



On the relationship between open cellular convective cloud patterns and the spatial distribution of precipitation

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Abstract. Precipitation is thought to be a necessary but insufficient condition for the transformation of stratocumulus-topped closed cellular convection to open cellular cumuli-form convection. Here we test the hypothesis that the spatial distribution of precipitation is a key element of the closed-to-open cell transition. A series of idealized 3-D simulations are conducted to evaluate the dependency of the transformation on the areal coverage of rain, and to explore the role of interactions between multiple rainy areas in the formation of the open cells. When rain is restricted to a small area, even substantial rain (order few mm day^{-1}) does not result in a transition. With increasing areal coverage of the rain, the transition becomes possible provided that the rain rate is sufficiently large. When multiple small rain regions interact with each other, the transition occurs and spreads over a wider area, provided that the distance between the rain regions is short. When the distance between the rain areas is large, the transition eventually occurs, albeit slowly. For much longer distances between rain regions the system is anticipated to remain in a closed-cell state. These results suggest a connection to the recently hypothesized remote control of open-cell formation. Finally it is shown that this transition occurs along a consistent path in the phase space of the mean vs. coefficient of variation of the liquid water path, droplet number and optical depth. This could be used as a diagnostic tool for global analyses of the statistics of closed- and open-cell occurrence and transitions between them.

1 Introduction

Low clouds cover a wide area of the Earth's atmosphere and have been the topic of decades of study via observation and numerical modeling. They represent a particularly interesting system comprising micro- to meso-scale interactions between microphysics, radiation and turbulence. Low cloud systems tend to organize into either closed or open cellular convection with cell sizes > 10 km. The transition from the mostly cloudy closed cellular convective state to the mostly clear open cellular state is interesting from a purely system dynamics point of view (e.g., Koren and Feingold, 2011; Feingold and Koren, 2013) as well as from the perspective of radiative forcing of the climate (e.g., Stevens et al., 2005).

The breakup of the stratocumulus-capped boundary layer into shallow cumulus clouds is a broad subject. At least three breakup mechanisms have been intensively studied: advection of a stratocumulus deck over a warmer ocean (e.g., Wyant et al., 1997; Pincus et al., 1997; Sandu and Stevens, 2011), cloud-top entrainment instability (e.g., Randall, 1980; Deardorff, 1980b; van der Dussen et al., 2014) and aerosol–cloud–precipitation interactions (e.g., Savijovic and Stevens, 2008; Wang and Feingold, 2009a; Feingold et al., 2010). The focus of this study is on stratocumulus cloud breakup through aerosol–cloud–precipitation interactions that promote transformation of closed- to open-cellular convection.

There is ample observational evidence that precipitation plays a key role in the transformation of the closed- to open-cellular state in regions that prefer the closed state (Stevens et al., 2005; vanZanten and Stevens, 2005; Comstock et al., 2005; Petters et al., 2006; Sharon et al., 2006; Wood et al., 2008, 2011a; Terai et al., 2014). For instance, Wood et al. (2011a) documented that for closed cells most of the drizzle evaporates below cloud base, but in open-cell regions a significant amount of precipitation reaches the surface. For the cases examined, the measured cloud-base rain rate was similar in magnitude for both closed and open cells. Surface precipitation is a mechanism for removing drops (and therefore aerosol particles) from the atmosphere, and significantly reduced cloud droplet and aerosol number concentrations are commonly measured within open cells.

Large-eddy simulation (LES) has successfully shown that in regions of high aerosol (or droplet) concentrations, closed cells are preferred while in low-concentration environments open cells are preferred, all else equal (Xue et al., 2008; Savic-Jovicic and Stevens, 2008; Wang and Feingold, 2009a). Follow-on modeling studies with LES, cloud system resolving model (CSRМ) and simple heuristic model have explored the relationship between the formation and maintenance of open cells and the control of aerosol/droplet number concentration on precipitation, entrainment and dynamical responses associated with convergence of precipitation-generated outflows (Wang and Feingold, 2009b; Wang et al., 2010; Feingold et al., 2010; Koren and Feingold, 2011; Kazil et al., 2011; Berner et al., 2011, 2013; Mechem et al., 2012; Ovchinnikov et al., 2013). Convergence of surface outflows originating from the downdrafts produced by precipitation in the walls of open cells results in updrafts that generate convection and eventually the intersection zones or walls of new open cells. This cycle of precipitation-generated interaction between outflows has been hypothesized as the main mechanism for open-cell transition as well as long-lived open cells (Wang and Feingold, 2009a; Feingold et al., 2010). A homogeneously distributed low aerosol/droplet number concentration generates widespread precipitation and has been the most frequently used method to initiate the closed- to open-cell transition.

Anecdotal evidence from research scientists and pilots suggests that strong precipitation may exist in closed cellular convection. Figure 1 shows radar reflectivity measured with the ship-mounted W- and C-band radars on 23 November 2008 between 07:00 and 09:00 UTC during the Variability of the American Monsoon Systems (VAMOS) Ocean Cloud Atmosphere Land Study Regional Experiment (VOCALS-REx; Wood et al., 2011b). The figure also shows the 4 km resolution infrared images observed from the Geostationary Operational Environmental Satellite (GOES). A threshold value of -15 dBZ is commonly used to indicate the existence of strong drizzle (-15 dBZ corresponds to roughly 0.1 mm d^{-1} rain rate; Comstock et al., 2004). Both radars detect persistent drizzle over the 2 h period. The W-band radar reflectiv-

ity implies that precipitation reaches the surface, with some rain almost certainly evaporating below cloud base; the C-band radar detects large rainy regions. However, there is no evidence of open cellular structure – at least at the 4 km spatial resolution, and over the range of the C-band radar (~ 100 km). Precipitation thus appears to be necessary, but not sufficient for open-cell transition (e.g., Feingold et al., 2010; Wood et al., 2011a). Feingold et al. (2010) hypothesized that the open-cell transition requires another element – sufficiently widespread precipitation. In the figure discussed above, the rainy areas seem to be widespread. Speculative reasons are that (1) the rain rate may not be strong enough; and (2) the spatial distribution of the rainy patches is not conducive to interactions between outflows that are a characteristic of open-cellular convection.

A few studies have considered aerosol gradients within horizontal domain lengths on the order of 25–100 km (Wang and Feingold, 2009b; Wang et al., 2010; Berner et al., 2013). The large characteristic length scale of organized convection requires domains much larger than those applied in typical LES. Wang and Feingold (2009b) showed that with a simulation initialized with an aerosol gradient, the outflow associated with the precipitation in open cells transported moisture toward non-precipitating closed cells. The resultant buildup of liquid water path (LWP) eventually initiated precipitation in the closed cells, with subsequent formation of open cells. Wang et al. (2010) investigated the initiation of open-cell formation with regions of lower aerosol as well as perturbations of moisture and temperature, and found that the latter initiated drizzle and open-cell formation in neighboring unperturbed closed cells instead of in the perturbed area. They referred to this as a remote control of open-cell formation.

In this study, we focus more closely on the importance of aerosol or precipitation gradients and the associated spatial distribution of rain on the selection of the cellular state of the system. A series of idealized numerical experiments is performed to evaluate the influence of these parameters on the structural transition from closed to open cells. Structure is defined by particular distributions of a few key parameters. A series of simulations is conducted to address the influence of both the formation and areal coverage of precipitation when the region of precipitation is concentrated at one location. Further numerical simulations are performed to understand the influence of multiple precipitation regions separated by some distance. A metric that quantifies the degree of open-cell transition is developed to aid in objective analysis.

The next section describes the numerical model and simulations. The metric for the open-cell transition is described in Sect. 3. Results are presented in Sect. 4. Additional discussions and conclusions are given in Sects. 5 and 6, respectively.

