



Emissions of nitrogen oxides from US urban areas: estimation from Ozone Monitoring Instrument retrievals for 2005–2014

Z. Lu¹, D. G. Streets¹, B. de Foy², L. N. Lamsal^{3,4}, B. N. Duncan⁴, and J. Xing⁵

¹Energy Systems Division, Argonne National Laboratory, Argonne, IL 60439, USA

²Department of Earth and Atmospheric Sciences, Saint Louis University, St. Louis, MO 63108, USA

³Goddard Earth Sciences Technology and Research, Universities Space Research Association, Columbia, MD 21046, USA

⁴NASA Goddard Space Flight Center, Greenbelt, MD 20771, USA

⁵US Environmental Protection Agency, Research Triangle Park, NC 27711, USA

Correspondence to: Z. Lu (zlu@anl.gov)

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Abstract. Satellite remote sensing of tropospheric nitrogen dioxide (NO₂) can provide valuable information for estimating surface nitrogen oxides (NO_x) emissions. Using an exponentially modified Gaussian (EMG) method and taking into account the effect of wind on observed NO₂ distributions, we estimate 3-year moving-average emissions of summertime NO_x from 35 US (United States) urban areas directly from NO₂ retrievals of the Ozone Monitoring Instrument (OMI) during 2005–2014. Following conclusions of previous studies that the EMG method provides robust and accurate emission estimates under strong-wind conditions, we derive top-down NO_x emissions from each urban area by applying the EMG method to OMI data with wind speeds greater than 3–5 m s⁻¹. Meanwhile, we find that OMI NO₂ observations under weak-wind conditions (i.e., < 3 m s⁻¹) are qualitatively better correlated to the surface NO_x source strength in comparison to all-wind OMI maps; therefore, we use them to calculate the satellite-observed NO₂ burdens of urban areas and compare with NO_x emission estimates. The EMG results show that OMI-derived NO_x emissions are highly correlated ($R > 0.93$) with weak-wind OMI NO₂ burdens as well as with bottom-up NO_x emission estimates over 35 urban areas, implying a linear response of the OMI observations to surface emissions under weak-wind conditions. The simultaneous EMG-obtained effective NO₂ lifetimes ($\sim 3.5 \pm 1.3$ h), however, are biased low in comparison to the summertime NO₂ chemical lifetimes. In general, isolated urban areas with NO_x emission intensities greater than ~ 2 Mg h⁻¹ produce statis-

tically significant weak-wind signals in 3-year average OMI data. From 2005 to 2014, we estimate that total OMI-derived NO_x emissions over all selected US urban areas decreased by 49 %, consistent with reductions of 43, 47, 49, and 44 % in the total bottom-up NO_x emissions, the sum of weak-wind OMI NO₂ columns, the total weak-wind OMI NO₂ burdens, and the averaged NO₂ concentrations, respectively, reflecting the success of NO_x control programs for both mobile sources and power plants. The decrease rates of these NO_x-related quantities are found to be faster (i.e., -6.8 to -9.3 % yr⁻¹) before 2010 and slower (i.e., -3.4 to -4.9 % yr⁻¹) after 2010. For individual urban areas, we calculate the R values of pair-wise trends among the OMI-derived and bottom-up NO_x emissions, the weak-wind OMI NO₂ burdens, and ground-based NO₂ measurements, and high correlations are found for all urban areas (median $R = 0.8$), particularly large ones (R up to 0.97). The results of the current work indicate that using the EMG method and considering the wind effect, the OMI data allow for the estimation of NO_x emissions from urban areas and the direct constraint of emission trends with reasonable accuracy.

1 Introduction

Nitrogen oxides (NO_x), the sum of nitrogen dioxide (NO₂) and nitric oxide (NO), is one of the six criteria pollutants identified by the US EPA (United States Environmental Pro-

tection Agency) under the requirement of the Clean Air Act. NO_x plays a crucial role in tropospheric chemistry processes such as the formation of ground-level ozone and secondary inorganic and organic aerosols and, thus, it is also linked with other criterion pollutants including ozone, particulate matter, carbon monoxide, and sulfur oxides. Therefore, NO_x is not only harmful to human health but also implicated in a number of environmental problems, such as acid rain, smog, eutrophication, and climate change. NO_x emissions come from both anthropogenic (e.g., man-made combustion of fossil fuels, biofuel, and biomass) and natural sources (e.g., lightning, microbial processes in soils, and wildfires). Bottom-up inventories of NO_x emissions can be quite uncertain because the emission factors of anthropogenic sources strongly depend on the fuel type, technology, and combustion condition, while natural sources are inherently difficult to quantify.

Due to the strong absorption of NO₂ molecules in the visible wavelength range of the spectrum, satellite instruments based on the principle of optical absorption spectroscopy serve as powerful tools to detect NO₂ signals from space (Martin, 2008 and references therein). The short lifetime of NO_x in the atmosphere leads to a close correlation between observed NO₂ columns and surface NO_x emission sources, implying the potential of space-borne instruments to aid in the estimation of NO_x emissions (Streets et al., 2013, 2014; and references therein). In the past 2 decades, satellite remote sensing of tropospheric NO₂ columns has been widely and successfully used to map the spatial distributions of NO₂ at local, regional, and global scales (e.g., Kim et al., 2009; Russell et al., 2010; Boersma et al., 2007, 2011; Martin et al., 2003), identify intensive point and area NO_x emission sources (e.g., Duncan et al., 2013; Kim et al., 2006; Lu and Streets, 2012; Streets et al., 2014; Wang et al., 2010; Zhang et al., 2009), and monitor diurnal/weekly/monthly/interannual variations of NO₂ (e.g., Hilboll et al., 2013; Hudman et al., 2010; Richter et al., 2005; Russell et al., 2012; Schneider et al., 2015; Tong et al., 2015; van der A et al., 2008) for both anthropogenic and natural sources.

In general, local, regional, and global NO_x emissions can be verified, estimated, and optimized by using forward and inverse modeling of satellite NO₂ columns (e.g., Boersma et al., 2005; Jaeglé et al., 2005; Kim et al., 2009; Martin et al., 2003; Wang et al., 2012). However, NO_x emissions and NO₂ lifetimes can also be determined directly by analyzing the downwind patterns of the satellite-observed NO₂ columns near the sources. Leue et al. (2001) used an exponential function to fit the downwind decay of GOME (Global Ozone Monitoring Experiment)-observed NO₂ columns at the eastern shore of the US and estimated the NO₂ lifetime by using the fitted *e*-folding distance and the averaged wind velocity. Kunhikrishnan et al. (2004) conducted a similar analysis over the Arabian Sea outflow region to estimate the regional NO_x lifetime for the Indian subcontinent. This method was revised by Beirle et al. (2004), who fitted the GOME-observed NO₂ columns across the shipping lane between Sri Lanka and

Indonesia with an exponentially modified Gaussian (EMG) function and derived the mean NO_x lifetime and the corresponding ship emissions for 1996–2001. Hereinafter, we call this approach the EMG method. By fitting the downwind line densities of the OMI (Ozone Monitoring Instrument)-observed NO₂ separately for eight wind directions, Beirle et al. (2011) further improved the EMG method and determined the average NO_x emissions and lifetimes simultaneously for nine worldwide megacities during 2005–2009. Using a similar method, Ialongo et al. (2014) estimated the average summertime NO_x emissions and lifetimes of three cities in the Baltic Sea region during 2005–2011. The EMG method and its variant versions have also been applied to the satellite observations of SO₂ to constrain SO₂ lifetimes and emissions from volcanoes (Beirle et al., 2014; Krotkov et al., 2010; Theys et al., 2013; Carn et al., 2013) and large anthropogenic point sources (Fioletov et al., 2015).

Recently, several studies discussed the applicability and reliability of the EMG method. Valin et al. (2013) suggested that the NO_x emissions and chemical lifetimes would be better quantified when winds are fast because the downwind NO₂ decay under this condition is dominated by chemical removal, not variability of the winds. Introducing the plume rotation technique, they inferred NO_x emissions of Riyadh from the OMI measurements with fast winds ($> 6.4 \text{ m s}^{-1}$) only and derived NO_x chemical lifetimes in slower wind conditions with the mass balance method. Additionally, de Foy et al. (2014) evaluated the performance of the EMG method using simulated column densities over a point source with known emissions under three chemical lifetime cases. They found that the EMG method generally provided reliable emission estimates at fast-wind-speed conditions ($> 3 \text{ m s}^{-1}$); however, the lifetime estimates were biased low and quite sensitive to the selection of the wind-speed cutoff and the accuracy of the plume rotation. This implies that, in practice, the EMG-derived lifetimes should not be treated as chemical lifetimes but rather as “effective lifetimes” that include the influences of chemical conversion, plume meandering, grid resolution, sampling issues, etc. (see also Fioletov et al., 2015; Ialongo et al., 2014). Nevertheless, the EMG method can provide quite accurate emission estimates if the issues of wind speed and direction are appropriately treated.

In this study, we use OMI NO₂ retrievals and an EMG method to estimate NO_x emissions from 35 major US urban areas during the OMI era of 2005–2014. Although there have been a number of studies reporting satellite observations of NO₂ over some US cities, they mainly focused on the interannual trends and/or monthly/weekly variations of the satellite signals themselves (van der A et al., 2008; Hilboll et al., 2013; Schneider et al., 2015; Russell et al., 2012; Kim et al., 2009) or the comparison of satellite observations with pre-existing emissions and/or surface measurement data sets (Tong et al., 2015; Lamsal et al., 2015). In this study, we use the EMG method to estimate NO_x emissions of nearly all major US cities directly from satellite NO₂ observations and

without using a chemical transport model. The prime motivation of this work is not to demonstrate the well-known dramatic decrease of urban NO₂ across the country (although we do have new findings by taking into account the wind effect) but to show the capability of the EMG method to provide direct and reliable estimates of urban NO_x emissions. The current work also differs from previous EMG-related studies (Beirle et al., 2004, 2011; de Foy et al., 2014, 2015; Ialongo et al., 2014; Valin et al., 2013), all of which used a multi-annual averaged satellite NO₂ map in the EMG fit and thus only obtained a long-term averaged emission estimate. This work, to our knowledge, is the first study to show that the EMG method can also provide estimates of emission trends with reasonable accuracy. The rest of the paper is organized as follows: Sect. 2 documents the methodology and data sets; Sect. 3 highlights the effect of wind on the OMI NO₂ observations (Sect. 3.1), presents the relationship between the EMG-derived NO_x emissions and the OMI NO₂ observations (Sect. 3.2), and compares the trends of various NO_x-related quantities (Sect. 3.3); and Sect. 4 summarizes the major findings of this work.

