



Linking climate and air quality over Europe: effects of meteorology on PM_{2.5} concentrations

A. G. Megaritis^{1,2}, C. Fountoukis², P. E. Charalampidis^{2,3}, H. A. C. Denier van der Gon⁴, C. Pilinis³, and S. N. Pandis^{1,2,5}

¹Department of Chemical Engineering, University of Patras, 26500 Patras, Greece

²Institute of Chemical Engineering Sciences, Foundation for Research and Technology Hellas (FORTH), 26504 Patras, Greece

³Department of Environment, University of the Aegean, University Hill, 81100, Mytilene, Greece

⁴Netherlands Organisation for Applied Scientific Research TNO, Princetonlaan 6, 3584 CB Utrecht, the Netherlands

⁵Department of Chemical Engineering, Carnegie Mellon University, Pittsburgh, PA 15213, USA

Correspondence to: S. N. Pandis (spyros@chemeng.upatras.gr)

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Abstract. The effects of various meteorological parameters such as temperature, wind speed, absolute humidity, precipitation and mixing height on PM_{2.5} concentrations over Europe were examined using a three-dimensional chemical transport model, PMCAMx-2008. Our simulations covered three periods, representative of different seasons (summer, winter, and fall). PM_{2.5} appears to be more sensitive to temperature changes compared to the other meteorological parameters in all seasons.

PM_{2.5} generally decreases as temperature increases, although the predicted changes vary significantly in space and time, ranging from $-700 \text{ ng m}^{-3} \text{ K}^{-1}$ ($-8 \% \text{ K}^{-1}$) to $300 \text{ ng m}^{-3} \text{ K}^{-1}$ ($7 \% \text{ K}^{-1}$). The predicted decreases of PM_{2.5} are mainly due to evaporation of ammonium nitrate, while the higher biogenic emissions and the accelerated gas-phase reaction rates increase the production of organic aerosol (OA) and sulfate, having the opposite effect on PM_{2.5}. The predicted responses of PM_{2.5} to absolute humidity are also quite variable, ranging from $-130 \text{ ng m}^{-3} \%^{-1}$ ($-1.6 \% \%^{-1}$) to $160 \text{ ng m}^{-3} \%^{-1}$ ($1.6 \% \%^{-1}$) dominated mainly by changes in inorganic PM_{2.5} species. An increase in absolute humidity favors the partitioning of nitrate to the aerosol phase and increases the average PM_{2.5} during summer and fall. Decreases in sulfate and sea salt levels govern the average PM_{2.5} response to humidity during winter. A decrease of wind speed (keeping the emissions constant) increases all PM_{2.5} species (on average $40 \text{ ng m}^{-3} \%^{-1}$) due

to changes in dispersion and dry deposition. The wind speed effects on sea salt emissions are significant for PM_{2.5} concentrations over water and in coastal areas. Increases in precipitation have a negative effect on PM_{2.5} (decreases up to $110 \text{ ng m}^{-3} \%^{-1}$) in all periods due to increases in wet deposition of PM_{2.5} species and their gas precursors. Changes in mixing height have the smallest effects (up to $35 \text{ ng m}^{-3} \%^{-1}$) on PM_{2.5}.

Regarding the relative importance of each of the meteorological parameters in a changed future climate, the projected changes in precipitation are expected to have the largest impact on PM_{2.5} levels during all periods (changes up to $2 \mu\text{g m}^{-3}$ in the fall). The expected effects in future PM_{2.5} levels due to wind speed changes are similar in all seasons and quite close to those resulting from future precipitation changes (up to $1.4 \mu\text{g m}^{-3}$). The expected increases in absolute humidity in the future can lead to large changes in PM_{2.5} levels (increases up to $2 \mu\text{g m}^{-3}$) mainly in the fall due to changes in particulate nitrate levels. Despite the high sensitivity of PM_{2.5} levels to temperature, the small expected increases of temperature in the future will lead to modest PM_{2.5} changes and will not dominate the overall change.

1 Introduction

Over the past decades, atmospheric particulate matter (PM) has received considerable attention due to its impact on human health, climate change, and visibility. In particular, fine particulate matter with an aerodynamic diameter less than 2.5 μm (PM_{2.5}) has detrimental effects on human health, as it is associated with increases in mortality, as well as respiratory and cardiovascular diseases (Schwartz et al., 1996; Bernard et al., 2001; Pope et al., 2009). PM_{2.5} has also been implicated in various air quality problems such as changes of the energy balance of the planet (IPCC, 2007), visibility reduction (Seinfeld and Pandis, 2006), and the formation of acid rain and acid fogs (Burtraw et al., 2007).

Concentrations of PM are strongly influenced by meteorology. For example, increasing temperature can lead to elevated sulfate concentrations due to the increased rate of SO₂ oxidation (Aw and Kleeman, 2003; Dawson et al., 2007; Jacob and Winner, 2009; Lecoœur and Seigneur, 2013). In contrast, semi-volatile organic and inorganic aerosols evaporate as temperature increases (Sheehan and Bowman, 2001; Dawson et al., 2007; Tsigaridis and Kanakidou, 2007; Jimenez-Guerrero et al., 2012). Temperature also has a significant indirect effect on secondary organic aerosol (SOA) concentrations. In a warmer climate, secondary organic aerosol can increase due to higher biogenic VOC (volatile organic compounds) emissions (Heald et al., 2008; Jacob and Winner, 2009). Changes in absolute humidity also affect PM_{2.5} levels. Increases in humidity favor nitric acid partitioning to the aerosol phase and therefore can lead to nitrate concentration increases (Dawson et al., 2007; Galindo et al., 2011; Lecoœur and Seigneur, 2013). Wet deposition is in most areas the major removal process for PM_{2.5}, hence changes in precipitation rates or the area extent of precipitation have a significant impact on aerosol concentrations (Dawson et al., 2007; Lecoœur and Seigneur, 2013). Changes in wind speed lead to changes in dispersion and transport as well as to changes in marine and desert aerosol production (Jacob and Winner, 2009; Aksoyoglu et al., 2011). Finally, mixing height determines to a large extent the dilution of primary and the formation of secondary pollutants (Jimenez-Guerrero et al., 2012; Pay et al., 2012).

Over the next decades, climate is expected to change and this change will influence PM_{2.5} concentrations. Based on IPCC projections for Europe (IPCC, 2013), temperature is expected to rise from 1 to 5.5 K over the next century. Emissions of biogenic VOCs are also expected to increase as temperature increases. Forkel and Knoche (2007) predicted a 30 % increase (locally up to 50 %) of biogenic VOCs emissions in Europe due to a predicted 1.7–2.4 °C temperature increase, under the IPCC IS92a scenario applied to the next 30 years. Higher temperatures in a future climate will also lead to increases in absolute humidity (IPCC, 2007). Precipitation is also expected to change over Europe in the future, having large spatial and seasonal variations. Based on

the IPCC A2 emission scenario, Räisänen et al. (2004) predicted an increase in mean winter precipitation in northern and central Europe (up to 50 %) and a substantial decrease in southern Europe in the next century. During summer, precipitation was projected to decrease throughout central and southern Europe. Similar projections for precipitation were also reported by other modeling studies (Giorgi and Meleux, 2007; Hedegaard et al., 2008; Kjellström et al., 2010). In addition, general circulation models (GCMs) and regional climate models (RCMs) predict changes in both rainfall intensity and frequency (Christensen and Christensen, 2004; Frei et al., 2006; Buonomo et al., 2007; Boe et al., 2009; Argüeso et al., 2012). Jacob and Winner (2009) suggested that the critical variable that affects PM concentrations is precipitation frequency rather than precipitation rate. Wind speed is also predicted to change in a future climate. Andersson and Engardt (2010) predicted increases in wind speed over northern Europe, and decreases in the southern regions. Similar projections for wind speed were reported by other model studies (Räisänen et al., 2004; Kjellström et al., 2010; Katragkou et al., 2011). Hedegaard et al. (2013) reported increasing mixing height in most of Europe under a future climate (by more than 100 m in southeastern Europe), but Jimenez-Guerrero et al. (2011) predicted an average decrease for most of continental Europe.

The impact of various climate scenarios on air quality over Europe as well as the correlation between meteorology and PM concentrations have been the subject of several studies (Koch et al., 2003; Heald et al., 2008; Hedegaard et al., 2008, 2013; Jacob and Winner, 2009; Redington et al., 2009; Roustan et al., 2010; Galindo et al., 2011; Im et al., 2012; Manders et al., 2012; Pay et al., 2012; Megaritis et al., 2013). Carvalho et al. (2010) applied a regional CTM, CHIMERE, over Europe with downscaled meteorology generated by a global GCM to study the impact of climate change on ozone and PM₁₀ levels, using the IPCC A2 scenario, which describes a very heterogeneous world, with continuously increasing population, self-reliance and preservation of local identities. Their predicted PM₁₀ concentration changes showed a strong spatial and temporal variability with increases over the continental regions and decreases over water. They concluded that the PM₁₀ response was mainly driven by changes in the boundary layer height and wind speed. Jimenez-Guerrero et al. (2012) used a regional modeling system, MM5-CHIMERE, over southwestern Europe in order to study how concentrations of air pollutants respond to a changing climate for 2100 under the IPCC A2 scenario. Their findings suggest that aerosol species are strongly influenced by the higher future temperatures. They predicted an increase of sulfate and secondary organic aerosols (SOA) due to faster reactions and higher emissions of biogenic VOCs, and a decrease of particulate nitrate. In a multi-year simulation (2000–2008), Lecoœur and Seigneur (2013) used a three-dimensional CTM, Polyphemus/Polair3D, to investigate the response of PM_{2.5} species to changes in meteorology. Their

results suggest that wind speed and precipitation have a strong negative effect on PM_{2.5} and its components (with sea salt being the only exception, for which a positive correlation with wind speed was predicted), while the response of PM_{2.5} to temperature changes varied significantly among the PM_{2.5} species considered. The negative response of PM_{2.5} to wind speed changes and the variable effects caused by changes in temperature were also reported by Aksoyoglu et al. (2011). Additional work has been conducted in several areas over the world, with the majority focused on the United States (Hogrefe et al., 2004; Racherla and Adams, 2006; Dawson et al., 2007, 2009; Tagaris et al., 2007, 2008; Zhang et al., 2008; Avise et al., 2009; Pye et al., 2009; Mahmud et al., 2010; Day and Pandis, 2011; Singh and Palazoglu, 2012; Tai et al., 2012; Jeong and Park, 2013). The predicted PM_{2.5} changes due to climate are quite variable in space and time, and there are often conflicting conclusions about the meteorological variables driving these changes.

Most of the earlier modeling studies have focused on the overall effect of future climate on PM_{2.5} concentrations. While this has provided valuable insights, it has often been difficult to quantify the effects of changes of individual meteorological parameters and processes. One of the few available studies has focused on the United States, studying the sensitivity of PM_{2.5} to different meteorological perturbations (Dawson et al., 2007). However, this study covered a relatively short simulation period, and it did not assess how important these meteorological changes are for individual processes that are related to the formation, transport and removal of PM_{2.5} components. In addition, to our knowledge, there has been little work trying to quantify how these individual processes (such as the partitioning of semi-volatile PM components, the marine aerosol production, etc.) can be affected by changes in meteorology and eventually, how sensitive PM_{2.5} is to these changes. The goal of this study is to conduct a detailed sensitivity analysis quantifying how changes in temperature, wind speed, absolute humidity, precipitation, and mixing height, and their subsequent effects on different processes, can influence fine particulate matter (PM_{2.5}) concentrations over Europe. Each of these parameters is studied separately so that the relative importance of each as well as the subsequent response of PM_{2.5} can be quantified. For this purpose we use a three-dimensional Chemical Transport Model (CTM), PMCAMx-2008, over Europe. PMCAMx-2008 implements a state-of-the-art organic module for organic aerosol (OA) modeling based on the volatility basis set framework (VBS) (Donahue et al., 2006), which has not been used in earlier versions of the model, as well as in earlier climate–air quality interactions studies. The model also uses updated inorganic aerosol modules for the simulation of inorganic PM species. In addition, we covered a 3-month-long simulation period, in order to obtain more representative results regarding the seasonal dependence of the PM_{2.5} response to changes in meteorology.