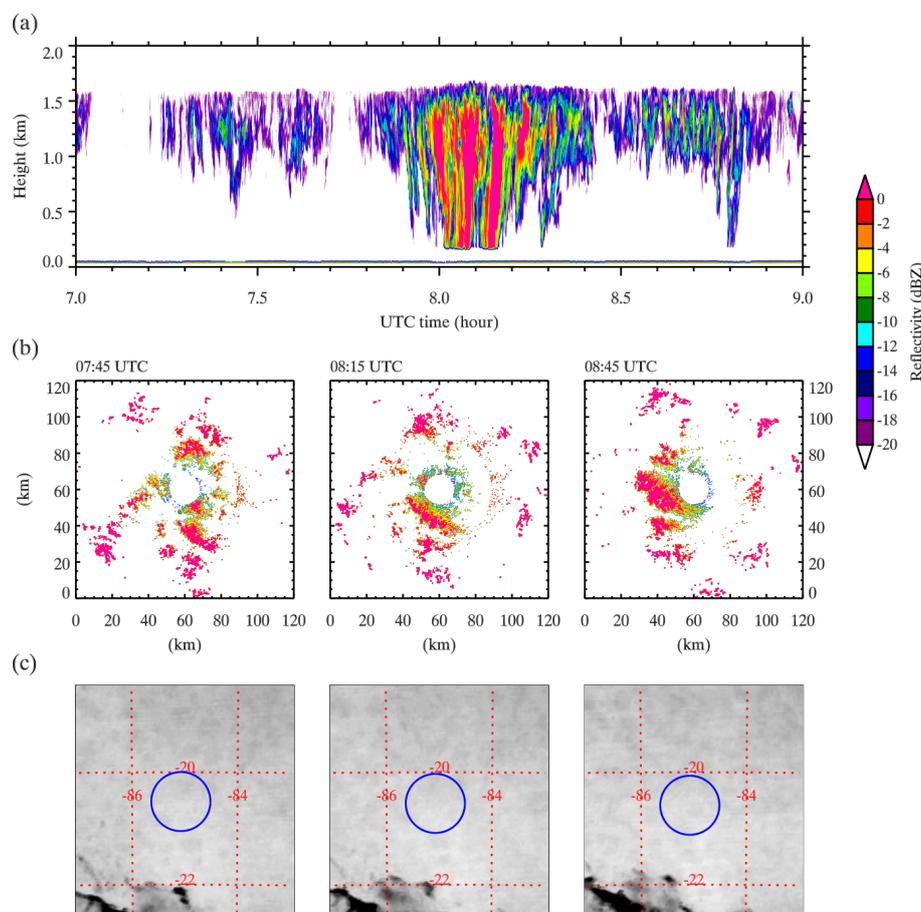


Figure 1. Ship- and satellite-based observations from VOCALS-REx between 07:00 and 09:00 UTC on 23 November 2008: (a) W-band radar reflectivity, (b) C-band radar reflectivity and (c) corresponding GOES infrared image (4 km resolution). The blue circle on each GOES image is the approximate range of the C-band radar.

2 Numerical model and simulations

2.1 Model

The System for Atmospheric Modeling (SAM; Khairoutdinov and Randall, 2003) is employed. SAM is formulated as an anelastic system, spatially discretized on the Arakawa C grid with a vertical height coordinate. The equations of motion are advanced with a third-order Adams–Bashforth scheme (Durran, 1991), and with a second-order center-difference scheme for advection. The subgrid-scale model is the 1.5-order turbulence kinetic energy scheme of Deardorff (1980). The bulk 2-moment microphysics scheme of Feingold et al. (1998) is used with a bulk sedimentation scheme following Morrison (2012), which is equivalent and more efficient than the earlier bin sedimentation scheme. Activation of aerosol is based on the ambient, calculated supersaturation, and the assumption of a lognormal aerosol size distribution with geometric-mean diameter and geometric standard deviation ($0.2\ \mu\text{m}$ and 1.5, respectively). Prognostic scalar variables are liquid water static energy, mixing ratios of water

vapor, cloud water and rain water, supersaturation, number concentrations of aerosol, cloud droplets and rain drops, and subgrid-scale turbulence kinetic energy. All scalars are transported with the monotonic fifth-order scheme of Yamaguchi et al. (2011).

The model configuration is based on the Global Energy and Water Exchanges Project (GEWEX) Cloud System Study (GCSS, currently known as the Global Atmospheric System Studies – GASS) LES intercomparison case of the Second Dynamics and Chemistry of Marine Stratocumulus (DYCOMS-II RF02; Ackerman et al., 2009) with several modifications. This is a nocturnal drizzling stratocumulus case. There are two frequently used observations for studies of open-cell transition with numerical simulations: DYCOMS-II RF02 (e.g., Savic-Jovicic and Stevens, 2008), and cases based on VOCALS-REx (e.g., Wang et al., 2010). The major difference between DYCOMS-II and VOCALS-REx is the depth of their boundary layers, approximately 800 m for DYCOMS-II RF02 and 1300 m for VOCALS-REx (Fig. 1). This difference manifests itself in both cloud depth and degree of vertical mixing, with VOCALS-REx boundary

layers supporting deeper clouds, more rain, and a stronger tendency to decouple. We note that Fig. 1 is associated with VOCALS-REx and should thus be viewed as an illustration of the central idea explored here, rather than a case for direct comparison with the current DYCOMS-II RF02 simulations. While it is interesting to explore the influence of boundary layer heights on the closed-to-open cell transition, we defer it for future research.

A horizontally square domain covers 51.2 km width with 200 m resolution and the depth of the domain is 1.6 km with 10 m vertical resolution. The lateral boundary conditions are doubly periodic. The time step is 1 s. The initial horizontal wind is set to 0 m s^{-1} in order to produce horizontally stationary cell patterns (Wang and Feingold, 2009a). For consistency, the Coriolis effect is not applied. No modifications are made to the standard GCSS DYCOMS-II RF02 forcings: large-scale subsidence is computed with the specified large-scale horizontal wind divergence of $3.75 \times 10^{-6} \text{ s}^{-1}$; surface sensible and latent heat fluxes are constant at 16 W m^{-2} and 93 W m^{-2} , respectively; friction velocity for surface momentum flux is constant and 0.25 m s^{-1} ; longwave radiative flux is computed with a simple longwave radiation scheme (Ackerman et al., 2009).

Simulation of mesoscale organization requires large domains, and in the current case, a large number of simulations. (The importance of large domains will become apparent.) The associated computational expense requires a CSR approach. The relatively coarse resolution, especially in the horizontal direction, and high grid aspect ratio (horizontal to vertical grid ratio is 20 : 1) in our configuration do not fit in the realm of LES. Earlier work by Wang and Feingold (2009a) demonstrated that mesoscale circulations simulated with CSR are qualitatively the same as their higher-resolution counterparts and therefore the current choice of grid size is not expected to change the qualitative nature of results to be presented herein.

SAM's standard horizontal mean statistics data (time series and profiles) are recorded every minute. Posteriori 2-D horizontal fields of LWP, vertically integrated droplet number concentration ($N_d \text{ mm}^{-2}$) and optical depth (τ) for total liquid water (i.e., the sum of cloud and rain water) as well as rain water path (RWP) and surface rain rate are output every minute. Albedo is calculated from τ with the two-stream approximation of Bohren (1987).

2.2 Numerical simulations

Three series of simulations are designed for this study. The first series, S1 uses a variety of a homogeneously distributed initial aerosol number concentrations (n_a), and are run for 12 h. Rain water is not allowed to precipitate over the first hour to allow turbulence to develop. S1 consists of 11 cases with different n_a ranging between 50 mg^{-1} and 250 mg^{-1} in increments of 20 mg^{-1} . Note that $1 \text{ mg}^{-1} = 1 \text{ cm}^{-3}$ at an air density of 1 kg m^{-3} . Each case is named with the initial n_a :

Table 1. List of cases for S2 simulations with a variety of single patches placed at the domain center. Each case branches off from N130 of S1 at 3 h, and is run for 9 h. Case name represents specified fractional area and n_t of the patch. F015-N030 uses a patch with area fraction of 0.15 and $n_t = 30 \text{ mg}^{-1}$.

Case	Fractional area	Radius (km)	n_t (mg^{-1})
F005-N0...	0.05	6.46	10, 30, 50, 70, 90
F010-N0...	0.10	9.13	10, 30, 50, 70, 90
F015-N0...	0.15	11.19	10, 30, 50, 70, 90
F020-N0...	0.20	12.92	10, 30, 50, 70, 90

e.g., N070 denotes the case with $n_a = 70 \text{ mg}^{-1}$. This series of simulations is the basis for development of the metric for closed- to open-cell transition.

