatures from simultaneous spectroscopic observations of 25 bands, *Atmos. Chem. Phys.*, 15, 3647–3669, doi:10.5194/acp-15-3647-2015, 2015.
- Ohoyama, H., Kasai, T., Yoshimura, Y., Kimura, H., and Kuwata, K.: Initial distribution of vibration of the OH radicals produced in the $\text{H} + \text{O}_3 \rightarrow \text{OH}(X^2\Pi_{1/2,3/2}) + \text{O}_2$ reaction.

- Chemiluminescence by a crossed beam technique, *Chem. Phys.*, 118, 263–266, doi:10.1016/0009-2614(85)85312-4, 1985.
- Patat, F., Moehler, S., O'Brien, K., Pompei, E., Bensby, T., Carraro, G., de Ugarte Postigo, A., Fox, A., Gavignaud, I., James, G., Korhonen, H., Ledoux, C., Randall, S., Sana, H., Smoker, J., Stefl, S., and Szeifert, T.: Optical atmospheric extinction over Cerro Paranal, *Astron. Astrophys.*, 527, doi:10.1051/0004-6361/201015537, 2011.
- Pendleton Jr., W., Espy, P., Baker, D., Steed, A., and Fetrow, M.: Observation of OH Meinel (7, 4) $P(N'' = 13)$ transitions in the night airglow, *J. Geophys. Res.*, 94, 505–510, doi:10.1029/JA094iA01p00505, 1989.
- Pendleton Jr., W. R., Espy, P. J., and Hammond, M. R.: Evidence for non-local-thermodynamic-equilibrium rotation in the OH nightglow, *J. Geophys. Res.*, 98, 11567–11579, doi:10.1029/93JA00740, 1993.
- Pendleton Jr., W. R., Baker, D. J., Reese, R. J., and O'Neil, R. R.: Decay of $O_2(a_1\Delta_g)$ in the evening twilight airglow: implications for the radiative lifetime, *Geophys. Res. Lett.*, 23, 1013–1016, doi:10.1029/96GL00946, 1996.
- Perminov, V. I. and Lipatov, K. V.: Empirical model of variations in the emission of the Infrared Atmospheric system of molecular oxygen: 3. temperature, *Geomagn. Aeron.*, 53, 389–396, doi:10.1134/S0016793213030158, 2013.
- Perminov, V. I. and Semenov, A. I.: Nonequilibrium of the rotational temperature of OH bands with high vibrational excitation, *Geomagn. Aeron.*, 32, 175–178, 1992.
- Perminov, V. I., Semenov, A. I., and Shefov, N. N.: On rotational temperature of the hydroxyl emission, *Geomagn. Aeron.*, 47, 756–763, doi:10.1134/S0016793207060084, 2007.
- Reisin, E. R., Scheer, J., Dyrland, M. E., Sigernes, F., Deehr, C. S., Schmidt, C., Höppner, K., Bittner, M., Ammosov, P. P., Gavrielyeva, G. A., Stegman, J., Perminov, V. I., Semenov, A. I., Knieling, P., Koppmann, R., Shiokawa, K., Lowe, R. P., López-González, M. J., Rodríguez, E., Zhao, Y., Taylor, M. J., Burity, R. A., Espy, P. J., French, W. J. R., Eichmann, K.-U., Burrows, J. P., and von Savigny, C.: Traveling planetary wave activity from mesopause region airglow temperatures determined by the Network for the Detection of Mesospheric Change (NDMC), *J. Atmos. Sol.-Terr. Phys.*, 119, 71–82, doi:10.1016/j.jastp.2014.07.002, 2014.
- Remsberg, E. E., Marshall, B. T., Garcia-Comas, M., Krueger, D., Lingenfelter, G. S., Martin-Torres, J., Mlynczak, M. G., Russell III, J. M., Smith, A. K., Zhao, Y., Brown, C., Gordley, L. L., Lopez-Gonzalez, M. J., Lopez-Puertas, M., She, C.-Y., Taylor, M. J., and Thompson, R. E.: Assessment of the quality of the version 1.07 temperature-versus-pressure profiles of the middle atmosphere from TIMED/SABER, *J. Geophys. Res.*, 113, D17101, doi:10.1029/2008JD010013, 2008.