2 Data and methodology

2.1 OMI NO₂ retrievals and processing

The OMI is an ultraviolet/visible nadir spectrometer onboard the National Aeronautics and Space Administration (NASA) Aura satellite, which was launched in a sun-synchronous ascending orbit at 705 km altitude in July 2004 (Levelt et al., 2006). It measures solar irradiance and earthshine radiance in the wavelength range of 270–500 nm and has been continuously providing aerosol and gaseous (including NO₂) column observations at approximately 13:45 local Equator-crossing time (LT) with nearly daily global coverage in the past decade. In this work, we use the version 2.0 product of the Dutch OMI NO₂ (DOMINO) tropospheric vertical column densities (TVCDs) developed at the Royal Netherlands Meteorological Institute (KNMI) for the years from 2005 to 2014 (Boersma et al., 2007, 2011). This version of the product uses an improved OMI NO₂ retrieval algorithm on the basis of better air mass factors, an a posteriori correction for across-track stripes, and high-resolution input profiles of the terrain height and the surface albedo climatology, and has been reported to be in better agreement with independent measurements and model simulations in comparison to the previous version (Boersma et al., 2011).

To increase the amount of valid OMI data, the filter criteria of the level-2 swath data were relaxed somewhat. We removed the daily pixel retrievals with solar zenith angle > 80°, cloud radiance fraction > 0.5, or surface albedo > 0.3. The largest five pixels at the swath edges (i.e., rows 1–5 and rows 56–60) were excluded to limit the across-track pixel width to < 70 km. Since June 2007, some row anomalies (RAs)

have developed on the OMI detectors and affected the data quality of a number of cross-track scenes (<http://www.knmi.nl/omi/research/product/rowanomaly-background.php>). The RAs change over time and we therefore dynamically removed the affected pixels based on the RA flags in the DOMINO product. We only used the summer half-year data (i.e., April–September) because the short NO_x lifetime in this period makes the relationship between NO_x emissions and satellite NO₂ observations more direct than in other months (e.g., Russell et al., 2012; Wang et al., 2012; Lu and Streets, 2012). For the OMI NO₂ maps of the entire domain of the continental US (e.g., Fig. 1a), all valid pixels were oversampled on a 2 km × 2 km grid to obtain detailed spatial distributions of NO₂ over hotspots (Lu et al., 2013; Russell et al., 2010; Fioletov et al., 2011, 2013; de Foy et al., 2009).

2.2 Selection of urban areas

Table 1 lists and Fig. 1a shows the locations of the urban areas selected in this work. We examined the top 50 urban areas in the US based on population size and the observed satellite NO₂ signals (Fig. 1a). We combined adjacent urban areas that are so close as to share the same NO₂ hotspot (e.g., Washington, DC, and Baltimore; and Los Angeles and Riverside) and omitted some urban areas where the NO₂ signals are not isolated due to the influence of large NO_x emitting sources nearby (e.g., Pittsburgh, Milwaukee, San Francisco). In total, 35 urban areas were selected for analysis and together they accounted for ~ 23 % of total NO_x emissions and ~ 50 % of the total urban population in the US during the period 2005–2014.

2.3 Wind fields

Wind information (including speed and direction) is crucial in exploring its influence on the OMI NO₂ observations and estimating the NO_x emissions with the EMG method described in Sect. 2.4. In this work, we use the gridded wind field data sets of the ERA-interim reanalysis at the resolution of 0.5° × 0.5° developed by the European Centre for Medium-range Weather Forecast (ECMWF; Dee et al., 2011). The ERA-interim reanalysis provides global wind fields for 60 vertical levels at four time steps per day (i.e., 00:00, 06:00, 12:00, 18:00 UTC) from 1979 to present on the N128 reduced Gaussian grid. Valin et al. (2013) expected that the EMG results would be insensitive to the choice of wind field data sets. This was confirmed by de Foy et al. (2015), who tested wind fields of both the ERA-interim reanalysis and the North American Regional Reanalysis (NARR) and obtained similar results in the EMG analysis. The NO_x emitted near the surface of the urban areas can undergo rapid vertical mixing, and we thus use the averaged wind fields of the bottom eight levels (i.e., from the surface to ~ 500 m), similar to the treatment of Beirle et al. (2011). We assume that the daily OMI NO₂ spatial pattern of a hotspot should not reflect

Table 1. Summary of the average bottom-up NO_x emissions, OMI-derived NO_x results, ground-based NO₂ measurements, and their linear trends for 35 US urban areas during the summer half-year (April–September) from 2006* to 2013*^a.

Urban areas	Latitude	Longitude	NEI emissions 2005–2014 (Mg h ⁻¹)	Results at high wind speeds (WS)			Results at low WS		Linear trends from 2006* to 2013* (% yr ⁻¹)				Mean R ^b
				Mean WS (m s ⁻¹)	OMI-derived emissions (Mg h ⁻¹)	Effective lifetime (h)	Mean WS (m s ⁻¹)	OMI burden (Mg)	NEI emissions	OMI-derived emissions	OMI burden at WS < 3 m s ⁻¹	AQS	
Atlanta, GA	33.74	-84.32	12.7	5.7	6.7 ± 2.8	4.3 ± 1.4	1.9	25.0 ± 10.1	-7.9	-7.8 ± 3.3	-15.3 ± 6.2	-5.0	0.80
Boston, MA	42.38	-71.02	10.3	6.1	10.9 ± 4.5	5.3 ± 1.7	1.9	29.5 ± 11.9	-6.1	-13.4 ± 5.6	-8.1 ± 3.3	-4.7	0.84
Charlotte, NC	35.34	-80.86	3.3	5.8	2.7 ± 1.1	4.0 ± 1.3	1.8	8.7 ± 3.5	-1.3	-7.9 ± 3.4	-12.9 ± 5.2	-6.1	0.54
Chicago, IL	41.78	-87.68	30.7	7.4	23.3 ± 9.7	6.1 ± 1.9	2.1	86.3 ± 34.8	-6.6	-7.8 ± 3.3	-5.7 ± 2.3	-6.8	0.95
Cincinnati, OH	39.12	-84.50	7.9	4.9	4.9 ± 2.0	5.6 ± 1.8	1.8	16.9 ± 6.8	-3.9	-8.5 ± 3.6	-6.0 ± 2.5	-6.0	0.73
Dallas, TX	32.86	-96.96	14.8	7.4	8.1 ± 3.4	3.3 ± 1.1	2.0	26.8 ± 10.8	-7.4	-12.2 ± 5.1	-5.3 ± 2.1	-7.3	0.86
Denver, CO	39.78	-105.04	10.0	6.0	12.1 ± 5.0	3.5 ± 1.1	1.8	21.7 ± 8.8	-2.4	-9.8 ± 4.1	-9.4 ± 3.8	-2.1	0.78
Detroit, MI	42.26	-83.12	26.1	6.4	18.7 ± 7.8	5.2 ± 1.7	2.0	57.7 ± 23.3	-6.6	-3.4 ± 1.5	-5.8 ± 2.3	-5.3	0.76
El Paso, TX	31.74	-106.38	2.2	6.7	3.2 ± 1.3	3.0 ± 0.9	1.9	7.2 ± 2.9	-3.3	-7.2 ± 1.6	-4.4 ± 1.8	-4.6	0.77
Houston, TX	29.82	-95.28	13.5	5.9	11.3 ± 4.7	4.1 ± 1.3	1.9	30.3 ± 12.2	-7.9	-5.2 ± 2.3	-5.1 ± 2.1	-4.8	0.79
Indianapolis, IN	39.80	-86.12	4.3	5.6	3.1 ± 1.3	4.2 ± 1.3	2.0	8.5 ± 3.4	-3.2	-5.7 ± 2.5	-6.5 ± 2.6	-7.8	0.86
Jacksonville, FL	30.40	-81.60	5.2	5.7	4.7 ± 2.0	2.5 ± 0.8	1.9	9.9 ± 4.0	-9.5	-6.3 ± 2.8	-9.5 ± 3.8	-2.9	0.80
Kansas City, MO	39.10	-94.56	10.2	6.6	5.1 ± 2.1	3.9 ± 1.2	1.9	14.3 ± 5.8	-4.4	-7.8 ± 3.3	-13.4 ± 5.4	-4.9	0.86
Las Vegas, NV	36.18	-115.14	6.1	6.4	6.7 ± 2.8	2.0 ± 0.7	1.9	11.2 ± 4.5	-3.3	-10.3 ± 4.4	-12.3 ± 5.0	-3.0	0.60
Los Angeles, CA	34.06	-117.92	40.1	3.7	40.0 ± 16.6	3.6 ± 1.2	2.0	124.4 ± 50.2	-10.7	-7.0 ± 2.9	-8.5 ± 3.4	-7.6	0.97
Louisville, KY	38.20	-85.74	6.3	5.6	2.5 ± 1.0	3.5 ± 1.1	1.9	8.1 ± 3.3	-7.6	-9.0 ± 3.8	-11.3 ± 4.6	-13.1	0.70
Memphis, TN	35.10	-90.04	4.4	5.9	1.5 ± 0.6	3.2 ± 1.0	1.9	3.4 ± 1.4	-7.3	-25.9 ± 10.8	-10.2 ± 4.1	-2.7	0.83
Miami, FL	26.02	-80.34	13.4	5.4	5.6 ± 2.3	5.0 ± 1.6	1.9	28.7 ± 11.6	-6.5	-10.2 ± 4.3	-4.5 ± 1.8	-9.3	0.80
Minneapolis, MN	44.96	-93.22	12.8	6.9	9.3 ± 3.9	2.7 ± 0.9	2.0	25.9 ± 10.5	-8.6	-12.4 ± 5.2	-11.3 ± 4.6	-11.0	0.89
Nashville, TN	36.14	-86.62	2.9	5.6	2.0 ± 0.8	2.8 ± 0.9	1.8	4.0 ± 1.6	-4.4	-13.8 ± 5.8	-14.9 ± 6.0	-6.9	0.78
New Orleans, LA	29.98	-90.22	7.2	5.3	3.6 ± 1.5	3.2 ± 1.0	1.8	6.0 ± 2.5	-5.2	-7.3 ± 3.2	-1.8 ± 1.1	-5.4	0.48
New York, NY	40.72	-73.80	43.2	5.3	50.7 ± 21.1	3.1 ± 1.0	1.9	128.1 ± 51.7	-6.3	-5.9 ± 2.5	-6.8 ± 2.8	-6.7	0.96
Philadelphia, PA	39.98	-75.16	17.8	5.2	23.3 ± 9.8	3.2 ± 1.0	1.9	53.0 ± 21.4	-7.2	-9.1 ± 4.0	-18.1 ± 7.3	-7.2	0.93
Phoenix, AZ	33.54	-112.00	10.8	5.4	12.2 ± 5.1	1.8 ± 0.6	1.7	21.1 ± 8.5	-4.7	-13.0 ± 5.5	-6.4 ± 2.6	-4.6	0.80
Portland, OR	45.44	-122.60	6.9	3.9	9.9 ± 4.1	1.2 ± 0.4	2.1	15.8 ± 6.4	-3.8	-5.0 ± 2.2	-11.6 ± 4.7	-8.1	0.91
Richmond, VA	37.42	-77.30	3.6	4.9	1.8 ± 0.7	3.5 ± 1.1	2.0	5.1 ± 2.1	-7.8	-5.7 ± 2.7	-14.7 ± 5.9	-9.4	0.68
Salt Lake City, UT	40.72	-111.92	3.6	4.8	8.2 ± 3.5	1.3 ± 0.4	1.8	14.3 ± 5.8	-3.7	-12.1 ± 5.4	-9.3 ± 3.8	-10.8	0.89
San Antonio, TX	29.56	-98.44	5.4	5.7	3.2 ± 1.4	2.1 ± 0.7	2.0	7.9 ± 3.2	-5.7	-10.2 ± 4.4	-8.2 ± 3.4	-1.6	0.75
San Diego, CA	32.66	-116.86	6.0	4.0	8.8 ± 3.7	3.1 ± 1.0	2.0	21.7 ± 8.8	-9.8	-6.3 ± 3.0	-4.2 ± 1.7	-7.8	0.91
Seattle, WA	47.42	-122.22	13.0	3.7	13.3 ± 5.7	3.4 ± 1.1	2.0	30.0 ± 12.1	-4.8	-6.5 ± 3.3	-4.3 ± 1.9	-5.6	0.80
St. Louis, MO	38.64	-90.32	11.0	5.2	4.9 ± 2.0	6.8 ± 2.1	1.9	15.8 ± 6.4	-0.7	-10.9 ± 4.5	-8.9 ± 3.6	-10.0	0.59
Tampa, FL	27.90	-82.42	8.5	5.6	6.9 ± 2.9	2.7 ± 0.9	1.8	14.3 ± 5.8	-9.5	-6.6 ± 2.8	-9.1 ± 3.7	-10.5	0.80
Tucson, AZ	32.24	-110.88	3.1	5.9	1.5 ± 0.6	3.6 ± 1.2	1.8	3.9 ± 1.6	-6.0	-6.2 ± 2.7	-4.0 ± 1.7	-7.4	0.59
Virginia Beach, VA	36.90	-76.32	6.1	6.2	4.6 ± 1.9	1.4 ± 0.4	2.0	7.3 ± 2.9	-8.7	-8.9 ± 3.7	-8.7 ± 3.5	-6.1	0.79
Washington, DC	39.20	-76.58	18.5	5.0	13.0 ± 5.5	4.7 ± 1.5	1.9	48.5 ± 19.6	-7.3	-10.2 ± 4.3	-6.9 ± 2.8	-6.2	0.92