A brief description of the PMCAMx-2008, along with the characteristics of its application in the European domain, is given in Sect. 2. The PMCAMx-2008 base-case predictions for PM_{2.5} concentrations and some information regarding the model evaluation are given in Sect. 3. The description of each sensitivity simulation conducted in this study as well as the predicted response of PM_{2.5} to these meteorological perturbations are presented in the next sections. Finally, the relative importance of the various meteorological parameters and the main conclusions are presented.

2 The PMCAMx-2008 CTM

2.1 Model description

PMCAMx-2008 (Fountoukis et al., 2011; Megaritis et al., 2013) uses the framework of the CAMx air quality model (Environ, 2003). The chemical mechanism used in this study to describe the gas-phase chemistry is based on the SAPRC99 mechanism (Environ, 2003; Carter, 2010) and includes 211 reactions of 56 gases and 18 free radicals. For the simulation of the aerosol species, the model uses three detailed modules: inorganic aerosol growth (Gaydos et al., 2003; Koo et al., 2003), aqueous phase chemistry (Fahey and Pandis, 2001) and SOA formation and growth (Murphy and Pandis, 2009). These modules employ a sectional approach using 10 aerosol size sections, spanning the diameter range from 40 nm to 40 μm. In this study, inorganic aerosol formation was simulated using the “bulk equilibrium approach”, where the bulk inorganic aerosol and gas phase are assumed to always be in equilibrium. The organic aerosol treatment in PMCAMx-2008 is based on the volatility basis set framework (Donahue et al., 2006; Stanier et al., 2008). Primary organic aerosol is assumed to be semi-volatile in PMCAMx-2008, while the model treats all organic species (primary and secondary) as chemically reactive. Further details regarding the simulation of inorganic and organic aerosol species in PMCAMx-2008 can be found in Fountoukis et al. (2011).

For the simulation of wet scavenging the model assumes that the scavenging rate within or below a cloud due to precipitation is equal to the product of the concentration of a pollutant and the respective scavenging coefficient (Seinfeld and Pandis, 2006). Dry deposition, for the gas-phase species, is simulated using the resistance model of Wesely (1989), while for aerosol species the PMCAMx-2008 uses the resistance approach of Slinn and Slinn (1980) as implemented in UAM-AERO (Kumar et al., 1996). More information about the simulation of removal processes can be found in Fountoukis et al. (2011) and Megaritis et al. (2013).

2.2 Model application

PMCAMx-2008 was set to simulate the atmosphere over Europe covering a 5400 × 5832 km region, with a 36 × 36 km resolution grid and 14 vertical layers extending up to

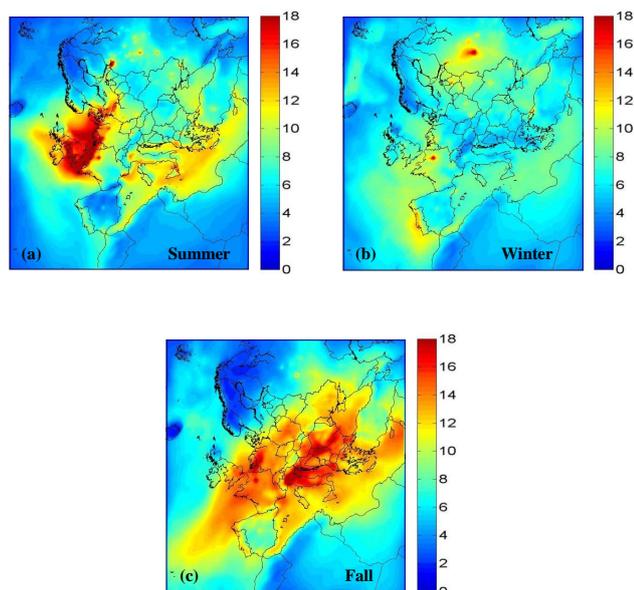


Figure 1. Predicted average base case PM_{2.5} ground-level concentrations ($\mu\text{g m}^{-3}$) during the modeled (a) summer, (b) winter and (c) fall periods.

approximately 6 km altitude. Three month-long periods, representative of different seasons (summer, winter, and fall) were simulated. The summer simulations were based on a hot late spring period (1–29 May 2008), the fall modeled period was from 15 September to 17 October 2008, while the winter simulation covered a cool late winter period (25 February–23 March 2009). The first 2 days from each simulation were used as model initialization days and were excluded from the analysis. All three periods showed a variety of meteorological conditions and pollution levels. The summer period was characterized by a blocking anticyclone (especially in the first half of May), leading to stable meteorological conditions and enhanced pollution over central Europe. In addition, high temperatures were observed in most of Europe (Pikridas et al., 2010; Hamburger et al., 2011; Poulain et al., 2011; Mensah et al., 2012), typical for summer conditions. Fall represented the transition from summer to winter with a moderate temperature (which was decreasing during this period), less stable atmospheric pressure and frequent precipitation events (EMEP, 2010; Poulain et al., 2011) while the winter period was characterized by low temperatures in most of Europe (Hildebrandt et al., 2010; Freney et al., 2011; Poulain et al., 2011; Mensah et al., 2012).

The necessary inputs to the model included emissions, meteorological conditions, land use data and initial and boundary conditions of the simulated PM species. Anthropogenic gas emissions included land as well as international shipping emissions and were developed by the TNO team as a continuation of the work in GEMS and MACC (Visschedijk et al., 2007; Denier van der Gon et al., 2010). Anthropogenic particulate organic and elemental carbon emissions were based

on the EUCAARI Pan-European Carbonaceous Aerosol Inventory (Kulmala et al., 2011). Biogenic emissions were produced by utilizing the MEGAN (Model of Emissions of Gases and Aerosols from Nature) model (Guenther et al., 2006). Sea salt emissions (O'Dowd et al., 2008) as well as wildfire emissions (Sofiev et al., 2009) were also included. Further details about the emissions data used in this study can be found in Fountoukis et al. (2011). The meteorological input into the model included hourly data of temperature, pressure, water vapor, clouds, rainfall, horizontal wind components and vertical diffusivity generated using the meteorological model WRF (Weather Research and Forecasting) (Skamarock et al., 2008). For the boundary conditions of the major PM species, we used the same concentrations as Fountoukis et al. (2011). The boundary conditions were chosen on the basis of measurements taken in sites close to the boundaries (e.g., Seinfeld and Pandis, 2006; Zhang et al., 2007).

3 Base case simulations and model evaluation

The predicted concentrations of total PM_{2.5} during the three modeled base case periods are presented in Fig. 1. During the summer period, the domain-average ground-level concentration of total PM_{2.5} is $7.7 \mu\text{g m}^{-3}$. Elevated PM_{2.5} concentrations are predicted in most of western Europe (up to $25 \mu\text{g m}^{-3}$), due mainly to high ammonium nitrate levels in this area (maximum concentration of $11 \mu\text{g m}^{-3}$). In central and northern Europe fine organic matter dominates with biogenic secondary OA and oxidized primary OA (including OA from intermediate volatile organic compounds, IVOCs) contributing the most (approximately 60 and 25 %, respectively). PM_{2.5} concentrations in these areas may exceed $20 \mu\text{g m}^{-3}$. Sulfate is predicted to be the dominant PM_{2.5} species over the eastern Mediterranean. The strong photochemical activity in this area favors the conversion of sulfur dioxide to sulfate and can partly explain the relatively high PM_{2.5} levels in this area (up to $15 \mu\text{g m}^{-3}$).

During the winter period the mean predicted ground-level concentration of total PM_{2.5} over the domain is $7.1 \mu\text{g m}^{-3}$. Sulfate and organics are predicted to be the major components, contributing approximately 25 and 28 % of total PM_{2.5} mass. Peak period-average concentrations of total PM_{2.5} (mostly over central and northern Europe) exceed $20 \mu\text{g m}^{-3}$, especially in areas with large industrial activity or large urban emissions. OA accounts for up to half of total PM_{2.5} in these areas with fresh primary OA being the dominant OA component.

In the fall period the model predicts an average total PM_{2.5} concentration of $8.3 \mu\text{g m}^{-3}$ over the domain. The elevated PM_{2.5} levels are due to a combination of high ammonium nitrate, sulfate, and organic aerosol. On a domain-average basis organic aerosol and sulfate are predicted to account for 28 % of total PM_{2.5} mass each, followed by ammonium (12 %), and nitrate (10 %). High levels of PM_{2.5} over the Balkans

and the Mediterranean (up to $22 \mu\text{g m}^{-3}$), are mainly due to high sulfate concentrations, while ammonium nitrate dominates over the western parts of the domain. The largest OA concentrations, with a peak of $6.8 \mu\text{g m}^{-3}$, are predicted in the Po Valley area. High OA levels are also predicted over the Balkans.

The model performance against high time resolution Aerosol Mass Spectrometer (AMS) measurements under the various meteorological conditions of the three simulated periods was encouraging. The May period was characterized by the highest average temperatures among the three simulated periods, the largest average mixing height (the average predicted mixing height was 550 m in May compared to 380 m in February/March and 440 m in September/October) and the lowest precipitation rates (Fig. S2 in the Supplement). The wind speed and absolute humidity were quite variable in space and time during all three periods. PM₁ organic aerosol was well predicted by the model during the photochemically active period of May (monthly average measured concentration: $3.3 \mu\text{g m}^{-3}$, PMCAMx-2008: $3 \mu\text{g m}^{-3}$). The observed data included ground measurements taken at four sites in Europe (Cabauw, Finokalia, Mace Head, and Melpitz) (Fountoukis et al., 2011). A similar model performance was found for the September/October period, with the model reproducing more than 74 % of the hourly averaged OA data within a factor of 2, with a fractional bias of -0.1 and a fractional error of 0.48. During the fall period, hourly AMS measurements were performed at several sites in Europe (Hyytiälä, k-Puszta, Melpitz, Puy de Dome, Payerne, Puijo, and Vavihill) during an EMEP intensive campaign. The largest discrepancies between model predictions and measurements for OA were found during the low-temperature period of February/March (monthly average measured concentration: $2.3 \mu\text{g m}^{-3}$, PMCAMx-2008: $1.1 \mu\text{g m}^{-3}$), however, this underprediction of OA was attributed to an underestimation of wood burning emissions (Fountoukis et al., 2014). The model performed equally well in reproducing observed PM₁ sulfate concentrations during all periods, regardless of the different meteorological conditions. The monthly average concentration of PM₁ sulfate predicted by the model in the four sites during summer was $2.9 \mu\text{g m}^{-3}$ compared to the measured value of $2.8 \mu\text{g m}^{-3}$. In agreement with observations, the model predicts the lowest sulfate concentrations during the winter period (monthly average measured concentration: $1.0 \mu\text{g m}^{-3}$, PMCAMx-2008: $0.9 \mu\text{g m}^{-3}$) and somewhat higher concentrations ($1.4 \mu\text{g m}^{-3}$ observed and $1.5 \mu\text{g m}^{-3}$ predicted) during the fall. The model also had a reasonable performance for PM₁ nitrate and PM₁ ammonium in most studied sites, with the exception of sites characterized by both high levels of sea salt and nitrate (e.g., Mace Head), where the model significantly overpredicted both fine nitrate and ammonium concentrations. These errors are mainly due to the assumption of bulk equilibrium that PMCAMx-2008 uses for the inorganic aerosol simulation (Trump et al., 2014). In general, the model performance

was found encouraging for all major PM₁ species under the variable meteorological conditions observed during the three seasons.

