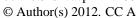
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(a)



# Peroxy radical chemistry and OH radical production during the NO<sub>3</sub>-initiated oxidation of isoprene

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**Abstract.** Peroxy radical reactions  $(RO_2 + RO_2)$  from the NO<sub>3</sub>-initiated oxidation of isoprene are studied with both gas chromatography and a chemical ionization mass spectrometry technique that allows for more specific speciation of products than in previous studies of this system. We find high nitrate yields ( $\sim 80\%$ ), consistent with other studies. We further see evidence of significant hydroxyl radical (OH) formation in this system, which we propose comes from  $RO_2 + HO_2$  reactions with a yield of  $\sim 38-58$  %. An additional OH source is the second generation oxidation of the nitrooxyhydroperoxide, which produces OH and a dinitrooxyepoxide with a yield of  $\sim 35$  %. The branching ratio of the radical propagating, carbonyl- and alcohol-forming, and organic peroxide-forming channels of the RO<sub>2</sub> + RO<sub>2</sub> reaction are found to be  $\sim$  18–38 %,  $\sim$  59–77 %, and  $\sim$  3–4 %, respectively. HO<sub>2</sub> formation in this system is lower than has been previously assumed. Addition of RO<sub>2</sub> to isoprene is suggested as a possible route to the formation of several isoprene C<sub>10</sub>-organic peroxide compounds (ROOR). The nitrooxy, allylic, and C<sub>5</sub> peroxy radicals present in this system exhibit different behavior than the limited suite of peroxy radicals that have been studied to date.

### 1 Introduction

The global emissions of isoprene (440–660 Tg yr<sup>-1</sup>) (Guenther et al., 2006) are larger than those of any other nonmethane hydrocarbon. Because of its high abundance and reactivity towards atmospheric radicals, isoprene plays a major role in the oxidative chemistry of the troposphere (Chameides et al., 1988; Williams et al., 1997; Roberts et al., 1998; Horowitz et al., 1998; Paulot et al., 2009a) and is an important precursor for secondary organic aerosol (SOA) (Claeys et al., 2004; Kroll et al., 2005, 2006; Surratt et al., 2006, 2010; Carlton et al., 2009).

Nitrate radicals (NO<sub>3</sub>), which form primarily from the reaction of NO2 and O3, are likely the dominant oxidant of isoprene at night when photochemical production of hydroxyl radicals (OH) ceases. Although nighttime isoprene emissions are negligible (Sharkey et al., 1996; Harley et al., 2004), isoprene emitted late in the day, as OH concentrations drop, remains in the nighttime atmosphere (Starn et al., 1998; Stroud et al., 2002; Warneke et al., 2004; Steinbacher et al., 2005; Brown et al., 2009). The rate constant for isoprene's reaction with NO<sub>3</sub> is  $\sim 50\,000$  times higher than that of its reaction

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with O<sub>3</sub>, the other major nighttime oxidant (Atkinson, 1997). Assuming an NO<sub>3</sub> mixing ratio of 10 ppt and an O<sub>3</sub> mixing ratio of 40 ppb, oxidation of isoprene by NO<sub>3</sub> will proceed more than an order of magnitude faster than that by O<sub>3</sub>. Mixing ratios of NO<sub>3</sub> in the nighttime continental boundary layer generally exceed 10 ppt, being in the range of 10–100 ppt (Platt and Janssen, 1995; Smith et al., 1995; Heintz et al., 1996; Carslaw et al., 1997), though concentrations on the order of several hundred ppt have been reported (Platt et al., 1981; von Friedeburg et al., 2002; Brown et al., 2006; Penkett et al., 2007).

During the day,  $NO_3$  is efficiently destroyed by photolysis and reaction with NO (Wayne et al., 1991), but significant daytime concentrations have been measured under conditions of sufficient  $O_x$  ( $O_x = O_3 + NO_2$ ) and low actinic flux.  $NO_3$  has been shown to reach concentrations of  $\sim 1$  ppt and be responsible for  $\sim 10$ % of total isoprene oxidation in the daytime under clouds or in a forest canopy (Brown et al., 2005; Forkel et al., 2006; Fuentes et al., 2007). In Houston, with large concentrations of both  $NO_x$  and  $O_3$ ,  $NO_3$  concentrations between 5–30 ppt in the hours before sunset have been measured (Geyer et al., 2003a).

The reaction of isoprene and NO<sub>3</sub> can be significant to atmospheric carbon and nitrogen budgets - and subsequently ozone formation – particularly on a regional scale. Globally, it is estimated that the isoprene + NO<sub>3</sub> reaction is responsible for  $\sim$  6–7 % of total isoprene oxidation (Horowitz et al., 2007; Ng et al., 2008) and  $\sim 15\%$  of oxidized nitrogen consumption (Brown et al., 2009). Field studies in the northeastern United States, which has a mix of NO<sub>x</sub> and isoprene sources, find that  $\sim$  22 % of isoprene oxidation in the residual daytime boundary layer,  $\sim 40$  % of isoprene oxidation in airmasses advected offshore within the marine boundary layer, and  $\sim 73$  % of NO<sub>3</sub> consumption can be attributed to this reaction (Warneke et al., 2004; Brown et al., 2009). In addition, the isoprene + NO<sub>3</sub> reaction is likely an important source of isoprene nitrates, which are significant NO<sub>x</sub>-reservoir compounds affecting regional ozone formation (von Kuhlmann et al., 2004; Fiore et al., 2005; Horowitz et al., 1998, 2007).

The oxidation mechanism and products of the isoprene + NO<sub>3</sub> reaction have been the subject of numerous studies (Jay and Stieglitz, 1989; Barnes et al., 1990; Skov et al., 1992; Kwok et al., 1996; Berndt and Boge, 1997; Suh et al., 2001; Zhang and Zhang, 2002; Fan and Zhang, 2004; Ng et al., 2008; Perring et al., 2009; Rollins et al., 2009). The initial step in the reaction is NO<sub>3</sub> addition to one of the double bonds, followed by addition of O<sub>2</sub> to make a nitrooxyalkyl peroxy radical (RO<sub>2</sub>). The RO<sub>2</sub> radicals then react with NO<sub>2</sub> (to make short-lived peroxynitrate compounds), NO<sub>3</sub>, HO<sub>2</sub>, NO<sub>2</sub>, or another RO<sub>2</sub>, leading to a variety of 1st generation products (Fig. 1). We neglect RO<sub>2</sub> reactions with NO as NO concentrations are generally very low at night in the remote environments where this reaction is most likely to occur (and low under our experimental conditions (Sect. 2) due to the rapid reaction  $NO_3 + NO \rightarrow 2NO_2$ ).

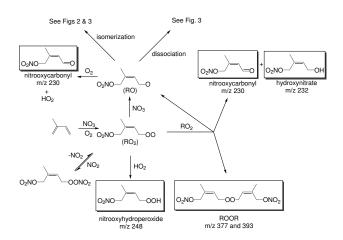


Fig. 1. Generalized reaction mechanism of the isoprene +  $NO_3$  reaction. Boxed compounds are detected by CIMS instrument as  $CF_3O^-$  adducts at the indicated m/z values.

In a previous study (Ng et al., 2008), we show that the SOA yield from the reaction of isoprene with NO $_3$  radicals is higher when experimental conditions favor RO $_2$  + RO $_2$  reactions over RO $_2$  + NO $_3$  reactions. This phenomenon is explained in part by the formation of low vapor pressure C $_{10}$ -organic peroxides (ROOR), a product channel that had previously been considered insignificant. In light of the potential importance of RO $_2$  + RO $_2$  reactions, we present here a detailed product study of the RO $_2$  + RO $_2$  reactions from the NO $_3$ -initiated oxidation of isoprene.

Our study also requires analysis of  $RO_2 + HO_2$  reactions, which inevitably occur in this system. Such reactions are generally considered to form peroxides (ROOH), but there is a growing body of work showing that, for certain  $RO_2$ , other product channels are significant, in particular the channel leading to the formation of hydroxyl radical (OH) (Hasson et al., 2004, 2012; Jenkin et al., 2007, 2008, 2010; Dillon and Crowley, 2008; Birdsall et al., 2010; Birdsall and Elrod, 2011). Since isoprene +  $NO_3$  reactions occur when there is limited photochemical production of OH, such a channel may play an important role in determining the oxidative capacity of the nighttime atmosphere.

### 2 Experimental

This work presents a detailed product study of the "excess isoprene" experiment discussed in Ng et al. (2008). The thermal decomposition of  $N_2O_5$  serves as the source of  $NO_3$  radicals.  $N_2O_5$  is synthesized by mixing streams of nitric oxide ( $\geq$  99.5%, Matheson Tri Gas) and ozone in a glass bulb, which forms  $N_2O_5$  via the following reactions (Davidson et al., 1978):

$$NO + O_3 \rightarrow NO_2 + O_2 \tag{R1}$$

$$NO_2 + O_3 \rightarrow NO_3 + O_2 \tag{R2}$$

$$NO_2 + NO_3 \leftrightarrow N_2O_5$$
 (R3)

Ozone is generated by flowing oxygen through an ozonizer (OREC V10-0); its mixing ratio is found to be  $\sim 2\,\%$  as measured by a UV/VIS spectrometer (Hewlett-Packard 8453). The flow rate of nitric oxide into the glass bulb is adjusted until the brown color in the bulb disappears. The  $N_2O_5$  is trapped for  $2\,h$  in an acetone-dry ice bath at approximately  $-80\,^{\circ}\text{C}$ , cold enough to trap  $N_2O_5$  but not  $O_3$ , as condensed  $O_3$  can explode upon warming. After synthesis, the bulb containing the  $N_2O_5$ , a white solid, is stored in a liquid nitrogen dewar.

Experiments are performed in the Caltech dual 28 m<sup>3</sup> Teflon chambers (Cocker et al., 2001; Keywood et al., 2004). O<sub>3</sub> (Horiba, APOA 360), NO and NO<sub>2</sub> (Horiba, APNA 360), and temperature and relative humidity (RH) (Vaisala, HMP 233) are continuously monitored. The chambers are maintained in the dark at room temperature ( $\sim 20$ – 21 °C) under dry conditions (RH < 10 %). Prior to an experiment, the chambers are continuously flushed for at least 24 h. The  $N_2O_5$  is removed from the liquid nitrogen and vaporizes into an evacuated 500 ml glass bulb, the pressure in which is continuously monitored by a capacitance manometer (MKS). Once a sufficient pressure of N<sub>2</sub>O<sub>5</sub> has been achieved in the bulb, the bulb's contents are flushed into the chamber with a  $51\,\mathrm{min}^{-1}$  air stream. After waiting  $\sim 1\,\mathrm{h}$  to allow the  $N_2O_5$  to become well-mixed in the chamber, a known volume of isoprene (Aldrich, 99 %) is injected into a glass bulb and flushed into the chamber with a 51 min<sup>-1</sup> dry air stream, which initiates the reaction.

The amount of isoprene added corresponds to a mixing ratio in the chamber of  $\sim 800\, ppb$ , while the  $N_2O_5$  concentration is  $\sim 150\, ppb$ . The large excess of hydrocarbon with respect to  $N_2O_5$  maximizes peroxy radical self- and cross-reactions and minimizes  $NO_3$  reactions with both peroxy radicals and stable first generation products (i.e., species other than isoprene). This excess is magnified by adding the hydrocarbon after the  $N_2O_5$  is well-mixed in the chamber: within the injected plume, hydrocarbon concentrations will be much greater than  $800\, ppb$ .

An Agilent 6890N gas chromatograph with flame ionization detector (GC-FID) measures isoprene and the oxidation products methyl vinyl ketone, methacrolein, and 3-methylfuran. The GC-FID, equipped with a bonded polystyrene-divinylbenzene based column (HP-Plot Q,  $15 \, \text{m} \times .53 \, \text{mm}$ ,  $40 \, \mu \text{m}$  thickness, J&W Scientific), is held at  $60 \, ^{\circ}\text{C}$  for  $0.5 \, \text{min}$ , then ramped at  $35 \, ^{\circ}\text{C} \, \text{min}^{-1}$  to  $200 \, ^{\circ}\text{C}$ , after which the temperature is held steady for  $3.5 \, \text{min}$ .

The other gas phase products reported here are monitored with a custom-modified Varian 1200 chemical ionization mass spectrometer (CIMS) (Ng et al., 2007; Paulot et al., 2009b; St. Clair et al., 2010), which selectively clusters CF<sub>3</sub>O<sup>-</sup> with compounds having a high fluoride affinity (e.g., acids, peroxides, and multifunctional nitrooxy- and hydroxy-compounds), forming ions detected at

m/z MW + 85 (Crounse et al., 2006). The quadrupole mass filter scans from m/z 50 to m/z 425, with a dwell time of 0.5 s per mass. The CIMS enables more specific speciation of organic nitrates than other techniques that have been employed to study the isoprene + NO<sub>3</sub> system: Fourier transform infrared (FT-IR) (Barnes et al., 1990; Skov et al., 1992; Berndt and Boge, 1997), thermal dissociation-laser induced fluorescence (TD-LIF) (Perring et al., 2009; Rollins et al., 2009), and proton transfer reaction mass spectrometry (PTR-MS) (Kwok et al., 1996; Perring et al., 2009; Rollins et al., 2009). FT-IR and TD-LIF measure the amount of a certain functionality (e.g., nitrates), but in complex mixtures it is difficult to distinguish compounds sharing a common functional group (e.g., nitrooxycarbonyls and hydroxynitrates). The PTR-MS allows for identification of individual compounds, but does so with significant fragmentation and water clustering, which leads to complex mass spectra and an increased probability of mass analogs. In contrast, the CIMS technique does not lead to significant fragmentation or water clustering under these experimental conditions, which simplifies interpretation of mass spectra.