The second series, S2 consists of 20 cases, which all branch off (i.e., restart) from N130 of S1 at 3 h and are run for a further 9 h (i.e., for a total of 12 h). As shown in the next section, N130 is the case with the smallest initial n_a among S1 that maintains closed cells for 12 h. A cylinder is placed vertically in the center of the horizontal domain at 3 h. This “patch” covers a specified horizontal fractional area. Inside the patch, in each grid box, the total number concentration, n_t , which is the sum of n_a and cloud water droplet number concentration (n_c), is changed to the specified value to initiate precipitation. Partitioning the specified n_t to n_a and n_c is done based on the number ratio of these two variables, i.e., $n_a = [n_a/(n_a + n_c)]n_t$ at 3 h. The list of cases is presented in Table 1. Each case is named according to the horizontal fractional area and n_t , e.g., F015-N030 denotes a patch fractional area of 0.15 and $n_t = 30 \text{ mg}^{-1}$.

The third series, S3 also branches off from N130 of S1 at 3 h, but uses a horizontal domain twice as large (102.4 km width). Restart data at 3 h of N130 are also prepared by running N130 on the larger domain. It is also run for a further 9 h. The simulations consists of four cases: one case uses a single patch, and the other three cases place one patch at the domain center and six patches at the vertices of the hexagon whose center is located at the domain center. The inter-patch distance, i.e., the shortest distance between two patches differs among these three cases. The total horizontal area fraction of patches is constant among the cases and is set to 0.04. The total number concentration of patches is also constant amongst the cases and is set to 30 mg^{-1} . The list of cases is shown in Table 2.

3 Metric for structural transition from closed to open cells

Albedo fields for N090, N110 and N130 for series S1 at 4, 8 and 12 h are presented in Fig. 2. It is clear that N090 transfers into an open-cell field, while N130 remains in the closed-cell state.

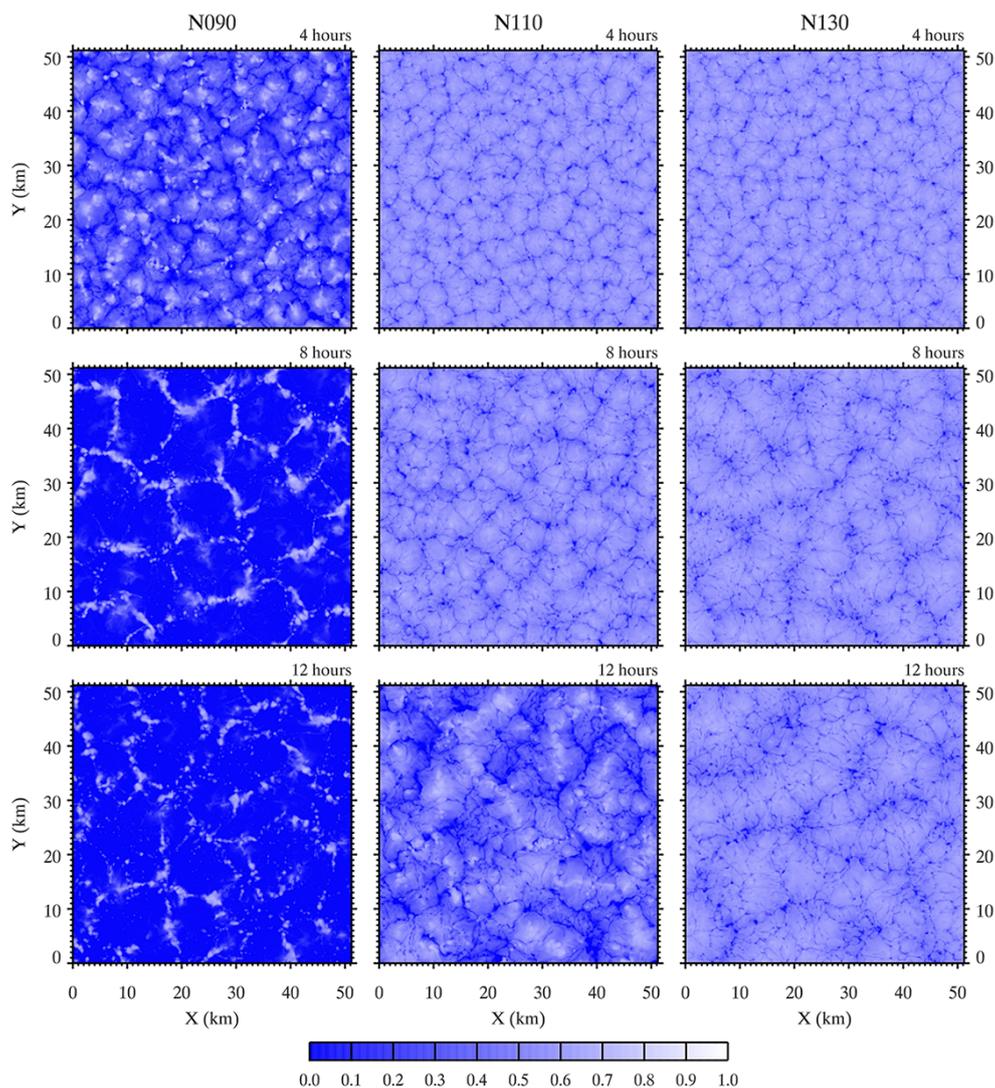


Figure 2. Time evolution of the computed albedo for N090, N110 and N130 of S1.

Domain mean statistics for all S1 cases are presented in Fig. 3. The closed-cell cases ($n_a \geq 130 \text{ mg}^{-1}$) maintain $\text{LWP} \gtrsim 100 \text{ g m}^{-2}$ and negligibly small RWP. For the last hour, well mixed profiles are present for these cases. Although very subtle, the larger initial n_a result in deeper boundary layers, higher cloud bases, and smaller liquid water content. This is related to entrainment resulting in faster evaporation for the higher n_c (and n_a) (Wang et al., 2003; Xue and Feingold, 2006; Bretherton et al., 2007; Xue et al., 2008; Hill et al., 2009).

The open-cell cases ($n_a \leq 90 \text{ mg}^{-1}$) show the runaway precipitation feedback discussed by Feingold and Kreidenweis (2002); LWP decreases rapidly while RWP increases rapidly to values of the order of 10 g m^{-2} , and precipitation rates exceeds 0.5 mm day^{-1} at the surface. For the last 2 h, the fields reach a steady state. The well-mixed profiles disappear for these cases and they have substantially lower plan-

etary boundary layer (PBL) height due to precipitation suppressing entrainment (Stevens et al., 1998). Profiles are in agreement with past studies (Savic-Jovicic and Stevens, 2008; Wang and Feingold, 2009a).

N110 is interesting in that it does not quite fit into either category. At 12 h, LWP is similar to the closed-cell cases, but in contrast RWP is close to the open-cell cases, and there is negligible domain mean surface precipitation. Profiles for the thermodynamic variables are very similar to the closed-cell cases but slightly less well-mixed. Turbulence profiles differ from the closed-cell cases, which suggests that N110 completes the 12 h simulation in the middle of the transition. Looking back at the albedo fields of N110 at 12 h (Fig. 2), bright cellular lines of clouds exist, as observed in N090 at 4 h (notwithstanding the different cell sizes). These bright lines do not exist in N130 for the duration of the simulation.