^a 2006* and 2013* denote the 3-year average of 2005–2007 and 2012–2014, respectively. ^b Average correlation coefficients (*R*) of pair-wise trends among the NEI NO_x emissions, the OMI-derived NO_x emissions, the OMI NO₂ burdens, and the AQS NO₂ measurements.

the wind strength and direction at the OMI overpass time of $\sim 13:45$ LT but the average wind fields a few hours before the satellite takes the measurement. For simplicity, we choose 12:00 LT as the time of the wind fields. The reasons are as follows: first, the choice of 12:00 LT of the ERA-interim re-analysis data set has been proven to reproduce the observed spatial transport pattern of the OMI NO₂ successfully at the daily level (Valin et al., 2013); second, if the average wind over the last 6 h is considered, the results have been reported to change only $\sim 10\%$ (Beirle et al., 2011). Consequently, daily wind speed and wind direction maps over the entire domain of the US were interpolated temporally at 12:00 LT and spatially on a 2 km \times 2 km grid in association with the over-sampled OMI NO₂ maps.

2.4 Exponentially modified Gaussian method

Beirle et al. (2011) presented a method using an EMG function to fit the downwind patterns of OMI NO₂ line densities separately for eight wind directions and simultaneously determined the NO_x emissions and lifetimes for nine megacities around the world. In this work, we follow a similar methodology but with a number of enhancements. For each urban area, we did not separate the OMI NO₂ measurements into different wind directions but rotated and overlapped the daily OMI NO₂ maps in the range of 300 km around the urban center (see Table 1 for the latitudes and

longitudes) to align all the wind directions at the urban center in the *x* direction (Valin et al., 2013; de Foy et al., 2014, 2015). This process increases the number of OMI samples, potentially increases the signal-to-noise ratio, and benefits the trend analysis using the EMG method. The wind-aligned OMI NO₂ maps were further reduced to one-dimensional line densities by integrating the NO₂ data in the across-wind direction over a maximum interval of ± 120 km (e.g., Chicago in Fig. 2). Depending on the size of the urban areas, smaller across-wind integration intervals down to ± 60 km were chosen for smaller NO₂ hotspots to minimize interference from background NO₂ and neighboring NO_x sources. The EMG model proposed by Beirle et al. (2011) was then used to fit the NO₂ line densities. As a function of the distance from the urban center *x*, the EMG model is expressed

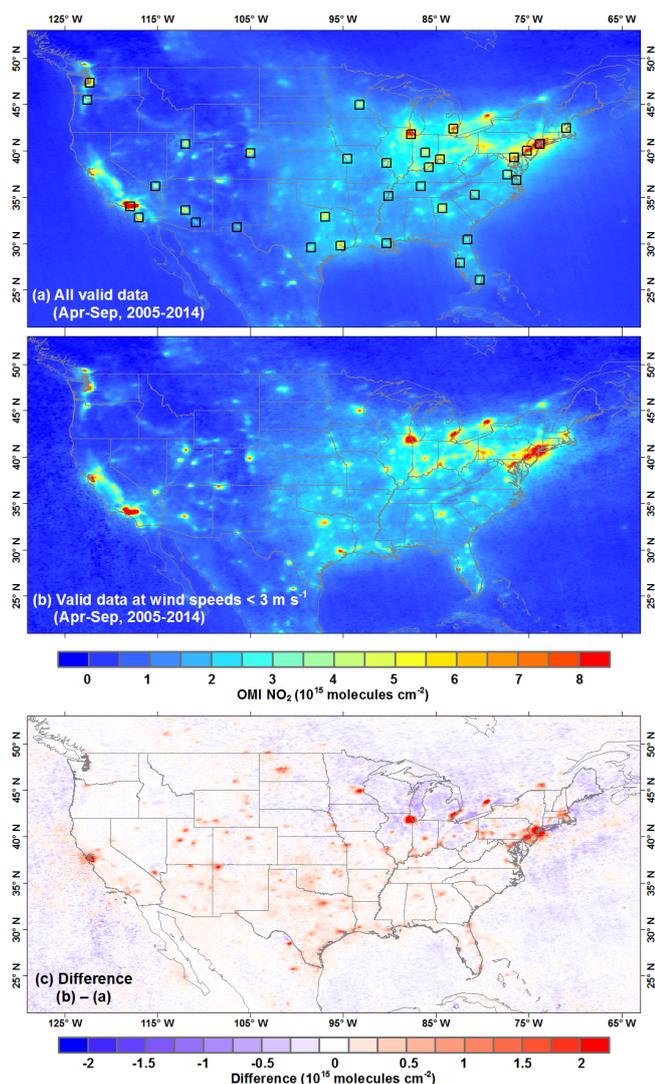


Figure 1. Average summer half-year (i.e., April–September) OMI NO₂ TVCDs over the continental US during 2005–2014: (a) all valid data were used, (b) only valid data with wind speeds $< 3 \text{ m s}^{-1}$ were used, and (c) the difference between (b) and (a). Squares in (a) indicate the urban areas selected in this work.

as Eqs. (1)–(3):

$$\begin{aligned} \text{OMI}_{\text{NO}_2, \text{line}}(x|\mu, \sigma, x_0, \alpha, B) &= \alpha \cdot f(x|\mu, \sigma, x_0) + B \\ &= \alpha \cdot [e(x|x_0, \mu) \otimes G(x|\sigma)] + B \end{aligned} \quad (1)$$

$$\begin{aligned} e(x|x_0, \mu) &= \exp\left(-\frac{x-\mu}{x_0}\right) \text{ for } x \geq \mu, \\ \text{otherwise } e(x|x_0, \mu) &= 0, \end{aligned} \quad (2)$$

$$G(x|\sigma) = \frac{1}{\sqrt{2\pi}\sigma} \exp\left(-\frac{x^2}{2\sigma^2}\right). \quad (3)$$

The estimation problem is nonlinear with five parameters to be determined (i.e., μ , σ , x_0 , α , and B). Mathematically, Eqs. (1)–(3) can be written as (Kalambet et al., 2011, and

references therein)

$$\begin{aligned} \text{OMI}_{\text{NO}_2, \text{line}}(x|\mu, \sigma, x_0, \alpha, B) \\ = \alpha \cdot \left[\frac{1}{x_0} \exp\left(\frac{\mu}{x_0} + \frac{\sigma^2}{2x_0^2} - \frac{x}{x_0}\right) \Phi\left(\frac{x-\mu}{\sigma} - \frac{\sigma}{x_0}\right) \right] + B, \end{aligned} \quad (4)$$

where x_0 in the exponential function $e(x)$ is the e -folding distance downwind, representing the length scale of the NO₂ decay; μ is the location of the apparent source relative to the city center; σ is the standard deviation of the Gaussian function $G(x)$, representing the Gaussian smoothing length scale; Φ is the cumulative distribution function; B is the offset factor representing the background NO₂; $f(x)$ is the convolution of $e(x)$ and $G(x)$; and α is the scale factor of $f(x)$. Since the integration of $f(x)$ equals one, the parameter α physically means the total number of NO₂ molecules observed near the hotspot, excluding the effect of background NO₂. α can be converted to mass units, representing the observed OMI NO₂ burden over the urban areas. Using the mean zonal wind speed w of the NO₂ line density domain, the mean effective NO₂ lifetime $\tau_{\text{effective}}$ and the mean NO_x emissions E can be calculated from the fitted parameters x_0 and α as

$$\tau_{\text{effective}} = x_0/w, \quad (5)$$

$$E = 1.32 \cdot \alpha / \tau_{\text{effective}} = 1.32 \cdot \alpha \cdot w / x_0, \quad (6)$$

where the factor of 1.32 is the mean NO_x/NO₂ ratio suggested by Beirle et al. (2011).