Because authentic standards for the major products are unavailable, we estimate the sensitivity of the CIMS to these products using the empirical method of Su and Chesnavich (1982). This method estimates the collision rate of CF<sub>3</sub>O<sup>-</sup> and an analyte based on the analyte's dipole moment and polarizability. We calculate the conformationally averaged dipole moment and polarizability of the analytes with the Spartan06 quantum package using molecular structures optimized with the B3LYP/6-31G(d) method. While this theoretical approach compares favorably with experimentally derived sensitivities for many compounds (Garden et al., 2009; Paulot et al., 2009b,a), it represents the largest source of uncertainty (±25 %) for the CIMS data.

### 3 Results and discussion

Because the isoprene + NO $_3$  reaction is rapid, the low time resolution of our measurements (one measurement every  $\sim 12\,\mathrm{min}$  for the GC-FID and  $\sim 8\,\mathrm{min}$  for the CIMS) allows us to determine only the final product distribution (Table 1). The molar yields in Table 1 vary slightly from those reported in Ng et al. (2008) due to refinements in the estimated CIMS sensitivity, but these changes do not significantly alter the conclusions drawn in our earlier work. Due to the computational cost of estimating the conformationally averaged dipole and polarizability of large molecules, we have assumed that the CIMS has the same sensitivity to all of the C9 and C10 compounds.

The only species for which we see time dependent signals are the ROOR  $C_{10}$ -organic peroxide compounds (CIMS m/z 332, 377, and 393), which reach peak signals 1–3 h after the reaction is initiated, followed by a slow decay. This behavior is likely because these compounds have low vapor

Table 1. Products detected by GC-FID and CIMS.

Compound	Method	m/z (CIMS)	Final concentration (ppb) <sup>a</sup>	Percent Yield (%) <sup>b</sup>
C <sub>4</sub> non-nitrate compounds				
MACR	GC-FID	_	3	2.3
MVK	GC-FID	_	6	4.7
C <sub>4</sub> -hydroxycarbonyl	CIMS	171	< 0.5	$\sim 0$
C <sub>5</sub> Nitrates				
C <sub>5</sub> -nitrooxycarbonyl	CIMS	230	45.7	35.6
C <sub>5</sub> -hydroxynitrate	CIMS	232	27.5	21.4
C <sub>5</sub> -nitrooxyhydroperoxide	CIMS	248	12.5	9.7
C <sub>4</sub> /C <sub>5</sub> Isomerized nitrates				
C <sub>5</sub> -nitrooxyhydroxycarbonyl	CIMS	246	5.5	4.3
C <sub>5</sub> -nitrooxydiol	CIMS	248	3.3	2.6
C <sub>5</sub> -nitrooxyhydroxyhydroperoxide	CIMS	264	2.1	1.6
C <sub>4</sub> -nitrooxycarbonyl	CIMS	216	0.6	0.5
C <sub>5</sub> Hydroxy compounds				
C <sub>5</sub> -hydroxycarbonyl	CIMS	185	2.6	2.0
C <sub>5</sub> -diol	CIMS	187	2.3	1.8
C <sub>5</sub> -hydroxyhydroperoxide	CIMS	203	4.2	3.3
C <sub>5</sub> Isomerized hydroxy compounds				
C <sub>5</sub> -dihydroxycarbonyl	CIMS	201	1.5	1.2
C <sub>5</sub> -triol	CIMS	203	1.3	1.0
C <sub>5</sub> -dihydroxyhydroperoxide	CIMS	219	<0.5	~0
Organic peroxides				
C <sub>10</sub> -dinitrooxy ROOR	CIMS	377	1.0	0.8
C <sub>10</sub> -isomerized dinitrooxy ROOR	CIMS	393	0.6	0.5
C <sub>10</sub> -nitrooxycarbonyl ROOR	CIMS	330	< 0.5	$\sim 0$
C <sub>10</sub> -hydroxynitrate ROOR	CIMS	332	0.6	0.5
C <sub>10</sub> -nitrooxyhydroperoxide ROOR	CIMS	348	< 0.5	$\sim 0$
C <sub>9</sub> -nitrooxycarbonyl ROOR	CIMS	316	<0.5	$\sim 0$
Other				
3-MF	GC-FID	_	4.5	3.5
hydroxyacetone	CIMS	159	0.5	0.4
hydrogen peroxide	CIMS	119	5.5	4.3
glycolaldehyde	CIMS	145	0.9	0.7
Total <sup>c</sup>			128.4	

<sup>&</sup>lt;sup>a</sup> Products with small but non-zero signals are noted as <0.5 ppb.

pressures and thus interact significantly with instrument tubing or condense into secondary organic aerosol ( $\sim 10\,\mu g\,m^{-3}$  of SOA forms rapidly in this experiment). For these compounds, the reported values are the peak mixing ratios seen during the experiment.

# 3.1 Nitrate yield

C<sub>5</sub>-nitrooxycarbonyls, hydroxynitrates, and nitrooxyhydroperoxides, the major products of the isoprene + NO<sub>3</sub> reaction, are detected by the CIMS at m/z 230, 232, and 248, respectively. In addition, we see compounds appearing at m/z 216, 246, and 264, which are consistent with nitrate

b Molar yield.

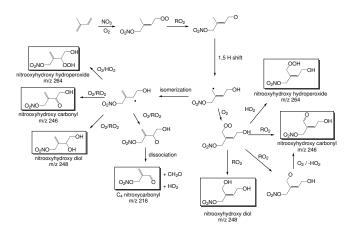
 $<sup>^{</sup>c}$  Sum of all products except hydrogen peroxide and minor signals.  $C_{10}$  compounds are counted twice as they comprise two isoprene molecules.

products resulting from the isomerization of the alkoxy (RO) radical originating from the  $\delta$ -nitrooxyperoxy radical formed by (1,4) or (4,1) addition (the notation (x,y) indicates  $NO_3$ addition to the x carbon and subsequent O2 addition to the y carbon) (Fig. 2). Previous studies have shown that (1,4) additions are dominant in this system (Skov et al., 1992; Berndt and Boge, 1997; Suh et al., 2001). Isomerization also leads to a nitrate product at m/z 248, the same mass as the nitrooxyhydroperoxide. To estimate the ratio of these two isobaric species, we assume that the alkoxy radical yield from RO<sub>2</sub> + RO<sub>2</sub> reactions is identical for both the non-isomerized and isomerized nitrooxyperoxy radical (the branching ratio of RO<sub>2</sub> + RO<sub>2</sub> is discussed further in Sect. 3.4). Finally, we see  $C_{10}$ -organic peroxides at m/z 332, 377, and 393 (further discussed in Sect. 3.6). Summing the concentrations of these nitrates (and noting that the ROOR compounds at m/z 377 and 393 sequester two nitrates), we find a total organic nitrate concentration of  $\sim 100$  ppb.

We can express the nitrate yield with respect to both reacted nitrogen or carbon. For the nitrogen-based yield, we divide the nitrate concentration by the amount of NO<sub>3</sub> radical consumed, which is equivalent to the loss of N2O5 during this reaction. Lacking a quantitative measurement of N<sub>2</sub>O<sub>5</sub>, we use the change in NO2 concentration after the addition of isoprene ( $\sim$  125 ppb) as a proxy. Every conversion of N<sub>2</sub>O<sub>5</sub> to NO<sub>3</sub> releases NO<sub>2</sub>, but the total change in NO<sub>2</sub> may be an overestimate of total NO3 reacted because NO2 can also be released in the formation of methyl vinyl ketone (MVK), methacrolein (MACR), 3-methylfuran (3-MF), and the C<sub>5</sub> hydroxycarbonyl (Fig. 3), though in Sect. 3.2 we discuss alternative formation pathways for these compounds. Subtracting these additional NO<sub>2</sub> sources to get a lower limit for NO<sub>3</sub> consumption leads to an NO<sub>3</sub> consumption range of 109-125 ppb and a corresponding nitrate yield of  $\sim 80-90\%$  (all percentage yields in this work are calculated on a molar basis).

This high yield suggests that the NO<sub>3</sub> radical reacts with isoprene predominantly, if not exclusively, via addition to a double bond. The CIMS does not see a detectable rise in HNO<sub>3</sub>, indicating that hydrogen abstraction is not a significant pathway for this reaction (our sensitivity to HNO<sub>3</sub>, however, is hampered by a large background – probably from impurities in the N2O5 or reaction of N2O5 with trace water on the surface of the chamber). Assuming most of the 16.1 ppb of MVK, MACR, 3-MF, and the C<sub>5</sub>-hydroxycarbonyl originates from nitrooxyperoxy radicals, we can account for  $\sim 100\%$  of the reacted NO<sub>3</sub>. Additionally, although our experimental design seeks to minimize reactions of NO<sub>3</sub> with species other than isoprene, there are possible (likely small) losses of NO<sub>3</sub> from reaction with other radicals or first generation products, or heterogeneously to the chamber walls or SOA.

The measured nitrate yield with respect to  $NO_3$  is consistent with the substantial yields determined by other studies:  $\sim 95\%$  (under NO-free conditions) (Berndt and Boge, 1997),



**Fig. 2.** Formation mechanism of compounds resulting from the isomerization of alkoxy radicals and measured by the CIMS at m/z 216, 246, 248, and 264. This figure assumes initial NO<sub>3</sub> attachment to the 1-carbon and formation of an (E)-δ-peroxy radical, but other isomers are possible.

 $57 \pm 11$  % (Perring et al., 2009), and  $70 \pm 8$  % (Rollins et al., 2009). Variance in yields with different experimental methods is not surprising because they depend on the relative concentrations of different radicals, as well as physical loss and mixing processes, which are unique to each work. Furthermore, the final product distribution is a strong function of the distribution of peroxy radical isomers:  $\delta$ -nitrooxyperoxy radicals tend to maintain their nitrate functionality (with the exception of the possible formation of hydroxycarbonyl or 3-MF), while  $\beta$ -nitrooxyperoxy radicals, if they become nitrooxyalkoxy radicals, are likely to lose the nitrate to form MVK or MACR (Vereecken and Peeters, 2009). Berndt and Boge (1997) and Peeters et al. (2009) suggest that peroxy radical isomers formed from isoprene oxidation are continuously interconverting. If this is true, the degree of interconversion is affected by the rate at which RO<sub>2</sub> become stable products relative to the interconversion rate, i.e., the magnitudes of  $k_1$  and  $k_2$  with respect to  $k_{\text{int}1}$  and  $k_{\text{int}2}$  in Fig. 4. These rates are specific to the unique experimental conditions of each study, such as temperature, pressure, the degree of mixing, and hydrocarbon and oxidant concentrations. Therefore, the distribution of isomers – which defines the final product distribution – may be sensitive to specific experimental conditions.

To calculate the nitrate yield with respect to carbon, we divide the concentration of nitrates by the amount of isoprene reacted. Because a portion of the isoprene reacts immediately upon introduction into the chamber, we do not know the exact starting isoprene concentration. Therefore, we assume that each of the products listed in Table 1 comes from one isoprene molecule, with the exception of the ROOR compounds (which comprise two isoprene molecules) and hydrogen peroxide (which comprises zero). This leads to an estimate of  $\sim 130\,\mathrm{ppb}$  of isoprene reacted, and a nitrate yield of  $\sim 80\,\%$ .

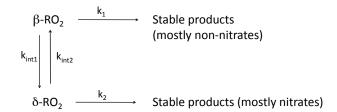
**Fig. 3.** Formation mechanisms of methyl vinyl ketone (a), methacrolein (b), 3-methylfuran (c), and hydroxycarbonyl (d), leading to release of NO<sub>2</sub>. The exact mechanism of 3-methylfuran formation is still uncertain (Francisco-Marquez et al., 2005).

As with the nitrogen-based yield, this result too is consistent with other studies:  $\sim$  80 % (Barnes et al., 1990),  $\sim$  90 % (Berndt and Boge, 1997),  $70\pm8$  % (Rollins et al., 2009), and  $65\pm12$  % (Perring et al., 2009).

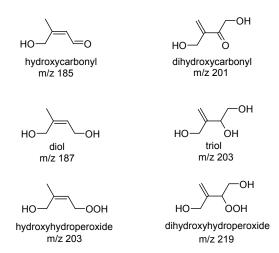
Much of the discrepancy between our estimates of isoprene and NO<sub>3</sub> consumption is likely due to our lack of an empirical calibration for the CIMS. Some of it, however, is due to an additional loss process of isoprene besides oxidation by NO<sub>3</sub>, which we discuss in the following section.