4 Sensitivity to meteorological variables

For each of the three modeled periods, we performed a suite of individual sensitivity simulations by perturbing various meteorological parameters, one at a time. The sensitivity tests included perturbations in temperature, wind speed, absolute humidity, precipitation rate, precipitating area and mixing height (Table 1).

Sensitivity to temperature was tested by performing four different simulations. The impact of temperature on biogenic emissions and PM_{2.5} levels was examined, using temperature-sensitive biogenic emissions produced by the MEGAN model, and based on an increase of 2 K. In this simulation the only change was on the biogenic emissions inventory. The temperatures used by the model (to simulate chemistry, thermodynamics, etc.) were those from the base case scenario. The effect of temperature on aerosol thermodynamics was tested in another simulation where we increased temperature by 2 K only for the modules of PMCAMx-2008 that simulate the partitioning of semi-volatile inorganic and organic PM_{2.5} species. Similar to the first simulation, temperatures for the other processes in PMCAMx and all the other meteorological parameters were the same as in the base case simulation. The third test studied the sensitivity of PM_{2.5} to the temperature dependence of the gas-phase reaction rates. The overall temperature effect on PM_{2.5} concentrations (also using temperature-dependent biogenic emissions) was studied in a different simulation where all surface and air temperatures were increased uniformly over the domain by 2 K, keeping all the other meteorological inputs constant.

The effect of wind speed on PM_{2.5} concentrations was studied by two different simulations. We first used a simplified scenario where horizontal wind speed was decreased uniformly over the entire domain by 10 %, keeping all other inputs constant. The vertical wind components were calculated from the perturbed horizontal wind speeds to ensure mass conservation. In this simulation the only changes were on the dispersion coefficients, as well as the transport (vertical velocity, advection, dilution) and removal processes (dry deposition rate), while sea salt emissions were kept constant as in the base case. In the second test, we examined the effect of wind speed on marine aerosol emissions, recalculating the corresponding emissions inventory for wind speeds decreased by 10 %. This simulation examines only changes in sea salt emissions, therefore wind speed and all other meteorological data used as input by the model were those of the base case scenario.

The effect of absolute humidity was tested based on a uniform increase of 5 % over the entire domain. Precipitation intensity was increased uniformly by 10 % to study its

Table 1. Description of performed sensitivity simulations.

Meteorological parameter	Change examined	Directly affected in simulation
Temperature	+2 K	Biogenic VOC emissions only
	+2 K	Organic and inorganic aerosol thermodynamics only
	+2 K	Gas-phase chemistry only
	+2 K	All temperature-dependent processes (including BVOC emissions)
Wind speed	−10 %	Turbulent dispersion coefficients, advection, dry deposition; emissions (including marine) were kept constant.
	−10 %	Marine aerosol emissions only
Absolute humidity	+5 %	Reaction rates with H ₂ O, aerosol thermodynamics.
Precipitation rate	+10 %	Wet deposition
Precipitation area	+10 %	Wet deposition
Mixing height	+1 model layer	Vertical dispersion

effects. Sensitivity to the spatial extent of precipitation was investigated in a simulation where the area undergoing precipitation was increased by +10 %. This was done by extending the existing precipitating area into non-precipitating but adjacent cells, which were chosen randomly. In addition, the sensitivity of PM_{2.5} to mixing height changes was examined in a simulation where the mixing height was increased by one model layer. This was done by changing the vertical diffusivity in only the layer immediately above the base case mixing height. The corresponding average change was an increase in mixing height by approximately 150 m.

Table 1 summarizes the sensitivity simulations performed in this study and the processes that were perturbed directly in each change. Initial and boundary conditions of the modeled PM species did not change compared to the baseline scenario, consistent over all tests. Emissions of all pollutants were also kept constant as in the base case conditions in all tests, except for the two simulations using temperature sensitive biogenic emissions and new sea salt emissions due to wind speed change. Previous work has shown that the processes that are perturbed in each of these sensitivity runs are well represented in the model. Inorganic aerosol thermodynamics are simulated through the ISORROPIA model (Karydis et al., 2010), while organic gas/aerosol partitioning is simulated with the state-of-the-art volatility basis set framework (Murphy and Pandis, 2009; Fountoukis et al., 2014). For the gas-phase chemistry, the chemical mechanism SAPRC99 is used, instead of the older Carbon Bond-IV, increasing the number of reactions from 100 to 211 (Tsimpidi et al., 2010). Biogenic VOC (BVOC) emissions as well as marine aerosol emissions are well represented, as described in Fountoukis et al. (2011).

5 Sensitivity to temperature

5.1 Temperature-dependent biogenic emissions

The predicted changes (sensitivity scenario – base case) in average ground-level concentrations of total PM_{2.5} due to higher biogenic emissions (based on a 2 K temperature increase) are shown in Fig. 2. During the modeled summer period, PM_{2.5} is predicted to increase by 10 ng m^{−3} K^{−1} (0.13 % K^{−1}) on a domain average basis, with a maximum increase of 250 ng m^{−3} K^{−1} (2 % K^{−1}) in France (Fig. 2a). This is mainly due to an OA increase as higher biogenic emissions lead to increases in biogenic SOA concentrations which account for almost 90 % of the OA increase. The increased biogenic VOCs, on the other hand, result in reductions of OH in several areas. The reduced OH levels slow down the gas-phase formation of sulfate (through SO₂ oxidation), and also lead to decreases of ammonium nitrate. This negative effect of increased biogenic VOCs on OH levels and hence on inorganics was also noted by Zhang et al. (2008). However, the predicted decreases of inorganic PM_{2.5} components are less than the increases of total OA, thus the net impact is an increase of total PM_{2.5} levels.

Biogenic emissions also have a positive effect on total PM_{2.5} concentrations during the modeled winter and fall periods. PM_{2.5} is predicted to increase throughout the domain by 10 ng m^{−3} K^{−1} (0.1 % K^{−1}) and 20.3 ng m^{−3} K^{−1} (0.25 % K^{−1}), on average, during the winter and fall, respectively (Fig. 2b, c). The predicted increases during the winter period can reach up to 130 ng m^{−3} K^{−1} (1 % K^{−1}) while those during fall are even higher (up to 200 ng m^{−3} K^{−1} or 1.5 % K^{−1}). Increases in OA levels dominate the response of total PM_{2.5}, while inorganic PM_{2.5} is less sensitive to biogenic emissions during these seasons.

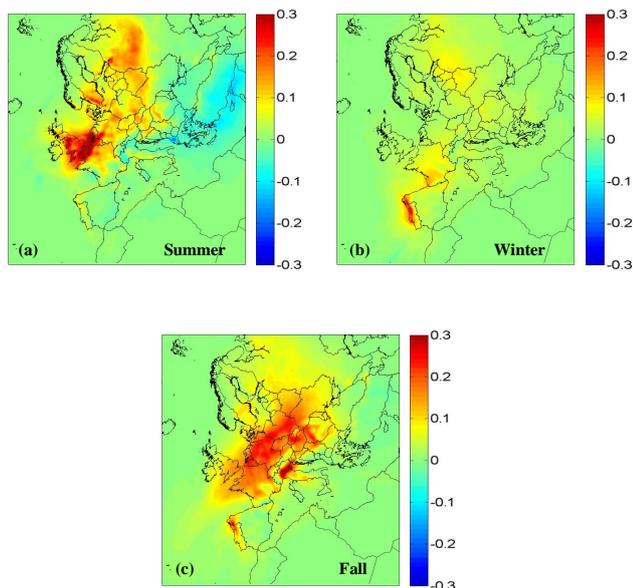


Figure 2. Predicted average change in ground-level concentrations ($\mu\text{g m}^{-3}$) of total PM_{2.5} due to changes on biogenic VOC emissions (based on a 2 K temperature increase) during the modeled (a) summer, (b) winter, and (c) fall periods. A positive value corresponds to an increase.

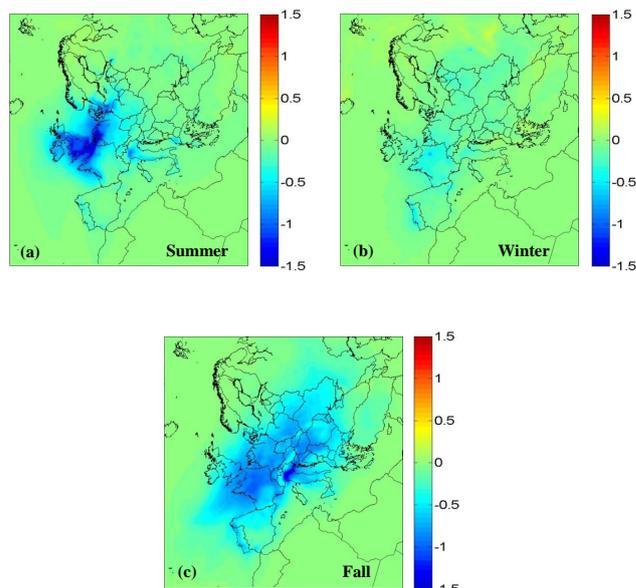


Figure 3. Predicted average change in ground-level concentrations ($\mu\text{g m}^{-3}$) of total PM_{2.5} due to changes on gas/aerosol partitioning (based on a 2 K temperature increase) during the modeled (a) summer, (b) winter, and (c) fall periods. A positive value corresponds to an increase.

5.2 Temperature effects on gas/aerosol partitioning

Increasing temperature by 2 K only for the partitioning of semi-volatile PM components has a significant effect on total PM_{2.5} levels in all three periods (Fig. 3). The predicted response of PM_{2.5} shows a strong spatial variability, as a result of competing changes in inorganic species concentrations and, to a lesser extent, in organic ones. In the modeled summer period, total PM_{2.5} concentrations decrease by $49 \text{ ng m}^{-3} \text{ K}^{-1}$ ($1 \% \text{ K}^{-1}$) on average, although the change is quite variable and ranges from -700 ng m^{-3} ($-5 \% \text{ K}^{-1}$) to 50 ng m^{-3} ($1.5 \% \text{ K}^{-1}$). The predicted PM_{2.5} decrease is largely due to significant decreases of ammonium nitrate. Rising temperature leads to increased volatilization of ammonium nitrate, which partitions to the gas phase (Seinfeld and Pandis, 2006). As a result, less ammonium nitrate exists in the particulate phase, leading to significant decreases of nitrate, which reach up to $600 \text{ ng m}^{-3} \text{ K}^{-1}$ ($14 \% \text{ K}^{-1}$). On the contrary, as particulate nitrate decreases, the cloud pH increases and the aqueous-phase formation of particulate sulfate accelerates. This complex effect of temperature changes on partitioning of semi-volatile inorganic PM_{2.5} is consistent with the results of other studies (e.g., Dawson et al., 2007; Aksoyoglu et al., 2011; Jimenez-Guerrero et al., 2012). OA is also sensitive to temperature, mainly due to changes in the levels of secondary OA components and, to a lesser extent, on primary OA. Higher temperature leads to evaporation of all OA components and subsequently to decreases of their levels. The sensitivity of OA to temperature,

as well as the increased gas-phase partitioning as temperature increases, have also been reported by earlier studies (Dawson et al., 2007; Megaritis et al., 2013).

During the modeled winter period, total PM_{2.5} also shows a negative response to temperature, with an average decrease of $25 \text{ ng m}^{-3} \text{ K}^{-1}$ ($0.4 \% \text{ K}^{-1}$) (Fig. 3b) over the domain. The predicted decrease of PM_{2.5} is significant in central Europe, due largely to decreases in nitrate and to a lesser extent in OA levels.

During the modeled fall period, total PM_{2.5} decreases by $88 \text{ ng m}^{-3} \text{ K}^{-1}$ ($1 \% \text{ K}^{-1}$) on average over the domain. Significant decreases are predicted mainly over the central and southwestern areas of the domain (Fig. 3c). Nitrate is significantly reduced (its predicted decreases exceed $10 \% \text{ K}^{-1}$), and along with total OA decreases, dominate the response of total PM_{2.5}, despite the predicted increases in sulfate levels.