We made additional treatments to the OMI NO₂ data when using the EMG method. For urban areas surrounded by significant NO_x emission sources, we discarded the OMI data with certain wind directions in the plume rotation process to limit the influence of surrounding sources on the wind-aligned OMI NO₂ line densities. For example, Washington, DC, is located $\sim 150 \text{ km}$ southwest of Philadelphia. On the one hand, the NO₂ of Philadelphia can be transported to Washington, DC, through northeasterly winds and affect the upwind pattern of the OMI line densities of Washington, DC. On the other hand, southwesterly winds can bring the NO₂ plume of Washington, DC, to Philadelphia and affect the downwind pattern of the line densities. In this case, daily OMI NO₂ maps with azimuths of $15\text{--}105^\circ$ (i.e., northeasterlies) and $195\text{--}285^\circ$ (i.e., southwesterlies) were excluded in the map rotation process. As discussed in detail in Sects. 3.1 and 3.2, the EMG method provides more accurate estimates of NO_x emissions for OMI NO₂ line densities obtained at stronger-wind-speed conditions (de Foy et al., 2014; Valin et al., 2013; Ialongo et al., 2014), while OMI NO₂ burdens under weak-wind conditions correlate better with NO_x emissions. We therefore divided the OMI observations into high- and low-wind-speed groups (e.g., Fig. 2a and b) and applied the EMG method to both groups. For the high-wind-speeds group, the OMI-derived NO_x emissions (E) and effective NO₂ lifetimes ($\tau_{\text{effective}}$) were estimated and the wind speeds were set to be above thresholds of $3\text{--}5 \text{ m s}^{-1}$ depending on

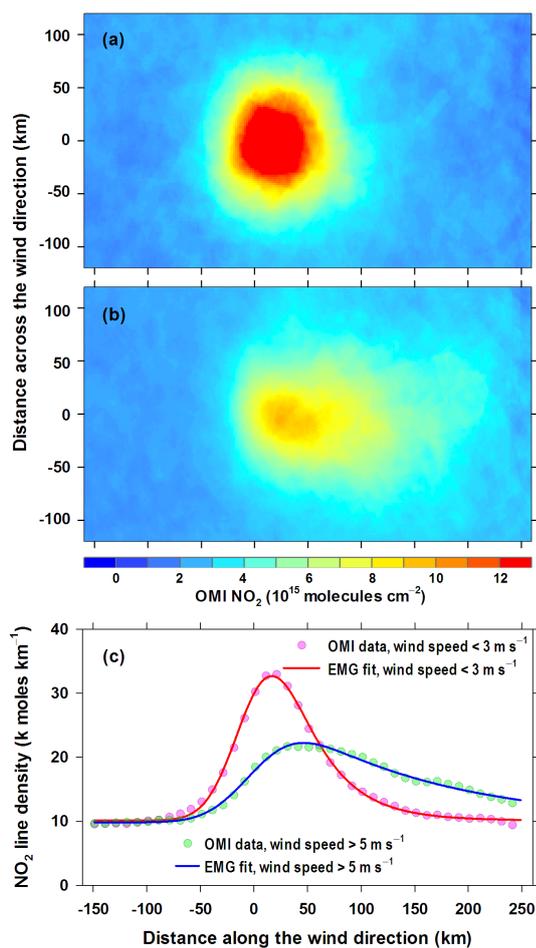


Figure 2. Wind-aligned OMI NO₂ TVCD maps at wind speeds (a) $< 3 \text{ m s}^{-1}$ and (b) $> 5 \text{ m s}^{-1}$ for Chicago in summer months (i.e., April–September) during 2005–2007. (c) OMI NO₂ line densities of (a) and (b) and the corresponding EMG fits. Line densities are from the integration of the NO₂ data in the across-wind direction.

the wind fields of each urban area. For the low-wind-speeds group, the criterion was set to be below 3 m s^{-1} for all investigated urban areas, and the OMI NO₂ burdens (α) under the slow wind condition were determined. To get reliable estimates through the EMG fit, we further combined all the valid data in 3 consecutive years in the analysis. Therefore, most results shown in this work are 3-year averages or 3-year moving trends. For simplicity, we add an asterisk to the middle year to represent the period of 3 years (e.g., 2006* denotes 2005–2007). Through the abovementioned treatments, there are at least 30 (up to ~ 250) valid OMI observations covering the line density domain of each urban area for both the high- and low-wind-speed cases in any 3 consecutive years during 2005–2014.

We follow the same method used by Beirle et al. (2011) to characterize the uncertainties of the estimates. Unless specified otherwise, the term “uncertainty” in this article refers to 1 standard deviation (± 1 SD) or the coefficient of variation

(CV, SD divided by the mean) expressed as a percentage. Total uncertainties of estimated NO_x emissions are the quadrature sum of the uncertainties in the NO_x / NO₂ ratio (10 %), DOMINO v2.0 TVCDs (25 %), EMG fitted results, the selected across-wind integration intervals for the line densities (10 %), and the wind fields (30 %; Beirle et al., 2011; Boersma et al., 2011). The latter four terms are taken into account for estimated NO₂ burdens, and the latter three are used in calculating the uncertainties of the effective NO₂ lifetimes.

2.5 Bottom-up NO_x emissions and ground-based NO₂ measurements

The OMI-derived NO_x emissions, NO₂ burdens, and their trends for the major US urban areas are compared with both bottom-up NO_x emissions and ground-based NO₂ measurements. The bottom-up NO_x emissions are based on the US EPA’s National Emission Inventory (NEI; <http://www.epa.gov/ttn/chief/eiinformation.html>). For each urban area, we grouped the counties covering the major urban extent and the major OMI NO₂ plume and treated the sum of NEI emissions of these counties as the bottom-up NO_x emissions of this urban area. The counties grouped for each urban area are listed in detail in Table S1 of the Supplement. NO_x emissions at the county level for years 2005, 2008, and 2011 were taken from the NEI directly, and emissions in other years were scaled on the basis of the NEI annual emission trends (<http://www.epa.gov/ttn/chief/trends/index.html>). We did not take into account NO_x emissions of natural sources such as open biomass burning, soil, and lightning, because they are negligible compared to anthropogenic emissions over urban areas in summer.

The ground-based NO₂ measurements are from the US EPA’s Air Quality System (AQS) database (Demerjian, 2000) acquired from the EPA’s AirData website (http://www.epa.gov/airdata/ad_data.html). We only chose monitoring sites spatially located in the NO₂ hotspots of the urban areas and temporally having continuous records in April–September from 2005 to 2014. A total of 110 qualified sites in 35 urban areas were selected and the detailed information is provided in Table S2. Since the OMI daily overpass time is at $\sim 13:45$ LT, for consistent comparison with the OMI-derived results, we only use hourly NO₂ measurements at 13:00 and 14:00 LT in the analysis. The 3-year average NO₂ concentration of an urban area is then determined from hourly measurements of all included sites.

3 Results and discussion

3.1 OMI NO₂ TVCDs over the continental US and the wind effects

Figure 1a shows the spatial distribution of average summer OMI NO₂ TVCDs over the US during 2005–2014 with

all valid OMI pixel data passing the criteria described in Sect. 2.1. The average TVCDs for the periods of 2005–2007 (i.e., 2006*) and 2012–2014 (i.e., 2013*) are shown in Fig. 3a and b, respectively. Obviously, the NO₂ signals of all selected urban areas are identifiable in these maps because the short lifetime of NO_x in the lower atmosphere makes the NO₂ TVCDs correlate closely with the surface NO_x emissions (e.g., Kim et al., 2006; Martin et al., 2003; Richter et al., 2005). In terms of the OMI-observed NO₂ trend during 2005–2014, a significant reduction in TVCDs of up to 50 % is observed in visible hotspots and an increase of up to $\sim 0.3 \times 10^{15}$ molecules cm⁻² is observed in rural areas, particularly in the central US (Fig. 3c). The hotspot reductions are caused by the technology improvement in the vehicle fleet for the urban areas (Dallmann and Harley, 2010; McDonald et al., 2012) and the mandatory implementation of emission control devices for power plants (Duncan et al., 2013; Kim et al., 2006). The rural increases are believed to be associated with the variations of soil NO_x emissions in recent years (Hudman et al., 2010; Russell et al., 2012). The above NO₂ trends over the US have been more or less reported in a number of previous studies (e.g., Kim et al., 2009; Russell et al., 2012; Tong et al., 2015), although we extended our analysis to the most recent year (i.e., 2014). It should be noted that these previous studies were all based on the satellite maps under all-wind conditions, while we will mainly discuss how the winds affect the satellite-observed NO₂ signals and trends over the urban areas in the following paragraphs.

The effect of different wind speeds on the patterns of the OMI-observed NO₂ columns was first shown by Valin et al. (2013) for Riyadh, Saudi Arabia. Here, we take Chicago as an example to demonstrate this effect. Figures 2a and b display the wind-aligned OMI NO₂ TVCD maps of Chicago when wind speeds are slow ($< 3 \text{ m s}^{-1}$) and fast ($> 5 \text{ m s}^{-1}$), respectively. The corresponding NO₂ line densities are shown in Fig. 2c. At low wind speeds, NO_x emissions accumulate and stagnate near the urban center, making the peak NO₂ columns about twice those observed at high wind speeds, once the background NO₂ is removed. In contrast, NO₂ plumes can be transported further in high-wind-speed conditions, increasing the downwind NO₂ columns at 250 km from the urban center by $\sim 0.9 \times 10^{15}$ molecules cm⁻². These results clearly indicate that the presence of winds, especially high-speed winds, affects the satellite NO₂ observations.

