Supporting Information for:

Towards a better understanding of the origins, the 3 chemical composition and the aging of the Oxygenated 4 study of **Mediterranean** Organic Aerosol: case a 5 **Industrialized Environment, Marseille** 6

7

Imad El Haddad ^{a,1,*}, Barbara D'Anna ^b, Brice Temime-Roussel ^a, Melanie Nicolas ^b, Antoinette Boreave ^b, Olivier Favez^{b,2}, Didier Voisin ^c, Jean Sciare ^d, Christian George ^b, Jean-Luc Jaffrezo ^c, Henri Wortham ^a, and Nicolas Marchand ^{a,*} 8

9

10 ^a Aix-Marseille Univ, Laboratoire Chimie Environnement, 13331, Marseille cedex 03, France

^b Université Lyon 1, Lyon, F-69626, France; CNRS, UMR5256, IRCELYON, Institut de Recherches sur la 11 12 Catalyse et l'Environnement de Lyon, Villeurbanne, F-69626, France

13 ^c Universités Joseph Fourier-Grenoble 1-CNRS, UMR 5183, Laboratoire de Glaciologie et Géophysique de

14 l'Environnement, Saint Martin d'Hères, 38402, France.

15 ^d Laboratoire des Sciences du Climat et de l'Environnement, CEA-CNRS-UVSQ-IPSL, Gif sur Yvette, 16 91191, France.

17 ¹Now at Laboratory of Atmospheric Chemistry, Paul Scherrer Institut, 5232 Villigen PSI, Switzerland.

²Now at INERIS, DRC/CARA/CIME, Parc Technologique Alata, BP2, Verneuil-en-Halatte, 60550, 18 19 France.

20 * Corresponding authors:

21 I. El Haddad – email: imad.el-haddad@psi.ch, phone: +41 5 63 10 27 85, fax: +41 5 63 10 45 25

22 N. Marchand – email: Nicolas.Marchand@univ-amu.fr, phone: +33 4 13 55 10 51, fax: +33 4 13 55 10 60

23

24 S.1 Synoptic air masses

Fig.S1 illustrates the air masses impacting the sampling site during the measurement period and shows that these air masses are mostly associated with long range transport from the Atlantic and the Mediterranean.



Fig.S1: 24h HYSPLIT air masses backward trajectories at 100m above sea level (Rolph, 2010)
illustrating the overall air masses circulation occurring during the entire measurement campaign.
Backward trajectories are confirmed by both MM5 modeling and local wind measurements.

32 S.2 Aerosol online chemical composition

33 The chemical composition of fine PM was measured in 2 min averages using a compact 34 time-of-flight (c-TOF, Tofwerk) Aerodyne Aerosol Mass Spectrometer (AMS, 35 Aerodyne). This instrument allows real-time measurements of PM_1 non-refractory 36 components (OA, NH₄, NO₃ and SO₄) combining thermal vaporization and electron 37 ionization (Drewnick et al., 2005). Aerosol size distribution (mobility diameters from 11 38 to 1083 nm), was investigated using a Scanning Mobility Particle Sizer (SMPS, L-DMA 39 and CPC5403, GRIMM). Semi-continuous hourly concentrations of elemental carbon 40 (EC) and organic carbon (OC) PM_{2.5} were obtained in the field from an OC/EC Sunset 41 field instrument (Sunset Laboratory, Forest Grove, OR, USA; Bae et al., 2004) running at 42 81 min⁻¹. AMS results are all corrected for the collection efficiency by using a common

- 43 factor of 0.65±0.14 estimated based on the comparison of total AMS measured mass and
- 44 SMPS + EC measured mass. Fig.S2 displays EC, OA, NH₄, NO₃ and SO₄ time series
- 45 recorded over the period of study.



1/7/08 2/7/08 3/7/08 4/7/08 5/7/08 6/7/08 7/7/08 8/7/08 9/7/08 10/7/08 11/7/08 12/7/08 13/7/08 14/7/08 15/7/08 Date and Time

46 47 **Fig.S2:** Time series of the main PM_1 components (EC, OA, NH_4 , NO_3 and SO_4) during the 48 period of study. AMS data are corrected for the collection efficiency using a common factor of 49 0.65 ± 0.14 . Due to technical issues, AMS measurements are not available between the 9th and 10^{th} 50 of July.

51 **S.3 Offline chemical analyses**

52 PM_{2.5} collected onto 150 mm-diameter filters was comprehensively characterized. 53 Technical description of the analysis techniques can be found in El Haddad et al., 2011a 54 and b. A brief outline of these measurements is included here.

EC/OC, ions, WSOC, HULIS_{WS} and elements: The carbonaceous content was analyzed 55 56 for EC and OC using a Thermo-Optical Transmission method on a Sunset Lab analyzer 57 (Birch and Cary, 1996), following both NIOSH (Schmid et al., 2001) and EUSAAR-2 58 (Cavalli et al., 2010) protocols. It is well established that different protocols result in very 59 different values for EC (Schmid et al., 2001). We based our analysis (i.e., Chemical Mass Balance analysis and multiple regression analysis, see below sections S.4 and S.9) on 60 61 concentrations determined following NIOSH protocol, as source profiles were 62 determined based on this protocol. Biases arising from discrepancies between the two protocols are all discussed in S.9.2 of the supporting material. 63

Sample fractions of 11.34 cm² taken from the sample filter were extracted into 15 mL 64 ultrapure Milli-Q water by 30 min short vortex agitation for the analyses of major ions 65 $(NH_4^+, SO_4^{2-}, NO_3^-)$, water-soluble organic carbon (WSOC) and water-soluble humic like 66 substances (HULIS). HULIS analysis was performed following the method described in 67

68 Baduel et al. (2009, 2010). This method involves extraction of HULIS by adsorption onto

- 69 DEAE resin (GE Healthcare®, HiTrapTM DEAE FF, 0.7 cm ID×2.5 cm length) and its
- round subsequent quantification with an OI Analytical 700 total organic carbon analyzer.
- 71 Finally, fifty elements were measured using ICP-MS (Agilent 7500ce) following the

complete dissolution of filter aliquots in a mixture of high-purity concentrated HF and

- HNO₃. Element concentrations were then calculated using the rock reference material BR
- 74 (Chauvel et al., 2010).

72

Radiocarbon measurements: Radiocarbon (¹⁴C) measurements were conducted on high 75 volume quartz filter fractions (~40 cm²) using ARTEMIS Accelerator Mass 76 77 Spectrometry. Each sample was first packed into a prefired quartz tube containing CuO 78 and Ag powder to be combusted at 850°C in a muffle furnace for 4 hours. Carbon dioxide 79 was collected and purified before its conversion into graphite by hydrogen reduction at 80 600°C using an Fe catalyst. The modern fraction (fm) was determined as the ratio of ¹⁴C/¹²C in aerosol sample to ¹⁴C/¹²C in the NBS Oxalic Acid standard (NIST-SRM-81 82 4990B).