### 3.2 Hydroxyl radical (OH) formation

The CIMS detects the formation of products at m/z 185, 187, 203, and 201, which are indicative of compounds at MW 100, 102, 118, and 116, respectively. These compounds are analogous to those depicted in Figs. 1 and 2, only with oxidation initiated by the hydroxyl radical (OH) (Fig. 5) (Surratt et al., 2010); the relative contribution of isobaric species is determined in the same manner as in Sec. 3.1. Some of the signal at m/z 201 may also be attributable to C<sub>5</sub>hydroperoxyaldehydes, which have recently been reported to result from OH oxidation (Crounse et al., 2011). Perring et al. (2009) report PTR-MS signals at m/z 101, 103, 119, and 117, which could be the protonated clusters of these compounds, though they attribute the latter three m/z to water clusters of other major product ions. Under the dry conditions of our experiment, however, we do not typically observe water clusters with, or significant fragmentation of, our product ions, so we are confident that the signals on the CIMS in fact represent hydroxy compounds. OH formation may also contribute to some or all of the MVK and MACR produced in our system, though it is likely that most of the 3-MF comes



**Fig. 4.** Schematic of the relationship between the interconversion of peroxy radical isomers and nitrate yields.



**Fig. 5.** Products detected by CIMS that may result from the OH-initiated oxidation of isoprene. Other isomers are possible.

from isoprene + NO<sub>3</sub> reactions because its yield in the isoprene + OH system is low (Ruppert and Becker, 2000; Paulot et al., 2009b).

We evaluate five possible routes to OH formation in our system: reactions of (i) O<sub>3</sub> and isoprene (Neeb and Moortgat, 1999), (ii) HO<sub>2</sub> and O<sub>3</sub> (Sinha et al., 1987), (iii) HO<sub>2</sub> and NO (Seeley et al., 1996), (iv) HO2 and NO3 (Mellouki et al., 1993), and (v) RO<sub>2</sub> and HO<sub>2</sub> (Hasson et al., 2004, 2005; Jenkin et al., 2007, 2008, 2010; Dillon and Crowley, 2008). Routes (i) and (ii) are unlikely to be significant sources of OH in our experiments. Not only does our O3 monitor not detect any ozone during the experiment (limit of detection  $\sim 2 \,\mathrm{ppb}$ ), but we also see no evidence in the CIMS data of significant organic acid or peroxide formation, which would result from the reaction of O<sub>3</sub> with isoprene (Hasson et al., 2001; Orzechowska and Paulson, 2005). Furthermore, for route (ii) to be feasible,  $HO_2 + O_3$  reactions  $(k = 1.9 \times 10^{-15} \text{ cm}^3 \text{ molec}^{-1} \text{ s}^{-1} \text{ at } 298 \text{ K}, \text{ Sander et al.},$ 2011) must be significantly faster than HO<sub>2</sub>+HO<sub>2</sub> reactions  $(k = 2.3 \times 10^{-12} \text{ cm}^3 \text{ molec}^{-1} \text{ s}^{-1} \text{ at 1 atm and 298 K},$ Sander et al., 2011), which produce ppb levels of H<sub>2</sub>O<sub>2</sub> in the system (Table 1). This would require O<sub>3</sub> to be more than three orders of magnitude more abundant than HO<sub>2</sub>, i.e.,

**Table 2.** Reactions considered for box model assessment of OH sources in the isoprene + NO<sub>3</sub> system.

No.	Reaction	Rate constant <sup>a</sup>	Source
1 <sup>b</sup>	$NO_3 + isoprene \rightarrow RO_2 + HO_2$	$6.6 \times 10^{-13}$	Atkinson (1997)
2	$RO_2 + RO_2 \rightarrow products$	$1 \times 10^{-13}$	Atkinson et al. (2006), and references therein
3	$RO_2 + NO_3 \rightarrow products$	$3 \times 10^{-12}$	Biggs et al. (1994); Daele et al. (1995);
			Canosa-Mas et al. (1996); Vaughan et al. (2006)
4	$RO_2 + HO_2 \rightarrow products$	$2.2 \times 10^{-11}$	Atkinson et al. (2006), and references therein
5	$RO_2 + NO_2 \rightarrow RO_2NO_2$	$5 \times 10^{-11}$	Sander et al. (2011), and references therein
6	$RO_2NO_2 \rightarrow RO_2 + NO_2$	20	Sander et al. (2011), and references therein
7	$NO_3 + HO_2 \rightarrow OH + NO_2 + O_2$	$3.5 \times 10^{-12}$	Sander et al. (2011), and references therein
8	$NO_2 + NO_3 \rightarrow N_2O_5$	$6.7 \times 10^{-12}$	Sander et al. (2011), and references therein
9	$N_2O_5 \rightarrow NO_2 + NO_3$	$2.2 \times 10^{-1}$	Sander et al. (2011), and references therein
10	$HO_2 + HO_2 \rightarrow H_2O_2$	$2.3 \times 10^{-12}$	Sander et al. (2011), and references therein
11	$HO_2 + NO_2 \rightarrow HO_2NO_2$	$2.8 \times 10^{-12}$	Sander et al. (2011), and references therein
12	$HO_2NO_2 \rightarrow HO_2 + NO_2$	$1.8 \times 10^{-1}$	Sander et al. (2011), and references therein
13	$NO_3 + NO_2 \rightarrow NO + NO_2 + O_2$	$6.6 \times 10^{-16}$	Sander et al. (2011), and references therein
14	$NO_3 + NO \rightarrow 2NO_2$	$2.6 \times 10^{-11}$	Sander et al. (2011), and references therein
15	$HO_2 + NO \rightarrow NO_2 + OH$	$8.0 \times 10^{-12}$	Sander et al. (2011), and references therein

<sup>&</sup>lt;sup>a</sup> At 1 atm and 298 K. Units are cm<sup>3</sup> molec<sup>-1</sup> s<sup>-1</sup>, except  $k_6$ ,  $k_9$  and  $k_{12}$ , which are s<sup>-1</sup>. Reaction rates involving RO<sub>2</sub> are approximated from values found in the literature.

at ppm levels that cannot come from trace contamination of the chamber.

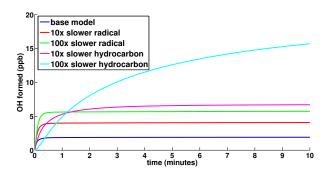
To examine the remaining hypotheses, we create a box model incorporating the major reactions in the system for developing a qualitative understanding of which processes may be important for the final product yield. Table 2 lists the parameters of this box model; for rate constants that have not been experimentally determined, we use estimates based on the rate constants of similar reactions found in the literature, but caution that the actual rate constants may differ significantly. Initial conditions reflect the nominal concentration of reagents in the chamber: [isoprene] = 800 ppb,  $[N_2O_5] = 125 \text{ ppb}$ , and  $[NO_2] = 50 \text{ ppb}$  (the NO<sub>2</sub> likely results from decomposition of N<sub>2</sub>O<sub>5</sub> prior to isoprene injection). In reality, though, the isoprene concentration is higher than 800 ppb during the reaction because of our injection method. As discussed later (Sect. 3.4), there are major uncertainties in the HO2 sources and magnitudes, so for the purposes of assessing possible OH sources, we assume as an upper limit that the formation rate of HO<sub>2</sub> is the same as that of RO<sub>2</sub> in Eq. (1) of Table 2; our final concentration of peroxides (i.e., [ROOH] +  $2 \times [H_2O_2]$ ) is  $\sim 29$  ppb, much less than the  $\sim 109-125$  ppb of RO<sub>2</sub> that is formed (Sect. 3.1), suggesting that the formation of HO<sub>2</sub> is significantly less than that of  $RO_2$ .

The box model shows that the NO levels in the chamber are too low to sustain substantial OH formation via route (iii). The  $NO_x$  monitor measures <1 ppb of NO throughout our experiment, and any NO that may exist prior to the experiment (or as a trace impurity in the  $N_2O_5$ ) reacts quickly

with NO<sub>3</sub> after N<sub>2</sub>O<sub>5</sub> injection; the NO lifetime is  $\sim 1$  s with our N<sub>2</sub>O<sub>5</sub> loading. Although NO may be generated as a minor channel of the NO<sub>2</sub> + NO<sub>3</sub> reaction, the rapid reaction of NO and NO<sub>3</sub> limits the steady state concentration of NO to  $<\sim 4$  ppt; at this concentration, NO cannot compete with other radicals reacting with HO<sub>2</sub> (i.e., RO<sub>2</sub>, HO<sub>2</sub>, NO<sub>3</sub>, and NO<sub>2</sub>). Therefore, HO<sub>2</sub> + NO is unlikely to contribute significantly to the  $\sim 12$ –21 ppb of OH that is formed in our system.

The box model also suggests that route (iv) is not feasible because of the substantial difference in the rates of the  $NO_3$  + isoprene and  $NO_3$  +  $HO_2$  reactions, both of which are well established experimentally. Under the base conditions of our box model in Table 2, which significantly overestimates the prevalence of HO<sub>2</sub> and underestimates the concentration of isoprene in the plume, less than 1 % of the NO<sub>3</sub> reacts with HO<sub>2</sub>, while 94 % reacts with isoprene and the rest with RO<sub>2</sub>. Therefore, while there is significant uncertainty with the  $RO_2 + HO_2$ ,  $RO_2 + RO_2$ , and  $RO_2 + NO_3$  rate constants, the frequency of the NO<sub>3</sub> + HO<sub>2</sub> reaction predicted by the model is very insensitive to these rates because NO<sub>3</sub> reactivity is dominated by its reaction with isoprene. Even if we favor  $NO_3 + HO_2$  reactions by reducing the  $RO_2 + HO_2$  and  $RO_2 + NO_3$  rate constants by a factor of 100, we only obtain  $\sim$  5 ppb of OH formation; in contrast, lowering the NO<sub>3</sub> + isoprene rate constant would lead to significantly more production of OH via NO<sub>3</sub> + HO<sub>2</sub> (Fig. 6). These simulations are consistent with the observation of Atkinson et al. (1988) during hydrocarbon + NO<sub>3</sub> kinetics studies that there is OH formation when slower reacting hydrocarbons are studied. The reaction of isoprene with NO3 is sufficiently fast under our

<sup>&</sup>lt;sup>b</sup> HO<sub>2</sub> yield is an upper limit to facilitate model analysis.



**Fig. 6.** Box model simulations for OH production in isoprene + NO<sub>3</sub> system. Blue: base case described in Table 2; Red: RO<sub>2</sub> + HO<sub>2</sub> and RO<sub>2</sub> + NO<sub>3</sub> rate constants reduced by factor of 10; Green: RO<sub>2</sub> + HO<sub>2</sub> and RO<sub>2</sub> + NO<sub>3</sub> rate constants reduced by factor of 100; Pink: isoprene + NO<sub>3</sub> rate constant reduced by factor of 10; Light Blue: isoprene + NO<sub>3</sub> rate constant reduced by factor of 100. Initial conditions: 150 ppb N<sub>2</sub>O<sub>5</sub>, 800 ppb isoprene, 50 ppb NO<sub>2</sub>.

experimental conditions, however, that such behavior should not occur.

We therefore suggest that formation of OH radicals most likely results from the reaction of  $RO_2$  and  $HO_2$  radicals. Quantifying the branching ratio of the  $RO_2 + HO_2$  reaction, however, is not trivial. There are four documented pathways for the  $RO_2 + HO_2$  reaction:

$$RO_2 + HO_2 \rightarrow ROOH + O_2 \tag{R4}$$

$$RO_2 + HO_2 \rightarrow ROH + O_3 \tag{R5}$$

$$RO_2 + HO_2 \rightarrow RO + OH + O_2 \tag{R6}$$

$$RO_2 + HO_2 \rightarrow R'CHO + H_2O + O_2 \tag{R7}$$

Channel (R4) can be quantified with CIMS measurements of peroxides. We neglect channel (R5), first because we don't see any evidence for ozone formation, and also because this channel is believed to proceed via a hydrotetroxide intermediate that only yields  $O_3$  if  $RO_2$  is an acylperoxy radical (RC(O)OO) (Hasson et al., 2005). To quantify channel (R6), we can use the sum of OH products as a tracer, but MVK, MACR, and the  $C_5$ -hydroxycarbonyl can come from either OH or  $NO_3$ , which leads to uncertainty in this quantity. Similarly, the nitrooxycarbonyl can come directly from channel (R7), indirectly from the RO formed in channel (R6), or from  $RO_2 + RO_2$ . Because multiple pathways share common products, and lacking more knowledge about these individual pathways, we cannot unambiguously constrain the  $RO_2 + HO_2$  branching ratios with the available data.

Recognizing the uncertainties, we estimate the OH yield from  $RO_2 + HO_2$  but emphasize that our assumptions and results must be verified by further studies. We assume channel (R7) is negligible, as well as OH from  $RO_2 + HO_2$  reactions where the  $RO_2$  originates from isoprene + OH (Paulot

et al., 2009a). We thus constrain the range of OH formation to 9–20.5 ppb, with the upper limit incorporating all the hydroxy products plus MVK and MACR, and the lower limit being the upper limit minus MVK, MACR, and the hydroxycarbonyl. We estimate channel (R4) by the concentration of the nitrooxyhydoperoxides at m/z 248 and m/z 264, so obtain a range for (R6)/[(R6)+(R4)] of between 9/(9+12.5+2.1) and 20.5/(20.5+12.5+2.1), or 38–58%. Because this analysis assumes that  $RO_2+HO_2$  reactions are the exclusive source of OH radicals and also ignores channel (R7), this yield should be considered an upper limit. Also, this yield is for the isomeric mix of  $RO_2$  in this system, which is dominated by  $RO_2$  from (1,4) additions, but also contains other isomers.