In practice, OMI NO₂ TVCD maps are averaged from valid pixel data with winds at different speeds from different directions (e.g., Figs. 1a; 3a, b); and, consequently, NO₂ signals near the NO_x emitting sources are smeared spatially. Figure 1b shows the summer mean NO₂ TVCDs over the US during 2005–2014 at wind speeds $< 3 \text{ m s}^{-1}$, and the maps for 2005–2007 and 2012–2014 are shown in Fig. 3d and e, respectively. Compared to NO₂ maps under the all-wind condition, NO₂ signals at low wind speeds are obviously higher over the urban areas (as well as in the big iso-

lated power plant areas) and lower in surrounding rural areas (Figs. 1c; 3g, h) and, consequently, more NO₂ hotspots are visible. In summary, satellite NO₂ maps for low-wind-speed conditions highlight the NO_x emission sources. This was also demonstrated in a recent study by Ialongo et al. (2014), who used the OMI pixels with wind speeds $< 5 \text{ m s}^{-1}$ to highlight the NO_x signals of three cities in the Baltic Sea region.

The effect of winds on satellite-observed NO₂ columns is not uniform but depends on the characteristics of the wind fields at each urban location. Figure 4 compares the OMI NO₂ TVCD maps for Chicago and Los Angeles under the all-wind and the weak-wind conditions. Compared to the significant differences in NO₂ columns over Chicago, the discrepancies between the weak-wind and the all-wind conditions over Los Angeles are nearly negligible. This is because the wind fields in these two cities are quite different. According to the ERA-interim reanalysis, the average wind speed of Los Angeles in summer during 2005–2014 was 2.4 m s^{-1} and $\sim 80 \%$ of the total valid OMI data were measured under wind speeds of $< 3 \text{ m s}^{-1}$. However, the average wind speed was 4.8 m s^{-1} for Chicago and the fraction of valid pixel data with wind speeds $< 3 \text{ m s}^{-1}$ was only $\sim 25 \%$. Figure 4 also shows that the OMI NO₂ signals observed at low wind speeds are better correlated with the NO_x emissions. Comparing OMI maps under the all-wind condition (Fig. 4a vs. 4d), Chicago seems to have significantly lower NO_x emissions than Los Angeles. However, on the basis of the NEI, average NO_x emissions from Chicago were about three-quarters of those from Los Angeles during 2005–2014 (Table 1). After removing the pixels with strong winds, the OMI NO₂ signals of Chicago and Los Angeles match the amounts of their NEI NO_x emissions much better.

For the reasons discussed above, in contrast to previous studies that use all-wind NO₂ maps, we utilize the OMI data under weak-wind conditions to calculate the satellite-observed NO₂ columns, burdens, and trends in the quantities in this work. The threshold of the wind speed was chosen to be 3 m s^{-1} for all the urban areas to ensure enough valid OMI samples (> 30 in 3 consecutive years). However, we do not discard the OMI data under the strong wind conditions, but use them with the EMG method to obtain “top-down” NO_x emissions (see Sect. 2.4 and the following sections). It should be noted that the presence of the strong winds may also change the observed NO₂ trends. Figure 3i shows the differences in OMI NO₂ changes over the US between the weak-wind and the all-wind conditions. Greater NO₂ reductions from 2005* to 2013* are observed under the weak-wind condition than under the all-wind condition over a number of cities, including Chicago, Minneapolis, New York, Las Vegas, and Cincinnati. This is possibly related to the nonlinearity of the NO_x chemistry. At fast wind speeds, the decreased NO₂ level over polluted urban areas may increase the NO_x lifetime so that the same reduction in NO_x emissions would lead to a smaller reduction in observed NO₂ columns compared to the slow-wind-speed conditions. For this reason, it

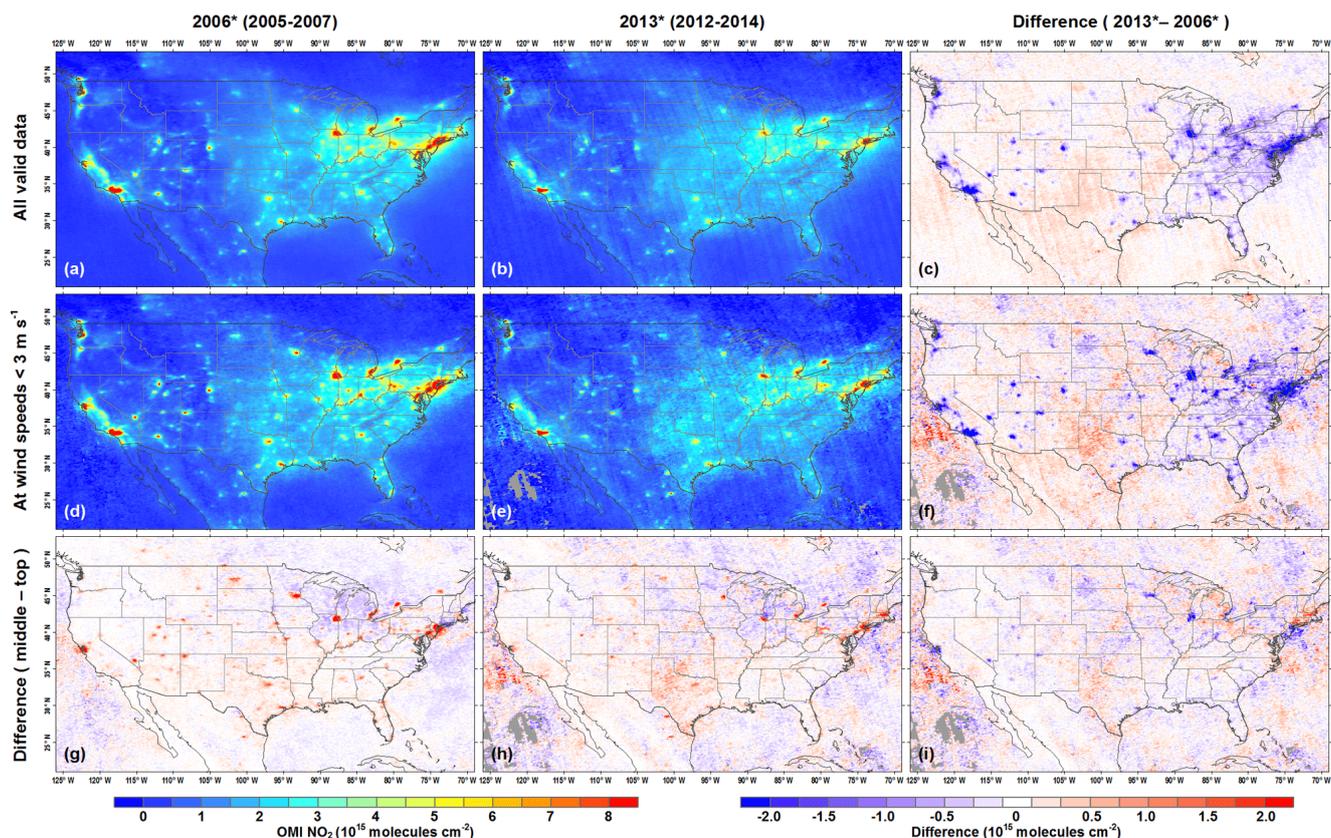


Figure 3. Average summer half-year (i.e., April to September) OMI NO₂ TVCDs over the continental US for (a, d) 2006* (i.e., 2005–2007) and (b, e) 2013* (i.e., 2012–2014): (a, b) all valid data were used, (d, e) only valid data with wind speeds $< 3 \text{ m s}^{-1}$ were used. The right column shows the differences in maps between the middle and the left column. The bottom row shows the differences in maps between the middle and the top row.

is expected that we would derive a higher rate of decline in OMI NO₂ columns over US urban areas than previous studies that have used all-wind OMI data.

3.2 NO_x emissions of US urban areas estimated from the OMI retrievals

As mentioned in Sect. 2.4, we use the EMG method to estimate NO_x emissions from US urban areas. In the original EMG method presented by Beirle et al. (2011), an average wind speed of at least 2 m s^{-1} was required for a target area to guarantee clear downwind outflow NO₂ patterns. However, Valin et al. (2013) pointed out that the variations of wind speed impact the nonlinear NO_x chemistry, and the NO₂ lifetime (and NO_x emissions) inferred from the average spatial pattern of the NO₂ plume is not necessarily equal to the average lifetime (and emissions). They restricted their analysis to OMI measurements made when winds were fast (i.e., $> 6.4 \text{ m s}^{-1}$) because under this condition the downwind decay of NO₂ is dominated by chemical removal, not variability of the winds. In fact, instead of using 6.4 m s^{-1} , the same emission estimates would be obtained if they had used

the OMI measurements with wind speed $> 5 \text{ m s}^{-1}$ (see Fig. 4 of Valin et al., 2013). Recently, de Foy et al. (2014) evaluated the EMG method using simulated column densities over an ideal point source with different chemical lifetimes and wind speeds. They found that the EMG method provided fairly robust and accurate emission estimates when wind speeds were $> 3 \text{ m s}^{-1}$. Furthermore, the EMG-obtained emissions seemed to be more accurate when wind speeds were higher, especially for case of the chemical lifetime $< 12 \text{ h}$ (see Table 2 of de Foy et al., 2014). In this work, we therefore apply the EMG method to the OMI line densities under strong-wind-speed conditions to estimate NO_x emissions. The criterion for the wind speed was set to be $> 5 \text{ m s}^{-1}$ and, if necessary, relaxed to 4 or 3 m s^{-1} to ensure at least 30 valid OMI samples in 3 consecutive years.