In order to account for the thermonuclear weapon tests of the late 1950s and early 1960s (Levin et al., 1985), the modern fraction (f_m) is divided by a ratio of 1.1 to get a corrected non-fossil fraction (f_{nf}) (Levin and Hesshaimer, 2000). This value is subsequently subtracted from one to obtain the fossil fraction (f_f) .

87 <u>Organic speciation:</u> A chemical derivatisation/gas chromatography-mass spectrometry
88 (GC-MS) approach was used to quantify primary and secondary organic markers,
89 including α-pinene oxidation products, a major focus of this study. The approach is fully
90 described in El Haddad et al., 2011b and will be only outlined in the following.

91 Prior to extraction, filters were spiked with known amounts of two isotope-labelled 92 standards: tetracosane-d50 and cholesterol-d6. Organic species were extracted from 93 filters with a dichloromethane/acetone mix (1/1 v:v) using an accelerated pressurized 94 solvent extraction device (ASE, Dionex 300). Extracts were then reduced to a volume of 95 500µL using a Turbo Vap II concentrator. The remainder was split into two fractions. 96 The first fraction was directly injected, whilst the second fraction was subjected to 97 derivation for 2 h at 70 °C, using N,O-Bis(trimethylsilyl)-trifluoroacetamide containing 98 10% trimethyl-chlorosilane ,before GC-MS analysis. The two fractions were analyzed

following the same GC-MS conditions: Aliquots of 2 μ L were analyzed using a Thermo Trace GC chromatograph interfaced to a Polaris Q ion trap mass spectrometer fitted with an external electron ionization source. The chromatographic separation was accomplished on a TR-5MS capillary column (Thermo Electron, 30 m × 0.25 mm i.d. × 0.25 μ m film thickness). Field blank filters were also treated with the same procedure and none of the target compounds were detected in these field blanks.

105 Primary organic markers, including n-alkanes, hopanes, polycyclic aromatic hydrocarbon 106 (PAH) and levoglucosan, were quantified and used as inputs in the CMB analysis to 107 apportion primary aerosol and VOC sources. A complete list of these primary organic 108 markers is included in El Haddad al. (2011a). α -pinene oxidation products quantified by 109 this method were used to estimate α -pinene SOA contributions. As described in El 110 Haddad al. (2011b), we identified and quantified 9 α -pinene SOA markers, whose 111 structures are presented in Fig.2. These included pinic (PA) and pinonic (PNA) acid, 112 which were identified and quantified using authentic standards. Seven other 113 multifunctional compounds (A1-A7), for which native standards are not available, were 114 tentatively identified by examining their retention times and MS characteristics (for more 115 details refer to El Haddad et al., 2011b). They included 3-hydroxyglutaric acid (A1), 3-(2-hydroxyethyl)-2,2-dimethylcyclobutane carboxylic 116 acid (A2), 3-hydroxy-4,4-117 dimethylglutaric acid (A3), 3-acetylglutaric acid (A4), 3-acetyladipic acid (A5), and 3-118 isopropylglutaric acid (A6) and 3-methyl-1,2,3-butanetricarboxylic (A7). These 119 compounds were quantified using the response factor of malic acid as a surrogate for all 120 of the compounds. Relative standard deviation of the concentrations of these species based on duplicate analysis was between 5 and 15%. 121

122 S.4 Comparison between offline and online measurements

The aim of this section is to evaluate biases and artefacts associated with the offline and online measurements of OA (e.g., AMS particle collection efficiency, adsorption artefacts onto filters). Fig.S3 conveys the comparison between AMS (PM_1) and filter based ($PM_{2.5}$) measurements for the two major aerosol components: SO₄ and OA. SO₄ is expected to primarily occur in the PM₁ fraction and to be mostly associated with ammonium sulfate and bisulfate (with very little influence from sea salt), and thus to be quantitatively analysed by the AMS. A very good agreement was observed between the AMS-SO₄ and the HiVol-SO₄ (s~1, i~0 and R²>0.9; Fig.S3), substantiating our AMS measurements and the particle collection efficiency factor, HiVol-OA was derived from OC concentrations measured from filter samples, corrected for differences in diameter cut-offs between the AMS and the HiVol sampler; **it constitutes our best estimate of offline** *PM*₁*OA*.



Fig.S3: Comparison between AMS and offline measurements for SO₄ (**A**) and OA (**B**). HiVol-OA was corrected for differences in the diameter cut-offs between the AMS and the HiVol sampler (see text); it refers to the PM₁ fraction. Also shown are the 1:1 line and the slopes (s), intercepts (i) and coefficients of determination (\mathbb{R}^2) obtained by linear fits of the data.

140 The calculation of PM1OA proceeded as follows: Based on size resolved EC/OC measurements performed on the LPI samples, the fraction of PM_1OC in $PM_{2.5}OC$ was 141 retrieved: $PM_1OC/PM_{25}OC = 0.82 \pm 0.06$. PM_1OC was then scaled by an average 142 143 OM/OC ratio of 1.67 \pm 0.05, obtained by comparing the AMS-OA to the LPI PM_1OC . The comparison between PM_1OA and AMS-OA shows that both fractions exhibit similar 144 variability (R^2 >0.7), with a slope close to 1. However, a negative intercept of -1.3±0.7 µg 145 m⁻³ was observed, implying that filter based measurements (PM_1OA) were systematically 146 associated with a positive bias of 1.3 µg m⁻³ engendered by adsorption artefacts onto 147 filter samples. As a result, filter based measurements tended to overestimate the absolute 148 concentrations of OA by up to 28%[†]. However, such artefacts would have only a minor 149

[†] The following estimation of adsorption artefacts onto HiVol filter samples (positive artefacts of 28%) is obtained by assuming first no volatilisation artefacts occurred during sampling with the LPI and second a quantitative transmission and evaporation of PM_1 organic particles in the AMS. Negative artefacts are common for sampling under low pressure and losses and slow vaporisation of large particles can be an issue in the AMS. Therefore, our estimation of positive artefacts onto the HiVol samples should be regarded as highest estimate.,

150 influence on our apportionments, provided that they evenly impacted the different 151 components of OA.







Fig.S4: Influence of varying FPEAK parameter on factors' mass spectra and time series for the 155 4-factor solution and at FPEAKs between 0 and 1.

To assess the robustness of the 4 factor solution, rotational ambiguity has been 156 157 investigated by varying FPEAK from -2 to 2 with 0.1 steps. Two main groups of 158 solutions were identified, the first one corresponding to FPEAK values below 0, for which unrealistic zero time series values are observed for LVOOA, and the other one corresponding to "FPEAK" above 0. Robust solutions were found for solutions at FPEAK between 0 to 1, with very little variability in the factors' time series and mass spectra (Fig.S4). The influence of the initial conditions seed (corresponding to pseudorandom starting-points of the PMF2 algorithm) ranging from 0 to 59 (with steps of 1) was also verified. No influence of different seed was observed, which provides evidence of the robustness of the chosen solution.