To our knowledge, this is the first study that has attempted to quantify the OH yield from RO<sub>2</sub> + HO<sub>2</sub> reactions involving the nitrooxyperoxy radicals in our system. Thus far, significant OH yields (15-80%) have been found for acylperoxy (RC(O)OO), methoxymethylperoxy (CH<sub>3</sub>OCH<sub>2</sub>OO), and  $\alpha$ carbonylperoxy (RC(O)CH<sub>2</sub>OO) radicals, and evidence for OH formation also exists for bicyclic hydroxyperoxy radicals derived from toluene; in contrast, alkylperoxy and hydroxyalkylperoxy radicals have exhibited minimal vields (Hasson et al., 2004, 2012; Jenkin et al., 2007, 2008, 2010; Dillon and Crowley, 2008; Birdsall et al., 2010; Birdsall and Elrod, 2011). For the peroxy radicals in this study, the high OH yields may result from the presence of the electronwithdrawing nitrooxy group conjugated through the double bond, which may stabilize (i.e., lower the enthalpy of) the alkoxy radical formed by the radical propagating channel (R6), thereby making this channel more thermodynamically favored.

# 3.2.1 OH formation from 2nd generation dinitrooxyepoxide formation

While this study focuses on the first generation products from the isoprene + NO<sub>3</sub> reaction, another nighttime source of OH in the atmosphere would be the further oxidation of the nitrooxyhydroperoxide, which can produce a dinitrooxyepoxide and OH (Paulot et al., 2009a). In another experiment described in detail in Ng et al. (2008), we first add 179 ppb of isoprene to the chamber followed by three additions of  $N_2O_5$  ( $\sim 120, 50, \text{ and } 210 \text{ ppb}$ ). After the first two additions, isoprene is completely consumed, so the third aliquot leads primarily to the formation of second generation products; some second generation products may be oxidized by this third addition, but the amount of N<sub>2</sub>O<sub>5</sub> added is similar to the concentration of first generation products (which is roughly equal to the starting isoprene concentration), so such tertiary chemistry is likely to be minimal. After this third addition, the nitrooxyhydroperoxide signal drops  $\sim$  6 ppb, while the signal for the dinitrooxyepoxide (at m/z 293) rises  $\sim 2.3$  ppb, indicating that the epoxide (and OH) yield from the NO<sub>3</sub> oxidation of the nitrooxyhydroperoxide is  $\sim 35\%$ , compared to  $\sim 75\%$  for OH oxidation of isoprene hydroxyhydroperoxides (Paulot et al., 2009a). The yield in the NO<sub>3</sub> system is likely lower because the dominant first generation peroxy radical is from the (1,4) addition of NO<sub>3</sub>. Therefore, to form an epoxide the second NO<sub>3</sub> must add to the 2-carbon, creating a secondary alkyl radical, whereas the more favored addition is likely to the 3-carbon creating a more stable tertiary alkyl radical (Fig. 7). In the OH system, (1,2) and (4,3) additions in the first oxidation step are most common (Paulot et al., 2009b), which means that the epoxide forming channel involves a second OH addition to the more favored 4or 1-carbon, respectively. Another factor which may contribute to a lower yield in our experiment is the presence of large amounts of nitric acid, which may increase heterogenous loss processes for the epoxide, either on the surface of the SOA or the chamber walls. Epoxides have been shown to undergo significant reactive uptake by acidic aerosol (Eddingsaas et al., 2010; Surratt et al., 2010).

# 3.3 $RO_2 + RO_2$ branching ratio

 $RO_2 + RO_2$  reactions have three product channels (Fig. 1):

$$RO_2 + RO_2 \rightarrow R'CHO + ROH + O_2$$
 (R8)

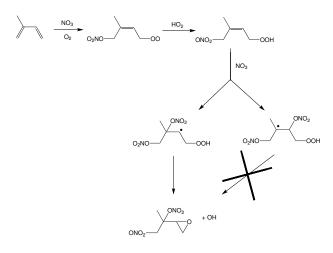
$$RO_2 + RO_2 \rightarrow ROOR + O_2 \tag{R9}$$

$$RO_2 + RO_2 \rightarrow 2RO + O_2 \tag{R10}$$

The C<sub>5</sub>-hydroxynitrate at m/z 232 comes exclusively from channel (R8), so we estimate that 55 ppb of RO<sub>2</sub> passes through this channel, leading to 27.5 ppb of hydroxynitrate and 27.5 ppb of the C<sub>5</sub>-nitrooxycarbonyl at m/z 230. Similarly, we see that  $\sim$  3 ppb of RO<sub>2</sub> follows channel (R9) to become ROOR (the sum of m/z 377 and 393, multiplied by 2 because each ROOR comprises two RO<sub>2</sub>), although some ROOR is likely in the aerosol phase as well, which we are unable to quantify.

For channel (R10), we assume that the nitrooxycarbonyl yield in excess of the hydroxynitrate yield arises from ROforming channels of either  $RO_2 + HO_2$  or  $RO_2 + RO_2$ . Thus, the total RO formation is estimated to be 33.5–45 ppb; the lower limit comprises the sum of the excess nitrooxycarbonyl, isomerized nitrates, and 3-MF, while the upper limit includes the lower limit plus MVK, MACR, and hydroxycarbonyl (assumed to come exclusively from isoprene+NO<sub>3</sub> reactions). Of the RO formation, 9-20.5 ppb comes from  $RO_2 + HO_2$  (Sect. 3.2), depending on the true provenance of MVK, MACR, and the hydroxycarbonyl. Taking into account the uncertainties involving RO, we get between 13 (i.e., 33.5–20.5) and 36 (i.e., 45–9) ppb of RO coming from  $RO_2 + RO_2$ , and summing up all the product channels yields a range of between 71 (i.e., 55+3+13) and 94 (i.e., 55 + 3 + 36) ppb of RO<sub>2</sub> undergoing RO<sub>2</sub> + RO<sub>2</sub> reactions.

With the above analysis (and neglecting possible  $RO_2 + RO_2$  reactions involving the hydroxyperoxy  $RO_2$ ), we



**Fig. 7.** Formation mechanism of dinitrooxyepoxide and hydroxyl radical from oxidation of nitrooxyhydroperoxide.

derive ranges for the  $RO_2 + RO_2$  branching ratio of 59–77 % (i.e., 55/94–55/71) for (R8), 3–4 % (i.e., 3/94–3/71) for (R9), 19–38 % (i.e., 13.5/71–36/94) for (R10).

To our knowledge, this is the first study analyzing the branching ratio of RO<sub>2</sub> + RO<sub>2</sub> reactions of isoprene nitrooxyperoxy radicals. For most peroxy radicals that have been studied, channel (R10) is typically more than 50%, while channel (R9) is generally considered negligible (Atkinson, 1997, and references therein). Ziemann (2002) proposes ROOR formation as the source of diacyl peroxides found in SOA from cycloalkene ozonolysis. Preliminary work in our laboratory has also detected ROOR compounds as products of RO<sub>2</sub> + RO<sub>2</sub> reactions from the NO<sub>3</sub>-initiated oxidation of 1,3-butadiene and 1,4-pentadiene (for the latter compound, ROOR is only apparent in the aerosol phase, M. N. Chan, personal communication, 2010), as well as the OH-initiated oxidation of 1,3-butadiene and isoprene. There remain many uncertainties regarding the mechanism of RO<sub>2</sub> + RO<sub>2</sub> reactions (Dibble, 2008), so it is difficult to assess whether reported ROOR formation (or lack thereof) is a result of the particular radicals studied or the analytical techniques employed to study their reaction. It is possible that the larger peroxy radicals we have studied are more likely to form ROOR than the smaller radicals that have been the subject of most previous work because larger RO<sub>2</sub> have more vibrational modes with which to distribute collisional energy and prevent breaking apart upon combination with another RO2 radical. Thus, ROOR formation from RO<sub>2</sub> + RO<sub>2</sub> reactions may be analogous to organic nitrate (RONO<sub>2</sub>) yield from RO<sub>2</sub> + NO reactions, which tends to increase with increasing RO<sub>2</sub> size (O'Brien et al., 1998; Arey et al., 2001; Matsunaga and Ziemann, 2009).

### 3.4 RO radical fate and HO<sub>2</sub> production

The fate of the alkoxy radical is important both for understanding the chamber studies and in nighttime chemistry as it leads to the production of  $HO_2$ . From the amount of excess hydroxycarbonyl formed, we estimate that 18 ppb of  $HO_2$  forms from  $O_2$  abstraction of RO (the formation of which we constrain to 33.5–45 ppb, Sect. 3.3), or 40–54 % of RO. This assumes that  $O_2$  abstraction from RO is the sole source of excess nitrooxycarbonyl, and that direct formation from  $RO_2 + HO_2$  reactions (R7) is negligible.

This is lower than the total sum of  $HO_2$  derived from peroxide measurements in our system (29 ppb, Sect. 3.2). Additional  $HO_2$  comes from  $O_2$  abstraction from minor alkoxy radicals (the isomerized nitrooxyalkoxy and hydroxyalkoxy radicals) and formation of the  $C_5$ -hydroperoxyaldehyde (Crounse et al., 2011), but it is not trivial to attempt an  $HO_2$  balance because of the uncertainties in both the CIMS calibration and the sources and sinks of  $HO_2$ .  $HO_2$  can come from the conversion of the nitrooxyalkoxy radical to a nitrooxycarbonyl, or MVK and MACR if the latter are from OH+isoprene; sinks of  $HO_2$  include peroxide formation,  $RO_2$ + $HO_2$  derived alkoxy radicals that do not undergo abstraction, and from nitrooxycarbonyls formed directly from  $RO_2$ + $HO_2$  via channel (R7).

Because the  $\delta$ -nitrooxyalkoxy radical, the dominant alkoxy radical in our system, can isomerize via a 1,5-H shift, the large HO<sub>2</sub> yield is somewhat surprising because isomerization reactions are typically faster than abstraction by O<sub>2</sub> (Atkinson and Arey, 2007). It is possible that the nitrooxy group limits isomerization when the  $\delta$ -nitrooxyalkoxy radical is in a Z conformation. According to the structure-activity relationship of Kwok and Atkinson (1995), H-abstraction from a carbon with an attached nitrooxy group is an order of magnitude slower than from a carbon with an attached methyl group. The nitrooxy group likely does not prevent isomerization – we see the analogous isomerized nitrates in experiments with 1,3-butadiene, for which the isomerization must abstract a hydrogen from the carbon  $\alpha$  to the nitrooxy group - but more study is required to elucidate the effect of the NO<sub>3</sub> group on isomerization rate. Large HO<sub>2</sub> yields would also be possible if  $O_2$  abstractions are faster for the alkoxy radicals in this system.

While our  $HO_2$  yield is higher than expected based on the alkoxy radical structure, it is lower than the value of 80 % of  $RO_2$  that has been used in modeling studies (Horowitz et al., 2007; Rollins et al., 2009). Therefore, models may overestimate the impact of isoprene +  $NO_3$  reactions on nighttime  $HO_x$  chemistry in this respect.

Each pathway to RO (e.g.,  $RO_2 + RO_2$ ,  $RO_2 + HO_2$ ,  $RO_2 + NO_3$ ) has a different heat of reaction, which may affect the RO fate (Berndt and Boge, 1997; Atkinson and Arey, 2007). Lacking any specific knowledge about how the RO fate in our system depends on the reaction enthalpy (and the values of the reaction enthalpies in our system), we have as-

sumed in the above analysis that every RO behaves the same, regardless of source.

# 3.5 Formation of $C_{10}$ -organic peroxides

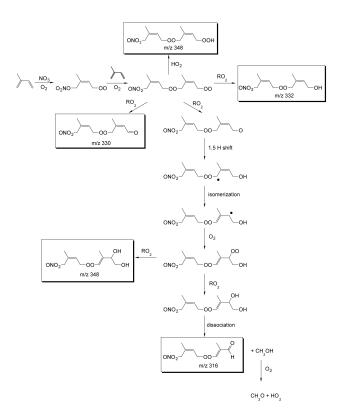
In Ng et al. (2008), we report the formation of ROOR  $C_{10}$ organic peroxide at m/z 377 and 393. Further examination of the CIMS data reveals other isoprene-based C<sub>10</sub>-organic peroxides. The most abundant of these, with a signal similar in magnitude to that of m/z 393, appears at m/z 332, which indicates a hydroxynitrate ROOR. One possible source for this compound is an  $RO_2 + RO_2$  reaction where one  $RO_2$  comes from NO<sub>3</sub> and the other from OH. The abundance of OH is much less than that of NO<sub>3</sub>, however, so it would be surprising if such a reaction would produce almost as much ROOR as the reaction between nitrooxyperoxy radicals. Alternatively, this compound may result from addition of an RO2 radical to isoprene, creating a C<sub>10</sub>-alkyl (and subsequently alkylperoxy) radical, which undergoes an RO<sub>2</sub> + RO<sub>2</sub> reaction to create the  $C_{10}$ -hydroxynitrate ROOR (Fig. 8). This mechanism for creating a C<sub>10</sub> RO<sub>2</sub> radical is analogous to the formation of bicyclic radicals by aromatic peroxy radicals (Atkinson and Arey, 2007). Small amounts of the corresponding nitrooxycarbonyl and nitrooxyhydroperoxide at m/z 330 and 348, respectively, are seen, as well as an apparent  $C_9$  compound at m/z 316 that can result from isomerization of a C<sub>10</sub>-nitrooxyalkoxy radical; isomerization may also lead to a diol at m/z 348, the same mass as the hydroperoxide.