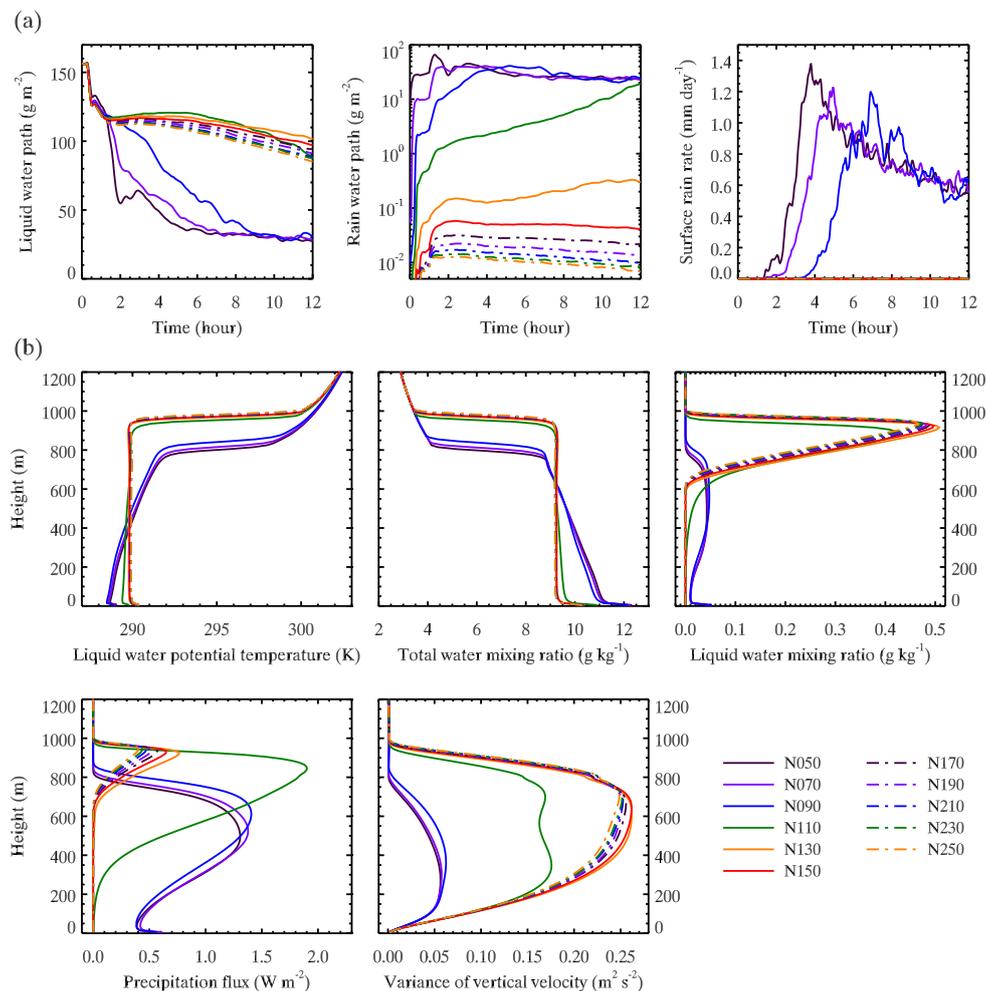


Figure 3. Domain mean statistics for S1. The vertical profiles are averaged over the last hour.

Wood and Hartmann (2006) developed a neural network scene classification method based on the distribution and power spectrum of LWP. This elaborate method classified a satellite image into one of four categories: no clear identifiable cellularity, closed cells, open cells and clouds containing cells without organization. We have developed an alternative scheme that yields information on the degree of transition between closed- and open-cell structures. In order to develop a metric for open-cell transition, the time evolution of the distributions of LWP, N_d (vertically integrated droplet number concentration) and τ are considered. To build a simple method, and possibly apply the metric to observational data in the future, we focus on cloud fields that can be measured or derived remotely, and avoid turbulence statistics. Figure 4a shows the LWP distribution for N090 between 2.5 and 4.5 h. It is roughly approximated by a unimodal distribution at 2.5 h; with time, the distribution develops bimodality but by 4.5 h it transitions back to a unimodal distribution with a mode at the smallest bin. These two unimodal structures were also shown in observations (Wood and Hart-

mann, 2006; Wood et al., 2011a). For a binned distribution, the mode M is defined as the value of the bin with the highest frequency (or count). Thus, the mode for LWP is quantified with $M = x_i$, where x_i is the value of the i th LWP bin. The time series of the mode based on LWP is shown in Fig. 4b for selected cases. For the open-cell cases (N070 and N090), there is a change from the mode associated with the closed-cellular state ($M \gtrsim 100 \text{ g m}^{-2}$) to the mode associated with the open-cellular state (the first bin, x_1 , i.e., $M = 4 \text{ g m}^{-2}$). The change of mode never takes place for the closed-cell case (N130). The mode change is in progress for N110 when the run ends. Thus, the differences between these two unimodal distributions reflect structural differences between closed and open cellular states.

Tracking the mode gives an indication of the timing of the transition to open cells. It cannot, however, distinguish between closed cells, open cells, or the point of transition.

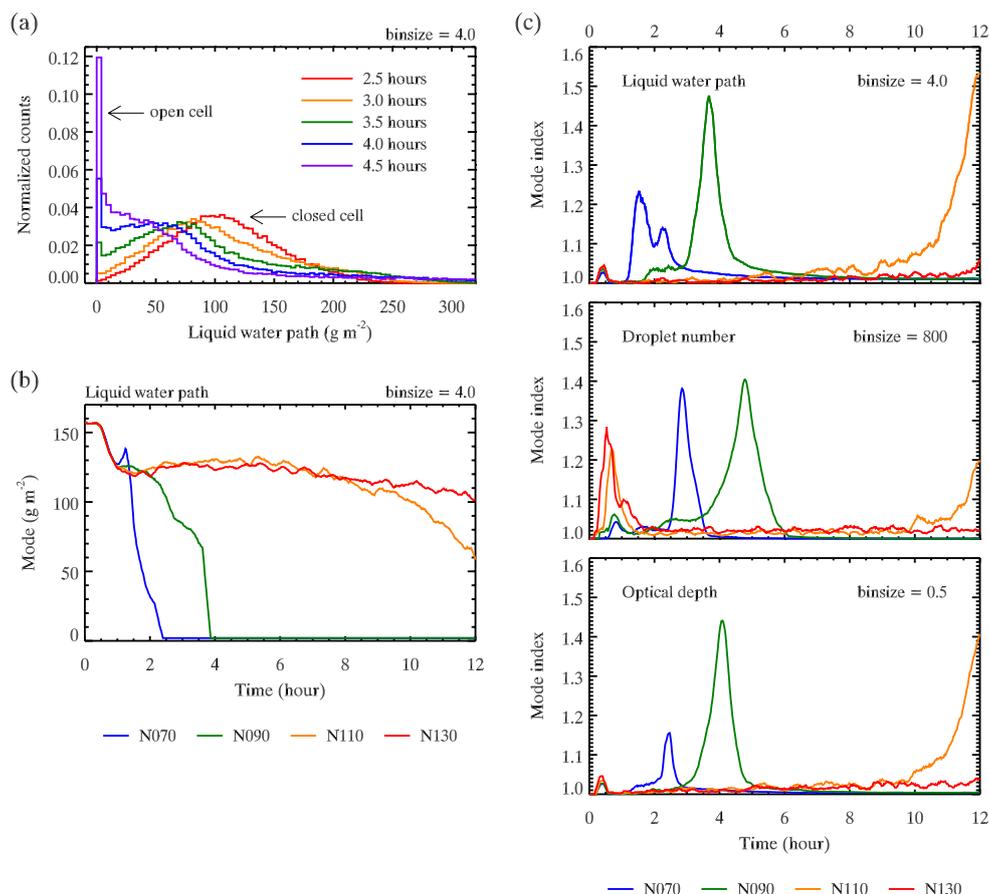


Figure 4. (a) Time evolution of the distribution of LWP for N090, (b) time series of the mode for the distribution of the LWP for selected S1 cases and (c) time series of the mode index based on LWP, N_d and τ . A running mean with a 15 min window is applied to the mode and mode index.

Therefore we introduce the mode index, defined as

$$I \equiv \frac{1}{2\max(f)} \sum_{i=1}^n |f_i - f_{i-1}|, \quad (1)$$

where f is the frequency and $f_0 = f_n = 0$ (boundary condition). By definition, $I = 1$ for a pure unimodal distribution, and $I = 2$ for a pure bimodal distribution.

Both mode and mode index are sensitive to bin size. Consider a histogram with a bin size that contains either 0 or 1 counts in each bin. For this case, the number of modes, N_M , equals the number of elements, and $I \geq 2$. (A special case is $I = 1$ when $f = 1$ for all bins except x_0 and x_n .) At the other extreme, $N_M = 1$ and $I = 1$ for a large enough bin size that contains all elements in a single bin. Considering these bounding cases, bin size is determined as follows: first the smallest bin size, which meets the condition that $I < 1.1$ at any time after spinup for N130 (a closed-cell case), is identified. With this bin size, the $\max(N_M)$ is identified for each case of S1 from the corresponding time series of N_M after spinup. The bin size is incrementally increased until $\max(N_M) \leq 2$ is satisfied for all cases. The bin sizes for

LWP, N_d and τ are 4 g m^{-2} , 800 mm^{-2} and 0.5, respectively. These are useful, ad hoc criteria and may need to be modified for different cases.

The time series of the mode indices for LWP, N_d and τ are presented in Fig. 4c for selected cases. After rain starts to reach the surface at 1 h, the mode indices for the open-cell cases (N070 and N090) pass a peak, which exceeds 1.1 for all variables. After 10 h, the mode indices for N110 increase to a value larger than 1.1. Since at 12 h the modes of N110 do not reach x_1 (Fig. 4b for LWP), the structure during transition is expected to be somewhat similar to the closed cells, as confirmed by Fig. 2. For the open-cell cases, there is a time lag for the appearance of the peak depending on the variable under consideration, i.e., LWP, N_d or τ : LWP comes first, followed by τ , and finally N_d . Thus the amount of condensate is a stronger indicator of the beginning of a transition than is a microphysical property such as N_d . The τ is a blend of condensate and microphysical properties and therefore appears second. This ordering may be partially related to the selected bin sizes and we do not place much emphasis on this result.