166 m/z
 167 Fig.S5: Factor spectral profiles derived from the 4 factor solution PMF2 analysis for FPEAK 0 and seed=0.

169 S.6 Chemical Mass Balance analysis

170 Available data used here also included source contributions to OC, apportioned using a 171 Chemical Mass Balance analysis (CMB) in conjunction with organic marker 172 concentrations, as fully described in El Haddad et al. (2011a). CMB model is based on 173 the mass conservation of individual organic markers. In the mass conservation equations, 174 known concentrations (C_{ik}) of specific markers of primary sources at receptor site *k* are 175 written as the product of known source profiles a_{ij} and unknown primary source 176 contributions s_{jk} (Watson et al., 1998) as expressed in equation 1:

$$C_{ik} = \sum_{j=1}^{m} a_{ij} s_{jk} \tag{1}$$

177 where *m* denotes the total number of emission sources and a_{ij} represents the fractional 178 abundances of chemical species in the source emissions, expressed as marker-to-OC 179 ratios. The set of linear equations generated by equation 1 is solved with an effective 180 variance weighted least squares method using the Environmental Protection agency EPA-181 CMB8.2 software.

- 182 Primary markers and source profiles selection is detailed in El Haddad et al., 2011a. 183 Primary markers include: levoglucosan as a specific marker for biomass burning 184 (BBOC), EC and three hopanes (i.e., 17(H),21(H)-norhopane, 17(H),21(H)-hopane and 185 22S,17(H), 21 (H)-homohopane) as key markers for vehicular emissions. In addition, a 186 series of C27-C32 n-alkanes was selected since this range demonstrates high odd-carbon 187 preference, specific to primary biogenic sources. In order to apportion industrial 188 emissions, four PAH (benzo[b,k]fluoranthene, benzo[e]pyrene, indeno[1,2,3-cd]pyrene, 189 and benzo[ghi] perylene), V, Ni and Pb were included as fitting species. Source profiles 190 comprise vehicular emissions derived from a tunnel study held in Marseille (El Haddad et 191 al., 2009), biomass burning emissions (Fine et al., 2002), vegetative detritus (Rogge et al., 192 1993a) and natural gas combustion (Rogge et al., 1993b). Three industrial-emission-193 related profiles were chosen, including metallurgical coke production (Weitkamp et al., 194 2005), HFO combustion/shipping (Agrawal et al., 2008), and steel manufacturing (Tsai et 195 al., 2007).
- In this study, emissions from the three industrial processes were lumped together under the term "industrial OA". Biomass burning, vegetative detritus and natural gas combustion contributed very little OC during the period of measurements (El Haddad et al., 2011a) and were not considered in the comparison between CMB and AMS/PMF results. CMB technique is unable to directly apportion secondary sources; however, the fraction of OC not attributed to primary sources is considered to be an upper limit estimate of secondary OC (SOC).
- In order to compare CMB and AMS/PMF results, primary OA associated with vehicular and industrial emissions were calculated by applying an OM-to-OC ratio of 1.2 (based on Aiken et al., 2008). SOA was considered as the difference between the total OA,

- determined by scaling the total OC by an OM-to-OC ratio of 1.67^{\ddagger} , and the primary OA.
- 207 An OM-to-OC ratio of 1.82 can be inferred for SOA (i.e. SOA-to-SOC), consistent with
- an overwhelmingly secondary origin of this fraction (Aiken et al., 2008).



209 S.7 Tracers' diurnal profiles

210Diurnal HourDiurnal Hour211Fig.S6: Tracers' diurnal profiles. Solid lines denote average profiles and shaded areas represent212[P25-P75] range.

213 S.8 HOA vs. EC

Fig.S7 presents a scatter plot of HOA vs. EC, from which it is possible to ascertain more thoroughly the sources of this fraction. As HOA and EC mainly arise from the same source (i.e., vehicular emissions (El Haddad et al., 2011a)), it was expected that the data point cluster around one line with a slope corresponding to the HOA/EC ratio at the point of emission. However, more scatter was observed with three different clear patterns:

 $^{^{\}ddagger}$ OM-to-OC ratio of 1.67 is calculated by comparing AMS OA with LPI OC measurements, see section S.3 and S.4.

219 Most of the data scattered around one line characterised by a ratio of HOA/EC of ~0.4, 220 representative of average vehicular emissions at typical ambient concentrations (see for 221 e.g. Chirico et al., 2011 and references therein). This is a clear indication that HOA was 222 mostly related to traffic. A second part of the data scattered around another line 223 characterised by a lower ratio of HOA/EC (<0.25), concomitant with the dilution of the 224 emissions as the boundary layer developed in the afternoon and with the enhancement of 225 the photochemical activity, which would increase the oxidation of HOA. The depletion of 226 traffic emission markers with respect to EC due to photochemistry was previously 227 demonstrated to occur during this field mission (El Haddad et al., 2011a). The third part 228 of the data presents more scattering, with higher HOA/EC ratios (around 0.75), occurring 229 mostly during meal hours, especially during the evening. This suggests that the HOA 230 factor was contaminated by cooking emissions that had a similar spectral profile as HOA. 231 Based on the comparison between EC and HOA, this contamination can be estimated as 20%, in agreement with the very low concentrations of cholesterol (0.13-3.32 ng m⁻³. El 232 233 Haddad et al., 2011a).







237 S.9.1 Multiple regression model

AMS/PMF apportionments and ¹⁴C measurements were combined using a multiple regression model to estimate the fossil and the non-fossil contributions to both SVOOA and LVOOA. It is worthwhile to note that such a combination is not straightforward, involving a certain number of assumptions that result in considerable uncertainties. Firstly, ¹⁴C measurements were conducted on $PM_{2.5}OC$ onto filter samples that are subjected to well-known but non systematic adsorption artefacts of gas phase organic compounds. In contrast, AMS provides real-time measurements of PM_IOA with little interference from gas phase organics. However, particle collection efficiency (CE) of the AMS, estimated in this dataset to be 0.65±0.14, can be highly variable and is dependent on the aerosol chemical nature and mixing state (Middlebrook et al., 2012).

Secondly, ¹⁴C measurements conducted in this study relate to the total carbon (TC) mass 248 249 that can be oxidized at 850 °C under oxygen, i.e., organic carbon (OC) and elemental 250 carbon (EC), whereas AMS quantifies OA that consists of OC and the associated 251 heteroatom (H, N, O, S ...). This fundamental difference engenders two major limitations for the assessment of fossil and non-fossil contributions of the OOA fractions. First, as 252 the TC apportioned by ¹⁴C measurements also included EC, assumptions related to the 253 254 origins of the latter must be made. Furthermore, since the separation between EC and OC 255 measured using the OC/EC instrument is method-dependent, biases associated with EC 256 determination can impact the estimations. Second, the AMS measurement of OA includes 257 heteroatom that can be unevenly distributed between the fossil carbon and the non-fossil 258 carbon. Here also, this distribution was not empirically accessible and assumptions have to be made to calculate our estimations. Further uncertainties can arise from PMF 259 calculations and residuals and from variability in the biomass ${}^{14}C/{}^{12}C$ ratio. The 260 assumptions made to achieve the apportionment are explicitly presented in this section 261 262 and the resulting biases and uncertainties are thoroughly discussed in section S.9.2.2.