We are not aware of any previous work that has examined external RO<sub>2</sub> addition to alkenes under atmospheric conditions, though it has been reported in both gas phase combustion (Osborne and Waddington, 1980; Stark and Waddington, 1995) and liquid phase studies (van Sickle et al., 1965a,b; Mayo, 1968; Simmons and van Sickle, 1973), producing both epoxides and polymeric peroxy radicals. This pathway may be another source of large peroxides in the ambient atmosphere.

### 4 Implications

The observed high nitrate yields, in general agreement with previous results, support the modeling results of Horowitz et al. (2007) that isoprene +  $NO_3$  reactions, while a minor sink of isoprene, are a substantial source of isoprene nitrates in the atmosphere. The formation and fate of these nitrates in turn significantly influences tropospheric  $NO_x$  and ozone.

Although we obtain similar nitrate yields relative to both reacted nitrogen and carbon, these two yields are fundamentally different quantities that coincidentally have similar magnitudes. The nitrogen-based yield (i.e., [nitrates]/[NO<sub>3</sub> consumed]) is non-unity due to competing reactions of NO<sub>3</sub> (e.g., reactions with RO<sub>2</sub>, HO<sub>2</sub>, and walls) and loss of the nitrate functionality by the initial isoprene-NO<sub>3</sub>



**Fig. 8.** Proposed formation mechanisms of products detected by CIMS at m/z 316, 330, 332, and 348. Other isomers are possible.

adduct (Fig. 3); in contrast, the carbon-based yield (i.e., [nitrates]/[isoprene reacted] or [nitrates]/[total carbon products]) is non-unity from both nitrate losses by the initial adduct and isoprene reactions with OH. For systems with extensive amounts of competing NO<sub>3</sub> sinks and/or OH formation, the nitrogen- and carbon-based yields may vary considerably. Thus, when applying experimental yields to atmospheric models, care must be taken to choose the appropriate value, as well as to consider the conditions under which those yields are obtained.

The yield of products produced from reaction of OH with isoprene is potentially very important for nighttime chemistry, particularly because we propose that the source of OH is from RO<sub>2</sub> + HO<sub>2</sub> reactions which likely dominate in the ambient environment (Beaver et al., 2012). Recent field studies suggest that the radical propagating channels of RO<sub>2</sub> + HO<sub>2</sub> reactions must be significant to explain observations (Thornton et al., 2002; Lelieveld et al., 2008). Previous studies of nighttime chemistry have only considered alkene ozonolysis and  $HO_2 + NO$ ,  $HO_2 + O_3$ , and  $NO_3 + HO_2$  reactions as sources of OH (Bey et al., 1997, 2001a,b; Harrison et al., 1998; Faloona et al., 2001; Golz et al., 2001; Geyer et al., 2003b; Ren et al., 2003; Geyer and Stutz, 2004; Vaughan et al., 2006). While a missing OH source may explain instances where models underestimate field measurements of OH (Faloona et al., 2001; Ren et al., 2003), Geyer et al. (2003b) overpredict measurements by a factor of two without such a source. Clearly, there remain many unresolved issues surrounding the abundance of oxidants in the nighttime atmosphere.

As demonstrated by Ng et al. (2008), while ROOR compounds are minor products of RO<sub>2</sub> + RO<sub>2</sub> reactions, they may be important in the atmosphere because they present a means of significantly increasing the mass of a molecule, thereby reducing its volatility and increasing its potential to form SOA. Currently, field measurements of SOA burdens often exceed those predicted by models (de Gouw et al., 2005; Heald et al., 2005; Johnson et al., 2005; Volkamer et al., 2006; Simpson et al., 2007), a discrepancy that may be explained by SOA formation pathways, such as ROOR, that are missing from models. In this work, we see evidence of an additional ROOR formation pathway, possibly the RO<sub>2</sub> addition to alkenes, that may be relevant to SOA formation. As SOA itself is a minor product of hydrocarbon oxidation (Donahue et al., 2009), pathways that are negligible in the context of gas phase oxidation mechanisms may in fact be important in the atmosphere if they represent efficient pathways to forming SOA.

The formation of ROOR compounds in the atmosphere, and the importance of  $RO_2 + RO_2$  reactions in general, is difficult to predict because of the large uncertainties in the rates of all the relevant competing pathways ( $RO_2 + RO_2$ ,  $RO_2 + NO_3$ ,  $RO_2 + NO$ ,  $RO_2 + HO_2$ ,  $RO_2 + alkene$ ) as well as the large variation in ambient mixing ratios of the relevant species. It is apparent, though, that  $RO_2 + RO_2$  reactions are most favored when the concentration of hydrocarbon is greater than that of oxidant. For the case of isoprene +  $NO_3$  reactions, this most likely would occur in the early evening, as OH concentrations drop and  $NO_3$  concentrations are still low, or during the daytime under clouds or in forest canopies, when isoprene concentrations are high and small amounts of  $NO_3$  can also exist.

Although the isoprene: NO<sub>3</sub> ratio in our experiment varies both temporally and spatially within the chamber, our box model (Sect. 3.2) shows that the ratios achieved during the experiment may be plausible in the ambient atmosphere. During the base case simulation, the peak NO<sub>3</sub> concentration is  $\sim 1.5$  ppb, which represents an isoprene:NO<sub>3</sub> ratio of  $\sim 500$ . If we increase the initial isoprene level to 3 ppm (to account for the fact that our concentrated plume contains more than 800 ppb isoprene) the peak NO<sub>3</sub> concentration is  $\sim 0.5$  ppb, for a ratio of  $\sim 6000$ . Thus, the isoprene:NO<sub>3</sub> ratios during the experiment, while variable, are likely on the order of  $\sim 1000$  when most of the reaction takes place. Apel et al. (2002) and Brown et al. (2009) find isoprene concentrations ~ 1 ppb at sunset in Northern Michigan and New England, respectively, so NO<sub>3</sub> levels of  $\sim$  1 ppt would yield similar ratios. Studies of daytime NO<sub>3</sub> oxidation find even higher ratios, with isoprene concentrations  $\sim 10$  ppb and NO<sub>3</sub> concentrations  $\sim 0.5$  ppt, or a ratio of  $\sim 20\,000$  (Brown et al., 2005; Fuentes et al., 2007). Thus, it is likely that our experimental conditions, which favor RO<sub>2</sub> + RO<sub>2</sub> reactions over  $RO_2 + NO_3$  reactions, can occur in the ambient atmosphere, but the ambient atmosphere will also likely have higher  $HO_2$  and NO levels (due to additional sources of these radicals such as soil emissions and pernitric acid decomposition) which will also influence the final product yield and must also be considered.

Situations favoring nighttime  $RO_2 + RO_2$ RO<sub>2</sub> + alkene) reactions may be more prevalent for monoterpenes and sesquiterpenes, which, unlike isoprene, may be emitted at night (Sakulyanontvittaya et al., 2008). Many of these compounds have exhibited high SOA yields in laboratory studies, though there are still many uncertainties in the SOA formation mechanism (Griffin et al., 1999; Hallquist et al., 1999; Spittler et al., 2006; Fry et al., 2009). Based on our experience with isoprene, the role of monoterpene and sesquiterpene RO<sub>2</sub> + RO<sub>2</sub> reactions in nighttime SOA formation is worthy of further study, particularly because the nitrooxy group is electron withdrawing, and the presence of electron withdrawing groups has been shown to significantly increase RO<sub>2</sub> + RO<sub>2</sub> reaction rates (Lightfoot et al., 1992). Resonance stabilization may also affect RO<sub>2</sub> reaction rates, though we know of no work that has examined this effect. As noted by other investigators (Brown et al., 2009; Fry et al., 2009), SOA formed from the reaction of biogenic compounds with NO<sub>3</sub> - an anthropogenic oxidant - is consistent the common finding that while SOA is largely composed of contemporary carbon, even in urban areas (Bench et al., 2007; Schichtel et al., 2008), its concentrations are correlated with anthropogenic emissions (de Gouw et al., 2005, 2008; Quinn et al., 2006; Sullivan et al., 2006; Weber et al., 2007).

Most of the RO<sub>2</sub> formed from isoprene + NO<sub>3</sub> reactions are primary radicals, however, whereas a significant amount of the RO2 derived from terpenes are likely to be secondary or tertiary. Primary RO<sub>2</sub> tend to undergo significantly faster RO<sub>2</sub> + RO<sub>2</sub> reactions than secondary or tertiary RO<sub>2</sub> (Lightfoot et al., 1992). Reactivity trends are less certain for  $RO_2 + NO$ ,  $RO_2 + NO_3$ , and  $RO_2 + HO_2$  reactions, but the variation in the available data is less pronounced than for RO<sub>2</sub> + RO<sub>2</sub> reactions (Lightfoot et al., 1992; Lesclaux, 1997; Wallington et al., 1997; Atkinson et al., 2006; Vaughan et al., 2006). So while terpenes are generally more reactive with NO3 than isoprene (i.e., have higher RO2 formation rates) (Atkinson and Arey, 2003), RO<sub>2</sub> + RO<sub>2</sub> reactions for these compounds may be less competitive than for isoprene under the same conditions because of the significantly lower RO<sub>2</sub> + RO<sub>2</sub> rate constants compared to competing reactions. On the other hand, if ROOR formation is analogous to RONO<sub>2</sub> formation from RO<sub>2</sub> + NO reactions, then the ROOR yield from terpene RO<sub>2</sub> + RO<sub>2</sub> reactions may be higher, particularly for secondary RO<sub>2</sub>. RONO<sub>2</sub> yields for secondary RO2 are about a factor of 2 higher than for primary or tertiary RO2, which have similar yields (Carter and Atkinson, 1985; Atkinson et al., 1987).

Although we have gained insights into the isoprene + NO<sub>3</sub> system in this work, chamber studies such as ours have limitations. There currently exists no stable precursor of NO<sub>3</sub> suitable for chamber studies, so chemical transformations occur the instant the NO<sub>3</sub> precursor and hydrocarbon meet; for isoprene, the chemistry occurs on a much faster timescale than the mixing. Therefore, because of our experimental conditions we are only able to do an end product analysis of our experiments, and are unable to perform kinetic modeling, which could provide deeper insights into the system. Furthermore, while we can constrain RO<sub>2</sub> reaction pathways in chamber studies of OH oxidation (i.e., RO<sub>2</sub> + NO for high  $NO_x$  conditions and  $RO_2 + HO_2$  for low  $NO_x$  conditions), this is currently not feasible for NO<sub>3</sub> chamber studies. Also, while the CIMS can speciate oxidation products with greater specificity than other techniques, the lack of commercial or easily synthesizable standards leads to uncertainties in product quantification.

Up to now, most studies relating to hydrocarbon oxidation mechanisms and kinetics have focused on ozone or the OH radical; increased transferring of the techniques employed in those studies to  $NO_3$  oxidation kinetics and mechanisms offers promise to significantly advance our understanding of nighttime atmospheric chemistry, but will require overcoming challenges such as reagent sythesis (including isomeric specificity), finding suitable radical precursors, and limiting secondary and competing reactions. Many of our results (e.g., OH yield from  $RO_2 + HO_2$ , RO yield from  $RO_2 + RO_2$ ,  $HO_2$  formation from RO, ROOR formation) differ from what is suggested by previous work on different – mostly small alkylperoxy, acylperoxy, or hydroxyalkylperoxy – systems. More studies focused on nitrooxy and allylic peroxy radicals, as well as larger peroxy radicals, are warranted.

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### References

Apel, E. C., Riemer, D. D., Hills, A., Baugh, W., Orlando, J., Faloona, I., Tan, D., Brune, W., Lamb, B., Westberg, H., Carroll, M. A., Thornberry, T., and Geron, C. D.: Measurement and interpretation of isoprene fluxes and isoprene, methacrolein, and methyl vinyl ketone mixing ratios at the PROPHET site during the 1998 Intensive, J. Geophys. Res., 107, 4034–4048, 2002.