The selection of the three different parameters is primarily to assess robustness of the transition metric.

With the mode and mode index, four states can be classified for each variable: closed cells, closed cells under transition, open cells under transition, and open cells. The four states are classified as follows:

- closed cells when $M \neq x_1$ and $I < 1.1$,
 closed cells under transition when $M \neq x_1$ and $I \geq 1.1$,
 open cells under transition when $M = x_1$ and $I \geq 1.1$, and
 open cells when $M = x_1$ and $I < 1.1$. (2)

The structural classification is performed for LWP, N_d and τ . For instance, for N110 at 12 h, the metric gives “closed cells under transition” for each variable (Fig. 4c). Although for four classified states and three variables, mathematically the total number of possible combination of states is 64, in reality there are far fewer possibilities since many combinations of states do not exist, e.g., closed-cell classification based on LWP and open-cell classification based on the other two variables.

4 Results

4.1 Single patch simulations

The time evolution of albedo for selected S2 cases is shown in Fig. 5. Based on the metrics, all 20 cases are categorized into only five combinations of classified states from closed to open cells. Unlike the open-cell cases for S1, there is no clear cellular pattern for the open-cell case (F015-N030) at 12 h, however the metric suggests that the structure strongly resembles that of open cells based on the distribution of LWP, N_d and τ . For the closed-cell case (F010-N090), the patch is completely swallowed by surrounding cloudy air. For the open-cell (left) and transition cases (three central columns), the divergence (or outflow from the domain center) created by the precipitation propagates outward and a low albedo area develops. That outflow propagates back toward the domain center once it reaches the horizontal boundary due to periodicity; this effect is examined later in Sect. 5.

Figure 6a shows the evolution of RWP between 6 and 10 h for F015-N030. The propagation of the RWP is approximately symmetrical since the patch is introduced at the domain center. As seen at 9 h, the distortion of the symmetry happens once the propagation reaches the domain boundary. Due to this quasi-symmetry, we calculate the radially averaged RWP, with the radius measured from the domain center, and plot this field as a modified Hovmöller diagram. Figure 6b shows the Hovmöller diagram of radially averaged RWP for three S2 cases. It captures the dynamical response of the generation of new raining areas due to convergence of outflows produced by precipitation. For F015-N030 and F010-N050, the onset of precipitation first creates a RWP

Table 2. List of cases for S3 simulations for multiple patches. Each case branches off from N130 of S1 at 3 h, performed with the large domain (102.4 km domain width), and is run for 9 h. The total area fraction covered by the patches is 0.04 and $n_t = 30 \text{ mg}^{-1}$. The inter-patch distance is the shortest distance between the sides of two patches.

Case	Formation	Radius (km)	Inter-patch distance (km)
SP	single patch	11.55	n/a
H1	hexagon (6 at vertices, 1 at center)	4.37	4.37
H2	hexagon (6 at vertices, 1 at center)	4.37	8.73
H3	hexagon (6 at vertices, 1 at center)	4.37	13.10

convergence region at approximately 4 km radius (slightly less than the radius of the patch) and 4.5 h, then the second RWP concentrated region forms near the domain center (i.e., 0 km radius) at 5.5 h. The second region persists longer than the first convergence region. The third region appears at approximately 16 km radius and 7.5 h after the second region has disappeared. After the third convergence, the perturbation reaches the horizontal boundary and proceeds inward to the domain center. Note that the closed-cell case (F010-N090) evolves in a similar way, albeit with slow propagation speed and weak magnitude, and it also manages to create the third precipitating region towards the end of the simulation.

Results of the metrics for the open-cell transition applied to the all S2 cases at 12 h are presented in Fig. 7. For a given fractional area, a larger gradient in n_t between the environment and the patch leads to the formation of open cells, and for a given gradient, a larger fractional area leads to the formation of open cells. Although the number of cases is limited, the results suggest that the gradient of n_t and the areal coverage of rain are of similar importance for the open-cell transition. To capture the features discussed for S1 (Fig. 4b and c), time series of the mode and mode index are shown in Fig. 8a for selected S2 cases. It is noted that (i) the case with an earlier LWP mode change also arrives earlier back to a mode index of 1.1 and (ii) because of the limited patch sizes, complete transition requires a long period of time – the fastest transition amongst the S2 cases (i.e., F020-N010, not shown in the figure) takes approximately 7 h based on these metrics. Transition times were much shorter in the S1 simulations (Fig. 4).

Figure 8b shows the time series of surface precipitation rate for domain mean and conditional mean. For the domain mean, the closed-cell case seems not to precipitate; its conditional average (over rain rates $> 0.1 \text{ mm day}^{-1}$), however, exceeds 2 mm day^{-1} , which suggests that the locally measured surface rain rate should not be considered a priori an indicator of open cells.

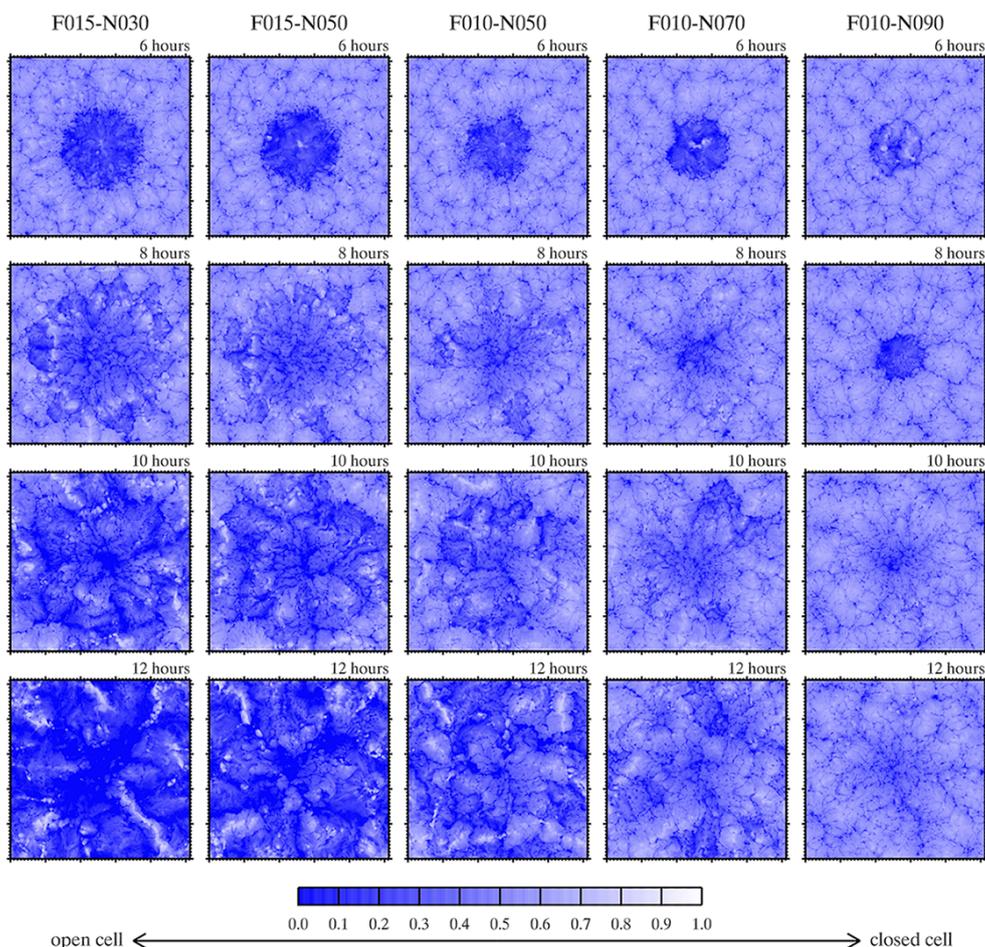


Figure 5. Time evolution of the computed albedo for selected S2 cases. Cases are placed so that the case with open cells is on the left side, the closed cell case is on the right side, and transition cases are in between. Categorization is based on the metric discussed in the text.