- Rezac, L., Kutepov, A., Russell III, J. M., Feofilov, A. G., Yue, J., and Goldberg, R. A.: Simultaneous retrieval of $T(p)$ and CO_2 VMR from two-channel non-LTE limb radiances and application to daytime SABER/TIMED measurements, *J. Atmos. Sol.-Terr. Phys.*, 130, 23–42, doi:10.1016/j.jastp.2015.05.004, 2015.
- Rothman, L. S., Gordon, I. E., Barbe, A., Chris Benner, D., Bernath, P. F., Birk, M., Boudon, V., Brown, L. R., Campargue, A., Champion, J.-P., Chance, K., Coudert, L. H., Dana, V., Devi, V. M., Fally, S., Flaud, J.-M., Gamache, R. R., Goldman, A., Jacquemart, D., Kleiner, I., Lacome, N., Lafferty, W. J., Mandin, J.-Y., Massie, S. T., Mikhailenko, S. N., Miller, C. E., Moazzen-Ahmadi, N., Naumenko, O. V., Nikitin, A. V., Orphal, J., Perevalov, V. I., Perrin, A., Predoi-Cross, A., Rinsland, C. P., Rotger, M., Šimečková, M., Smith, M. A. H., Sung, K., Tashkun, S. A., Tennyson, J., Toth, R. A., Vandaele, A. C., and Vander Auwera, J.: The HITRAN 2008 molecular spectroscopic database, *J. Quant. Spectrosc. Ra.*, 110, 533–572, doi:10.1016/j.jqsrt.2009.02.013, 2009.
- Rothman, L. S., Gordon, I. E., Babikov, Y., Barbe, A., Chris Benner, D., Bernath, P. F., Birk, M., Bizzocchi, L., Boudon, V., Brown, L. R., Campargue, A., Chance, K., Cohen, E. A., Coudert, L. H., Devi, V. M., Drouin, B. J., Fayt, A., Flaud, J.-M., Gamache, R. R., Harrison, J. J., Hartmann, J.-M., Hill, C., Hodges, J. T., Jacquemart, D., Jolly, A., Lamouroux, J., Le Roy, R. J., Li, G., Long, D. A., Lyulin, O. M., Mackie, C. J., Massie, S. T., Mikhailenko, S., Müller, H. S. P., Naumenko, O. V., Nikitin, A. V., Orphal, J., Perevalov, V., Perrin, A., Polovtseva, E. R., Richard, C., Smith, M. A. H., Starikova, E., Sung, K., Tashkun, S., Tennyson, J., Toon, G. C., Tyuterev, V. G., and Wagner, G.: The HITRAN2012 molecular spectroscopic database, *J. Quant. Spectrosc. Ra.*, 130, 4–50, doi:10.1016/j.jqsrt.2013.07.002, 2013.
- Rousselot, P., Lidman, C., Cuby, J.-G., Moreels, G., and Monnet, G.: Night-sky spectral atlas of OH emission lines in the near-infrared, *Astron. Astrophys.*, 354, 1134–1150, 2000.
- Russell III, J. M., Mlynczak, M. G., Gordley, L. L., Tansock Jr., J. J., and Esplin, R. W.: Overview of the SABER experiment and preliminary calibration results, *SPIE Proc. Ser.*, 3756, 277–288, doi:10.1117/12.366382, 1999.
- Scheer, J. and Reisin, E. R.: Refinements of a classical technique of airglow spectroscopy, *Adv. Space Res.*, 27, 1153–1158, doi:10.1016/S0273-1177(01)00150-8, 2001.