Again, we use the example of Chicago to demonstrate our analytical procedure. Figure 2b shows the wind-aligned OMI NO₂ TVCDs at wind speeds $> 5 \text{ m s}^{-1}$ for Chicago during 2005–2007 (i.e., the year 2006*). The NO₂ line densities and the corresponding EMG fit are shown in Fig. 2c. Clearly, the EMG fit reproduces the NO₂ pattern along the wind direction very well. The fitted e -folding distance x_0 , background

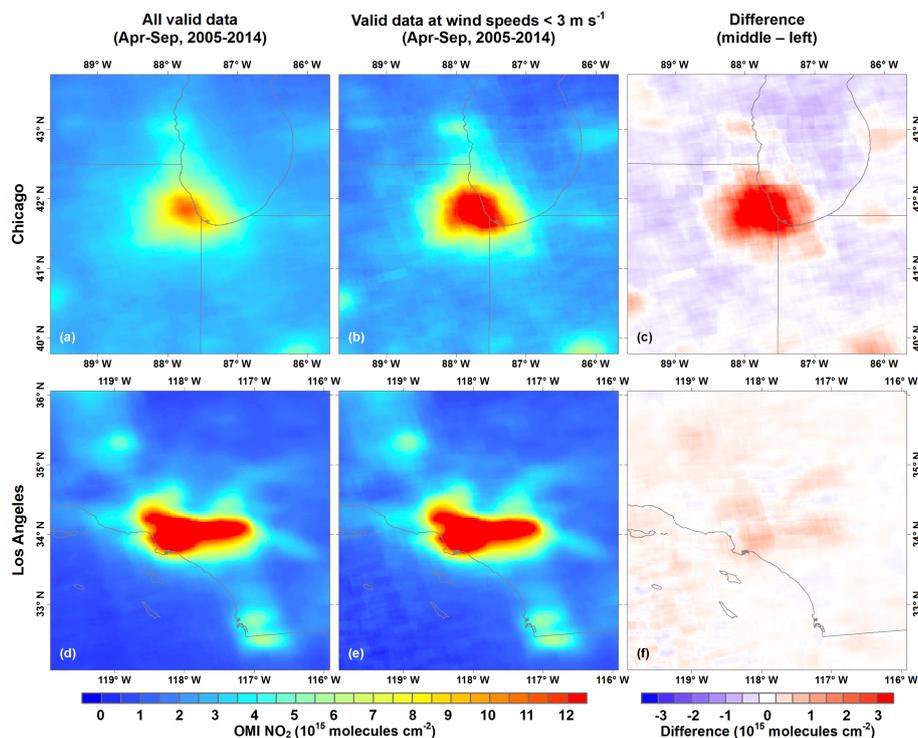


Figure 4. Average summer half-year (i.e., April–September) OMI NO₂ TVCDs over (a, b) Chicago and (d, e) Los Angeles during 2005–2014: (a, d) all valid data were used, (b, e) only valid data with wind speeds < 3 m s^{−1} were used, and (c, f) the difference between the middle and the left column.

B , and the burden α are 144 km, 9.83×10^3 mol km^{−1}, and 2.74×10^6 mol, respectively. The average wind speed w of valid OMI pixels over the studied domain is 7.3 m s^{−1}, so that the effective NO₂ lifetime $\tau_{\text{effective}}$ and the NO_x emissions E are determined to be 5.5 h and 30 Mg h^{−1} through Eqs. (5) and (6), respectively. We also use the EMG method to fit the NO₂ line densities at wind speeds < 3 m s^{−1} (see Fig. 2 and Sect. 3.1), and the OMI NO₂ burden under the weak-wind condition is estimated to be 98 Mg. The same analysis is conducted for all the 3 consecutive years during 2005–2014 and the 3-year moving NO₂ and NO_x trends are summarized in Fig. 5. Results show that the four NO_x-related trends in Chicago correlate with each other very well from 2006* to 2013* ($R > 0.89$). The linear annual average decreasing rates (AADRs) of the OMI-derived NO_x emissions, the OMI-observed NO₂ burdens at slow winds, the NEI NO_x emission estimates, and the ground-based NO₂ measurements are -7.8 , -5.7 , -6.6 , and -6.8 % yr^{−1}, respectively. The AADRs of the OMI-observed NO₂ burden is greater than previously reported values of -3.9 to -5.4 % yr^{−1} that use the all-wind satellite NO₂ maps (Lamsal et al., 2015; Hilboll et al., 2013; Schneider et al., 2015; Russell et al., 2012) but close to those of the “top-down” and “bottom-up” emissions as well as surface measurements. This further implies that the presence of strong winds changes the observed NO₂ trends,

and trends in NO₂ columns obtained at slow winds may better reflect the real bottom-up NO_x emission trends.

The above analysis procedure was applied to all 35 selected US urban areas and the average NO_x-related quantities and linear trends for the entire period of 2006*–2013* are summarized in Table 1. We focus on the relationship between NO_x emissions and NO₂ burdens in this section and discuss the trends in the next section. Figures 6a and b show the scatter plots of both the NEI and the OMI-derived NO_x emissions against the OMI NO₂ burdens under slow-wind conditions. Each point in the scatter plots represents a 3-year moving average for an urban area. Clearly, there is good agreement between NO_x emissions and OMI NO₂ burdens ($R > 0.95$), implying a linear response of the OMI observations to the surface emissions under the weak-wind condition. Generally, urban areas with NO_x emission intensities higher than ~ 2 Mg h^{−1} produce statistically significant OMI burdens and can be analyzed using the method described in this work.

Figure 6c shows the comparison between the OMI-derived and the NEI NO_x emissions for all the selected urban areas. Good agreement was also found between the top-down and the bottom-up emission estimates ($R = 0.94$). The slope of the linear fit indicates that the NEI emissions are on average ~ 6 % higher than the OMI-derived ones. Besides the uncertainties of both estimates, any remaining discrepancies can

Table 2. Summary of NO_x-related trends over all selected US urban areas during 2006*–2013*^a.

	2006*–2013*	2006*–2010*	2010*–2013*
Sum of OMI NO ₂ columns under winds < 3 m s ⁻¹	− 6.9 % yr ⁻¹	− 9.0 % yr ⁻¹	− 3.9 % yr ⁻¹
Total NEI NO _x emissions	− 6.2 % yr ⁻¹	− 6.8 % yr ⁻¹	− 4.9 % yr ⁻¹
Total OMI-derived NO _x emissions	− 7.4 % yr ⁻¹	− 8.7 % yr ⁻¹	− 3.4 % yr ⁻¹
Total OMI NO ₂ burdens under winds < 3 m s ⁻¹	− 7.3 % yr ⁻¹	− 9.3 % yr ⁻¹	− 4.6 % yr ⁻¹
Average NO ₂ concentrations	− 6.3 % yr ⁻¹	− 7.2 % yr ⁻¹	− 4.6 % yr ⁻¹

^a 2006*, 2010*, and 2013* denote the 3-year average of 2005–2007, 2009–2011, and 2012–2014, respectively.

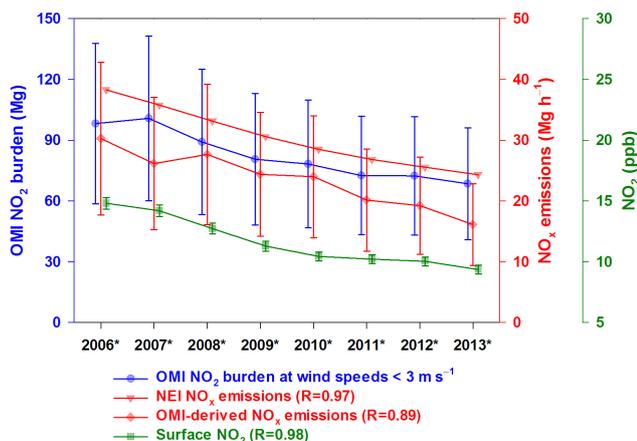


Figure 5. Interannual trends of NEI NO_x emissions, the OMI-derived summertime (April–September) NO_x emissions, the OMI-observed summertime NO₂ burdens at low-wind-speed (< 3 m s⁻¹) conditions, and the average summertime NO₂ concentrations at 13:00–14:00 LT in Chicago during 2006*–2013*. Error bars express the ±1 SD of the annually estimated results. *R* values shown are the correlation coefficients with the OMI-observed NO₂ burdens.

probably be attributed to three factors. First, the NEI NO_x emissions of an urban area are based on the sum of all emissions in counties covering the major urban extent and the major OMI NO₂ plume. Since the outer boundary of the urban area is often somewhat larger than its OMI signals, the NEI values may include more emissions. Second, the OMI-derived NO_x emissions are for the summer half-year, while we did not take into account the seasonality of the NEI emissions. Generally, NO_x emissions of urban areas are lower in summer than in winter because of the relatively low vehicle emissions on warm days and the higher rates of operation of NO_x control devices in some power plants during the ozone season (e.g., Xing et al., 2013; Duncan et al., 2013). Third, a typical NO_x/NO₂ ratio of 1.32 at noon was used in the determination of NO_x emissions, but this scale factor may vary in urban areas depending on the local NO_x chemistry.

In addition to the NO_x emissions, the EMG fits for the OMI line densities at strong winds also yield instantaneous daytime (or, more precisely, 13:00–14:00 LT) lifetimes of

NO₂ for selected urban areas. The model evaluation by de Foy et al. (2014) showed that the EMG method provides accurate estimates of the chemical lifetimes (τ_{chemical}) if the plumes are uniformly transported at fast winds (i.e., 5 m s⁻¹). However, influenced by the inaccurate plume rotation and the use of the satellite data at relatively slow wind speeds, they found the resulting lifetimes were always biased low and could not be treated as the true τ_{chemical} . Here, we call this term the effective lifetime ($\tau_{\text{effective}}$), which can be considered as a combination of τ_{chemical} and an extra lifetime term of τ_{extra} related to the influences of plume meandering, grid resolution, and sampling issues (i.e., lifetimes are combined inversely as shown in Eq. 7; de Foy et al., 2014, 2015). As summarized in Table 1, the estimated $\tau_{\text{effective}}$ values were in the range of 1.2–6.8 h with a mean of $\sim 3.5 \pm 1.3$ h for all studied urban areas during 2005–2014. They are biased low in comparison to the expected summertime NO₂ τ_{chemical} of ~ 7 h estimated for a broader region in the eastern United States (Lamsal et al., 2010) confirming the findings by de Foy et al. (2014), but are consistent with previously reported summertime NO₂ lifetimes of 1–7 h examined for plumes over urban areas (Beirle et al., 2011; Dommen et al., 1999; Ialongo et al., 2014; Nunnermacker et al., 1998; Spicer, 1982), power plants (Fioletov et al., 2015; Nunnermacker et al., 2000; Sillman, 2000), and open biomass burning (Alvarado et al., 2010; Mebust et al., 2011).