5.3 Temperature-dependent gas-phase reaction rates

Changes in gas-phase reaction rates, due to temperature changes, could also affect total PM_{2.5} levels (Dawson et al., 2007). At higher temperatures, gas-phase reactions will accelerate (Dawson et al., 2007; Jacob and Winner, 2009; Day and Pandis, 2011; Im et al., 2011). In all three modeled periods, PM_{2.5} is predicted to increase due to the combined increases on the individual PM_{2.5} components. In the modeled summer period, PM_{2.5} concentrations are predicted to increase by $26 \text{ ng m}^{-3} \text{ K}^{-1}$ ($0.3 \% \text{ K}^{-1}$) on a domain average basis. The effect is stronger over continental Europe, where

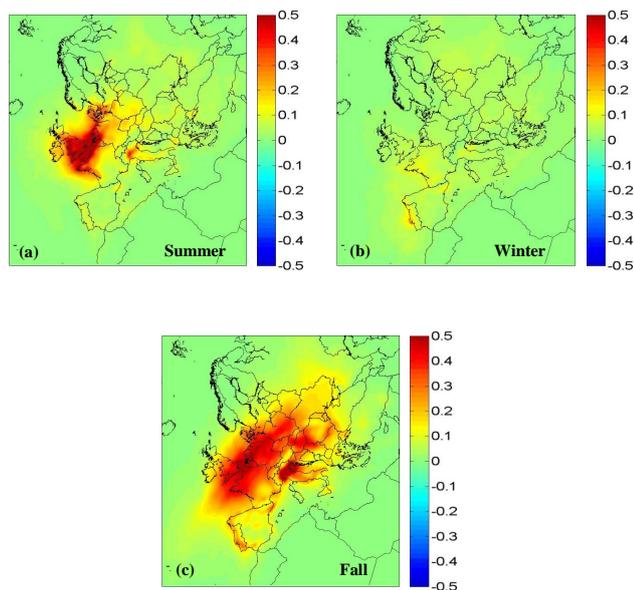


Figure 4. Predicted average change in ground-level concentrations ($\mu\text{g m}^{-3}$) of total PM_{2.5} due to changes on gas-phase reaction rates (based on a 2 K temperature increase) during the modeled (a) summer, (b) winter, and (c) fall periods. A positive value corresponds to an increase.

PM_{2.5} increases by $50 \text{ ng m}^{-3} \text{ K}^{-1}$ ($0.8 \% \text{ K}^{-1}$) on average, while in some areas in western Europe, increases in PM_{2.5} reach up to $400 \text{ ng m}^{-3} \text{ K}^{-1}$ ($2 \% \text{ K}^{-1}$) (Fig. 4a). The predicted response of total PM_{2.5} is mainly driven by increases of nitrate levels (approximately 45 % of the PM_{2.5} increase is due to nitrate), followed by increases in OA (largely attributed to secondary OA) and sulfate.

The lower oxidant availability during the winter leads to a lower increase of PM_{2.5} compared to summertime (Fig. 4b). Over continental Europe, the predicted increases are higher, up to $120 \text{ ng m}^{-3} \text{ K}^{-1}$ ($1 \% \text{ K}^{-1}$). Changes in organics and nitrate dominate (each of these two components accounts for around 40 % of the PM_{2.5} increase), while increases in sulfate tend to be rather small.

The effects are higher during the modeled fall period (an average increase of $47 \text{ ng m}^{-3} \text{ K}^{-1}$ or $0.6 \% \text{ K}^{-1}$ over the domain). The largest changes are in central and western Europe (Fig. 4c). Increases of fine particulate nitrate and organics are driving the PM_{2.5} response, while there are moderate increases in sulfate.

5.4 Overall temperature effects

An increase in temperature by 2 K is predicted to have a negative effect on average PM_{2.5} levels for all three modeled periods. On a domain average basis, PM_{2.5} decreases by $25 \text{ ng m}^{-3} \text{ K}^{-1}$ ($0.3 \% \text{ K}^{-1}$) in the summer, $7 \text{ ng m}^{-3} \text{ K}^{-1}$ ($0.1 \% \text{ K}^{-1}$) in the winter and $33 \text{ ng m}^{-3} \text{ K}^{-1}$ ($0.4 \% \text{ K}^{-1}$) in the modeled fall period. However, the overall effect

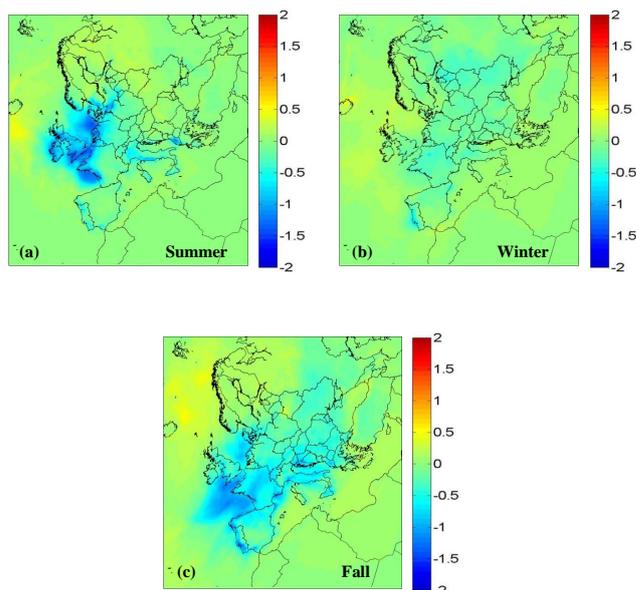


Figure 5. Predicted average change in ground-level concentrations ($\mu\text{g m}^{-3}$) of total PM_{2.5} due to an overall temperature increase by 2 K during the modeled (a) summer, (b) winter, and (c) fall periods. A positive value corresponds to an increase.

of temperature on PM_{2.5} levels is quite variable in space and time (Fig. 5) due to the different effects on the individual processes, as well as the competing responses of the PM_{2.5} species. The predicted changes on PM_{2.5} concentrations range from $-720 \text{ ng m}^{-3} \text{ K}^{-1}$ ($-8 \% \text{ K}^{-1}$) to $280 \text{ ng m}^{-3} \text{ K}^{-1}$ ($7 \% \text{ K}^{-1}$). Over continental Europe, PM_{2.5} changes are dominated by decreases in nitrate, which are mainly due to the evaporation of ammonium nitrate, leading to a reduction of average nitrate levels by 18 %. On the contrary, in several parts of the domain, the higher biogenic VOC emissions and the increased rate of SO₂ oxidation enhance the production of OA and sulfate, respectively. These increases can reach up to $225 \text{ ng m}^{-3} \text{ K}^{-1}$ ($7 \% \text{ K}^{-1}$) for sulfate and up to $190 \text{ ng m}^{-3} \text{ K}^{-1}$ ($4 \% \text{ K}^{-1}$) for OA. These results support the findings from previous studies that suggest the competing effects of temperature among the different processes and PM_{2.5} species (Dawson et al., 2007; Heald et al., 2008; Jacob and Winner, 2009; Jimenez-Guerrero et al., 2012). Summarizing, the semi-volatile PM_{2.5} evaporation appears to dominate and determine the overall PM_{2.5} response to temperature changes over Europe, during all seasons. The average changes in PM_{2.5} are higher during the fall.

6 Wind speed

Decreasing wind speed by 10 %, without any change on sea salt emissions (as well as on emissions from other sources), affects all PM_{2.5} components, resulting in increases of their

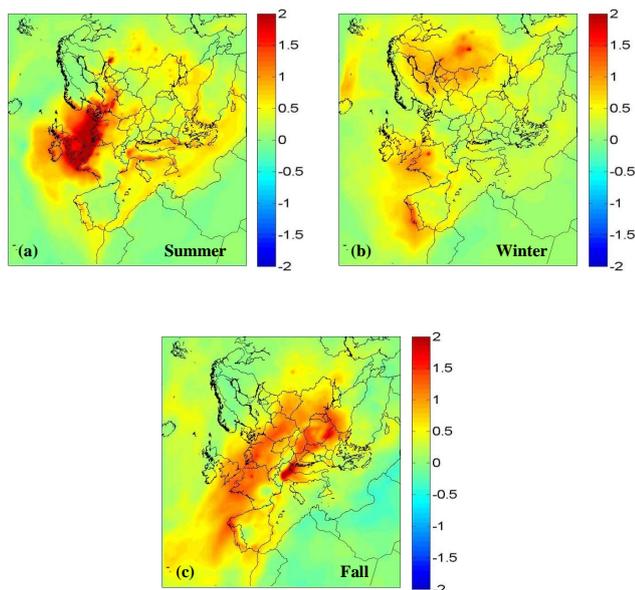


Figure 6. Predicted average change in ground-level concentrations ($\mu\text{g m}^{-3}$) of total PM_{2.5} due to a 10 % decrease in wind speed during the modeled (a) summer, (b) winter, and (c) fall periods. A positive value corresponds to an increase.

levels in all three modeled periods (Fig. 6). During summer, total PM_{2.5} is predicted to increase by $41 \text{ ng m}^{-3} \%^{-1}$ ($0.6 \% \%^{-1}$) on average over the entire domain (Fig. 6a). The effects of wind speed were found to be highest in the more polluted and populated areas of the domain. For example, in western Europe, a high ammonium nitrate area during summer, total PM_{2.5} increases up to $340 \text{ ng m}^{-3} \%^{-1}$ ($1.5 \% \%^{-1}$), driven mainly by increases of nitrate. Decreases in wind speed affect advection, dispersion and mixing as well as lead to changes in dry deposition. Approximately 7–13 % less PM_{2.5} is dry deposited due to the simulated 10 % reduction in wind speed. The effects of wind speed on total PM_{2.5} levels are similar during the other two periods. During winter, PM_{2.5} increases by $36 \text{ ng m}^{-3} \%^{-1}$ ($0.5 \% \%^{-1}$) on average over the domain. Significant increases are found mainly over northern Europe as well as in central and southwestern Europe (Fig. 6b), mainly due to increases of total OA and sulfate. In the modeled fall period PM_{2.5} shows a similar sensitivity. On a domain average basis, PM_{2.5} increases by $38 \text{ ng m}^{-3} \%^{-1}$ ($0.5 \% \%^{-1}$) with a maximum in central Europe (Fig. 6c). The predicted PM_{2.5} response is driven mainly by increases in particulate nitrate (it accounts for approximately 40 % of total PM_{2.5} increase) and to a lesser extent in ammonium, sulfate and organics.

Our results, regarding the PM_{2.5} response to wind speed, are consistent with those by Dawson et al. (2007), who found a PM_{2.5} sensitivity to wind equal to $0.77 \% \%^{-1}$ during summer and $0.56 \% \%^{-1}$ during winter in the eastern US. This negative effect of wind speed on PM_{2.5} has also been reported in earlier modeling studies over Europe (Carvalho

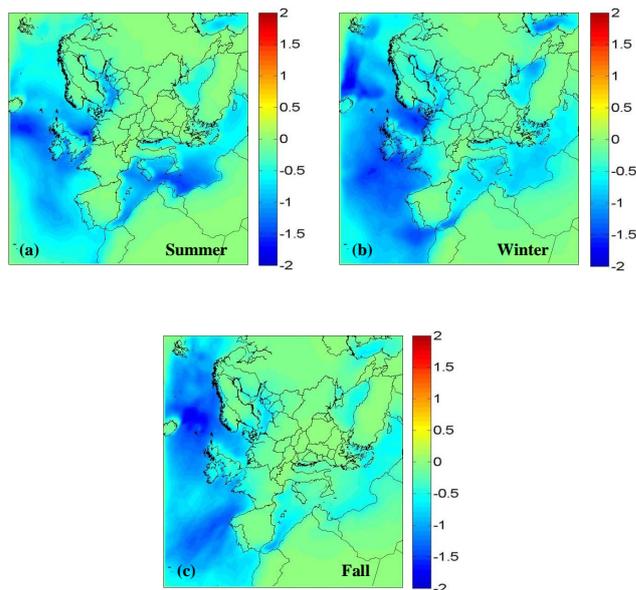


Figure 7. Predicted average change in ground-level concentrations ($\mu\text{g m}^{-3}$) of total PM_{2.5} due to changes on sea salt emissions (based on a 10 % decrease in wind speed) during the modeled (a) summer, (b) winter, and (c) fall period. A positive value corresponds to an increase.

et al., 2010; Aksoyoglu et al., 2011; Lecoer and Seigneur, 2013).