- Arey, J., Aschmann, S. M., Kwok, E. S. C., and Atkinson, R.: Alkyl nitrate, hydroxyalkyl nitrate, and hydroxycarbonyl formation from the  $NO_x$ -air photooxidations of  $C_5$ - $C_8$  n-alkanes, J. Phys. Chem. A, 105, 1020–1027, 2001.
- Atkinson, R.: Gas-phase tropospheric chemistry of volatile organic compounds: 1. Alkanes and alkenes, J. Phys. Chem. Ref. Data, 26, 215–290, 1997.
- Atkinson, R. and Arey, J.: Gas-phase tropospheric chemistry of biogenic volatile organic compounds: a review, Atmos. Environ., 37, Supplement 2, 197–219, 2003.
- Atkinson, R. and Arey, J.: Mechanisms of the gas-phase reactions of aromatic hydrocarbons and PAHs with OH and NO<sub>3</sub> radicals, Polycyclic Aromat. Compd., 27, 15–40, 2007.
- Atkinson, R., Aschmann, S., and Winer, A.: Alkyl nitrate formation from the reaction of a series of branched RO<sub>2</sub> radicals with NO as a function of temperature and pressure, J. Atmos. Chem., 5, 91–102, 1987.
- Atkinson, R., Aschmann, S. M., and Pitts, J. N.: Rate constants for the gas-phase reactions of the NO<sub>3</sub> radical with a series of organic compounds at 296 +/- 2 K, J. Phys. Chem., 92, 3454–3457, 1988.
- Atkinson, R., Baulch, D. L., Cox, R. A., Crowley, J. N., Hampson, R. F., Hynes, R. G., Jenkin, M. E., Rossi, M. J., Troe, J., and IUPAC Subcommittee: Evaluated kinetic and photochemical data for atmospheric chemistry: Volume II gas phase reactions of organic species, Atmos. Chem. Phys., 6, 3625–4055, doi:10.5194/acp-6-3625-2006, 2006.
- Barnes, I., Bastian, V., Becker, K. H., and Tong, Z.: Kinetics and products of the reactions of NO<sub>3</sub> with monoalkenes, dialkenes, and monoterpenes, J. Phys. Chem., 94, 2413–2419, 1990.
- Beaver, M. R., Clair, J. M. St., Paulot, F., Spencer, K. M., Crounse, J. D., LaFranchi, B. W., Min, K. E., Pusede, S. E., Wooldridge, P. J., Schade, G. W., Park, C., Cohen, R. C., and Wennberg, P. O.: Importance of biogenic precursors to the budget of organic nitrates: observations of multifunctional organic nitrates by CIMS and TD-LIF during BEARPEX 2009, Atmos. Chem. Phys., 12, 5773–5785, doi:10.5194/acp-12-5773-2012, 2012.
- Bench, G., Fallon, S., Schichtel, B., Malm, W., and McDade, C.: Relative contributions of fossil and contemporary carbon sources to PM 2.5 aerosols at nine Interagency Monitoring for Protection of Visual Environments (IMPROVE) network sites, J. Geophys. Res.-Atmos., 112, D10205, doi:10.1029/2006JD007708, 2007.
- Berndt, T. and Boge, O.: Gas-phase reaction of NO<sub>3</sub> radicals with isoprene: A kinetic and mechanistic study, Int. J. Chem. Kinet., 29, 755–765, 1997.
- Bey, I., Aumont, B., and Toupance, G.: The nighttime production of OH radicals in the continental troposphere, Geophys. Res. Lett., 24, 1067–1070, 1997.
- Bey, I., Aumont, B., and Toupance, G.: A modeling study of the nighttime radical chemistry in the lower continental troposphere 2. Origin and evolution of HO<sub>x</sub>, J. Geophys. Res.-Atmos., 106, 9991–10001, 2001a.
- Bey, I., Jacob, D., Yantosca, R., Logan, J., Field, B., Fiore, A., Li, Q., Liu, H., Mickley, L., and Schultz, M.: Global modeling of tropospheric chemistry with assimilated meteorology: Model description and evaluation, J. Geophys. Res., 106, 23073–23096, 2001b
- Biggs, P., Canosa-Mas, C. E., Fracheboud, J., Shallcross, D. E., and Wayne, R. P.: Investigation into the kinetics and mechanism of

- the reaction of NO<sub>3</sub> with CH<sub>3</sub> and CH<sub>3</sub>O at 298 K between 0.6 and 8.5 Torr: is there a chain decomposition mechanism in operation?, J. Chem. Soc. Faraday Trans., 90, 1197–1204, 1994.
- Birdsall, A. W., Andreoni, J. F., and Elrod, M. J.: Investigation of the role of bicyclic peroxy radicals in the oxidation mechanism of toluene, J. Phys. Chem. A, 114, 39, 10655–10633, 2010.
- Birdsall, A. W. and Elrod, M. J.: Comprehensive NO-dependent study of the products of the oxidation of atmospherically relevant aromatic compounds, J. Phys. Chem. A, 115, 21, 5397– 5407, 2011.
- Brown, S. S., Osthoff, H. D., Stark, H., Dube, W. P., Ryerson, T. B.,
  Warneke, C., de Gouw, J. A., Wollny, A. G., Parrish, D. D.,
  Fehsenfeld, F. C., and Ravishankara, A. R.: Aircraft observations of daytime NO<sub>3</sub> and N<sub>2</sub>O<sub>5</sub> and their implications for tropospheric chemistry, J. Photoch. Photobio. A, 176, 270–278, 2005.
- Brown, S. S., Ryerson, T. B., Wollny, A. G., Brock, C. A., Peltier, R., Sullivan, A. P., Weber, R. J., Dube, W. P., Trainer, M., Meagher, J. F., Fehsenfeld, F. C., and Ravishankara, A. R.: Variability in nocturnal nitrogen oxide processing and its role in regional air quality, Science, 311, 67–70, 2006.
- Brown, S. S., deGouw, J. A., Warneke, C., Ryerson, T. B., Dubé, W. P., Atlas, E., Weber, R. J., Peltier, R. E., Neuman, J. A., Roberts, J. M., Swanson, A., Flocke, F., McKeen, S. A., Brioude, J., Sommariva, R., Trainer, M., Fehsenfeld, F. C., and Ravishankara, A. R.: Nocturnal isoprene oxidation over the Northeast United States in summer and its impact on reactive nitrogen partitioning and secondary organic aerosol, Atmos. Chem. Phys., 9, 3027–3042, doi:10.5194/acp-9-3027-2009, 2009.
- Canosa-Mas, C. E., King, M. D., Lopez, R., Percival, C. J., Wayne, R. P., Shallcross, D. E., Pyle, J. A., and Daele, V.: Is the reaction between CH<sub>3</sub>C(O)O<sub>2</sub> and NO<sub>3</sub> important in the night-time troposphere?, J. Chem. Soc. Faraday Trans., 92, 2211–2222, 1996.
- Carlton, A. G., Wiedinmyer, C., and Kroll, J. H.: A review of Secondary Organic Aerosol (SOA) formation from isoprene, Atmos. Chem. Phys., 9, 4987–5005, doi:10.5194/acp-9-4987-2009, 2009.
- Carslaw, N., Carpenter, L. J., Plane, J. M. C., Allan, B. J., Burgess, R. A., Clemitshaw, K. C., Coe, H., and Penkett, S. A.: Simultaneous observations of nitrate and peroxy radicals in the marine boundary layer, J. Geophys. Res.-Atmos., 102, 18917–18933, 1997.
- Carter, W. and Atkinson, R.: Atmospheric chemistry of alkanes, J. Atmos. Chem., 3, 377–405, 1985.
- Chameides, W. L., Lindsay, R. W., Richardson, J., and Kiang, C. S.: The role of biogenic hydrocarbons in urban photochemical smog
  – Atlanta as a case study, Science, 241, 1473–1475, 1988.
- Claeys, M., Graham, B., Vas, G., Wang, W., Vermeylen, R., Pashynska, V., Cafmeyer, J., Guyon, P., Andreae, M. O., Artaxo, P., and Maenhaut, W.: Formation of secondary organic aerosols through photooxidation of isoprene, Science, 303, 1173–1176, 2004.
- Cocker, D. R., Flagan, R. C., and Seinfeld, J. H.: State-of-the-art chamber facility for studying atmospheric aerosol chemistry, Environ. Sci. Technol., 35, 2594–2601, 2001.
- Crounse, J. D., McKinney, K. A., Kwan, A. J., and Wennberg, P. O.: Measurement of gas-phase hydroperoxides by chemical ionization mass spectrometry, Anal. Chem., 78, 6726–6732, 2006.
- Crounse, J. D., Paulot, F., Kjaergaard, H. G., and Wennberg, P. O.: Peroxy radical isomerization in the oxidation of isoprene, Phys. Chem. Chem. Phys., 13, 13607–13613, 2011.

- Daele, V., Laverdet, G., Le Bras, G., and Poulet, G.: Kinetics of the reactions CH<sub>3</sub>O + NO, CH<sub>3</sub>O + NO<sub>3</sub>, and CH<sub>3</sub>O<sub>2</sub> + NO<sub>3</sub>, J. Phys. Chem., 99, 1470–1477, 1995.
- Davidson, J. A., Viggiano, A. A., Howard, C. J., Dotan, I., Fehsenfeld, F. C., Albritton, D. L., and Ferguson, E. E.: Rate constants for reactions of O<sub>2</sub><sup>+</sup>, NO<sub>2</sub><sup>+</sup>, NO<sup>+</sup>, H<sub>3</sub>O<sup>+</sup>, CO<sub>3</sub><sup>-</sup>, NO<sub>2</sub><sup>-</sup>, and halide ions with N<sub>2</sub>O<sub>5</sub> at 300 K, J. Chem. Phys., 68, 2085–2087, 1978.
- de Gouw, J. A., Middlebrook, A. M., Warneke, C., Goldan, P. D., Kuster, W. C., Roberts, J. M., Fehsenfeld, F. C., Worsnop, D. R., Canagaratna, M. R., Pszenny, A. A. P., Keene, W. C., Marchewka, M., Bertman, S. B., and Bates, T. S.: Budget of organic carbon in a polluted atmosphere: Results from the New England Air Quality Study in 2002, J. Geophys. Res.-Atmos., 110, D16305, doi:10.1029/2004JD005623, 2005.
- de Gouw, J. A., Brock, C. A., Atlas, E. L., Bates, T. S., Fehsenfeld, F. C., Goldan, P. D., Holloway, J. S., Kuster, W. C., Lerner, B. M., Matthew, B. M., Middlebrook, A. M., Onasch, T. B., Peltier, R. E., Quinn, P. K., Senff, C. J., Stohl, A., Sullivan, A. P., Trainer, M., Warneke, C., Weber, R. J., and Williams, E. J.: Sources of particulate matter in the northeastern United States in summer: 1. Direct emissions and secondary formation of organic matter in urban plumes, J. Geophys. Res.-Atmos., 113, D08301, doi:10.1029/2007JD009243, 2008.
- Dibble, T. S.: Failures and limitations of quantum chemistry for two key problems in the atmospheric chemistry of peroxy radicals, Atmos. Environ., 42, 5837–5848, 2008.
- Dillon, T. J. and Crowley, J. N.: Direct detection of OH formation in the reactions of HO<sub>2</sub> with CH<sub>3</sub>C(O)O<sub>2</sub> and other substituted peroxy radicals, Atmos. Chem. Phys., 8, 4877–4889, doi:10.5194/acp-8-4877-2008, 2008.
- Donahue, N. M., Robinson, A. L., and Pandis, S. N.: Atmospheric organic particulate matter: From smoke to secondary organic aerosol, Atmos. Environ., 43, 94–106, 2009.
- Eddingsaas, N. C., VanderVelde, D. G., and Wennberg, P. O.: Kinetics and products of the acid-catalyzed ring-opening of atmospherically relevant butyl epoxy alcohols, J. Phys. Chem. A, 114, 8106–8113, 2010.
- Faloona, I., Tan, D., Brune, W., Hurst, J., Barket, D., Couch, T. L., Shepson, P., Apel, E., Riemer, D., Thornberry, T., Carroll, M. A., Sillman, S., Keeler, G. J., Sagady, J., Hooper, D., and Paterson, K.: Nighttime observations of anomalously high levels of hydroxyl radicals above a deciduous forest canopy, J. Geophys. Res.-Atmos., 106, 24315–24333, 2001.
- Fan, J. W. and Zhang, R. Y.: Atmospheric oxidation mechanism of isoprene, Environ. Chem., 1, 140–149, 2004.
- Fiore, A. M., Horowitz, L. W., Purves, D. W., Levy, H., Evans, M. J., Wang, Y. X., Li, Q. B., and Yantosca, R. M.: Evaluating the contribution of changes in isoprene emissions to surface ozone trends over the eastern United States, J. Geophys. Res.-Atmos., 110, D12303, doi:10.1029/2004JD005485, 2005.
- Forkel, R., Klemm, O., Graus, M., Rappengluck, B., Stockwell, W. R., Grabmer, W., Held, A., Hansel, A., and Steinbrecher, R.: Trace gas exchange and gas phase chemistry in a Norway spruce forest: A study with a coupled 1-dimensional canopy atmospheric chemistry emission model, Atmos. Environ., 40, S28–S42, 2006.
- Francisco-Marquez, M., Alvarez-Idaboy, J. R., Galano, A., and Vivier-Bunge, A.: A possible mechanism for furan formation in the tropospheric oxidation of dienes, Environ. Sci. Technol., 39,

