1 SUPPORTING INFORMATION

2	On the spatial distribution and evolution of ultrafine particles in Barcelona
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Fig. S1. Monitoring sites in Barcelona, Spain



2 Fig. S2. Temporal trends of the particle number concentrations measured during the SAPUSS





Fig. S3. Temporal trends of the aerosol size distributions during the SAPUSS project at four
different monitoring sites.





Fig. S4. Vertical profiles of RH, temperature, wind direction, wind speed and air pressure obtained with radio sounding balloons performed twice per day (at 12:00 UTC and 00:00 UTC) at the UB site throughout the SAPUSS campaign (64 balloons in total). Enhanced in colour are the profiles obtained during the regional (blue), urban (red) and regional background only (green) nucleation events. Wind data for day 17/10/10 were not available.







NOAA HYSPLIT MODEL Backward trajectories ending at 1800 UTC 17 Oct 10 GDAS Meteorological Data





Fig. S5. Diurnal profile of the boundary layer and air mass back trajectories during the (a)
regional nucleation event on 25th September 2010, (b) regional background nucleation on 17th
October 2010 and (c) the urban nucleation event on 5th October 2010.



Fig. S6. Aerosol size distributions shown during the urban nucleation event (05/10/2010),
starting at about 15:00 in the city (shown for simplicity as an average between RS and UB),
moving to the urban background TC and reaching the regional background site (RB) at about
18:00 as a plume of ultrafine particles grown to about 50 nm (SMPS size distributions shown
also in Fig. 6).





- 1 Fig. S7. Linear correlation between BC (ng m^{-3}) and N (cm⁻³) at four different monitoring
- 2 sites for each hour of the day.



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5 **Fig. S8**. Gradient and intercept binned at six hour intervals obtained from the linear 6 correlations of Fig. S6.

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Fig S9. Percentage of primary non volatile traffic particles (% Traffic) at the four different
monitoring sites.

1 Calculation of the Condensation Sink

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3 The condensation sink (CS) describes how rapidly condensable vapour molecules will 4 condense on the existing aerosol. Specifically this quantity describes the loss rate of 5 molecules with diameter d_p , diffusion coefficient D, and mean free path λ_v onto a distribution 6 $n(d_p)$ (or N_i in the discrete case) of existing particles and as such, can be obtained from 7 integrating over the particle size spectrum (Dal Maso *et al.* 2002).

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9

$$CS = 2\pi D \int_{0}^{\infty} d_{p} \beta_{M} (d_{p}) n(d_{p}) dd_{p}$$

$$= 2\pi D \sum_{i}^{o} \beta_{M_{i}} d_{p,i} N_{i}$$
(1)

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12 The transitional correction factor β_{M} can be expressed as (Fuchs *et al.* 1971)

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14
$$\beta_{M} = \frac{Kn+1}{0.377Kn+1+\frac{4}{3}\alpha^{-1}Kn^{2}+\frac{4}{3}\alpha^{-1}Kn}$$
(2)

15

16 Where Kn is the Knutson number $\text{Kn} = 2\lambda_v/d_p$ and α is the mass accommodation coefficient. 17

18 In Eq. (2), α is assumed to be 1 while Kn can be expressed in terms of particle diameter and 19 the mean free path of vapour molecules (Pirjola *et al.* 1999) as

20

21
$$Kn = \frac{2\lambda_{\nu}}{d_{p}}$$
(3)

1 The mean free path λ_v is pressure and temperature dependent and can be determined from the 2 following formula from (Willeke 1976)

3

$$\lambda_{\nu} = \lambda_r \left(\frac{101}{P}\right) \left(\frac{T}{293}\right) \left(\frac{1 + \frac{110}{293}}{1 + \frac{110}{T}}\right) \tag{4}$$

5

4

6 Where P is in kPa and T in K.

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9 Dal Maso, M., M. Kulmala, et al. (2002). "Condensation and coagulation sinks and formation
10 of nucleation mode particles in coastal and boreal forest boundary layers." J. Geophys.
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12 Fuchs, N. A. and A. G. Sutugin (1971). <u>Highly Dispersed Aerosols</u>. New York, Pergamon.

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