Wind effects on sea salt emissions

The predicted changes (sensitivity scenario – base case) in average ground-level concentrations of PM_{2.5} using a new sea salt emission inventory (based on a 10 % decrease of wind speed) are shown in Fig. 7. As expected, lower sea salt emissions result in lower PM_{2.5} concentrations in all modeled periods, especially over water and in coastal areas. The predicted PM_{2.5} response is, as expected, not uniform throughout the domain. During the modeled summer period, the predicted PM_{2.5} decrease exceeds $60 \text{ ng m}^{-3} \%^{-1}$ (or $0.5 \% \%^{-1}$), and may reach up to $170 \text{ ng m}^{-3} \%^{-1}$ ($0.9 \% \%^{-1}$), mainly due to decreases in particulate sodium and chloride. The predicted decreases are even larger (up to $200 \text{ ng m}^{-3} \%^{-1}$ or $2.7 \% \%^{-1}$) during the winter modeled period, as sea salt emissions and the accompanying concentrations of particulate sodium chloride were higher, while similar results have been obtained for the fall. Over continental Europe, the effects on PM_{2.5} levels due to lower marine aerosol emissions are small. PM_{2.5} is also reduced, however the predicted decrease does not exceed $20 \text{ ng m}^{-3} \%^{-1}$ ($0.1 \% \%^{-1}$) in all three periods.

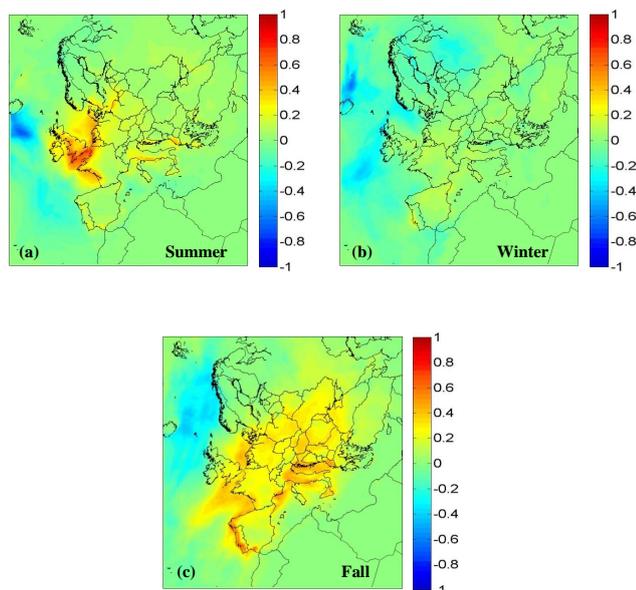


Figure 8. Predicted average change in ground-level concentrations ($\mu\text{g m}^{-3}$) of total PM_{2.5} due to a 5 % increase in absolute humidity during the modeled (a) summer, (b) winter, and (c) fall periods. A positive value corresponds to an increase.

7 Effects of absolute humidity

Changes in absolute humidity affect total PM_{2.5} concentrations, however its predicted response varies significantly in space (Fig. 8) due to the competing changes among PM_{2.5} species. In the modeled summer period, increases of absolute humidity by 5 % result in an average increase of total PM_{2.5} by $8 \text{ ng m}^{-3} \%^{-1}$ ($0.2 \% \%^{-1}$) over the entire domain. This is consistent with the Dawson et al. (2007) study for the eastern US, who reported a $20 \text{ ng m}^{-3} \%^{-1}$ increase in summer PM_{2.5} levels due to increases in absolute humidity by 5–20 %. The highest changes are predicted in western Europe (Fig. 8a) as a result of significant increases in nitrate. Increases in relative humidity shift the equilibrium of the ammonia–nitric acid system toward the particles (Seinfeld and Pandis, 2006). As absolute humidity increases by 5 %, approximately 15 % more HNO₃ is predicted to move to the aerosol phase, leading to higher particulate nitrate concentrations. These changes in nitrate, along with increases in ammonium and OA, are driving the PM_{2.5} response over land. On the contrary, over the ocean, total PM_{2.5} decreases as humidity increases, due mainly to changes in sulfate and sodium chloride. The negative response of PM_{2.5} in this area arises from increases in the size of the particles and accelerated dry deposition (in all modeled periods a 5 % increase in absolute humidity resulted in a 9–15 % increase in dry deposited mass of sulfate, sodium, and chloride). Absolute humidity also has a positive effect on PM_{2.5} levels during the modeled fall period. Significant increases are predicted in most areas of continental Europe (up to $130 \text{ ng m}^{-3} \%^{-1}$

or $1 \% \%^{-1}$) (Fig. 8c), mainly due to significant increases in particulate nitrate (approximately 65 % of the PM_{2.5} increase) while over the ocean, total PM_{2.5} decreases. The predicted increases of nitrate along with the increase in ammonium and total OA exceed the decreases in sulfate and sea salt, thus the net impact on total PM_{2.5} is an average increase of $11.5 \text{ ng m}^{-3} \%^{-1}$ ($0.2 \% \%^{-1}$).

In the modeled winter period, the predicted response of total PM_{2.5} to absolute humidity differs. In spite of the increase in nitrate concentrations, the predicted decreases in fine particulate sulfate and sea salt aerosol dominate and determine the response of total PM_{2.5} (Fig. 8b). On a domain average basis, the net effect of absolute humidity on PM_{2.5} is a decrease by $7.5 \text{ ng m}^{-3} \%^{-1}$ ($0.2 \% \%^{-1}$), while the predicted concentration changes range from $-130 \text{ ng m}^{-3} \%^{-1}$ ($-1.6 \% \%^{-1}$) to $44 \text{ ng m}^{-3} \%^{-1}$ ($0.5 \% \%^{-1}$).

8 Precipitation

8.1 Precipitation rate

The effect of the precipitation rate on PM_{2.5} concentrations is similar during all the modeled periods. The predicted response of average ground-level PM_{2.5} concentrations after a 10 % increase in precipitation rate (without changing the precipitation area) is shown in Fig. S1 in the Supplement. As it is expected, increases in precipitation rate accelerate the wet removal of PM_{2.5} species and their gas precursors and consequently result in decreases of their concentrations. In this simulation we predict a 2–4 % increase in PM_{2.5} wet deposited mass as well as a 5–12 % increase in the wet deposition of PM_{2.5} gas precursors due to a 10 % increase in precipitation rate.

During the modeled summer period, total PM_{2.5} is predicted to decrease as precipitation increases, by $13 \text{ ng m}^{-3} \%^{-1}$ ($0.2 \% \%^{-1}$) on average. Precipitation affects all the individual PM_{2.5} species, leading to reductions of their levels in most areas of the domain (Fig. S1a in the Supplement). Over the western parts of the domain, total PM_{2.5} is reduced up to $110 \text{ ng m}^{-3} \%^{-1}$ ($1.8 \% \%^{-1}$). However, even in areas with little rainfall during this period (e.g. eastern Mediterranean) (Fig. S2 in the Supplement), total PM_{2.5} also decreased, indicating that changes due to precipitation in upwind areas can affect the levels of PM_{2.5} over downwind areas. Similar effects are predicted during the other two periods. PM_{2.5} is also reduced as precipitation rate increases, having an average decrease of $0.2 \% \%^{-1}$ in both periods. The predicted effects are strongest in areas receiving moderate or little precipitation. This negative correlation has been also pointed out in earlier studies (Hedegaard et al., 2008; Jacob and Winner, 2009; Jimenez-Guerrero et al., 2012; Manders et al., 2012; Lecoeur and Seigneur, 2013). Dawson et al. (2007) predicted

quite similar sensitivities for total PM_{2.5} during summer (approximately 0.2 % %⁻¹).

8.2 Precipitation area

The predicted reduction of total PM_{2.5} for a 10 % increase in the spatial extent of precipitation covers a significant portion of Europe, during all periods (Fig. S3 in the Supplement). During summer the predicted reduction of PM_{2.5} reaches a maximum of 19 ng m⁻³ %⁻¹ (0.3 % %⁻¹), with an average sensitivity of 8 ng m⁻³ %⁻¹ (0.1 % %⁻¹). The predicted reductions arise mainly from the increases in PM_{2.5} wet deposited mass (approximately 2–5 %). The predicted effect is quite similar during the winter period, while in the modeled fall period the predicted response of total PM_{2.5} is a little higher, 13 ng m⁻³ %⁻¹ (0.16 % %⁻¹) on average. Our results support the conclusion that not only the precipitation intensity, but the area undergoing precipitation as well, can affect total PM_{2.5} concentrations (Lecoeur and Seigneur, 2013).

9 Mixing height

The predicted simulation-averaged changes in PM_{2.5} due to an increase in mixing height (by approximately 150 m) are shown in Fig. S4 in the Supplement. As expected, increases in mixing height affect all the individual PM_{2.5} components resulting in decreases in their concentrations during all modeled periods. In the summer, the average total PM_{2.5} concentrations decrease by 3.5 ng m⁻³ %⁻¹ (or 0.05 % %⁻¹). Similar effects on PM_{2.5} levels are also predicted for the other two periods. The effect of mixing height is strongest over polluted areas, where the predicted reduction of total PM_{2.5} can exceed 35 ng m⁻³ %⁻¹ (0.8 % %⁻¹) (over western Europe, during the modeled summer period). Our results are consistent with those by Dawson et al. (2007), who predicted a PM_{2.5} sensitivity to mixing height equal to 0.08 % %⁻¹ during summer and 0.05 % %⁻¹ during winter in the eastern US.

10 Relative importance of meteorological parameters

In order to evaluate the relative importance of the various meteorological parameters, we estimated the potential effects that each of them may have on total PM_{2.5} concentrations in a future climate. Our estimates were based on the predicted average PM_{2.5} sensitivities to the meteorological perturbations (Fig. 9) and the projected future changes for each parameter. Figure 9 summarizes the sensitivity distributions of average PM_{2.5} in the different locations of the domain to the various meteorological parameters. The projected meteorological changes are shown in Table S1 in the Supplement. According to the different IPCC (2013) scenarios, the average temperature in Europe is expected to increase over the next century from 1 to 5.5 K. Projections for wind speed and precipitation in Europe vary significantly in space. Based on the

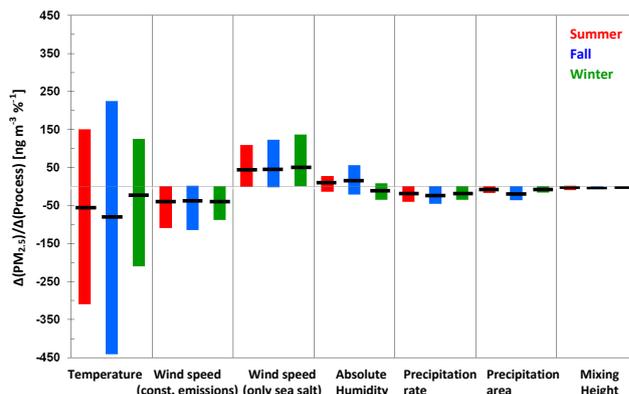


Figure 9. Predicted simulation-average sensitivities of total PM_{2.5} to changes in temperature, wind speed, sea salt emissions, absolute humidity, precipitation rate, precipitating area, and mixing height, during the three modeled periods. Each bar shows the range between the 10th and 90th percentiles. The black line in each bar shows the mean PM_{2.5} sensitivity over the domain.