- 8797-8802, 2005.
- Fry, J. L., Kiendler-Scharr, A., Rollins, A. W., Wooldridge, P. J., Brown, S. S., Fuchs, H., Dubé, W., Mensah, A., dal Maso, M., Tillmann, R., Dorn, H.-P., Brauers, T., and Cohen, R. C.: Organic nitrate and secondary organic aerosol yield from NO<sub>3</sub> oxidation of β-pinene evaluated using a gas-phase kinetics/aerosol partitioning model, Atmos. Chem. Phys., 9, 1431–1449, doi:10.5194/acp-9-1431-2009, 2009.
- Fuentes, J. D., Wang, D., Bowling, D. R., Potosnak, M., Monson, R. K., Goliff, W. S., and Stockwell, W. R.: Biogenic hydrocarbon chemistry within and above a mixed deciduous forest, J. Atmos. Chem., 56, 165–185, 2007.
- Garden, A. L., Paulot, F., Crounse, J. D., Maxwell-Cameron, I. J., Wennberg, P. O., and Kjaergaard, H. G.: Calculation of conformationally weighted dipole moments useful in ion-molecule collision rate estimates, Chem. Phys. Lett., 474, 45–50, 2009.
- Geyer, A. and Stutz, J.: Vertical profiles of  $NO_3$ ,  $N_2O_5$ ,  $O_3$ , and  $NO_x$  in the nocturnal boundary layer: 2. Model studies on the altitude dependence of composition and chemistry, J. Geophys. Res.-Atmos., 109, D12307, doi:10.1029/2003JD004211, 2004.
- Geyer, A., Alicke, B., Ackermann, R., Martinez, M., Harder, H., Brune, W., di Carlo, P., Williams, E., Jobson, T., Hall, S., Shetter, R., and Stutz, J.: Direct observations of daytime NO<sub>3</sub>: Implications for urban boundary layer chemistry, J. Geophys. Res.-Atmos., 108, 4368, doi:10.1029/2002JD002967, 2003a.
- Geyer, A., Bachmann, K., Hofzumahaus, A., Holland, F., Konrad, S., Klupfel, T., Patz, H. W., Perner, D., Mihelcic, D., Schafer, H. J., Volz-Thomas, A., and Platt, U.: Nighttime formation of peroxy and hydroxyl radicals during the BERLIOZ campaign: Observations and modeling studies, J. Geophys. Res.-Atmos., 108, 8249, doi:10.1029/2001JD000656, 2003b.
- Golz, C., Senzig, J., and Platt, U.: NO<sub>3</sub>-initiated oxidation of biogenic hydrocarbons, Chemosphere, 3, 339–352, 2001.
- Griffin, R. J., Cocker, D. R., Flagan, R. C., and Seinfeld, J. H.: Organic aerosol formation from the oxidation of biogenic hydrocarbons, J. Geophys. Res.-Atmos., 104, 3555–3567, 1999.
- Guenther, A., Karl, T., Harley, P., Wiedinmyer, C., Palmer, P. I., and Geron, C.: Estimates of global terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols from Nature), Atmos. Chem. Phys., 6, 3181–3210, doi:10.5194/acp-6-3181-2006, 2006.
- Hallquist, M., Wangberg, I., Ljungstrom, E., Barnes, I., and Becker, K. H.: Aerosol and product yields from NO<sub>3</sub> radical-initiated oxidation of selected monoterpenes, Environ. Sci. Technol., 33, 553–559, 1999.
- Harley, P., Vasconcellos, P., Vierling, L., Pinheiro, C. C. D., Greenberg, J., Guenther, A., Klinger, L., De Almeida, S. S., Neill, D., Baker, T., Phillips, O., and Malhi, Y.: Variation in potential for isoprene emissions among Neotropical forest sites, Global Change Biol., 10, 630–650, 2004.
- Harrison, R. M., Shi, J. P., and Grenfell, J. L.: Novel nighttime free radical chemistry in severe nitrogen dioxide pollution episodes, Atmos. Environ., 32, 2769–2774, 1998.
- Hasson, A. S., Ho, A. W., Kuwata, K. T., and Paulson, S. E.: Production of stabilized Criegee intermediates and peroxides in the gas phase ozonolysis of alkenes 2. Asymmetric and biogenic alkenes, J. Geophys. Res.-Atmos., 106, 34143–34153, 2001.
- Hasson, A. S., Tyndall, G. S., and Orlando, J. J.: A product yield study of the reaction of HO<sub>2</sub> radicals with ethyl peroxy

- $(C_2H_5O_2)$ , acetyl peroxy  $(CH_3C(O)O_2)$ , and acetonyl peroxy  $(CH_3C(O)CH_2O_2)$  radicals, J. Phys. Chem. A, 108, 5979–5989, 2004.
- Hasson, A. S., Kuwata, K. T., Arroyo, M. C., and Petersen, E. B.: Theoretical studies of the reaction of hydroperoxy radicals (HO<sub>2</sub>) with ethyl peroxy (CH<sub>3</sub>CH<sub>2</sub>O<sub>2</sub>), acetyl peroxy (CH<sub>3</sub>C(O)O<sub>2</sub>) and acetonyl peroxy (CH<sub>3</sub>C(O)CH<sub>2</sub>O<sub>2</sub>) radicals, J. Photoch. Photobio. A, 176, 218–230, 2005.
- Hasson, A. S., Tyndall, G. S., Orlando, J. J., Singh, S., Hernandez, S. Q., Campbell, S., and Ibarra, Y: Branching ratios for the reaction of selected carbonyl-containing peroxy radicals with hydroperoxy radicals, J. Phys. Chem. A, 116, 6264–6281, 2012.
- Heald, C. L., Jacob, D. J., Park, R. J., Russell, L. M., Huebert, B. J., Seinfeld, J. H., Liao, H., and Weber, R. J.: A large organic aerosol source in the free troposphere missing from current models, Geophys. Res. Lett., 32, L18809, doi:10.1029/2005GL023831, 2005.
- Heintz, F., Platt, U., Flentje, H., and Dubois, R.: Long-term observation of nitrate radicals at the tor station, Kap Arkona (Rugen), J. Geophys. Res.-Atmos., 101, 22891–22910, 1996.
- Horowitz, L. W., Liang, J. Y., Gardner, G. M., and Jacob, D. J.: Export of reactive nitrogen from North America during summertime: Sensitivity to hydrocarbon chemistry, J. Geophys. Res.-Atmos., 103, 13451–13476, 1998.
- Horowitz, L. W., Fiore, A. M., Milly, G. P., Cohen, R. C., Perring, A., Wooldridge, P. J., Hess, P. G., Emmons, L. K., and Lamarque, J. F.: Observational constraints on the chemistry of isoprene nitrates over the eastern United States, J. Geophys. Res.-Atmos., 112, D12S08, doi:10.1029/2006JD007747, 2007.
- Jay, K. and Stieglitz, L.: The gas-phase addition of NO<sub>x</sub> to olefins, Chemosphere, 19, 1939–1950, 1989.
- Jenkin, M. E., Hurley, M. D., and Wallington, T. J.: Investigation of the radical product channel of the CH<sub>3</sub>C(O)O<sub>2</sub>+HO<sub>2</sub> reaction in the gas phase, Phys. Chem. Chem. Phys., 9, 3149–3162, 2007.
- Jenkin, M. E., Hurley, M. D., and Wallington, T. J.: Investigation of the radical product channel of the CH<sub>3</sub>C(O)CH<sub>2</sub>O<sub>2</sub>+HO<sub>2</sub> reaction in the gas phase, Phys. Chem. Chem. Phys., 10, 4274–4280, 2008.
- Jenkin, M. E., Hurley, M. A., and Wallington, T. J.: Investigation of the radical product channel of the  $CH_3OCH_2O_2 + HO_2$  reaction in the gas phase, J. Phys. Chem. A, 114, 408–416, 2010.
- Johnson, D., Jenkin, M. E., Wirtz, K., and Martin-Reviejo, M.: Simulating the formation of secondary organic aerosol from the photooxidation of aromatic hydrocarbons, Environ. Chem., 2, 35–48, 2005.
- Keywood, M. D., Varutbangkul, V., Bahreini, R., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol formation from the ozonolysis of cycloalkenes and related compounds, Environ. Sci. Technol., 38, 4157–4164, 2004.
- Kroll, J. H., Ng, N. L., Murphy, S. M., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol formation from isoprene photooxidation under high-NO<sub>x</sub> conditions, Geophys. Res. Lett., 32, L18808, doi:10.1029/2005GL023637, 2005.
- Kroll, J. H., Ng, N. L., Murphy, S. M., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol formation from isoprene photooxidation, Environ. Sci. Technol., 40, 1869–1877, 2006.
- Kwok, E. S. C. and Atkinson, R.: Estimation of hydroxyl radical reaction rate constants for gas-phase organic compounds using a structure-activity relationship – an update, Atmos. Environ., 29, 1685–1695, 1995.

- Kwok, E. S. C., Aschmann, S. M., Arey, J., and Atkinson, R.: Product formation from the reaction of the NO<sub>3</sub> radical with isoprene and rate constants for the reactions of methacrolein and methyl vinyl ketone with the NO<sub>3</sub> radical, Int. J. Chem. Kinet., 28, 925–934, 1996.
- Lelieveld, J., Butler, T. M., Crowley, J. N., Dillon, T. J., Fischer, H., Ganzeveld, L., Harder, H., Lawrence, M. G., Martinez, M., Taraborrelli, D., and Williams, J.: Atmospheric oxidation capacity sustained by a tropical forest, Nature, 452, 737–740, 2008.
- Lesclaux, R.: Combination of peroxyl radicals in the gas phase, in: Peroxyl Radicals, edited by: Alfassi, Z., 81–112, Wiley, 1997.
- Lightfoot, P. D., Cox, R. A., Crowley, J. N., Destriau, M., Hayman, G. D., Jenkin, M. E., Moortgat, G. K., and Zabel, F.: Organic peroxy radicals – kinetics, spectroscopy and tropospheric chemistry, Atmos. Environ. A-Gen., 26, 1805–1961, 1992.
- Matsunaga, A. and Ziemann, P. J.: Yields of beta-hydroxynitrates and dihydroxynitrates in aerosol formed from OH radical-initiated reactions of linear alkenes in the presence of NO<sub>x</sub>, J. Phys. Chem. A, 113, 599–606, 2009.
- Mayo, F. R.: Free radical autoxidations of hydrocarbons, Acc. Chem. Res., 1, 193–201, 1968.
- Mellouki, A., Talukdar, R. K., Bopegedera, A., and Howard, C. J.: Study of the kinetics of the reactions of NO<sub>3</sub> with HO<sub>2</sub> and OH, Int. J. Chem. Kinet., 25, 25–39, 1993.
- Neeb, P. and Moortgat, G. K.: Formation of OH radicals in the gas-phase reaction of propene, isobutene, and isoprene with O<sub>3</sub>: Yields and mechanistic implications, J. Phys. Chem. A, 103, 9003–9012, 1999.
- Ng, N. L., Chhabra, P. S., Chan, A. W. H., Surratt, J. D., Kroll, J. H., Kwan, A. J., McCabe, D. C., Wennberg, P. O., Sorooshian, A., Murphy, S. M., Dalleska, N. F., Flagan, R. C., and Seinfeld, J. H.: Effect of NO<sub>x</sub> level on secondary organic aerosol (SOA) formation from the photooxidation of terpenes, Atmos. Chem. Phys., 7, 5159–5174, doi:10.5194/acp-7-5159-2007, 2007.
- Ng, N. L., Kwan, A. J., Surratt, J. D., Chan, A. W. H., Chhabra, P. S., Sorooshian, A., Pye, H. O. T., Crounse, J. D., Wennberg, P. O., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol (SOA) formation from reaction of isoprene with nitrate radicals (NO<sub>3</sub>), Atmos. Chem. Phys., 8, 4117–4140, doi:10.5194/acp-8-4117-2008, 2008.
- O'Brien, J. M., Czuba, E., Hastie, D. R., Francisco, J. S., and Shepson, P. B.: Determination of the hydroxy nitrate yields from the reaction of C<sub>2</sub>-C<sub>6</sub> alkenes with OH in the presence of NO, J. Phys. Chem. A, 102, 8903–8908, 1998.
- Orzechowska, G. E. and Paulson, S. E.: Photochemical sources of organic acids. 1. Reaction of ozone with isoprene, propene, and 2-butenes under dry and humid conditions using SPME, J. Phys. Chem. A, 109, 5358–5365, 2005.
- Osborne, D. A. and Waddington, D. J.: Reactions of oxygenated radicals in the gas phase: 7. Reactions of methylperoxyl radicals and alkenes, J. Chem. Soc. Perk. T., 2, 925–930, 1980.
- Paulot, F., Crounse, J. D., Kjaergaard, H. G., Kurten, A., St Clair, J. M., Seinfeld, J. H., and Wennberg, P. O.: Unexpected epoxide formation in the gas-phase photooxidation of isoprene, Science, 325, 730–733, 2009a.
- Paulot, F., Crounse, J. D., Kjaergaard, H. G., Kroll, J. H., Seinfeld, J. H., and Wennberg, P. O.: Isoprene photooxidation: new insights into the production of acids and organic nitrates, Atmos. Chem. Phys., 9, 1479–1501, doi:10.5194/acp-9-1479-2009, 2009b.