- Schmidt, C., Höppner, K., and Bittner, M.: A ground-based spectrometer equipped with an InGaAs array for routine observations of OH(3-1) rotational temperatures in the mesopause region, *J. Atmos. Sol.-Terr. Phys.*, 102, 125–139, doi:10.1016/j.jastp.2013.05.001, 2013.
- Semenov, A. I., Shefov, N. N., and Medvedeva, I. V.: Empirical model of variations in the continuum emission in the upper atmosphere. 1. Intensity, *Geomagn. Aeron.*, 54, 488–499, doi:10.1134/S0016793214040148, 2014.
- Shepherd, M. G., Evans, W. F. J., Hernandez, G., Offermann, D., and Takahashi, H.: Global variability of mesospheric temperature: mean temperature field, *J. Geophys. Res.*, 109, D24117, doi:10.1029/2004JD005054, 2004.
- Shiokawa, K., Otsuka, Y., Suzuki, S., Katoh, T., Katoh, Y., Satoh, M., Ogawa, T., Takahashi, H., Gobbi, D., Nakamura, T., Williams, B. P., She, C.-Y., Taguchi, M., and Shimomai, T.: Development of airglow temperature photometers with cooled-CCD detectors, *Earth Planets Space*, 59, 585–599, doi:10.1186/BF03352721, 2007.
- Smette, A., Sana, H., Noll, S., Horst, H., Kausch, W., Kimeswenger, S., Barden, M., Szyszka, C., Jones, A. M., Galenne, A., Vinther, J., Ballester, P., and Taylor, J.: Molecfit: a general tool for telluric absorption correction. I. Method and application to ESO instruments, *Astron. Astrophys.*, 576, doi:10.1051/0004-6361/201423932, 2015.
- Smith, A. K.: Global dynamics of the MLT, *Surv. Geophys.*, 33, 1177–1230, doi:10.1007/s10712-012-9196-9, 2012.

- Takahashi, H., Clemesha, B. R., and Batista, P. P.: Predominant semi-annual oscillation of the upper mesospheric airglow intensities and temperatures in the equatorial region, *J. Atmos. Terr. Phys.*, 57, 407–414, 1995.
- Takahashi, H., Gobbi, D., Batista, P. P., Melo, S. M. L., Teixeira, N. R., and Buriti, R. A.: Dynamical influence on the equatorial airglow observed from the south american sector, *Adv. Space Res.*, 2, 817–825, doi:10.1016/S0273-1177(97)00680-7, 1998.
- Vallance Jones, A. and Gattinger, R. L.: The seasonal variation and excitation mechanism of the $1.58\mu^1\Delta_g\text{-}^3\Sigma_g^-$ twilight airglow band, *Planet. Space Sci.*, 11, 961–974, doi:10.1016/0032-0633(63)90123-5, 1963.
- van der Loo, M. P. J. and Groenenboom, G. C.: Theoretical transition probabilities for the OH Meinel system, *J. Chem. Phys.*, 126, 114314, doi:10.1063/1.2646859, 2007.
- van der Loo, M. P. J. and Groenenboom, G. C.: Erratum: “Theoretical transition probabilities for the OH Meinel system” [*J. Chem. Phys.* 126, 114314], *J. Chem. Phys.*, 128, 159902, doi:10.1063/1.2899016, 2008.
- van Rhijn, P. J.: On the brightness of the sky at night and the total amount of starlight, *Publ. Astr. Lab. Groningen*, 31, 1–83, 1921.