IPCC SRES A2 scenario, wind speed is predicted to change from –10 to 10 %, while precipitation could change from –40 to 40 %. In our calculations, the changes in precipitation area and intensity were chosen to represent future projections for total precipitation over Europe. For mixing height, a potential range of changes was assumed, based on the estimates of Hedegaard et al. (2013). For this first-order estimate, we assumed the same meteorological changes for all seasons.

In all three periods, PM_{2.5} appears to be more sensitive to temperature changes compared to the rest meteorological parameters (Fig. 9). On average, PM_{2.5} shows a negative sensitivity to temperature changes, which is higher during fall compared to the other periods (Table S1 in the Supplement). However, the predicted PM_{2.5} sensitivities to temperature are spatially and temporally variable as a result of the different effects among the individual processes and the different responses of the PM_{2.5} species. During all seasons, the increased volatilization of ammonium nitrate dominates, causing large decreases in PM_{2.5} with increasing temperature. The negative predicted sensitivities reach up to 440 ng m⁻³ %⁻¹ in the fall and 310 ng m⁻³ %⁻¹ in the summer period (lower during winter) (Fig. 9). At the same time, the increasing temperatures lead to higher biogenic VOC emissions and accelerate the gas-phase chemical reactions. PM_{2.5} also shows a strong sensitivity to wind speed and its accompanying effects on the marine aerosol production. However, the predicted changes are somewhat lower compared to the PM_{2.5} sensitivities to temperature (Fig. 9). The sensitivity is similar in all seasons, and ranges from –115 ng m⁻³ %⁻¹ (due to changes in wind speed, without any change in the emissions) to 132 ng m⁻³ %⁻¹ (due to the effects of wind speed on sea salt emissions). PM_{2.5} appears to be less sensitive to absolute humidity changes. In all periods, PM_{2.5} concentrations respond differently to

absolute humidity, due to the competing effects between the individual PM_{2.5} species (e.g., increases in nitrate, decreases in sulfate), thus the average sensitivity does not exceed 12 ng m⁻³ %⁻¹ and the largest PM_{2.5} sensitivities are close to 55 ng m⁻³ %⁻¹. Changes in precipitation result in negative sensitivities for PM_{2.5} levels, which are comparable to those of absolute humidity, while mixing height seems to have a relatively small effect on average PM_{2.5} levels.

In a future climate, the projected changes in precipitation are expected to have the largest impact on PM_{2.5} levels during all periods (Fig. 10). These ranges were estimated simply by multiplying the expected changes with the calculated sensitivities and therefore include both the variability in space and the variability in the parameter change itself. PM_{2.5} concentrations could potentially change by several μg m⁻³ (up to approximately 2 μg m⁻³ during the fall period), with changes in precipitation intensity being rather more important than changes in precipitating area. Extrapolating from 10 % (Table 1) to 40 % (Table S1 in the Supplement) change in precipitation in order to calculate the expected concentration change adds some uncertainty due to possible non-linearities induced from a possible saturation of wet deposition to further increase of precipitation. Wind speed and absolute humidity may also lead to appreciable changes in future PM_{2.5} levels. The expected effects on PM_{2.5} due to changes in wind speed as well as its accompanying effects on the marine aerosol production are similar in all three periods and quite close to those resulting from future precipitation changes (up to 1.4 μg m⁻³). In addition, absolute humidity could potentially lead to large changes in PM_{2.5} mainly during the fall period (increases up to 2 μg m⁻³). The increased particulate nitrate levels, as higher absolute humidity favors its partitioning, are dominant during this period, causing large increases in PM_{2.5}. In the other two periods, the expected changes in PM_{2.5} are smaller due to the competing responses among the individual PM_{2.5} species. Temperature is expected to have a lower impact on future PM_{2.5} compared to the rest of the meteorological parameters, in all seasons. The expected PM_{2.5} concentration changes range from -1.1 to 0.5 μg m⁻³, driven from the offsetting effects of increased nitrate volatilization, higher biogenic VOC emissions and accelerated gas-phase chemistry. Mixing height is expected to have a relatively small impact on PM_{2.5} levels in a future climate.

11 Conclusions

Climate affects air quality through a complex web of interactions, starting with changes in the major meteorological variables like temperature, wind speed, absolute humidity, precipitation intensity, precipitation area, mixing height, etc. and progressing through changes in pollutant concentrations, formation and removal rates. In this study, we used a detailed three-dimensional CTM, PMCAMx-2008, to quantify the in-

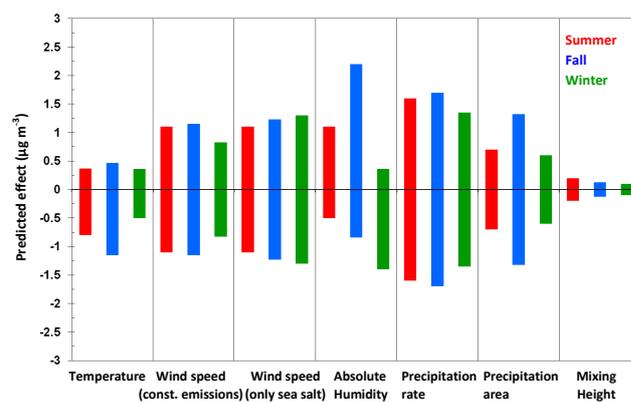


Figure 10. Expected PM_{2.5} concentration changes due to projected changes in temperature, wind speed, sea salt emissions, absolute humidity, precipitation rate, precipitating area, and mixing height in the future, during the three modeled periods.

dividual effects of the major meteorological parameters on the concentration and composition of PM_{2.5} over Europe.

Precipitation is expected to have the largest impact on PM_{2.5} concentrations under a changed future climate. In all periods, PM_{2.5} shows a negative sensitivity to precipitation, driven mainly by the accelerated wet deposition of PM_{2.5} species and their gas precursors. The average PM_{2.5} sensitivity is quite similar during all seasons (an approximate decrease of 15 ng m⁻³ %⁻¹) and also taking account the significant projected precipitation changes, PM_{2.5} concentrations could potentially change by several μg m⁻³ (up to 2 μg m⁻³ in the fall) in the future.

Wind speed can also have appreciable effects on future PM_{2.5} levels due to changes in dispersion and transport, dry deposition and marine aerosol production. The projected changes in wind speed over Europe in the future are expected to change PM_{2.5} levels up to 1.4 μg m⁻³.

Changes in absolute humidity influence mainly the inorganic PM_{2.5} species, resulting in competing responses. An increase in absolute humidity favors the partitioning of nitrate to the aerosol phase and leads to higher particulate levels. During the fall period, this effect dominates the overall PM_{2.5} response, and as absolute humidity is expected to rise in the future, it could lead to large increases of PM_{2.5} (up to 2 μg m⁻³). On the contrary, the increase in absolute humidity could lead to decreases in sulfate, and sea salt levels due to the increase in the size of the particles and the accelerated dry deposition. These negative effects may, to some extent, offset the predicted increases in nitrate, thus during summer and winter the expected changes in future PM_{2.5} due to absolute humidity are smaller.

Temperature is expected to have a lower average impact on future PM_{2.5} levels compared to the rest of the meteorological parameters due to the competing effects among the individual processes and the different responses of the PM_{2.5} species. The evaporation of semi-volatile PM_{2.5} species is

found to be the dominant process and determines to a large extent the PM_{2.5} response to temperature changes over Europe, during all seasons. Significant effects are predicted mainly on particulate ammonium nitrate, as the increase in temperature reduces its concentration levels up to 15 % K⁻¹. Especially during fall, the predicted reduction of nitrate drives the overall PM_{2.5} response, and as temperature is expected to rise in a future climate, could potentially lead to decreases in PM_{2.5} levels up to 1.1 μg m⁻³. However, as temperature increases, biogenic VOC emissions are expected to increase and gas-phase chemical reactions will accelerate, which will offset to some extent the reductions of PM_{2.5}, leading to even smaller changes in future PM_{2.5} levels during the summer and winter period. PM_{2.5} concentrations generally decrease as mixing height increases. However the predicted effects are not as significant as those of the other parameters for the average PM_{2.5} levels, due to the importance of secondary PM_{2.5} components that have a strong regional character.

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References

- Aksoyoglu, S., Keller, J., Barmpadimos, I., Oderbolz, D., Lanz, V. A., Prévôt, A. S. H., and Baltensperger, U.: Aerosol modelling in Europe with a focus on Switzerland during summer and winter episodes, *Atmos. Chem. Phys.*, 11, 7355–7373, doi:10.5194/acp-11-7355-2011, 2011.
- Andersson, C. and Engardt, M.: European ozone in a future climate: Importance of changes in dry deposition and isoprene emissions, *J. Geophys. Res.*, 115, D02303, doi:10.1029/2008JD011690, 2010.
- Argüeso, D., Hidalgo-Munoz, J. M., Gamiz-Fortiz, S. R., Esteban-Parra, M. J., and Castro-Diez, Y.: High-resolution projections of mean and extreme precipitation over Spain using the WRF model (2070–2099 versus 1970–1999), *J. Geophys. Res.*, 117, D12108, doi:10.1029/2011JD017399, 2012.
- Avise, J., Chen, J., Lamb, B., Wiedinmyer, C., Guenther, A., Salathé, E., and Mass, C.: Attribution of projected changes in summertime US ozone and PM_{2.5} concentrations to global changes, *Atmos. Chem. Phys.*, 9, 1111–1124, doi:10.5194/acp-9-1111-2009, 2009.
- Aw, J. and Kleeman, M. J.: Evaluating the first-order effect of inter-annual temperature variability on urban air pollution, *J. Geophys. Res.*, 108, 4365, doi:10.1029/2002JD002688, 2003.
- Bernard, S. M., Samet, J. M., Grambsch, A., Ebi, K. L., and Romieu, I.: The potential impact of climate variability and change on air pollution-related health effects in the United States, *Environ. Health Perspect.*, 109, Suppl. 2, 199–209, 2001.
- Boe, J., Terray, L., Cassou, C., and Najac, J.: Uncertainties in European summer precipitation changes: role of large scale circulation, *Clim. Dynam.*, 33, 265–276, 2009.
- Buonomo, E., Jones, R., Huntingford, C., and Hannaford, J.: On the robustness of changes in extreme precipitation over Europe from two high resolution climate change simulations, *Q. J. Roy. Meteor. Soc.*, 133, 65–81, 2007.
- Burtraw, D., Krupnick, A., Mansur, E., Austin, D., and Farrell, D.: Costs and benefits of reducing air pollutants related to acid rain, *Contemp. Econ. Policy*, 16, 379–400, doi:10.1111/j.1465-7287.1998.tb00527.x, 2007.
- Carter, W. P. L.: Programs and Files Implementing the SAPRC-99 Mechanism and its Associates Emissions Processing Procedures for Models-3 and Other Regional Models: <http://www.cert.ucr.edu/~carter/SAPRC99/> (last access: 15 April 2014), 2010.
- Carvalho, A., Monteiro, A., Solman, S., Miranda, A. I., and Borgego, C.: Climate-driven changes in air quality over Europe by the end of the 21st century, with special reference to Portugal, *Environ. Sci. Policy*, 13, 445–458, 2010.
- Christensen, O. B. and Christensen, J. H.: Intensification of extreme European summer precipitation in a warmer climate, *Global Planet. Change*, 44, 107–117, 2004.
- Dawson, J. P., Adams, P. J., and Pandis, S. N.: Sensitivity of PM_{2.5} to climate in the Eastern US: a modeling case study, *Atmos. Chem. Phys.*, 7, 4295–4309, doi:10.5194/acp-7-4295-2007, 2007.
- Dawson, J. P., Racherla, P. N., Lynn, B. H., Adams, P. J., and Pandis, S. N.: Impacts of climate change on regional and urban air quality in the eastern United States: Role of meteorology, *J. Geophys. Res.*, 114, D05308, doi:10.1029/2008JD009849, 2009.
- Day, M. C. and Pandis S. N.: Predicted changes in summertime organic aerosol concentrations due to increased temperature, *Atmos. Environ.*, 45, 6546–6556, 2011.
- Denier van der Gon, H. A. C., Visschedijk, A., van der Brugh, H., and Droge, R.: A high resolution European emission data base for the year 2005, TNO report TNO- 34-UT-2010-01895 RPTML, Netherlands Organisation for Applied Scientific Research TNO, Utrecht, the Netherlands, 2010.
- Donahue, N. M., Robinson, A. L., Stanier, C. O., and Pandis, S. N.: Coupled partitioning, dilution, and chemical aging of semivolatile organics, *Environ. Sci. Technol.*, 40, 2635–2643, 2006.
- EMEP: Transboundary acidification, eutrophication and ground level ozone in Europe in 2008, EMEP status report, Oslo, Norway, 2010.
- ENVIRON: User's Guide to the Comprehensive Air Quality Model with Extensions (CAMx), Version 4.02, Report, ENVIRON Int. Corp., Novato, Calif, available at: <http://www.camx.com> (last access: 15 April 2014), 2003.
- Fahey, K. and Pandis, S. N.: Optimizing model performance: variable size resolution in cloud chemistry modeling, *Atmos. Environ.*, 35, 4471–4478, 2001.
- Forkel, R., and Knoche, R.: Nested regional climate-chemistry simulations for central Europe, *C. R. Geosci.*, 339, 734–746, 2007.