The procedure went as follows: First, in order to estimate the fossil and non-fossil fractions of OC, EC was assumed to be entirely related to fossil carbon (assumption founded on Chemical Mass Balance calculations reported in El Haddad et al., 2011a). This is described in equations (2) and (3):

$$OC_f = TC_f - EC_f = TC_f - EC \left(EC_f = EC \parallel EC_{nf} = 0 \right)$$
(2)

$$OC_{nf} = TC_{nf} - EC_{nf} = TC_{nf} \left(EC_f = EC \parallel EC_{nf} = 0 \right)$$
(3)

267 Where TC_f , OC_f and EC_f correspond to the fossil TC, OC and EC, respectively, and 268 TC_{nf} , OC_{nf} and EC_{nf} to the non-fossil TC, OC and EC, respectively.

- 269 The second step involved the transition from $PM_{2.5}OC$ measured on filter samples to
- 270 PM_1OA determined by the AMS. This conversion was achieved for each of the samples
- 271 by scaling OC_f and OC_{nf} to fossil $OA(OA_f)$ and non-fossil $OA(OA_{nf})$, respectively, by
- 272 a factor α :

$$OA_{f} = \alpha \times OC_{f} \left(\alpha = PM_{1}OA / PM_{2.5}OC \right)$$
⁽⁴⁾

$$OA_{nf} = \alpha \times OC_{nf} \left(\alpha = PM_1 OA / PM_{2.5} OC \right)$$
⁽⁵⁾

where α is the ratio between AMS PM_1OA and filter $PM_{2.5}OC$. α is variable depending on the considered sample, but has an average of 0.92±0.21. The assumption underlying this scaling is that $PM_{2.5}OC$ and PM_1OA were associated with the same proportions of fossil and non-fossil mass (i.e. the ratio fossil/total was the same for $PM_{2.5}OC$ and PM_1OA). Biases arising form this assumption are discussed in the section S.9.2.2.

 OA_f and OA_{nf} obtained in equations 4 and 5 can be accordingly expressed as a linear 278 279 combination of the AMS/PMF OA fractions derived from fossil $(fOA_i: POA_f, SVOOA_f \text{ and } LVOOA_f)$ and non-fossil $(nfOA_i: POA_{nf}, SVOOA_{nf} \text{ and } LVOOA_f)$ 280 LVOOA_{nf}) sources, respectively: 281

$$OA_f = \sum_{i}^{l} fOA_i = POA_f + SVOOA_f + LVOOA_f$$
(6)

$$OA_{nf} = \sum_{i}^{m} nfOA_{i} = POA_{nf} + SVOOA_{nf} + LVOOA_{nf}$$
(7)

where *l* and *m* are the total numbers of fOA_i and $nfOA_i$ fractions, respectively. As fOA_i and $nfOA_i$ are not directly accessible, equations 6 and 7 can be written in terms of the OA factors (OA_i) determined by AMS/PMF analysis and the respective share of fossil and non-fossil fractions to these factors such that

$$OA_{f} = \sum_{i}^{l} a_{i} \times OA_{i} = a_{1} \times POA + a_{2} \times SVOOA + a_{3} \times LVOOA \left(a_{i} = \frac{fOA_{i}}{fOA_{i} + nfOA_{i}}\right)$$
(8)

$$OA_{nf} = \sum_{i}^{m} b_{i} \times OA_{i} = b_{1} \times POA + b_{2} \times SVOOA + b_{3} \times LVOOA \left(b_{i} = \frac{nfOA_{i}}{fOA_{i} + nfOA_{i}} \right)$$
(9)

286 where a_i and b_i denote the relative share of fossil and non-fossil fractions to OA_i factors, respectively $(a_1 \text{ for } POA_f, a_2 \text{ for } SVOOA_f, a_3 \text{ for } LVOOA_f, b_1 \text{ for } POA_{nf}, b_2 \text{ for } POA_{nf}, b_1 \text{ for } POA_{nf}, b_2 \text{ for } POA_{nf}, b_2 \text{ for } POA_{nf}, b_1 \text{ for } POA_{nf}, b_2 \text{ for } POA_{nf}, b_2 \text{ for } POA_{nf}, b_1 \text{ for } POA_{nf}, b_2 \text{ for } POA_{nf}, b_2 \text{ for } POA_{nf}, b_1 \text{ for } POA_{nf}, b_2 \text{ for } POA_{nf}, b_3 \text{ for } POA_{nf}, b_4 \text{ for$ 287 $SVOOA_{nf}$ and b_3 for $LVOOA_{nf}$). This system of linear equations can be visualised as 288 289 the following matrix equation:

$$\begin{bmatrix} POA & SVOOA & LVOOA & 0 & 0 & 0 \\ 0 & 0 & 0 & POA & SVOOA & LVOOA \\ 1 & 0 & 0 & 1 & 0 & 0 \\ 0 & 1 & 0 & 0 & 1 & 0 \\ 0 & 0 & 1 & 0 & 0 & 1 \end{bmatrix} \times \begin{bmatrix} a_1 \\ a_2 \\ a_3 \\ b_1 \\ b_2 \\ b_3 \end{bmatrix} = \begin{bmatrix} OA_f \\ OA_{nf} \\ 1 \\ 1 \\ 1 \end{bmatrix}$$
(10)

П

290 As POA was assumed be strictly related fossil to to sources (i.e. $POA_f = POA = HOA + F4$), the parameter a_1 can be assumed to equal 1, implying 291 that b_1 equals 0 (i.e. $POA_{nf} = 0$, the sensitivity of the results to this assumption is 292 293 assessed in section S.9.2.2). Equation 11 can be then simplified as follows:

$$\begin{bmatrix} SVOOA \ LVOOA \ 0 \ 0 \end{bmatrix} \times \begin{bmatrix} a_2 \\ a_3 \\ b_2 \\ 0 \ 1 \ 0 \ 1 \end{bmatrix} = \begin{bmatrix} SOA_f \\ SOA_{nf} \\ 1 \\ 1 \end{bmatrix}$$
(11)

294 With SOA_f and SOA_{nf} denoting the fossil and non-fossil fractions of SOA, respectively, calculated as: $SOA_f = OA_f - POA_f$ and $SOA_{nf} = OA_{nf} - POA_{nf}$. The equations of the 295 296 resulting linear system (equation 11) are not independent and thus the system has no 297 solution. For that reason, a multiple linear regression analysis was applied instead to 298 solve equation 11, which found the average values for a_i and b_i that fit best the equation, 299 in the sense of solving the quadratic minimization problem. In equation 11 AMS/PMF OA_i vectors were included as independent variables and OA_f and OA_{nf} as 300 301 dependant variables.