- Peeters, J., Nguyen, T. L., and Vereecken, L.: HO<sub>x</sub> radical regeneration in the oxidation of isoprene, Phys. Chem. Chem. Phys., 11, 5935–5939, 2009.
- Penkett, S. A., Burgess, R. A., Coe, H., Coll, I., Hov, O., Lindskog, A., Schmidbauer, N., Solberg, S., Roemer, M., Thijsse, T., Beck, J., and Reeves, C. E.: Evidence for large average concentrations of the nitrate radical (NO<sub>3</sub>) in Western Europe from the HANSA hydrocarbon database, Atmos. Environ., 41, 3465–3478, 2007.
- Perring, A. E., Wisthaler, A., Graus, M., Wooldridge, P. J., Lockwood, A. L., Mielke, L. H., Shepson, P. B., Hansel, A., and Cohen, R. C.: A product study of the isoprene+NO<sub>3</sub> reaction, Atmos. Chem. Phys., 9, 4945–4956, doi:10.5194/acp-9-4945-2009, 2009.
- Platt, U. and Janssen, C.: Observation and role of the free radicals NO<sub>3</sub>, ClO, BrO and IO in the troposphere, Faraday Discuss., 100, 175–198, 1995.
- Platt, U., Perner, D., Schroder, J., Kessler, C., and Toennissen, A.: The diurnal variation of NO<sub>3</sub>, J. Geophys. Res.-Oc. Atm., 86, 1965–1970, 1981.
- Quinn, P. K., Bates, T. S., Coffman, D., Onasch, T. B., Worsnop, D., Baynard, T., de Gouw, J. A., Goldan, P. D., Kuster, W. C., Williams, E., Roberts, J. M., Lerner, B., Stohl, A., Pettersson, A., and Lovejoy, E. R.: Impacts of sources and aging on submicrometer aerosol properties in the marine boundary layer across the Gulf of Maine, J. Geophys. Res., 111, D23S36, doi:10.1029/2006JD007582, 2006.
- Ren, X. R., Harder, H., Martinez, M., Lesher, R. L., Oliger, A., Shirley, T., Adams, J., Simpas, J. B., and Brune, W. H.: HO<sub>x</sub> concentrations and OH reactivity observations in New York City during PMTACS-NY2001, Atmos. Environ., 37, 3627–3637, 2003.
- Roberts, J. M., Williams, J., Baumann, K., Buhr, M. P., Goldan, P. D., Holloway, J., Hubler, G., Kuster, W. C., McKeen, S. A., Ryerson, T. B., Trainer, M., Williams, E. J., Fehsenfeld, F. C., Bertman, S. B., Nouaime, G., Seaver, C., Grodzinsky, G., Rodgers, M., and Young, V. L.: Measurements of PAN, PPN, and MPAN made during the 1994 and 1995 Nashville Intensives of the Southern Oxidant Study: Implications for regional ozone production from biogenic hydrocarbons, J. Geophys. Res.-Atmos., 103, 22473–22490, 1998.
- Rollins, A. W., Kiendler-Scharr, A., Fry, J. L., Brauers, T., Brown, S. S., Dorn, H.-P., Dubé, W. P., Fuchs, H., Mensah, A., Mentel, T. F., Rohrer, F., Tillmann, R., Wegener, R., Wooldridge, P. J., and Cohen, R. C.: Isoprene oxidation by nitrate radical: alkyl nitrate and secondary organic aerosol yields, Atmos. Chem. Phys., 9, 6685–6703, doi:10.5194/acp-9-6685-2009, 2009.
- Ruppert, L. and Becker, K. H.: A product study of the OH radicalinitiated oidation of isoprene: formation of C<sub>5</sub>-unsaturated diols, Atmos. Environ., 34, 1529–1542, 2000.
- Sakulyanontvittaya, T., Duhl, T., Wiedinmyer, C., Helmig, D., Matsunaga, S., Potosnak, M., Milford, J., and Guenther, A.: Monoterpene and sesquiterpene emission estimates for the United States, Environ. Sci. Technol., 42, 1623–1629, 2008.
- Sander, S. P., Abbatt, J., Barker, J. R., Burkholder, J. B., Friedl, R. R., Golden, D. M., Huie, R. E., Kolb, C. E., Kurylo, M. J., Moortgat, G. K., Orkin, V. L., and Wine, P. H.: Chemical kinetics and photochemical data for use in atmospheric studies: Evaluation Number 17, Jet Propulsion Laboratory, 2011.
- Schichtel, B., Malm, W., Bench, G., Fallon, S., McDade, C., Chow, J., and Watson, J.: Fossil and contemporary fine particulate car-

- bon fractions at 12 rural and urban sites in the United States, J. Geophys. Res., 113, D02311, doi:10.1029/2007JD008605, 2008.
- Seeley, J. V., Meads, R. F., Elrod, M. J., and Molina, M. J.: Temperature and pressure dependence of the rate constant for the HO<sub>2</sub> + NO reaction, J. Phys. Chem., 100, 4026–4031, 1996.
- Sharkey, T. D., Singsaas, E. L., Vanderveer, P. J., and Geron, C.: Field measurements of isoprene emission from trees in response to temperature and light, Tree Physiol., 16, 649–654, 1996.
- Simmons, K. E. and van Sickle, D.: Addition abstraction competition of acylperoxy radicals reacting with alkenes cooxidation of cyclohexene and valeraldehyde, J. Am. Chem. Soc., 95, 7759–7763, 1973.
- Simpson, D., Yttri, K. E., Klimont, Z., Kupiainen, K., Caseiro, A., Gelencser, A., Pio, C., Puxbaum, H., and Legrand, M.: Modeling carbonaceous aerosol over Europe: Analysis of the CARBOSOL and EMEP EC/OC campaigns, J. Geophys. Res.-Atmos., 112, D23S14, doi:10.1029/2006JD008158, 2007.
- Sinha, A., Lovejoy, E. R., and Howard, C. J.: Kinetic study of the reaction of HO<sub>2</sub> with ozone, J. Chem. Phys., 87, 2122–2128, 1987.
- Skov, H., Hjorth, J., Lohse, C., Jensen, N. R., and Restelli, G.: Products and mechanisms of the reactions of the nitrate radical (NO<sub>3</sub>) with isoprene, 1,3-butadiene and 2,3-dimethyl-1,3-butadiene in air, Atmos. Environ. A-Gen., 26, 2771–2783, 1992.
- Smith, N., Plane, J. M. C., Nien, C. F., and Solomon, P. A.: Nighttime radical chemistry in the San Joaquin Valley, Atmos. Environ., 29, 2887–2897, 1995.
- Spittler, M., Barnes, I., Bejan, I., Brockmann, K. J., Benter, T., and Wirtz, K.: Reactions of NO<sub>3</sub> radicals with limonene and alphapinene: Product and SOA formation, Atmos. Environ., 40, S116– S127, 2006.
- Stark, M. S. and Waddington, D. J.: Oxidation of propene in the gas-phase, Int. J. Chem. Kinet., 27, 123–151, 1995.
- Starn, T. K., Shepson, P. B., Bertman, S. B., Riemer, D. D., Zika, R. G., and Olszyna, K.: Nighttime isoprene chemistry at an urban-impacted forest site, J. Geophys. Res.-Atmos., 103, 22437–22447, 1998.
- St. Clair, J. M., McCabe, D. C., Crounse, J. D., Steiner, U., and Wennberg, P. O.: Chemical ionization tandem mass spectrometer for the in situ measurement of methyl hydrogen peroxide, Rev. Sci. Instrum., 81, 094102, doi:10.1063/1.3480552, 2010.
- Steinbacher, M., Dommen, J., Ordonez, C., Reimann, S., Gruebler, F., Staehelin, J., Andreani-Aksoyoglu, S., and Prevot, A. S. H.: Volatile organic compounds in the Po Basin. part B: Biogenic VOCs, J. Atmos. Chem., 51, 293–315, 2005.
- Stroud, C. A., Roberts, J. M., Williams, E. J., Hereid, D., Angevine, W. M., Fehsenfeld, F. C., Wisthaler, A., Hansel, A., Martinez-Harder, M., Harder, H., Brune, W. H., Hoenninger, G., Stutz, J., and White, A. B.: Nighttime isoprene trends at an urban forested site during the 1999 Southern Oxidant Study, J. Geophys. Res.-Atmos., 107, 4291, doi:10.1029/2001JD000959, 2002.
- Su, T. and Chesnavich, W. J.: Parameterization of the ion-polar molecule collision rate constant by trajectory calculations, J. Chem. Phys., 76, 5183–5185, 1982.
- Suh, I., Lei, W. F., and Zhang, R. Y.: Experimental and theoretical studies of isoprene reaction with NO<sub>3</sub>, J. Phys. Chem. A, 105, 6471–6478, 2001.
- Sullivan, A. P., Peltier, R. E., Brock, C. A., de Gouw, J. A., Holloway, J. S., Warneke, C., Wollny, A. G., and Weber, R. J.: Airborne measurements of carbonaceous aerosol soluble in water

- over northeastern United States: Method development and an investigation into water-soluble organic carbon sources, J. Geophys. Res.-Atmos., 111, D23S46, doi:10.1029/2006JD007072, 2006.
- Surratt, J. D., Murphy, S. M., Kroll, J. H., Ng, N. L., Hildebrandt, L., Sorooshian, A., Szmigielski, R., Vermeylen, R., Maenhaut, W., Claeys, M., Flagan, R. C., and Seinfeld, J. H.: Chemical composition of secondary organic aerosol formed from the photooxidation of isoprene, J. Phys. Chem. A, 110, 9665–9690, 2006.
- Surratt, J. D., Chan, A. W. H., Eddingsaas, N. C., Chan, M. N., Loza, C. L., Kwan, A. J., Hersey, S. P., Flagan, R. C., Wennberg, P. O., and Seinfeld, J. H.: Reactive intermediates revealed in secondary organic aerosol formation from isoprene, P. Natl. Acad. Sci. USA, 107, 6640–6645, 2010.
- Thornton, J. A., Wooldridge, P. J., Cohen, R. C., Martinez, M., Harder, H., Brune, W. H., Williams, E. J., Roberts, J. M., Fehsenfeld, F. C., Hall, S. R., Shetter, R. E., Wert, B. P., and Fried, A.: Ozone production rates as a function of  $NO_X$  abundances and  $HO_X$  production rates in the Nashville urban plume, J. Geophys. Res.-Atmos., 107, 4146, doi:10.1029/2001JD000932, 2002.
- van Sickle, D., Mayo, F. R., and Arluck, R. M.: Liquid-phase oxidation of cyclopentene, J. Amer. Chem. Soc., 87, 4832–4837, 1965a
- van Sickle, D., Mayo, F. R., and Arluck, R. M.: Liquid-phase oxidations of cyclic alkenes, J. Amer. Chem. Soc., 87, 4824–4832, 1965b.
- Vaughan, S., Canosa-Mas, C. E., Pfrang, C., Shallcross, D. E., Watson, L., and Wayne, R. P.: Kinetic studies of reactions of the nitrate radical (NO<sub>3</sub>) with peroxy radicals (RO<sub>2</sub>): an indirect source of OH at night?, Phys. Chem. Chem. Phys., 8, 3749–3760, 2006
- Vereecken, L. and Peeters, J.: Decomposition of substituted alkoxy radicals-part I: a generalized structure-activity relationship for reaction barrier heights, Phys. Chem. Chem. Phys., 11, 9062– 9074, 2009.
- Volkamer, R., Jimenez, J. L., San Martini, F., Dzepina, K., Zhang, Q., Salcedo, D., Molina, L. T., Worsnop, D. R., and Molina, M. J.: Secondary organic aerosol formation from anthropogenic air pollution: Rapid and higher than expected, Geophys. Res. Lett., 33, L17811, doi:10.1029/2006GL026899, 2006.
- von Friedeburg, C., Wagner, T., Geyer, A., Kaiser, N., Vogel, B., Vogel, H., and Platt, U.: Derivation of tropospheric NO<sub>3</sub> profiles using off-axis differential optical absorption spectroscopy measurements during sunrise and comparison with simulations, J. Geophys. Res.-Atmos., 107, 4168, doi:10.1029/2001JD000481, 2002.

- von Kuhlmann, R., Lawrence, M. G., Pöschl, U., and Crutzen, P. J.: Sensitivities in global scale modeling of isoprene, Atmos. Chem. Phys., 4, 1–17, doi:10.5194/acp-4-1-2004, 2004.
- Wallington, T., Nielsen, O., and Sehested, J.: Reactions of organic peroxy radicals in the gas phase, in: Peroxyl Radicals, edited by: Alfassi, Z., 113–172, Wiley, 1997.
- Warneke, C., de Gouw, J. A., Goldan, P. D., Kuster, W. C., Williams, E. J., Lerner, B. M., Jakoubek, R., Brown, S. S., Stark, H., Aldener, M., Ravishankara, A. R., Roberts, J. M., Marchewka, M., Bertman, S., Sueper, D. T., McKeen, S. A., Meagher, J. F., and Fehsenfeld, F. C.: Comparison of daytime and nighttime oxidation of biogenic and anthropogenic VOCs along the New England coast in summer during New England Air Quality Study 2002, J. Geophys. Res.-Atmos., 109, D10309, doi:10.1029/2003JD004424, 2004.
- Wayne, R. P., Barnes, I., Biggs, P., Burrows, J. P., Canosamas, C. E., Hjorth, J., Lebras, G., Moortgat, G. K., Perner, D., Poulet, G., Restelli, G., and Sidebottom, H.: The nitrate radical physics, chemistry, and the atmosphere, Atmos. Environ. A-Gen., 25, 1–203, 1991.
- Weber, R. J., Sullivan, A. P., Peltier, R. E., Russell, A., Yan, B., Zheng, M., de Gouw, J., Warneke, C., Brock, C., Holloway, J. S., Atlas, E. L., and Edgerton, E.: A study of secondary organic aerosol formation in the anthropogenic-influenced southeastern United States, J. Geophys. Res.-Atmos., 112, D13302, doi:10.1029/2007JD008408, 2007.
- Williams, J., Roberts, J. M., Fehsenfeld, F. C., Bertman, S. B., Buhr, M. P., Goldan, P. D., Hubler, G., Kuster, W. C., Ryerson, T. B., Trainer, M., and Young, V.: Regional ozone from biogenic hydrocarbons deduced from airborne measurements of PAN, PPN, and MPAN, Geophys. Res. Lett., 24, 1099–1102, 1997.
- Zhang, D. and Zhang, R. Y.: Unimolecular decomposition of nitrooxyalkyl radicals from NO<sub>3</sub>-isoprene reaction, J. Chem. Phys., 116, 9721–9728, 2002.
- Ziemann, P.: Evidence for low-volatility diacyl peroxides as a nucleating agent and major component of aerosol formed from reactions of  $O_3$  with cyclohexene and homologous compounds, J. Phys. Chem. A, 106, 4390–4402, 2002.