- Fountoukis, C., Racherla, P. N., Denier van der Gon, H. A. C., Polymeneas, P., Charalampidis, P. E., Pilinis, C., Wiedensohler, A., Dall'Osto, M., O'Dowd, C., and Pandis, S. N.: Evaluation of a three-dimensional chemical transport model (PMCAMx) in the European domain during the EUCAARI May 2008 campaign, *Atmos. Chem. Phys.*, 11, 10331–10347, doi:10.5194/acp-11-10331-2011, 2011.
- Fountoukis, C., Megaritis, A. G., Skyllakou, K., Charalampidis, P. E., Pilinis, C., Denier van der Gon, H. A. C., Crippa, M., Canonaco, F., Mohr, C., Prévôt, A. S. H., Allan, J. D., Poulain, L., Petäjä, T., Tiitta, P., Carbone, S., Kiendler-Scharr, A., Nemitz, E., O'Dowd, C., Swietlicki, E., and Pandis, S. N.: Organic aerosol concentration and composition over Europe: insights from comparison of regional model predictions with aerosol mass spectrometer factor analysis, *Atmos. Chem. Phys. Discuss.*, 14, 7597–7635, doi:10.5194/acpd-14-7597-2014, 2014.
- Frei, C., Scholl, R., Fukutome, S., Schmidli, J., and Vidale, P. L.: Future change of precipitation extremes in Europe: Intercomparison of scenarios from regional climate models, *J. Geophys. Res.*, 111, D06105, doi:10.1029/2005JD005965, 2006.
- Freney, E. J., Sellegri, K., Canonaco, F., Boulon, J., Hervo, M., Weigel, R., Pichon, J. M., Colomb, A., Prévôt, A. S. H., and Laj, P.: Seasonal variations in aerosol particle composition at the puy-de-Dôme research station in France, *Atmos. Chem. Phys.*, 11, 13047–13059, doi:10.5194/acp-11-13047-2011, 2011.
- Galindo, N., Varea, M., Gil-Molto, J., Yubero, E., and Nicolas, J.: The influence of meteorology on particulate matter concentrations at an urban Mediterranean location, *Water Air Soil Pollut.*, 215, 365–372, 2011.
- Gaydos, T., Koo, B., and Pandis, S. N.: Development and application of an efficient moving sectional approach for the solution of the atmospheric aerosol condensation/evaporation equations, *Atmos. Environ.*, 37, 3303–3316, 2003.
- Giorgi, F. and Meleux, F.: Modeling the regional effects of climate change on air quality, *C. R. Geosci.*, 339, 721–733, 2007.
- Guenther, A., Karl, T., Harley, P., Wiedinmyer, C., Palmer, P. I., and Geron, C.: Estimates of global terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols from Nature), *Atmos. Chem. Phys.*, 6, 3181–3210, doi:10.5194/acp-6-3181-2006, 2006.
- Hamburger, T., McMeeking, G., Minikin, A., Birmili, W., Dall'Osto, M., O'Dowd, C., Flentje, H., Henzing, B., Junninen, H., Kristensson, A., de Leeuw, G., Stohl, A., Burkhardt, J. F., Coe, H., Krejci, R., and Petzold, A.: Overview of the synoptic and pollution situation over Europe during the EUCAARI-LONGREX field campaign, *Atmos. Chem. Phys.*, 11, 1065–1082, doi:10.5194/acp-11-1065-2011, 2011.
- Heald, C. L., Henze, D. K., Horowitz, L. W., Feddema, J., Lamarque, J.-F., Guenther, A., Hess, P. G., Vitt, F., Seinfeld, J. H., Goldstein, A. H., and Fung, I.: Predicted change in global secondary aerosol concentrations in response to future climate, emissions and land use change, *J. Geophys. Res.*, 113, D05211, doi:10.1029/2007JD009092, 2008.
- Hedegaard, G. B., Brandt, J., Christensen, J. H., Frohn, L. M., Geels, C., Hansen, K. M., and Stendel, M.: Impacts of climate change on air pollution levels in the Northern Hemisphere with special focus on Europe and the Arctic, *Atmos. Chem. Phys.*, 8, 3337–3367, doi:10.5194/acp-8-3337-2008, 2008.
- Hedegaard, G. B., Christensen, J. H., and Brandt, J.: The relative importance of impacts from climate change vs. emissions change on air pollution levels in the 21st century, *Atmos. Chem. Phys.*, 13, 3569–3585, doi:10.5194/acp-13-3569-2013, 2013.
- Hildebrandt, L., Kostenidou, E., Mihalopoulos, N., Worsnop, D. R., Donahue, N. M., and Pandis, S. N.: Formation of highly oxygenated organic aerosol in the atmosphere: Insights from the Finokalia Aerosol Measurement Experiments, *Geophys. Res. Lett.*, 37, L23801, doi:10.1029/2010GL045193, 2010.
- Hogrefe, C., Lynn, B., Civerolo, K., Ku, J.-Y., Rosenthal, J., Rosenzweig, C., Goldberg, R., Gaffin, S., Knowlton, K., and Kinney, P. L.: Simulating changes in regional air pollution over the eastern United States due to changes in global and regional climate and emissions, *J. Geophys. Res.*, 109, D22301, doi:10.1029/2004JD004690, 2004.
- Im, U., Markakis, K., Poupkou, A., Melas, D., Unal, A., Gerasopoulos, E., Daskalakis, N., Kindap, T., and Kanakidou, M.: The impact of temperature changes on summer time ozone and its precursors in the Eastern Mediterranean, *Atmos. Chem. Phys.*, 11, 3847–3864, doi:10.5194/acp-11-3847-2011, 2011.
- Im, U., Markakis, K., Kocak, M., Gerasopoulos, E., Daskalakis, N., Mihalopoulos, N., Poupkou, A., Kindap, T., Unal, A., and Kanakidou, M.: Summertime aerosol chemical composition in the Eastern Mediterranean and its sensitivity to temperature, *Atmos. Environ.*, 50, 164–173, 2012.
- Intergovernmental Panel on Climate Change (IPCC): Fourth Assessment Report: Summary for Policymakers, Geneva, Switzerland, 2007.
- Intergovernmental Panel on Climate Change (IPCC): Summary for Policymakers, in: *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, edited by: Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M., Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013.
- Jacob, D. J. and Winner, D. A.: Effect of climate change on air quality, *Atmos. Environ.*, 43, 51–63, 2009.
- Jeong, J. I. and Park, R. J.: Effects of the meteorological variability on regional air quality in East Asia, *Atmos. Environ.*, 69, 46–55, 2013.
- Jimenez-Guerrero, P., Gomez-Navarro, J. J., Jerez, S., Lorente-Plazas, R., Garcia-Valero, J. A., and Montavez, J. P.: Isolating the effects of climate change in the variation of secondary inorganic aerosols (SIA) in Europe for the 21st century (1991–2100), *Atmos. Environ.*, 45, 1059–1063, 2011.
- Jimenez-Guerrero, P., Montavez, J. P., Gomez-Navarro, J. J., Jerez, S., and Lorente-Plazas, R.: Impacts of climate change on ground level gas-phase pollutants and aerosols in the Iberian Peninsula for the late XXI century, *Atmos. Environ.*, 55, 483–495, 2012.
- Karydis, V. A., Tsimpidi, A. P., Fountoukis, C., Nenes, A., Zavala, M., Lei, W., Molina, L. T., and Pandis, S. N.: Simulating the fine and coarse inorganic particulate matter concentrations in a polluted megacity, *Atmos. Environ.*, 44, 608–620, 2010.
- Katragkou, E., Zanis, P., Kioutsioukis, I., Tegoulas, I., Melas, D., Kruger, B. C., and Coppola, E.: Future climate change impacts on summer surface ozone from regional climate-air quality simulations over Europe, *J. Geophys. Res.*, 116, D22307, doi:10.1029/2011JD015899, 2011.