It should be noted that the apportionment procedure followed here is not unique. As this 302 303 study mainly focuses on AMS measurements, the chosen method is AMS data oriented in 304 that the resulting apportionments would exhibit the same variability as the AMS/PMF factors (e.g. $SVOOA_f + SVOOA_{nf} = SVOOA$). This is the result of the equation 11, in which AMS/PMF OA_i vectors were chosen as independent variables. As this equation is a self-consistent system (i.e. $OA_f + OA_{nf} - POA = SVOOA + LVOOA$), another approach is also possible; by considering OA_f and OA_{nf} as independent variables, this second approach yields the same average results but orients the variability towards ¹⁴C measurements.

311 **S.9.2** Output quality control, uncertainty assessments and potential biases

312 S.9.2.1 Output quality control and residual analyses

313 One of the major drawbacks of the multiple regression analysis applied here is that it 314 considers a constant contribution of fossil and non-fossil sources to each of the 315 OA_i factors (i.e., constant a_i and b_i ratios), while these contributions may significantly 316 vary over the course of the measurements. Accordingly, a_i and b_i ratios should be 317 regarded as average contributions of fossil and non-fossil sources to OA_i . These ratios are 318 reported in Tab.S1.

Tab.S1: a_i and b_i ratios for the POA,SVOOA and LVOOA fractions.

	POA	SVOOA	LVOOA
a_i	1.0	0.33±0.11	0.082 ± 0.085
b_i	0.0	0.67 ± 0.11	0.92 ± 0.08



319measured f [ug m⁻]measured nf [ug m⁻] $0.92 \times PM_{2.5}OC+EC [ug m⁻]320Fig.S8: Scatter plots of modelled vs. measured data for the fossil (A) and non-fossil (B) fractions.321Measured fossil fraction=EC+0.92 xOC_f, modelled fossil fraction=EC+HOA+F4+322SVOOA_f+LVOOA_f, Measured non-fossil fraction=0.92 xOC_f, and modelled fossil fraction=323SVOOA_{nf}+LVOOA_{nf}. The 0.92 value is the average value of the factor <math>\alpha$ used in equations 6 and3247. The comparison between filter measurements and AMS measurements is shown in panel C.

For all panels, the slope of the linear regression (s), its intercept (i) and its coefficient of determination (R) are also indicated (n=28 samples for each plot).

327 Fig.S8 compares the measured and the modelled concentrations for the total fossil and 328 non-fossil fractions. It shows that the model captured quite well the amounts and the 329 variability of the measured concentrations, especially in the case of the fossil fraction 330 (Fig.S8a). In the case of the non-fossil fraction, the model tended to slightly 331 underestimate (overestimate) the measured levels at low (high) concentrations (Fig.S8b). 332 It should be noted though that most of the variability observed in panels A and B of 333 Fig.S8 arose from discrepancies between the AMS PM_1 OA and filter-based $PM_{2.5}$ OC, as 334 shown in panel C of the same figure. Differences between the 2 measurement techniques 335 were accounted for in equations 4 and 5 prior to the multiple regression analyses by the 336 coefficient α that encompasses various conversion factors. Overall, these comparisons validated the representativeness of a_i and b_i obtained in the multiple regression analyses. 337

338



339

Fig.S9: Residuals' normal distributions (modelled – measured) derived from the multiple regression approach applied above for the fossil (**A**) and the non-fossil (**B**) fractions. Residuals are fitted using a Gaussian fit, from which the mean (\overline{X}) and the standard deviation (σ) are calculated for both fractions.

Assumptions underlying the residuals' distributions were examined for the fossil and non-fossil fractions in Fig.S9. Residuals followed normal distributions with mean values statistically equal to zero, implying that errors are homoscedastic (variance = 0) and are not correlated. From Fig.S9, it is possible to estimate the uncertainties related to the total fossil and non-fossil OA fractions. Fossil and non-fossil OA were accordingly estimated to contribute $1.52\pm0.31 \ \mu g \ m^{-3}$ (implying 20% of errors) and $2.52\pm0.78 \ \mu g \ m^{-3}$ (implying 31% of errors), respectively. These uncertainties included: (1) measurement differences between filter-based PM_{2.5} TC and AMS PM₁ OA + EC (Fig.S8c) and (2) variability in $a_i \ and \ b_i \ obtained$ in the multiple regression analyses (see the related uncertainties in Tab.S1).

354 S.9.2.2 General assessment of uncertainties and biases

355 It is worthwhile to note that a great part of the uncertainties assessed for the absolute 356 concentrations of fossil and non-fossil OA arose from discrepancies between AMS and 357 filter measurements and hence is not representative of the statistical significance of each 358 of the fractions. The statistical significance of the relative contributions of HOA, F4, 359 SVOOA_f, LVOOA_f, SVOOA_{nf} and LVOOA_{nf} were assessed through a sensitivity test 360 using a random selection technique. Inputs to the calculation are the PMF factor mass concentrations, ¹⁴C data, and OC/EC measurements. The calculation was performed 361 based on equations (2-11) and proceeded as follows: 362

- For each of the input parameters a range was assigned, within which these can
 vary (see Tab.S2). The criteria on which we based our assessment of these ranges
 are developed below, in Appendix A.
- o The parameters were then allowed to randomly vary within the range 366 predetermined in the previous step, assuming a normal distribution. This approach 367 368 is somewhat similar to Monte Carlo calculations and allows vast numbers of 369 combinations of input parameters to be computed (McKay et al., 1979). A Monte Carlo simulation would involve testing all possible combinations of input 370 371 parameters, which would be prohibitive in terms of processing time. In contrast, 372 random sampling is much more effective and for our purposes provided 373 essentially the same results as a full Monte Carlo analysis (McKay et al., 1979).
- o Following the approach described above, 50 sets of parameters were generated randomly and used subsequently in the equations 2-11 to calculate the inputs for the multiple linear regression analysis (i.e. SOA_f , SOA_{nf} , SVOOA and LVOOA). This provided for each set of parameters average values for a_i and b_i plus the corresponding uncertainties.

o For each set of the coefficients a_i and b_i previously generated, the average values of these coefficients were varied assuming a binomial distribution, derived based on the corresponding uncertainties provided by the multiple linear regression analyses. In this step and for each set of parameters, ten values were generated for a_i and b_i and used to compute the contributions of SVOOA_f, LVOOA_f, SVOOA_{nf} and LVOOA_{nf}. This gave in total 500 different solutions.

A great advantage of this approach is that combinations of parameters which are very unlikely (e.g., that only the minimum-possible values from each parameter were used) will represent only a small percentage of the output. The obtained 500 solutions are presented graphically as a probability density (frequency distribution) of possible solutions to the source apportionment problem we have set up, as shown in Fig.7 in the manuscript.