It should be noted that the slope of the regression lines of ~ 2.8 h in Fig. 6a and b is also a time term. It can be considered an average timescale of the OMI-observed NO₂ residency over the emission sources under the slow-wind condition. We therefore name it the residence lifetime $\tau_{\text{residence}}$ as suggested by de Foy et al. (2014). In addition to $\tau_{\text{effective}}$, $\tau_{\text{residence}}$ includes the influences of NO₂ physical dispersion in the atmosphere and can be calculated approximately as

$$\frac{1}{\tau_{\text{residence}}} = \frac{1}{\tau_{\text{dispersion}}} + \frac{1}{\tau_{\text{effective}}} = \frac{1}{\tau_{\text{dispersion}}} + \frac{1}{\tau_{\text{chemical}}} + \frac{1}{\tau_{\text{extra}}}, \quad (7)$$

where $\tau_{\text{dispersion}}$ is the physical dispersion timescale. For the slow-wind-speeds condition, the average dimension of the fitting domain downwind from the urban center was 150 km and the average wind speed was 2 m s⁻¹. Hence, $\tau_{\text{dispersion}}$ was about 21 h and the average $\tau_{\text{residence}}$ for all the urban areas was estimated to be ~ 3 h using Eq. (7), assuming that

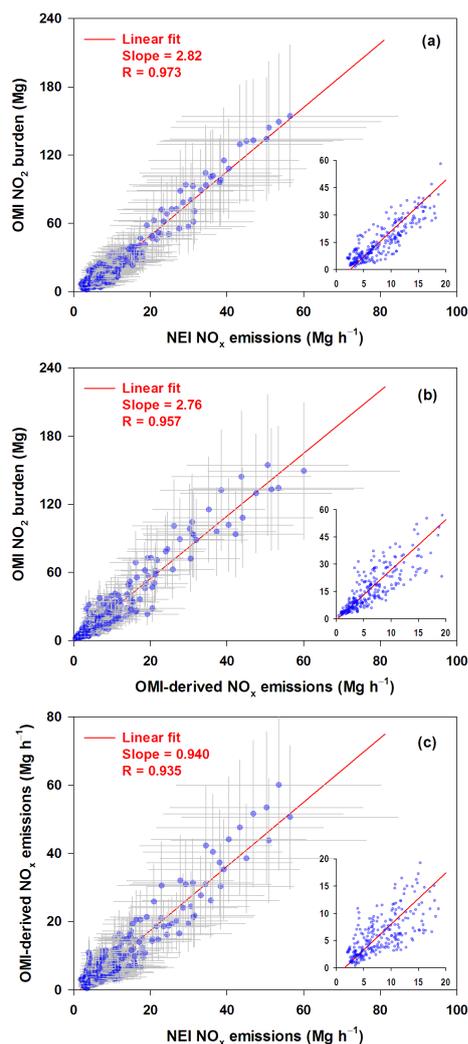


Figure 6. Scatter plots of (a) OMI-observed NO₂ burdens at low-wind-speed ($< 3 \text{ m s}^{-1}$) condition against NEI NO_x emissions, (b) OMI-observed NO₂ burdens against OMI-derived NO_x emissions, and (c) OMI-derived NO_x emissions against NEI NO_x emissions for 35 selected US urban areas during 2005–2014. Each point represents a 3-year result for an urban area. Error bars express the ± 1 SD uncertainties. Uncertainties of NEI emissions are set to be 50 % according to the expert judgment. The inset figures are the enlarged views of points with emissions $< 20 \text{ Mg h}^{-1}$.

$\tau_{\text{effective}}$ did not change significantly with wind speed. This $\tau_{\text{residence}}$ estimation is close to the ones derived directly from Fig. 6a and b (i.e., 2.8 h).

3.3 NO₂ and NO_x trends of US urban areas during 2005–2014

The linear trends of the NEI NO_x emissions, OMI-derived NO_x emissions, OMI burdens under the weak-wind condition, and the AQS NO₂ measurements for all 35 selected urban areas from 2006* to 2013* are summarized in Table 1.

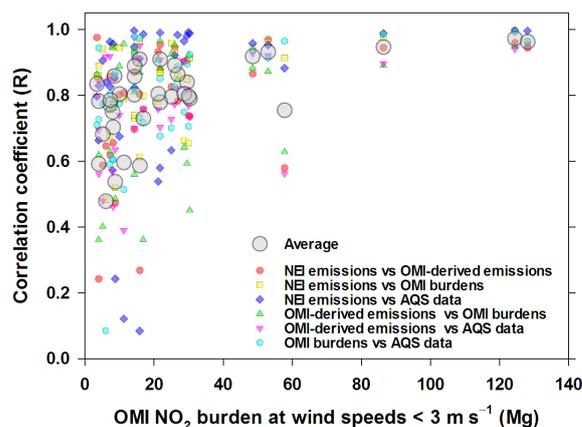


Figure 7. Correlation coefficients of pair-wise trends among the NEI NO_x emissions, the OMI-derived NO_x emissions, the OMI NO₂ burdens at wind speeds $< 3 \text{ m s}^{-1}$, and the AQS NO₂ measurements against the mean OMI NO₂ burdens under the weak-wind-speed condition ($< 3 \text{ m s}^{-1}$) for all selected urban areas during 2006*–2013*. Each large gray circle represents the average of the six correlation coefficients for an urban area.

We have calculated the correlation coefficients of pair-wise trends among these four NO_x-related quantities for each area (see Table 1) and plotted them against the average OMI-observed NO₂ burdens in Fig. 7. Significant reductions in NO₂ and NO_x have occurred in US urban areas. The average percentage reductions among all the studied urban areas from 2006* to 2013* were -34 ± 12 , -46 ± 13 , -45 ± 15 , and -37 ± 12 % for the NEI NO_x emissions, OMI-derived NO_x emissions, OMI-observed NO₂ burdens, and surface NO₂ measurements, respectively. In general, the time series of the four NO_x-related quantities correlate with each other very well in most of the areas. As shown in Fig. 7 and the last column of Table 1, 20 out of the 35 urban areas have an average correlation coefficient > 0.8 , and only 6 areas have an average correlation coefficient < 0.7 . The NO_x-related trends are in better agreement with each other for the larger OMI NO₂ hotspots such as New York, Los Angeles, Chicago, Philadelphia, and Washington, DC (mean $R > 0.92$). The poorest correlation among the four NO_x-related series is observed in New Orleans (mean $R = 0.48$), where NO_x emissions are close to the lowest detection limit of the EMG method we suggested in Sect. 3.2 ($\sim 2 \text{ Mg h}^{-1}$).

The differences in trends of the four NO_x-related quantities in individual urban areas can be attributed to the following reasons. For the OMI-derived NO₂ and NO_x emissions, we have discussed previously that the selection of the wind-speed group and inaccuracy in the wind rotation affects the observed NO₂ trends and the EMG fitted results. Moreover, the EMG method is best suited to point sources; however, urban NO_x emissions are area sources, and the size and shape of the urban area may introduce additional uncertainty to the EMG results. For the NEI emissions, although

NO_x emissions from power plants are measured directly using the continuous emissions monitoring system (CEMS), emissions from other sources (e.g., mobile emissions) are still estimated using bottom-up approaches, which have significant uncertainties inherent in the emission factors and the emissions models that are used (US EPA, 1996). For the AQS data, NO₂ measurements at a limited number of monitoring sites can be readily influenced by nearby emission sources and thus may sometimes reflect localized trends rather than urban-scale trends (e.g., Lamsal et al., 2015). Last but not least, there are spatial and temporal mismatches among emissions, OMI observations, and AQS data (e.g., Tong et al., 2015; Bechle et al., 2013). Spatially, OMI provides measurements of tropospheric NO₂ column densities, AQS data are so-called “nose-level” NO₂ concentrations, while emissions are NO_x masses directly discharged into the atmosphere at a variety of heights above the surface. Temporally, the NEI emissions are annual estimates, the OMI data were restricted to the summer half-year and have gone through a series of filtering processes to remove unreliable pixels, and although we restricted our analysis to the hourly NO₂ measurements close to the OMI overpass time, all AQS measurements at the chosen sites in April–September were used for the trend comparison.

Despite the trend discrepancies caused by these various factors in individual urban areas, we expect the trends of the total (or averaged) NO_x emissions, columns, burdens, and concentrations across all areas to be robust and to reflect the urban NO_x situation at the national level. Figure 8 shows the sum of 3-year averaged OMI NO₂ columns under the weak-wind-speed condition for all urban areas as a function of the distance from the urban centers. Clearly, the sum of OMI signals over the hotspot centers was continuously decreasing during 2006*–2013*. Based on the 2013*/2006* ratio, OMI NO₂ columns over US urban areas decreased by 40 % with an AADR of $-6.9\% \text{ yr}^{-1}$. The 3-year moving trends of the total NEI NO_x emissions, OMI-derived NO_x emissions, OMI-observed NO₂ burdens under slow-wind-speed conditions, and the area-weighted average NO₂ concentrations for all areas are shown in Fig. 9. The four NO_x or NO₂ trends are in excellent agreement with each other ($R > 0.99$). From 2006* to 2013*, total NEI NO_x emissions, OMI-derived NO_x emissions, OMI-observed NO₂ burdens, and the average NO₂ concentrations decreased by 36, 42, 41, and 37 % with AADRs of -6.2 , -7.4 , -7.3 , and $-6.3\% \text{ yr}^{-1}$, respectively (Table 2). The satellite-observed NO₂ rates of decrease obtained in this work are greater than previously reported values. For example, using the OMI BEHR (Berkeley High-Resolution) retrievals, Russell et al. (2012) detected consistent decreases in NO₂ columns (AADR of $-6.2\% \text{ yr}^{-1}$) over 47 US cities during 2005–2011; Tong et al. (2015) examined the OMI NO₂ columns over eight large US cities and found an average AADR of $-6.0\% \text{ yr}^{-1}$ for 2005–2012; with the newly developed NASA OMI product (version 2.2), Lamsal et al. (2015) quantified the average de-