- Kjellström, E., Nikulin, G., Hansson, U., Strandberg, G., and Ullerstig, A.: 21st century changes in the European climate: uncertainties derived from an ensemble of regional climate model simulations, *Tellus A*, 63, 24–40, 2010.
- Koch, D., Park, J., and Del Genio, A.: Clouds and sulfate are anti-correlated: a new diagnostic for global sulfur models, *J. Geophys. Res.-Atmos.*, 108, 4781, doi:10.1029/2003JD003621, 2003.
- Koo, B., Pandis, S. N., and Ansari, A.: Integrated approaches to modeling the organic and inorganic atmospheric aerosol components, *Atmos. Environ.*, 37, 4757–4768, 2003.
- Kulmala, M., Asmi, A., Lappalainen, H. K., Baltensperger, U., Brenguier, J.-L., Facchini, M. C., Hansson, H.-C., Hov, Ø., O’Dowd, C. D., Pöschl, U., Wiedensohler, A., Boers, R., Boucher, O., de Leeuw, G., Denier van der Gon, H. A. C., Feichter, J., Krejci, R., Laj, P., Lihavainen, H., Lohmann, U., McFiggans, G., Mentel, T., Pilinis, C., Riipinen, I., Schulz, M., Stohl, A., Swietlicki, E., Vignati, E., Alves, C., Amann, M., Ammann, M., Arabas, S., Artaxo, P., Baars, H., Beddows, D. C. S., Bergström, R., Beukes, J. P., Bilde, M., Burkhardt, J. F., Canonaco, F., Clegg, S. L., Coe, H., Crumeyrolle, S., D’Anna, B., Decesari, S., Gilardoni, S., Fischer, M., Fjaeraa, A. M., Fountoukis, C., George, C., Gomes, L., Halloran, P., Hamburger, T., Harrison, R. M., Herrmann, H., Hoffmann, T., Hoose, C., Hu, M., Hyvärinen, A., Hörrak, U., Iinuma, Y., Iversen, T., Josipovic, M., Kanakidou, M., Kiendler-Scharr, A., Kirkevåg, A., Kiss, G., Klimont, Z., Kolmonen, P., Komppula, M., Kristjánsson, J.-E., Laakso, L., Laaksonen, A., Labonnote, L., Lanz, V. A., Lehtinen, K. E. J., Rizzo, L. V., Makkonen, R., Manninen, H. E., McMeeking, G., Merikanto, J., Minikin, A., Mirme, S., Morgan, W. T., Nemitz, E., O’Donnell, D., Panwar, T. S., Pawlowska, H., Petzold, A., Pienaar, J. J., Pio, C., Plass-Duelmer, C., Prévôt, A. S. H., Pryor, S., Reddington, C. L., Roberts, G., Rosenfeld, D., Schwarz, J., Seland, Ø., Sellegri, K., Shen, X. J., Shiraiwa, M., Siebert, H., Sierau, B., Simpson, D., Sun, J. Y., Topping, D., Tunved, P., Vaattovaara, P., Vakkari, V., Veefkind, J. P., Visschedijk, A., Vuollekoski, H., Vuolo, R., Wehner, B., Wildt, J., Woodward, S., Worsnop, D. R., van Zadelhoff, G.-J., Zardini, A. A., Zhang, K., van Zyl, P. G., Kerminen, V.-M., S Carslaw, K., and Pandis, S. N.: General overview: European Integrated project on Aerosol Cloud Climate and Air Quality Interactions (EUCAARI) – integrating aerosol research from nano to global scales, *Atmos. Chem. Phys.*, 11, 13061–13143, doi:10.5194/acp-11-13061-2011, 2011.
- Kumar, N., Lurmann, F. W., Wexler, A. S., Pandis, S., and Seinfeld, J. H.: Development and application of a three dimensional aerosol model. Presented at the AWMA Specialty Conference on Computing in Environmental Resource Management, Research Triangle Park, North Carolina, USA, 2–4 December, 1996.
- Lecœur, È. and Seigneur, C.: Dynamic evaluation of a multi-year model simulation of particulate matter concentrations over Europe, *Atmos. Chem. Phys.*, 13, 4319–4337, doi:10.5194/acp-13-4319-2013, 2013.
- Mahmud, A., Hixson, M., Hu, J., Zhao, Z., Chen, S.-H., and Kleeman, M. J.: Climate impact on airborne particulate matter concentrations in California using seven year analysis periods, *Atmos. Chem. Phys.*, 10, 11097–11114, doi:10.5194/acp-10-11097-2010, 2010.
- Manders, A. M. M., van Meijgaard, E., Mues, A. C., Kranenburg, R., van Ulft, L. H., and Schaap, M.: The impact of differences in large-scale circulation output from climate models on the regional modeling of ozone and PM, *Atmos. Chem. Phys.*, 12, 9441–9458, doi:10.5194/acp-12-9441-2012, 2012.
- Megaritis, A. G., Fountoukis, C., Charalampidis, P. E., Pilinis, C., and Pandis, S. N.: Response of fine particulate matter concentrations to changes of emissions and temperature in Europe, *Atmos. Chem. Phys.*, 13, 3423–3443, doi:10.5194/acp-13-3423-2013, 2013.
- Mensah, A. A., Holzinger, R., Otjes, R., Trimborn, A., Mentel, Th. F., ten Brink, H., Henzing, B., and Kiendler-Scharr, A.: Aerosol chemical composition at Cabauw, The Netherlands as observed in two intensive periods in May 2008 and March 2009, *Atmos. Chem. Phys.*, 12, 4723–4742, doi:10.5194/acp-12-4723-2012, 2012.
- Murphy, B. N. and Pandis, S. N.: Simulating the formation of semivolatile primary and secondary organic aerosol in a regional chemical transport model, *Environ. Sci. Technol.*, 43, 4722–4728, 2009.
- O’Dowd, C. D., Langmann, B., Varghese, S., Scannell, C., Ceburnis, D., and Facchini, M. C.: A combined organic-inorganic sea-spray source function, *Geophys. Res. Lett.*, 35, L01801, doi:10.1029/2007GL030331, 2008.
- Pay, M. T., Jimenez-Guerrero, P., and Baldasano, J. M.: Assessing sensitivity regimes of secondary inorganic aerosol formation in Europe with the CALIOPE-EU modeling system, *Atmos. Environ.*, 51, 146–164, 2012.
- Pikridas, M., Bougiatioti, A., Hildebrandt, L., Engelhart, G. J., Kostenidou, E., Mohr, C., Prévôt, A. S. H., Kouvarakis, G., Zarmas, P., Burkhardt, J. F., Lee, B.-H., Psichoudaki, M., Michalopoulos, N., Pilinis, C., Stohl, A., Baltensperger, U., Kulmala, M., and Pandis, S. N.: The Finokalia Aerosol Measurement Experiment – 2008 (FAME-08): an overview, *Atmos. Chem. Phys.*, 10, 6793–6806, doi:10.5194/acp-10-6793-2010, 2010.
- Pope, C. A. I., Ezzati, M., and Dockery, D. W.: Fine-particulate air pollution and life expectancy in the United States, *New Engl. J. Med.*, 360, 376–386, 2009.
- Poulain, L., Spindler, G., Birmili, W., Plass-Dülmer, C., Wiedensohler, A., and Herrmann, H.: Seasonal and diurnal variations of particulate nitrate and organic matter at the IfT research station Melpitz, *Atmos. Chem. Phys.*, 11, 12579–12599, doi:10.5194/acp-11-12579-2011, 2011.
- Pye, H. O. T., Liao, H., Wu, S., Mickley, L. J., Jacob, D. J., Henze, D. K., and Seinfeld, J. H.: Effect of changes in climate and emissions on future sulfate-nitrate-ammonium aerosol levels in the United States, *J. Geophys. Res.*, 114, D01205, doi:10.1029/2008JD010701, 2009.
- Racherla, P. N. and Adams, P. J.: Sensitivity of global tropospheric ozone and fine particulate matter concentrations to climate change, *J. Geophys. Res.*, 111, D24103, doi:10.1029/2005JD006939, 2006.
- Räisänen, J., Hansson, U., Ullerstig, A., Döschner, R., Graham, L. P., Jones, C., Meier, H. E. M., Samuelsson, P., and Willen, U.: European climate in the late twenty-first century: regional simulations with two driving global models and two forcing scenarios, *Clim. Dynam.*, 22, 13–31, 2004.
- Redington, A. L., Derwent, R. G., Witham, C. S., and Manning, A. J.: Sensitivity of modeled sulphate and nitrate aerosol to cloud, pH, and ammonia emissions, *Atmos. Environ.*, 43, 3227–3234, 2009.

- Roustan, Y., Sartelet, K. N., Tombette, M., Debry, E., and Sportisse, B.: Simulation of aerosols and gas-phase species over Europe with the POLYPHEMUS system. Part II: Model sensitivity analysis for 2001, *Atmos. Environ.*, 44, 4219–4229, 2010.
- Schwartz, J., Dockery, D. W., and Neas, L. M.: Is daily mortality associated specifically with fine particles?, *J. Air Waste Manage. Assoc.*, 46, 927–939, 1996.
- Seinfeld, J. H. and Pandis, S. N.: *Atmospheric Chemistry and Physics: From Air Pollution to Climate Change*. 2nd ed.; John Wiley and Sons, Hoboken, NJ, 2006.
- Sheehan, P. E. and Bowman, F. M.: Estimated effects of temperature on secondary organic aerosol concentrations, *Environ. Sci. Technol.*, 35, 2129–2135, 2001.
- Singh, A. and Palazoglu, A.: Climatic variability and its influence on ozone and PM pollution in 6 non-attainment regions in the United States, *Atmos. Environ.*, 51, 212–224, 2012.
- Skamarock, W. C., Klemp, J. B., Dudhia, J., Gill, D. O., Barker, D. M., Wang, W., and Powers, J. G.: A Description of the Advanced Research WRF Version 3. NCAR Technical Note, available at: http://www.mmm.ucar.edu/wrf/users/docs/arw_v3.pdf (last access: 15 April 2014), 2008.
- Slinn, S. A. and Slinn, W. G. N.: Predictions for particle deposition on natural waters, *Atmos. Environ.*, 24, 1013–1016, 1980.
- Sofiev, M., Vankevich, R., Lotjonen, M., Prank, M., Petukhov, V., Ermakova, T., Koskinen, J., and Kukkonen, J.: An operational system for the assimilation of the satellite information on wildland fires for the needs of air quality modelling and forecasting, *Atmos. Chem. Phys.*, 9, 6833–6847, doi:10.5194/acp-9-6833-2009, 2009.
- Stanier, C. O., Donahue, N. M., and Pandis, S. N.: Parameterization of secondary organic aerosol mass fraction from smog chamber data, *Atmos. Environ.*, 42, 2276–2299, 2008.
- Tagaris, E., Manomaiphiboon, K., Liao, K. J., Leung, L. R., Woo, J. H., He, S., Amar, P., and Russell, A. G.: Impacts of global climate change and emissions on regional ozone and fine particulate matter concentrations over the United States, *J. Geophys. Res.*, 112, D14312, doi:10.1029/2006JD008262, 2007.
- Tagaris, E., Liao, K.-J., Manomaiphiboon, K., He, S., Woo, J.-H., Amar, P., and Russell, A. G.: The role of climate and emission changes in future air quality over southern Canada and northern Mexico, *Atmos. Chem. Phys.*, 8, 3973–3983, doi:10.5194/acp-8-3973-2008, 2008.
- Tai, A. P. K., Mickley, L. J., and Jacob, D. J.: Impact of 2000–2050 climate change on fine particulate matter (PM_{2.5}) air quality inferred from a multi-model analysis of meteorological modes, *Atmos. Chem. Phys.*, 12, 11329–11337, doi:10.5194/acp-12-11329-2012, 2012.
- Trump, E. R., Fountoukis, C., Donahue, N. M., and Pandis, S. N.: Improvement of simulation of fine inorganic PM levels through better descriptions of coarse particle chemistry, *Atmos. Environ.* in review, 2014.
- Tsigaridis, K. and Kanakidou, M.: Secondary organic aerosol importance in the future atmosphere, *Atmos. Environ.*, 41, 4682–4692, 2007.
- Tsimpidi, A. P., Karydis, V. A., Zavala, M., Lei, W., Molina, L., Ulbrich, I. M., Jimenez, J. L., and Pandis, S. N.: Evaluation of the volatility basis-set approach for the simulation of organic aerosol formation in the Mexico City metropolitan area, *Atmos. Chem. Phys.*, 10, 525–546, doi:10.5194/acp-10-525-2010, 2010.
- Visschedijk, A. J. H., Zandveld, P., and Denier van der Gon, H. A. C.: TNO Report 2007 A-R0233/B: A high resolution gridded European emission database for the EU integrated project GEMS, Organization for Applied Scientific Research, Apeldoorn, the Netherlands, 2007.
- Wesely, M. L.: Parameterization of Surface Resistances to Gaseous Dry Deposition in Regional-Scale Numerical Models, *Atmos. Environ.*, 23, 1293–1304, 1989.
- Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I., Alfarra, M. R., Takami, A., Middlebrook, A. M., Sun, Y. L., Dzepina, K., Dunlea, E., Docherty, K., De-Carlo, P., Salcedo, D., Onasch, T. B., Jayne, J. T., Miyoshi, T., Shimojo, A., Hatakeyama, N., Takegawa, N., Kondo, Y., Schneider, J., Drewnick, F., Weimer, S., Demerjian, K. L., Williams, P. I., Bower, K. N., Bahreini, R., Cottrell, L., Griffin, R. J., Rautiainen, J., and Worsnop, D. R.: Ubiquity and dominance of oxygenated species in organic aerosols in anthropogenically influenced Northern Hemisphere midlatitudes, *Geophys. Res. Lett.*, 34, L13801, doi:10.1029/2007GL029979, 2007.
- Zhang, Y., Hu, X.-M., Leung, L. R., and Gustafson Jr., W. I.: Impacts of regional climate change on biogenic emissions and air quality, *J. Geophys. Res.*, 113, D18310, doi:10.1029/2008JD009965, 2008.