391 This analysis provided strong support for our results, allowing the assessment of the 392 uncertainties underlying our measurements and assumptions and offering a measure of 393 our ability to separate the different components (statistical significance of each 394 component). It showed that the uncertainties of our estimations depend on the component 395 considered. Depending strictly on the PMF analysis errors, the uncertainties associated 396 with the contributions of POA (HOA and F4) were less than 10%. Conversely, for OOA 397 components the uncertainties were less homogenous. For non-fossil OOAs the 398 uncertainties were around 10% as these were well resolved by the regression model, 399 whereas for fossil OOAs uncertainties are higher (\sim 36% and \sim 58% for SVOOA_f and LVOOA_f, respectively), as these were poorly resolved by the regression model and 400 401 strongly dependant on the EC measurements and the assumptions made for POA. All 6 402 fractions were statistically significant with contributions higher than 0 (Z equal 23, 7.4, 403 2.8, 1.7, 9.2, 12 for HOA, F4, SVOOA_f, LVOOA_f, SVOOA_{nf} and LVOOA_{nf} respectively, 404 with Z=average/uncertainty).

405 Additionally, the sensitivity test offered the assessment of the biases on the 406 apportionments presented in the manuscript. The sensitivity test results suggest that we 407 might underestimate the contributions of F4, HOA, $SVOOA_{nf}$ and $LVOOA_{f}$ by 9%, 6%, 408 5% and 23%, respectively and overestimate the $SVOOA_{f}$ and $LVOOA_{nf}$ by 34% and 5%, 409 respectively. The main conclusion to be derived from this analysis is the robustness of the 410 results presented and discussed in the manuscript. For example, OOA_{nf} was clearly the 411 biggest contributor to OA. It is also clear that LVOOA was derived predominately from 412 non-fossil precursors (LVOOA_{nf} /LVOOA=89±7%), whereas SVOOA included a larger 413 fraction of fossil SOA (SVOOA_{nf}/SVOOA=75±8%). Given the wide range of 414 uncertainties used in the sensitivity test, these results demonstrate that, in general, we can 415 clearly identify the contribution from different components.





418 **Fig.S10:** 2D-framework for α -pinene SOA aging adapted from **Jimenez et al.**, 2009, representing 419 OA oxidation state (approximated by O:C) vs. OA volatility (log₁₀(C*) at 298 K). The ambient 420 OOA factors are represented in this 2D space by the 2 green squares, with LV-OOA being less 421 volatile and more oxidized than SV-OOA (Jimenez et al., 2009). First generation products from 422 α -pinene (yellow pentagon) + ozone reaction are distributed according to the blue contour. 423 Products derived from subsequent OH oxidation (1.5 OH lifetimes) of first generation products 424 are represented with purple contour (Jimenez et al., 2009). This oxidation reproduces a substantial 425 shift toward ambient LV-OOA volatility and oxidation state. We added on this 2D space α -pinene 426 first and subsequent generation oxidation products measured in this study by GC/MS (PNA, PA

427 and A1-A7). The volatilities of these compounds were calculated using the approach proposed by 428 Donahue et al., 2011. O:C ratios of pure SV-OOA and LV-OOA retrieved by AMS/PMF2 429 analysis were calculated following the parameterization proposed in Aiken et al. (2009) and 430 indicated by the orange and dark red lines, respectively. Following the same methodology, the 431 range of O:C ratios (0.48<O:C_{OOA}<0.72) of total OOA encountered during the measurement 432 period was determined and indicated by the dotted area.

434 Appendix A: Calculations of the different parameters entered as inputs in the sensitivity test

435 For the parameters in Tab.S2, the ranges were established as follows:

436 • For EC/OC measurements, the range was designed to encompass biases and uncertainties 437 associated with the separation between EC and OC. This range was bounded by 438 measurements determined following NIOSH and EUSAAR2 protocols, respectively. For 439 EC and OC, a constant bias between the 2 protocols was determined to be $40\pm8\%$ and 440 $6\pm5\%$.

- 441 The average uncertainty for the discrimination between fossil and non-fossil TC was 4%,
- 442 including uncertainties in 14 C measurements and errors due to the correction for 14 C 443 inputs from the bomb testing.

An assumption made in equations 4 and 5 relates to the origin of EC, which was
estimated to only pertain to the fossil fraction. This assumption would bias high the
contributions of fossil sources to the secondary OC fractions. As there was little influence
from biomass burning, we assumed an upper limit contribution of non-fossil sources to
EC to be 15%, (based on Minguillón et al., 2011 and references), and varied this
contribution between 0 and 15%.

E *	Parameters	Variables	Low	High	Remarks	
2, 3	OC_{f}, OC_{nf}	OC/EC	NIOSH	EUSAAR2		
		F_{f}	$0.96 \times F_{nf}$	$1.04 \times F_{nf}$	Uncertainties on measurements of ¹⁴ C in TC	
		EC_{f}	$0.85 \times EC$	EC	Origin of EC: $EC = EC_f + EC_{nf}$	
4, 5	$\alpha = PM_1OA/PM_{2.5}OC$ see Tab.S3	α_1^{OC}	0.76	0.88	Diameter cut-offs: $\alpha_1^{OC} = PM_1OC/PM_{2.5}OC$	
		a_2^{OC}	CI(-)	CI(+)	Positive artefacts based on Fig.S3: CI(-) and CI(+) are the upper and lower bounds of the confident interval on the linear regression	
8,9	AMS/PMF2 OA	Factors	FPEAK0	FPEAK1	AMS/PMF2 results obtained for FPEAKs between 0 and 1.	
10,11	POA _f , POA _{nf}	$a_1 \times POA$	0.75× <i>HOA</i> + <i>F</i> 4	0.9× <i>HOA</i> + <i>F</i> 4	$a_1 \times POA$ is the fraction of fossil POA. Its uncertainty is constrained based on Fig.S7	

Tab.S2: Ranges [Low, High] of the different parameters entered as inputs to the uncertainty calculation. E* denotes equations 2-11.