- Vernet, J., Dekker, D’Odorico, S., Kaper, L., Kjaergaard, P., Hammer, F., Randich, S., Zerbi, F., Groot, P. M., Hjorth, J., Guinouard, I., Navarro, R., Adolfe, T., Albers, P. W., Amans, J.-P., Andersen, J. J., Andersen, M. I., Binetruy, P., Bristow, P., Castillo, R., Chemla, F., Christensen, L., Conconi, P., Conzelmann, R., Dam, J., De Caprio, V., De Ugarte Postigo, A., Delabre, B., Di Marcantonio, P., Downing, M., Elswijk, E., Finger, G., Fischer, G., Flores, H., François, P., Goldoni, P., Guglielmi, L., Haignon, R., Hanenbourg, H., Hendriks, I., Horrobin, M., Horville, D., Jessen, N. C., Kerber, F., Kern, L., Kieckbusch, M., Kleszcz, P., Klougart, J., Kragt, J., Larsen, H. H., Lizon, J.-L., Lucuix, C., Mainieri, V., Manuputy, R., Martayan, C., Mason, E., Mazzoleni, R., Michaelsen, N., Modigliani, A., Moehler, S., Møller, P., Norup Sørensen, A., Nørregaard, P., Péroux, C., Patat, F., Pena, E., Pragt, J., Reinero, C., Rigal, F., Riva, M., Roelfsema, R., Royer, F., Sacco, G., Santin, P., Schoenmaker, T., Spano, P., Sweers, E., Ter Horst, R., Tintori, M., Tromp, N., van Dael, P., van der Vliet, H., Venema, L., Vidali, M., Vinther, J., Vola, P., Winters, R., Wistisen, D., Wulterkens, G., and Zacchei, A.: X-shooter, the new wide band intermediate resolution spectrograph at the ESO Very Large Telescope, *Astron. Astrophys.*, 536, doi:10.1051/0004-6361/201117752, 2011.
- von Savigny, C.: Variability of OH(3-1) emission altitude from 2003 to 2011: long-term stability and universality of the emission rate–altitude relationship, *J. Atmos. Sol.-Terr. Phys.*, 127, 120–128, doi:10.1016/j.jastp.2015.02.001, 2015.
- von Savigny, C., McDade, I. C., Eichmann, K.-U., and Burrows, J. P.: On the dependence of the OH* Meinel emission altitude on vibrational level: SCIAMACHY observations and model simulations, *Atmos. Chem. Phys.*, 12, 8813–8828, doi:10.5194/acp-12-8813-2012, 2012.
- von Zahn, U., Höffner, J., Eska, V., and Alpers, M.: The mesopause altitude: only two distinctive levels worldwide?, *Geophys. Res. Lett.*, 23, 3231–3234, doi:10.1029/96GL03041, 1996.
- Wallace, L., Hinkle, K. H., Livingston, W. C., and Davis, S. P.: An optical and near-infrared (2958–9250 Å) solar flux atlas, *Astrophys. J. Suppl. S.*, 195, doi:10.1088/0067-0049/195/1/6, 2011.
- Washenfelder, R. A., Toon, G. C., Blavier, J.-F., Yang, Z., Allen, N. T., Wennberg, P. O., Vay, S. A., Matross, D. M., and Daube, B. C.: Carbon dioxide column abundances at the Wisconsin Tall Tower site, *J. Geophys. Res.*, 111, D22305, doi:10.1029/2006JD007154, 2006.
- Watanabe, T., Nakamura, M., and Ogawa, T.: Rocket measurements of O₂ atmospheric and OH Meinel bands in the airglow, *J. Geophys. Res.*, 86, 5768–5774, doi:10.1029/JA086iA07p05768, 1981.
- Xu, J., Smith, A. K., Jiang, G., Gao, H., Wei, Y., Mlynczak, M. G., and Russell III, J. M.: Strong longitudinal variations in the OH nightglow, *J. Geophys. Res.*, 37, L21801, doi:10.1029/2010GL043972, 2010.
- Xu, J., Gao, H., Smith, A. K., and Zhu, Y.: Using TIMED/SABER nightglow observations to investigate hydroxyl emission mechanisms in the mesopause region, *J. Geophys. Res.*, 117, D02301, doi:10.1029/2011JD016342, 2012.
- Yee, J.-H., Crowley, G., Roble, R. G., Skinner, W. R., Burrage, M. D., and Hays, P. B.: Global simulations and observations of O(¹S), O₂(¹Σ) and OH mesospheric nightglow emissions, *J. Geophys. Res.*, 102, 19949–19968, doi:10.1029/96JA01833, 1997.