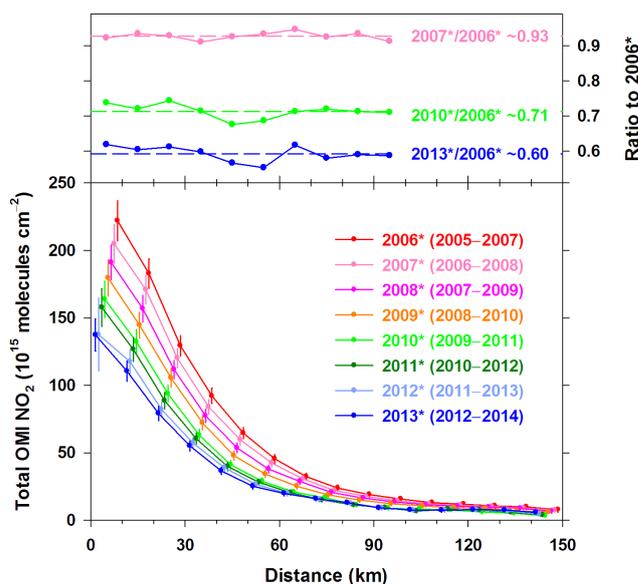


Figure 8. The sum of 3-year averaged OMI NO₂ TVCDs under the weak-wind-speed condition for 35 selected US urban areas as a function of the distance from the urban centers during 2006*–2013*. The background NO₂ of urban areas was removed. Error bars express the 95 % confidence intervals of the mean. The 2007*/2006*, 2010*/2006*, and 2013*/2006* ratios are shown at the top.

creasing rate of NO₂ columns in 20 major US cities from 2005 to 2013 to be $-5.8\% \text{ yr}^{-1}$; and Schneider et al. (2015) used the data from the SCIAMACHY (Scanning Imaging Absorption Spectrometer for Atmospheric Chartography) instrument onboard the Envisat platform and observed decreasing tropospheric NO₂ columns on the order of $-5.8\% \text{ yr}^{-1}$ over nine large urban agglomerations in the US for the period of 2002–2012. Although these previous studies differ in a number of aspects such as satellite data used (i.e., instruments/retrievals/products), time period studied (i.e., summer months or all months), urban areas selected, domain size chosen for each area, trend calculation method used, etc., they derived similar average NO₂ column decrease rates of -5.8 to $-6.2\% \text{ yr}^{-1}$ for US cities since ~ 2005 . This implies that the differences mentioned above may have minor influence on the overall trend analysis results at the country or regional level. However, we obtain a significantly greater column decrease rate of $-7.3\% \text{ yr}^{-1}$ in this work. As discussed in Sect. 3.1, the fact that all these previous studies used all-wind satellite NO₂ maps while we used weak-wind OMI data is the major reason for such discrepancy.

In previous studies such as Russell et al. (2012), Tong et al. (2015), and Lamsal et al. (2015), OMI NO₂ reduction rates were observed to be moderate ($\sim -7\% \text{ yr}^{-1}$), larger ($\sim -9\% \text{ yr}^{-1}$), and smaller ($\sim -3\% \text{ yr}^{-1}$) during the periods of 2005–2007, 2008–2009, and after 2010, respectively, over the US urban areas. The reason for these changes of

pace of the reduction was attributed in these previous studies to the combined effects of the gradually installed NO_x control devices in power plants, transformation to a less-polluting vehicle fleet, the economic recession that happened in 2008, and the slow recovery of the US economy after 2008. In this work, we found similar trends. As shown in Figs. 8 and 9, the sum of OMI columns, the total NEI NO_x emissions, OMI-derived NO_x emissions, OMI NO₂ burdens, and the average NO₂ concentrations over selected urban areas decreased at rates of -6.8 to -9.3% yr⁻¹ during 2006*–2010*, and -3.4 to -4.6% yr⁻¹ during 2010*–2013* (Table 2). We did not observe a greater decreasing rate during the economic recession period, probably because we used 3-year moving trends which smooth the short-term changes. Extrapolating the trends to the years of 2005 and 2014 with AADRs of earlier and later periods, respectively, we estimate that the above five NO_x-related quantities decreased by approximately 47, 43, 49, 49, and 44 %, respectively, during the whole period of 2005–2014.

Although satellite NO₂ column changes cannot be translated to NO_x emission changes directly, due to the nonlinear feedback of NO_x emissions on NO_x chemistry (Lamsal et al., 2011; Lu and Streets, 2012), we indeed obtained similar reductions in total NO_x emissions and total OMI NO₂ observations over all the selected urban areas in the US. Lamsal et al. (2011) used a dimensionless factor of β to express the relationship between changes in NO_x emissions and changes in NO₂ TVCDs:

$$\beta = \frac{\Delta E/E}{\Delta \text{TVCD}/\text{TVCD}}. \quad (8)$$

Generally, β is greater than one in clean regions, because increased NO_x emissions under the low NO₂ condition promotes the generation of OH radicals and thus decreases the NO_x lifetime, while β is less than one in polluted regions since an increase in NO_x emissions consumes OH radicals and increases the NO_x lifetime. On the basis of the monthly global gridded β calculated by Lamsal et al. (2011), the average β over the 35 selected urban areas during April–September was 1.03 ± 0.21 (bounding values 0.74–1.52). This result partially explains why we observed similar trends in total NO_x emissions and total OMI NO₂ columns in this work. It should be noted that we only discuss the overall atmospheric characteristics over all urban areas here. Individual areas may have β values significantly greater or smaller than one, reflecting the local sensitivity of changes in OMI NO₂ columns to NO_x emissions.

4 Summary and conclusions

In the present work, we use the satellite observations of NO₂ vertical columns from the OMI instrument to quantify the summer half-year (i.e., April–September) NO_x emissions and emission trends of 35 selected US urban areas during

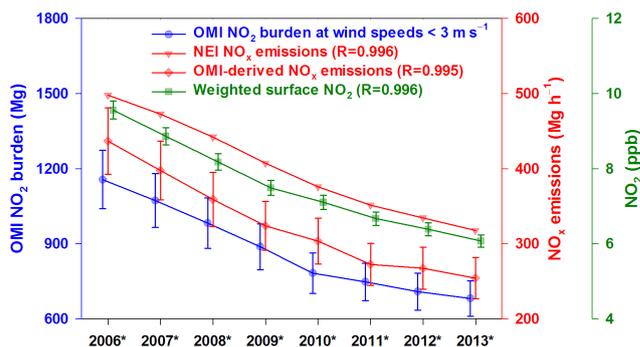


Figure 9. The 3-year moving trends of the total NEI NO_x emissions, the total OMI-derived NO_x emissions, the total OMI-observed NO₂ burdens under the weak-wind-speed condition, and the area-weighted average AQS surface NO₂ measurements for all selected urban areas during 2006*–2013*. Error bars express the ± 1 SD of the estimates. R values shown are the correlation coefficients with the OMI-observed NO₂ burdens.

2005–2014. To refine the analysis, we first explore the impact of winds on the satellite NO₂ observations. Significant differences are found between the OMI NO₂ maps averaged from all valid data and those from data with slow wind speeds only, and such differences are not uniform across all urban areas but depend on local meteorological conditions. Compared to NO₂ maps under all-wind conditions, the satellite-observed NO₂ signals at wind speeds $< 3 \text{ m s}^{-1}$ are significantly higher over the urban areas and lower in surrounding rural areas and are better correlated to the amounts of surface NO_x emissions. We observe greater NO₂ column reductions over a number of selected cities from 2006* (i.e., 2005–2007) to 2013* (i.e., 2012–2014) under the weak-wind condition than under the all-wind condition, implying that the effect of winds should be taken into account when comparing the trends of NO_x emissions and satellite NO₂ observations.

Noticing the importance of wind speed, we divide the OMI observations around each urban area into fast-wind-speed (> 3 to 5 m s^{-1}) and slow-wind-speed ($< 3 \text{ m s}^{-1}$) groups. Daily OMI NO₂ data of each wind-speed group are rotated and oversampled to generate wind-aligned OMI NO₂ maps, the along-wind line densities of which are further fitted by an EMG function. For each urban area in any 3 consecutive years during 2005–2014, we derive the corresponding NO_x emissions and effective NO₂ lifetimes from the EMG fits of the fast-wind-speed groups and the OMI NO₂ burdens from the slow-wind-speed groups. We find good linear agreement ($R > 0.93$) among NEI NO_x emissions, OMI-derived NO_x emissions, and OMI NO₂ burdens, implying the possibility of using the satellite NO₂ observations under the weak-wind condition to constrain the surface NO_x emissions directly. The simultaneously obtained effective NO₂ lifetimes ($\sim 3.5 \pm 1.3 \text{ h}$) are biased low in comparison to the summertime NO₂ chemical lifetime of $\sim 7 \text{ h}$, reflecting the influences

of plume meandering and the coarse sampling resolution on the EMG-fitted results.

Finally, we quantify the NO_x reductions in selected US urban areas and compare the trends of satellite observations with those of bottom-up emissions and ground-based measurements. We find that the time series of the NO_x-related quantities correlate with each other very well in most US urban areas, especially for large cities. Due to the successful control of NO_x emissions in both the power and transportation sectors, the total NEI NO_x emissions, the total OMI-derived NO_x emissions, the sum of OMI NO₂ columns (under the weak-winds condition), the total OMI NO₂ burdens (under the weak-winds condition), and the average measured NO₂ concentrations for all US urban areas decreased by 43, 49, 47, 49, and 44 %, respectively, from 2005 to 2014. Reductions of these five NO_x-related quantities were rapid, at rates of -6.8 to -9.3 % yr⁻¹, before 2010 and slowed down to rates of -3.4 to -4.9 % yr⁻¹ in recent years. Generally, the annual average rates of decrease of OMI NO₂ observations obtained in this work are greater than previously reported values derived from the all-wind satellite maps, further demonstrating the importance of considering winds.

We have shown that using the EMG method, the OMI has the capability to estimate NO_x emissions from urban areas directly and constrain their trends with reasonable accuracy. These OMI-derived emissions can provide independent and valuable information to policy makers and researchers in verifying the bottom-up emission estimates and inspecting the effectiveness of emission control measures, especially for areas without complete surface monitoring networks and lacking well-established emission inventories. We also show that a comprehensive and integrated analysis of satellite observations, ground measurements, and bottom-up emissions can provide a better understanding of the true NO_x situation in a given area. Furthermore, the method described in this work can be applied to the near-future satellite missions such as NASA's Tropospheric Emissions: Monitoring of Pollution (TEMPO; Chance et al., 2013) and the European Space Agency's (ESA) Tropospheric Monitoring Instrument (TROPOMI; Veefkind et al., 2012). With the improved temporal and/or spatial resolution offered by these missions, the diurnal variations of NO_x emissions and emissions from smaller sources are likely to be able to be inferred.

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