4.2 Multiple patch simulations

To explore the effect of multiple rain regions on the open-cell transition, the S3 simulations (Sect. 2, Table 2) are now discussed. The difference in the model configurations compared to the other simulations is that the horizontal domain size is extended to $102.4 \times 102.4 \text{ km}^2$. This large domain is required to fit distributed patches in the domain, as well as to avoid a bias caused by the periodic boundary conditions. As noted earlier, the outflow sometimes reaches the domain edge before the simulation ends and propagates inward to the domain, which essentially “contaminates” the flow. Note that n_t of patches is 30 mg^{-1} for all cases.

Figure 9 shows the evolution of the albedo for all S3 cases. The metrics applied for a $51.2 \times 51.2 \text{ km}^2$ subdomain and the full domain ($102.4 \times 102.4 \text{ km}^2$) at 12 h are also shown in schematic form. SP reaches an open-cell state based on the metric applied to the subdomain, which is anticipated since the areal fraction of 0.04 for the full domain is equivalent to that of 0.16 for the domain with 51.2 km width. Simulation F015-N030 of S2 (single patch covering 0.15 area fraction

for the $51.2 \times 51.2 \text{ km}^2$ domain) also achieves an open-cell state (Fig. 7). The transition occurs via interactions between precipitation-generated outflows associated with the patches in H1, H2 and H3; these patches are not able to transform to open cells without these interactions. This is anticipated since each patch only covers an areal fraction of approximately 0.006, which translates to 0.023 for the domain with 51.2 km width, and as shown above, F005-N030 of S2 (single patch covering 0.05 area fraction for the $51.2 \times 51.2 \text{ km}^2$ domain) remains in the closed-cell state (Fig. 7). The transition for H3 is the slowest amongst the cases due to the largest distance between patches. This is because the outflows have to travel the longest distance to converge with other outflows. The open areas of H3 are completely filled by surrounding cloudy air, but by 12 h open areas have appeared. For much larger inter-patch distances (all else equal) the outflows are not able to create strong enough updrafts at the convergence zones, the transition comes to a halt, and the system returns to the closed-cell state.

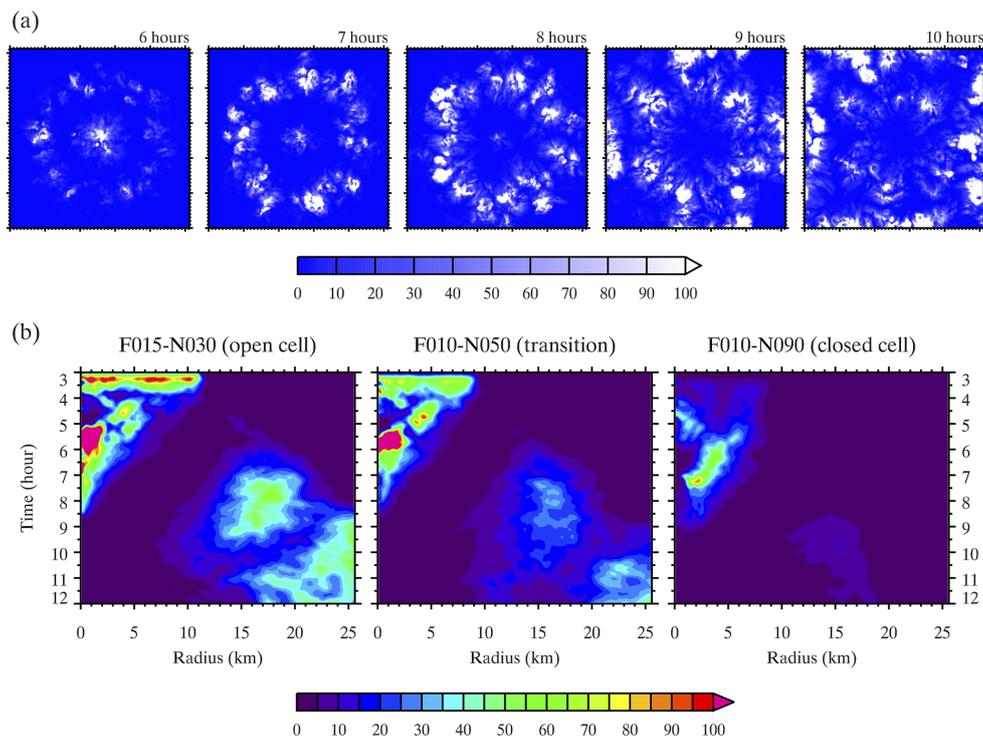


Figure 6. (a) Time evolution of RWP (g m^{-2}) between 6 and 10 h for F015-N030. (b) Hovmöller diagram of the radially averaged RWP for three S2 cases. A radius of 0 km represents the domain center.

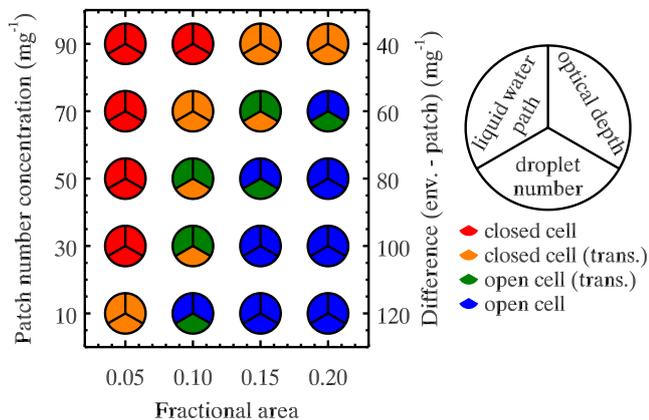


Figure 7. Metric for the open-cell transition for all S2 cases at 12 h. In the diagram, each circle represents one case, and each part of the circle indicates either LWP, N_d or τ , and is colored according to the classified state.

How widely the rainy areas spread for a given duration is determined by the distance between the outer edge of the precipitating area (the “rim”) and the domain boundary (hereafter, rim–boundary distance) as well as the inter-patch distance. The effect appears in the differences among the metrics in Fig. 9. For the rim–boundary distance, $\text{SP} > \text{H1} > \text{H2} > \text{H3}$, and for the inter-patch distance, $\text{H1} < \text{H2} < \text{H3}$. SP initiates precipitation within the patch,

which produces relatively strong dynamical feedback within the subdomain, but the case requires time to propagate rain regions to the outer reaches of the domain because of the large rim–boundary distance. On the other hand, the rim–boundary and inter-patch distances for H1 are optimal for spreading rain regions outward through dynamical response. Relative to SP, it does, however, require more time to fully transform the H1 subdomain into an open-cell state.

The dynamical feedback is shown in Fig. 10a in the form of a Hovmöller diagram of RWP. The wavelike pattern for SP is different from that for H1. SP creates a relatively short-lived convergence at the domain center at approximately 6.5 h, and later a strong domain center convergence at approximately 11.5 h after the inflow from the rain area located at approximately 15 km radius arrives at the domain center. H1 shows convergence at approximately half of the distance between patches (6.5 km radius) at 5 h following the initial rain event; then the large RWP associated with strong convergence at the domain center forms at approximately 6 h and spreads outward. H2 and H3 exhibit a similar evolution, but with slower phase speed since the inter-patch distance is longer. Figure 10b shows that strong, coherent surface rain occurs when the RWP convergence exists, and that H1 produces surface precipitation of $> 4.5 \text{ mm d}^{-1}$ while SP and all but two cases of S2 (Fig. 8b) do not produce surface rain $> 3.5 \text{ mm d}^{-1}$. H2 and H3 produce surface precipitation $> 4.5 \text{ mm d}^{-1}$ (not shown). These strong precipitation events