450 • The conversion from $PM_{2.5}OC$ measured on filter samples to PM_1OA determined by the 451 AMS was performed in equations 2 and 3, using a common factor α (with $\alpha =$ 452 $PM_1OA/PM_{2.5}OC$) for both fossil and non-fossil OC. This factor encompassed three key 453 corrections related to differences between the two measurement techniques, including 454 differences in diameter cut-offs between AMS and filter sampling (referred to as $PM_1OC/PM_{2.5}OC$ ratio), the adsorption artefacts on the filters, and the OM/OC ratio. The assumption underlying the $PM_{2.5}OC$ to PM_1OA conversion is that the aggregate of the aforementioned corrections was similar for both fossil and non-fossil OC and well represented by α . Indeed, these corrections can vary greatly between the primary and the secondary fractions, which exhibited variable contributions to the fossil and non-fossil OC. To address the biases resulting from the application of a single conversion factor α , the latter is deconvolved in equation (12) into several factors, such that

$$PM_1OA = \alpha \times PM_{2.5}OC = \sum_{i}^{n} \left(\prod_{j}^{p} \alpha_{ij}\right) \times OC_i$$
(12)

In this equation, α_{ij} denotes the factor used for a conversion (j) applied to an OC_i 462 fraction. n is the total number of OC fractions; in our case, it was limited to two fractions 463 representing the primary and the secondary OC. p is the total number of α_i conversions 464 applied to OC_i ; in our case, p was equal to 3, accounting for the $PM_1OC/PM_{2.5}OC$ 465 ratio (i=1), for the sampling artefacts (i=2) and for the OM/OC ratio (i=3). The following 466 is an example performed for the campaign average value, representing the matrix of 467 α_{ij} factors for primary and secondary OC (Tab.S3). Similar calculations were performed 468 for all of the data set to complete the sensitivity test. In this calculation, α_{ii} were 469 measured for the total OC $\left(\alpha_{i}^{OC}\right)$, assumed for the primary OC $\left(\alpha_{i}^{POC}\right)$, and inferred 470 for secondary OC $\left(\alpha_{i}^{SOC}\right)$. The different conversions include the following: 471

472 - α_1^{OC} denotes the $PM_1OC/PM_{2.5}OC$ ratio, estimated using size resolved OC 473 measurements, i.e., 0.82±0.06%. POC was assumed to pertain entirely to the PM₁ 474 fraction ($\alpha_1^{POC} = 1$), resulting in an α_1^{SOC} of 0.77 (i.e. 77% of the PM_{2.5} SOC are included 475 in the PM₁ fraction).

476 - α_2^{OC} is the ratio allowing the correction for the sampling artefacts retrieved from 477 Fig.S3, with an average value of 0.72. Artefacts were assumed to be evenly distributed 478 between the primary and the secondary fractions, i.e. $\alpha_2^{POC} = \alpha_2^{SOC}$ (Tab.S3). 479 - α_3^{OC} is the OM/OC ratio, obtained from the comparison between PM₁ AMS and LPI 480 measurements. An average α_3^{OC} value of 1.67 was found, and assuming an α_3^{POC} value 481 of 1.2 for primary OC a value of 1.81 can be inferred for the α_3^{SOC} .

482

Tab.S3: α_{ij} factors estimated for POC, SOC and total OC fractions

SOC and total OC fractions.							
$lpha_{ij}$	$\alpha_{_{j}}^{POC}$	α_{j}^{SOC}	α_{j}^{OC}				
α_{i1}	1.0	0.77	0.82				
α_{i2}	0.72	0.72	0.72				
α_{i3}	1.20	1.81	1.67				
$\prod_{j}^{p} lpha_{ij}$	0.84	0.98	0.95				

483 From α_{ij} matrix, overall conversion factors $\prod_{j}^{p} \alpha_{ij}$ can be inferred for POC, SOC and total

484 OC, which were equal to 0.84, 0.98 and 0.95, respectively. Estimated $\prod_{j}^{p} \alpha_{j}^{OC}$ (0.95) is

485 comparable to the average α empirically determined and used in equations (4) and (5) to 486 convert from $PM_{2.5}OC$ measured on filter samples to PM_1OA determined by the 487 AMS ($\alpha = 0.92 \pm 0.21$).

488 Using $\prod_{j=1}^{p} \alpha_{j}^{POC}$ and $\prod_{j=1}^{p} \alpha_{j}^{SOC}$ obtained above, one can apply different conversion factors

489 to POA and SOA. In the sensitivity test such a calculation has been made by considering 490 a range of α_1^{OC} and α_2^{OC} (Tab.S3).

In the calculation of equations 10 and 11, we considered for the apportionments the
AMS/PMF2 results, including *POA*, *SVOOA* and *LVOOA*. One approach to assess the
uncertainties on the AMS/PMF2 apportionments consists of varying FPEAK within a
reasonable range, in our case between FPEAK =0 and FPEAK =1.

In the calculation of equations 10 and 11, we considered that POA pertains only to the
fossil fraction. However, we observed evidence of inputs from cooking (a non fossil
primary source) to the HOA factor. To take this observation into account in the

- 498 uncertainty calculations, we considered that these cooking inputs contributed between
- 499 10% and 25% of total HOA, based on Fig.S7.

503 References

- Agrawal, H., Welch, W. A., Miller, J. W., and Cocker, D. R.: Emission measurements from a crude oil
 tanker at sea, Environ. Sci. Technol., 42, 7098–7103, 2008.
- Aiken, A. C., DeCarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty, K. S., Ulbrich, I. M.,
 Mohr, C., Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q., Trimborn, A., Northway, M., Ziemann, P.
 J., Canagaratna, M. R., Onasch, T. B., Alfarra, M. R., Prevot, A. S. H., Dommen, J., Duplissy, J.,
 Metzger, A., Baltensperger, U., and Jimenez, J. L.: O/C and OM/OC Ratios of Primary, Secondary,
 and Ambient Organic Aerosols with HighResolution Time-of-Flight Aerosol Mass Spectrometry,
 Environ. Sci. Technol., 42, 4478–4485, doi:10.1021/es703009q, 2008.
- Baduel, C., Voisin, D., and Jaffrezo, J. L.: Comparison of analytical methods for HULIS measurements in
 atmospheric particles, Atmos. Chem. Phys., 9, 5949-5962, doi:10.5194/acp-9-5949-2009, 2009.
- Baduel, C., Voisin, D., and Jaffrezo, J. L.: Seasonal variations of concentrations and optical properties of
 water soluble HULIS collected in urban environments, Atmos. Chem. Phys., 10, 4085–4095,
 doi:10.5194/acp-10-4085-2010, 2010.
- Bae, M.-S., Schauer J.J., DeMinter, J.T., Turner, J.R., Smith, D. and Cary, R.A.: Validation of a semicontinuous instrument for elemental carbon and organic carbon using a thermal-optical method, Atmos.
 Environ., 38, 2885–2893, 2004.
- 520 Cavalli, F., Viana, M., Yttri, K. E., Genberg, J., and Putaud, J.-P.: Toward a standardised thermal-optical
 521 protocol for measuring atmospheric organic and elemental carbon: the EUSAAR protocol, Atmos.
 522 Meas. Tech., 3, 79–89, doi:10.5194/amt-3-79-2010, 2010.
- 523 Chauvel, C., Bureau, S., and Poggi, C.: Comprehensive chemical and isotopic analyses of basalt and
 524 sediment reference materials, Geostand. Geoanalyt. Res., 35, 125–143, doi:10.1111/j.1751525 908X.2010.00086., 2010.
- 526 Chirico, R., A.S.H. Prevot, P.F. DeCarlo, M.F. Heringa, R. Richter, E. Weingartner, and Urs Baltensperger.
 527 Aerosol and Trace Gas Vehicle Emission Factors Measured in a Tunnel Using an Aerosol Mass
 528 Spectrometer and other On-line Instrumentation. Atmospheric Environment, 45(13), 2182-2192,
 529 2011.
- Donahue, N. M., Kroll, J. H., Pandis, S. N., Robinson, A. L.: A two-dimensional volatility basis set Part
 2: Diagnostics of organic-aerosol evolution, Atmos. Chem. Phys. Discuss., 11, 24883–24931,
 doi:10.5194/acpd-11-24883-201, 2011.
- Drewnick, F., Hings, S. S., DeCarlo, P., Jayne, J.T., Gonin, M., Fuhrer, K., Weimer, S., Jimenez, J.L.,
 Demerjian, K.L., Borrmann, S., Worsnop, D.R.: A new time-of-flight aerosol mass spectrometer
 (TOF-AMS) Instrument description and first field deployment, Aerosol Sci. Tech., 39, 637–658,
 2005.

- El Haddad, I., Marchand, N., Dron, J., Temime-Roussel, B., Quivet, E., Wortham, H., Jaffrezo, J. L.,
 Baduel, C., Voisin, D., Besombes, J. L., and Gille, G.: Comprehensive primary particulate organic
 characterization of vehicular exhaust emissions in France, Atmos. Environ., 43, 6190–6198, 2009.
- El Haddad, I., Marchand, N., Wortham, H., Piot, C., Besombes, J. -L., Cozic, J., Chauvel, C., Armengaud,
 A., Robin, D., Jaffrezo. J. -L., Primary sources of PM2.5 particles in an industrial Mediterranean
 city, Marseille, Atmospheric Chemistry and Physics, 11, 2039–2058, 2011a.
- El Haddad, I., Marchand, N., Temime-Roussel, B., Wortham, H., Piot, C., Besombes, J. -L., Baduel, C.,
 Voisin, D., Armengaud, A., Jaffrezo, J. -L., Insights into the secondary fraction of the organic
 aerosol in a Mediterranean urban area: Marseille, Atmospheric Chemistry and Physics, 11, 2059–
 2079, 2011b.
- Fine, P. M., Cass, G. R., and Simoneit, B. R. T.: Chemical Characterization of Fine Particle Emissions from
 the Fireplace Combustion of Woods Grown in the Southern United States, Environ. Sci. Technol.,
 36, 1442–1451, doi:10.1021/es0108988, 2002.
- 550 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo, P. 551 F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., 552 Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, 553 J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., 554 Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J. A., Onasch, T. B., 555 Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., 556 Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., 557 Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., 558 Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., 559 Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of organic aerosols in the atmosphere, 560 Science, 326, 1525-1529, 2009.
- Levin, I., Kromer, B., Schochfischer, H., Bruns, M., Munnich, M., Berdau, D., Vogel, J. C., and Munnich,
 K. O.: 25 Years of Tropospheric C-14 Observations in Central-Europe, Radiocarbon, 27, 1-19, 1985.
- Levin, I., and Hesshaimer, V.: Radiocarbon A unique tracer of global carbon cycle dynamics,
 Radiocarbon, 42, 69-80, 2000.
- 565 Minguillón, M. C., Perron, N., Querol, X., Szidat, S., Fahrni, S. M., Alastuey, A., Jimenez, J. L., Mohr, C.,
- 566 Ortega, A. M., Day, D. A., Lanz, V. A., Wacker, L., Reche, C., Cusack, M., Amato, F., Kiss, G., Hoffer,
- 567 A., Decesari, S., Moretti, F., Hillamo, R., Teinilä, K., Seco, R., Peñuelas, J., Metzger, A., Schallhart, S.,
- 568 Müller, M., Hansel, A., Burkhart, J. F., Baltensperger, U., and Prévôt, A. S. H.: Fossil versus
- 569 contemporary sources of fine elemental and organic carbonaceous particulate matter during the DAURE

- 570 campaign in Northeast Spain, Atmos. Chem. Phys., 11, 12067-12084, doi:10.5194/acp-11-12067-2011,
- 571 2011.
- 572 McKay, M. D., Beckman ,R. J. and Conover, W.J.: A Comparison of Three Methods for Selecting Values
- 573 of Input Variables in the Analysis of Output from a Computer Code, Technometrics, 21, 239-245, 1979.
- 574 Middlebrook, A.M., Bahreini, R., Jimenez, J.L. and Canagaratna M.R.: Evaluation of Composition-
- 575 Dependent Collection Efficiencies for the Aerodyne Aerosol Mass Spectrometer using Field Data.
- 576 Aerosol Science and Technology, DOI:10.1080/02786826.2011.620041, 46, 258–271, 2012
- Rolph, G. D.: Real-time Environmental Applications and Displays Ystem (READY) Website
 (http://ready.arl.noaa.gov). NOAA Air Resources Laboratory, Silver Spring, MD, 2010.
- Rogge, W. F., Hildemann, L. M., Mazurek, M. A., Cass, G. R., and Simoneit, B. R. T.: Sources of Fine
 Organic Aerosol .4. Particulate Abrasion Products from Leaf Surfaces of Urban Plants, Environ. Sci.
 Technol., 27, 2700–2711, 1993a.
- Rogge, W. F., Hildemann, L. M., Mazurek, M. A., Cass, G. R., and Simoneit, B. R. T.: Sources of Fine
 Organic Aerosol.5. NaturalGas Home Appliances, Environ. Sci. Technol., 27, 2736–2744, 1993b.
- 584 Schmid, H., Laskus, L., Jurgen Abraham, H., Baltensperger, U., Lavanchy, V., Bizjak, M., Burba, P.,
- 585 Cachier, H., Crow, D., Chow, J., Gnauk, T., Even, A., ten Brink, H. M., Giesen, K.-P., Hitzenberger, R.,
- 586 Hueglin, C., Maenhaut, W., Pio, C., Carvalho, A., Putaud, J.-P., Toom-Sauntry, D., and Puxbaum, H.:
- 587 Results of the "carbon conference" international aerosol carbon round robin test stage I, Atmos.
- 588 Environ., 35, 2111–2121, 2001.
- 589 Tsai, J.-H., Lin, K.-H., Chen, C.-Y., Ding, J.-Y., Choa, C.-G., and Chiang, H.-L.: Chemical constituents in
- 590 particulate emissions from an integrated iron and steel facility, J. Hazard. Mater., 147, 111–119, 2007.
- Watson, J. G., Robinson, N. F., Fujita, E. M., Chow, J. C., Pace, T. G., Lewis, C., and Coulter, T.: CMB8
 Applications and Validation Protocol for PM2.5 and VOCs, US EPA, USA, 1998.
- 593 Weitkamp, E. A., Lipsky, E. M., Pancras, P. J., Ondov, J. M., Polidori, A., Turpin, B. J., and Robinson, A.
- 594 L.: Fine particle emission profile for a large coke production facility based on highly time-resolved
- fence line measurements, Atmos. Environ., 39, 6719–6733, 2005.