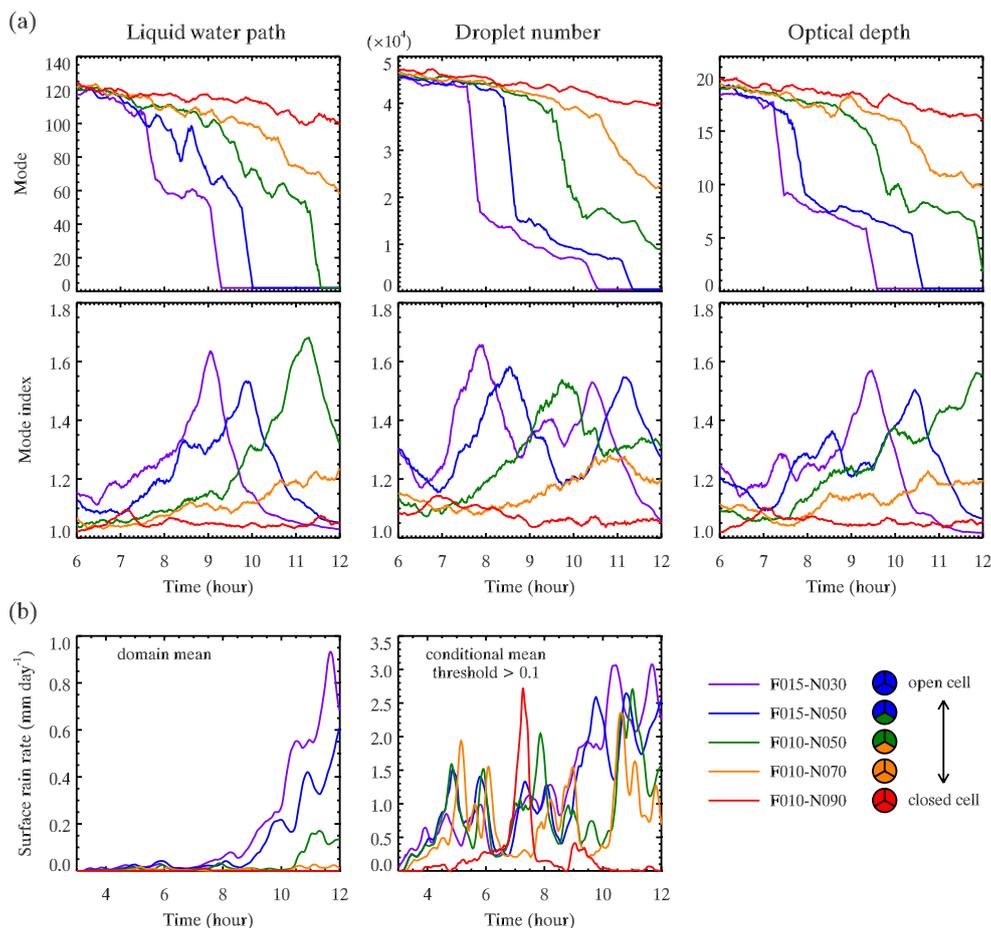


Figure 8. Time series for S2 simulations: (a) mode and mode index for LWP (g m^{-2}), N_d (mm^{-2}) and τ from 6 to 12 h and (b) domain mean and conditional mean (grid with $> 0.1 \text{ mm day}^{-1}$) of surface rain rate for the selected cases. A 15 min running mean is applied. For reference, $N_d = 1 \times 10^4 \text{ mm}^{-2}$ is equivalent to a local concentration of 50 cm^{-3} for a cloud depth of 200 m.

for the multipatch cases occur at the large RWP convergence zones at the domain center, e.g., at 7 h for H1. SP has the largest convergence at 11.5 h and the surface rain rate is at its maximum. The rain rate difference between these convergence zones can be shown to be related to the distance to the domain center from the previous convergence, which is the source of outflow traveling toward the domain center. Although the surface rain rates for the previous convergence for H1 and SP are similar to each other and approximately 2 mm d^{-1} , the previous convergence for H1 is located much closer to the domain center than that for SP, allowing the outflow to generate stronger updrafts.

5 Predictable transition pathway?

Berner et al. (2013) used multi-day, 2-D CSRM simulations to study the slow manifold path toward a shallow broken, or deeper well-mixed stratocumulus capped PBL with adjustment timescales of several days. In the spirit of synthesizing model output of system evolution, we consider a phase

diagram for the mean and coefficient of variation, c_v , (i.e., standard deviation divided by the mean) for LWP, N_d and τ . Relationships between the mean and variance of cloud fields are especially useful for boundary-layer parameterizations employed in climate models. We show here that in this phase-space the different model configurations exhibit a predictable evolutionary pathway when perturbed by low aerosol concentration patches. Figure 11 shows the phase diagram for the mean and c_v , for LWP, N_d and τ for all cases of S1 (Fig. 11a), S2 (Fig. 11b) and S3 (Fig. 11c) for the entire duration of the simulations (12 h for S1, 9 h for S2 and S3). The calculation of c_v and mean for S3 is performed on the $51.2 \times 51.2 \text{ km}^2$ subdomain.

For S1 (Fig. 11a), c_v for the closed-cell cases ($n_a > 130 \text{ mg}^{-1}$, red colored) remains small while the mean decreases in response to a deepening of the boundary layer, and entrainment reducing the cloud water (Fig. 3). For the open-cell cases ($n_a < 90 \text{ mg}^{-1}$, blue colored), the pathways converge to one curve, characterized by a decreasing mean and an increasing c_v . The transition case ($n_a = 110 \text{ mg}^{-1}$, green

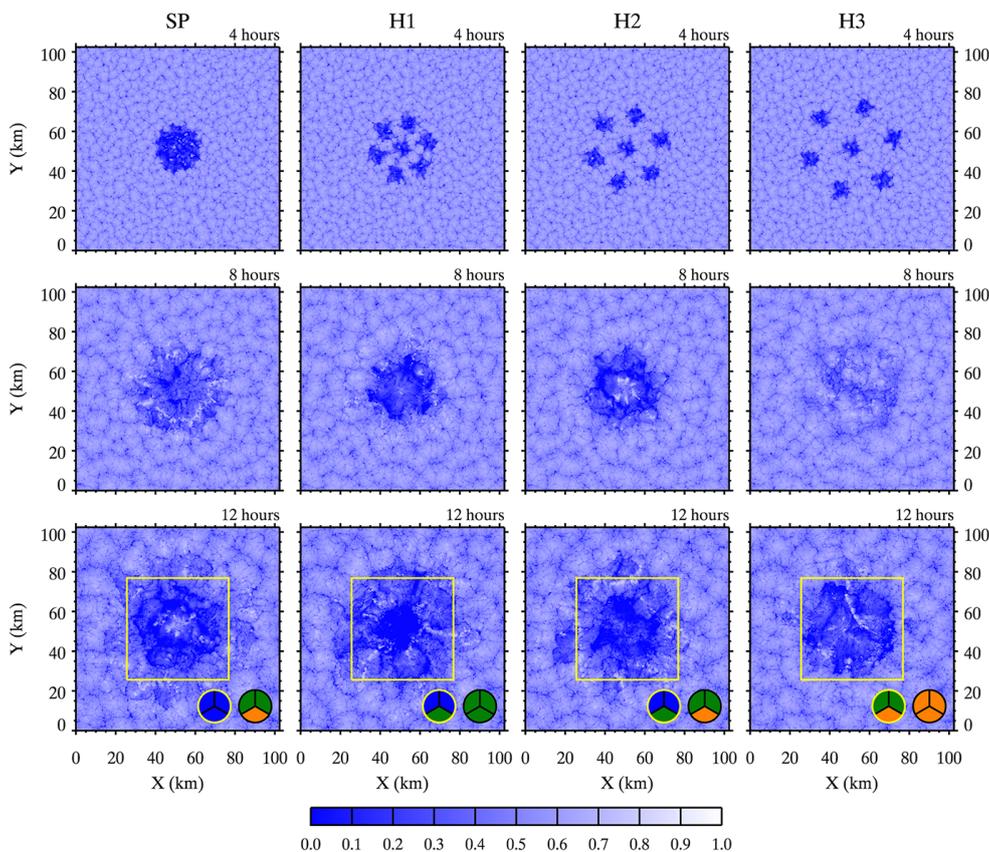


Figure 9. Time evolution of the computed albedo for S3 simulations. For the albedo at 12 h, the metric calculations for the open-cell transition are applied to two squares: 51.2 km width (yellow) and the entire domain (102.4 km width). The same colors are used for the square and the circumference of the schematic metric circles.

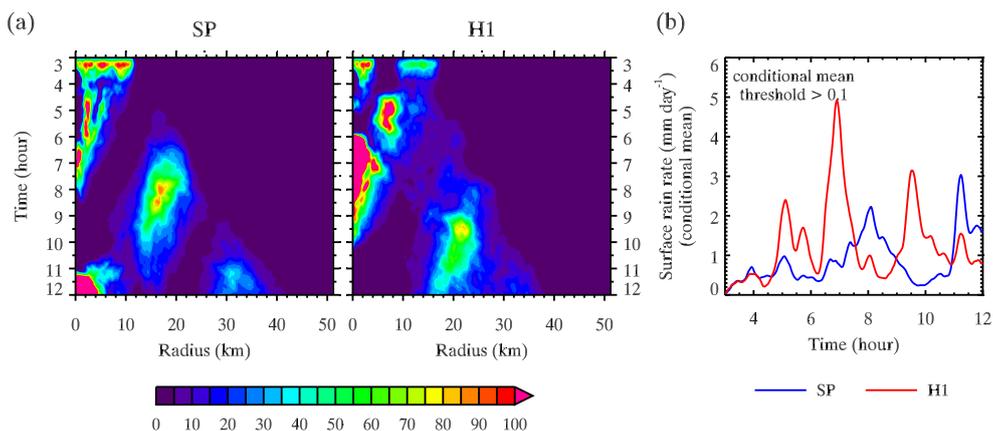


Figure 10. (a) Hovmöller diagram of the radially averaged RWP (g m^{-2}) for SP and H1. (b) Conditional mean surface rain rate for SP and H1 (grid points with rain rate $> 0.1 \text{ mm day}^{-1}$ averaged over the domain). A 15 min running mean is applied.

colored) also merges to the open-cell phase curve. It is noted that the rate of c_v increase for open-cell cases is larger for smaller initial n_a (not shown). It is also worth noting that the phase diagrams for the skewness vs. the mean also form a single curve (not shown).

The phase curves for S2 (Fig. 11b) exhibit similar behavior. They form a characteristic curve, which is independent of the area fraction or patch-environment gradient in n_t . More interestingly, the phase curve for LWP for S2 overlaps that of S1. The behavior for N_d and τ is somewhat different. Re-

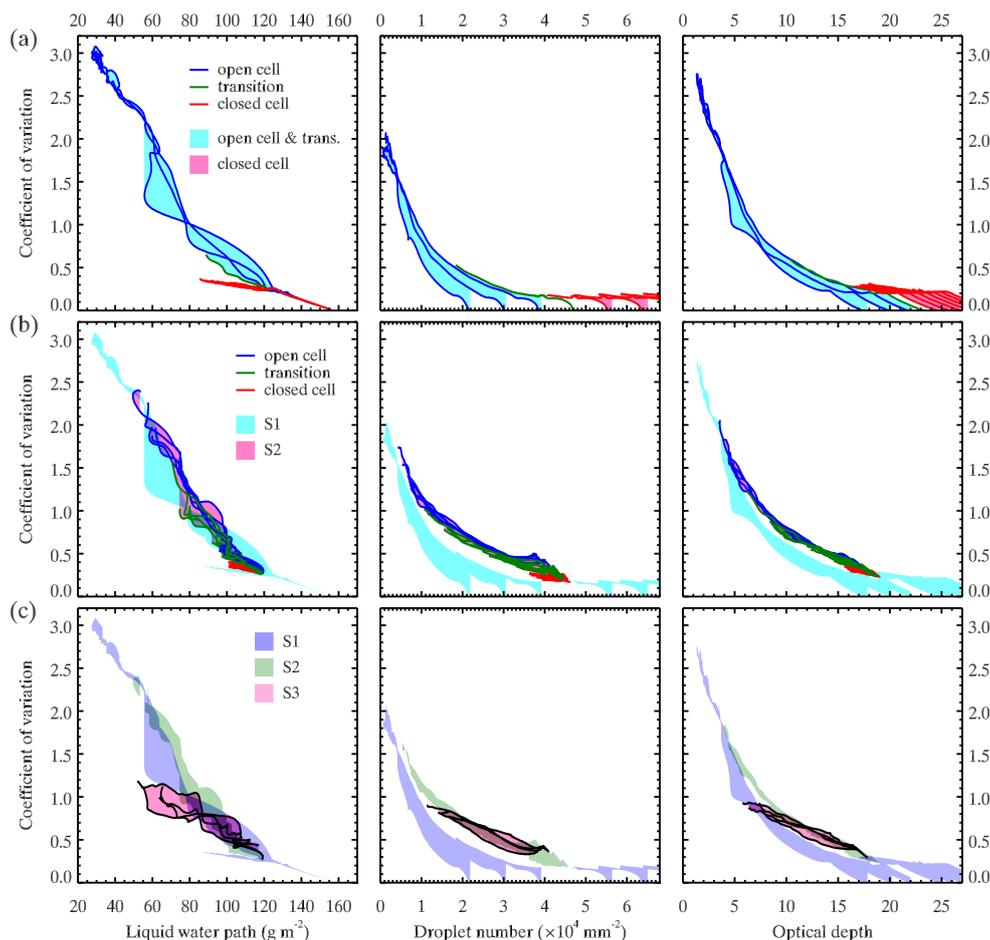


Figure 11. (a) Phase diagram of c_v of LWP, N_d , and τ for all S1 cases. The two shadings cover the range of c_v for open-cell and transition cases and for closed-cell cases. (b) Same as (a) but for all S2 cases. The two shadings are for S1 and S2. (c) Same as (a) but for all S3 cases. The S3 analysis is applied for the subdomain with 51.2 km width. The three shadings are for S1, S2 and S3. All shadings are translucent. A 15 min running mean is applied.

call that, for S2, when a patch is introduced, the liquid water mass mixing ratio is not altered but the number concentration is. This leads to initially higher c_v for N_d or τ as a result of both larger standard deviation and smaller mean compared with S1, and a trajectory that lies above that of S1. At higher c_v the S2 phase curves for N_d and τ merge to those of S1.

The effect of domain size can be seen in Fig. 11c. All S3 cases again form one phase curve, which starts at approximately the same mean and c_v as in S2, but deviates later. Since F015-N030 of S2 and SP of S3 only differ by 0.01 area fraction for the 51.2 km domain width, without the bias arising from the domain size the shadings for S2 and S3 should overlap each other for the entire duration of the simulation. One can see here that the bias associated with periodic boundary conditions and the associated premature inward propagation of the outflow generated by precipitation hastens the transition from closed to open cell.

6 Conclusions

The main goal of this work is to test the hypothesis posed by Feingold et al. (2010) that precipitation is not a sufficient condition for the formation of open cells in conditions conducive to closed-cellular convection, but that the spatial distribution of precipitation plays a role. To explore the effect of the spatial distribution of precipitation on the formation of open cells, a series of idealized simulations are performed by inserting single or multiple patches with reduced n_t into the closed-cell field. A simple metric for closed- to open-cell transition is developed; the metric could be applied to satellite-based observations of cloud fields, especially geostationary satellites.

For the simulations with single patches (S2), relatively strong, localized precipitation events (order 2 mm day^{-1}) are insufficient to initiate transition to open cells. In general, the transformation to open cells occurs for sufficiently small n_t and/or large patch area fraction. Both of these factors appear

to be of similar importance. For the simulations with multiple patches (S3), the inter-patch distance selects the resulting state for patches that alone are too weak to initiate the open-cell transformation. When patches are close to one another, interactions between precipitation-driven outflows bolster the transition to open cells, whereas when they are separated by large enough distances, the transition is less effective. Large enough distances of separation can be viewed as localized raining areas which, based on the S2 simulations, unless strong and large enough, will not transition the system to open cells.

Modified Hovmöller diagrams of the radially averaged RWP show wavelike patterns, with rainy regions migrating back and forth between the domain center and the outer reaches of the domain. This wavelike behavior is associated with what Wang et al. (2010) called a remote control of open-cell formation; perturbations or existing open cells generate moisture convergence into the surrounding unperturbed closed-cell region, initiate precipitation there, and result in open-cell formation. Furthermore, multiple patches can accomplish this remote control when one patch is not able to achieve this. We further anticipate that (i) if the perturbation is strong enough, the remote control probably propagates more effectively, forming open cells without additional support, and (ii) if multiple remote controls occur simultaneously, interaction of these remote controls could enhance the formation of open cells over a wider area.

Following the idea of slow manifolds of Bretherton et al. (2010), we show that the phase diagrams of the coefficient of variation and mean cloud field properties of LWP, N_d and τ follow one phase trajectory for given large-scale meteorological conditions. The extent to which the phase trajectory is consistent for a range of different meteorological states is left for future research.

In summary, we have shown here that strong, localized surface rain (order 2 mm day^{-1}) is not a sufficient condition for a transition from the closed- to the open-cell state. Both the gradient of n_t and the areal coverage of rain are of equal importance for initiating this transition. Interacting patches, if close enough to one another, also facilitate the transition. Further research will use a combination of satellite-based cloud retrievals and ship-based radar data to test this hypothesis, as well as the existence of preferred variance vs. mean phase trajectories.

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